

Heterogeneous core-mantle boundary heat flux in thermo-chemical core convection

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Abstract

Thermal coupling between convection in Earth's mantle and core was proposed to explain asymmetric features of the geomagnetic field during the history of the Earth. The coupling is caused by laterally varying heat transport from the core to the mantle, induced by lateral temperature gradients in the lower-most mantle. Clues for the temperature gradients were found by seismic tomography.

This numerical study aims to explore the influence of these non uniform boundary conditions, as compared to uniform heat flux and isothermal boundary conditions on thermo-chemical core convection and some distinct dynamo properties.

In our model the heat flux pattern is modeled by a single spherical harmonic of degree and order 2. This setting conserves equatorial symmetry, but imposes azimuthal heat flux gradients.

Today, convection in the Earth's core is assumed to be driven predominantly by a combination of thermal and compositional buoyancy sources located at the inner-core boundary. Thermal and compositional diffusivities differ by orders of magnitude. The resulting differences in the dynamical behavior of the two components demand an approach with distinct transport equations and boundary conditions for temperature and chemical concentration. Simulations for five different ratios of thermal and chemical driving are made.

We observe that fixed flux conditions promote larger flow scales and an increase of mean kinetic energy densities. The imposed flux pattern locks the outer core flow to the mantle and therefore breaks its azimuthal symmetry, even for relatively low thermal forcing ratios of 20 %. Despite of the symmetry breaking, stable and dipolar dynamos can be maintained due to the partly chemical forcing with its uniform boundary conditions. Additionally, the chemical component partly adopts to the geometry of the heat flux pattern because advective transport of concentration is more effective in regions of increased heat flux.

1 Introduction and Motivation

1.1 A thermo-chemical approach to dynamo modeling

1.2 Why use heterogeneous heat flux boundary conditions?

2 Modeling core convection and dynamo action

Rotating convection and dynamo action in Earth's core is a topic that has been extensively studied (Chandrasekhar, 1961). There exist various ideas of how to formulate a set of equations that depicts all relevant physical processes and that is still as simple and therefore numerically economical as possible. Computation time is the limiting factor when it comes to the question how realistic models of the inner core can be. The gap between the relevant physical parameters expected for the Earth and the parameters that are in range of numerical modeling is still big. It cannot be expected that this gap will be closed only with the help of the increasing computational resources that will become available within the next years. Alternatives to waiting for larger computers that allow to explore earth-like parameters have to be found. Asymptotic models are one possibility already revealing promising results (REFS ???).

A question that is closely related to the numerical costs and therefore to the accessible parameter range is the choice of the geometry. Most models are either Cartesian with periodic boundary conditions (BC) or spherical shell models. The latter are of course more realistic for the Earth but numerically more costly and therefore even less earth-like with regard to computational feasible parameters. In this work, a spherical model is used in order to be as earth-like as possible in a geometrical way. As a trade-off, parameter regimes in which Cartesian models could advance, are unreachable here.

We model the electrically conducting liquid outer core. It is enclosed by the inner-core boundary (ICB) at the bottom and the core-mantle boundary (CMB) at the top. Convection is driven by destabilizing thermal and compositional gradients across the sphere. For the compositional component, fixed chemical concentrations at the ICB and the CMB are imposed (Dirichlet BC). The thermal forcing is maintained by introducing a fixed in - and outflux of heat (Neumann BC). The heat flux at the CMB is laterally heterogeneous, i.e., there exist regions of higher and lower heat flux than lateral average. The inner core and mantle are assumed to be insulating and therefore have no influence on the evolution of magnetic fields.

In the following all relevant equations are introduced. There exist numerous detailed derivations so that this description will be held relatively short. The according references will be given in each section.

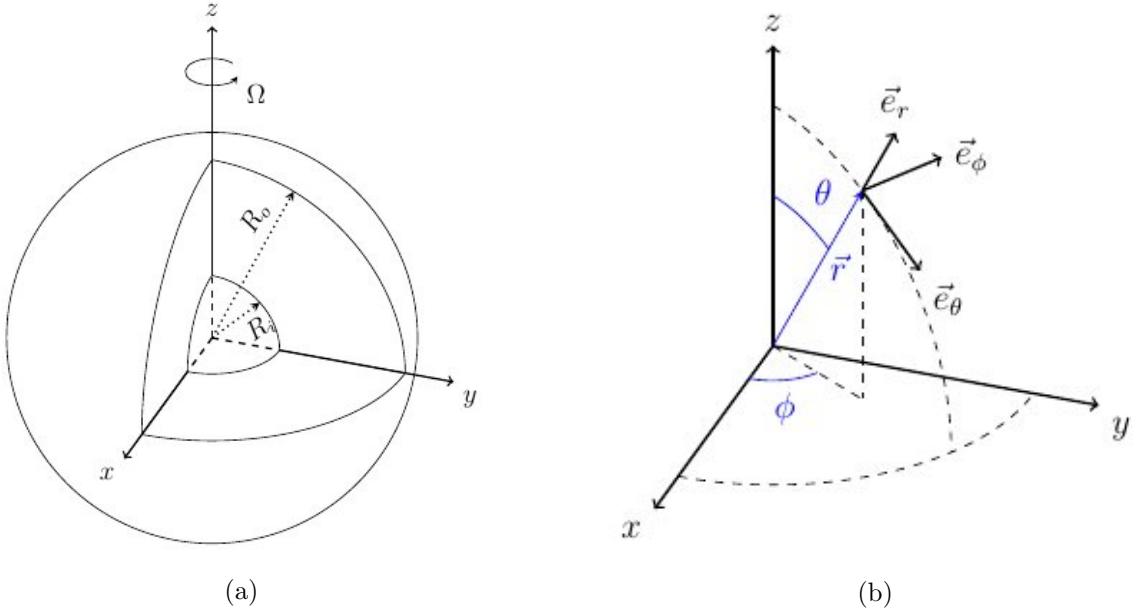


Figure 2.1: (a) Sketch of the liquid outer core. It is bounded by the two spheres of radii R_i and R_o . The rotation axis is the cartesian z-axis. (b) This work uses spherical coordinates with unit vectors e_r , e_Φ and e_θ (Trümper, 2014).

2.1 Frame of reference

The liquid outer core (LOC) is modeled in a spherical shell with an inner radius R_i and an outer radius R_o . It is bounded by the ICB at the bottom the CMB at the top. The shell thickness is chosen according to what is expected for today's state of the Earth and defined over the ratio between R_i and r_o : $a = R_i / R_o = 0.35$.

The LOC is constantly rotating about the z-axis of a Cartesian system. The angular velocity Ω is invariant in time. The effect of the rotation on the frame of reference will be discussed when it comes to the equation of motion. In the course of this study, it is quite nearby to chose spherical coordinates (??).

2.2 Equation of continuity

Per definition, the mass \mathcal{M} of a material volume $\mathcal{V}(t)$ with a density ρ , moving with a velocity $\mathbf{u}(\mathbf{r}, t)$ in a fluid is conserved:

$$\frac{d\mathcal{M}(\mathcal{V})}{dt} = \frac{d}{dt} \int_{\mathcal{V}(t)} \rho(\mathbf{r}, t) d^3r = 0 \quad (2.1)$$

Using Reynold's Transport Theorem and then applying Gauss's Theorem, this yields

$$\int_{\mathcal{V}(t)} \frac{\partial \rho}{\partial t} d^3r + \oint_{\partial \mathcal{V}(t)} \rho \mathbf{u} \cdot d\mathbf{s} = \int_{\mathcal{V}(t)} \left[\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) \right] d^3r = 0. \quad (2.2)$$

Since this has to hold for all possible material volumes \mathcal{V} , one gets

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \mathbf{u}) = 0, \quad (2.3)$$

the general form of the equation of continuity. The introduction of the *material derivative* $\frac{D}{Dt} = \frac{\partial}{\partial t} + \mathbf{u} \cdot \nabla$ suggests another useful formulation:

$$\frac{\partial \rho}{\partial t} + \rho \nabla \cdot \mathbf{u} + \mathbf{u} \cdot \nabla \rho = \frac{D\rho}{Dt} + \rho \nabla \cdot \mathbf{u} = 0. \quad (2.4)$$

2.3 Equation of momentum

In an inertial, non rotating frame of reference, the change of momentum of a material volume \mathcal{V} can be written as

$$\begin{aligned} \frac{d}{dt} \int_{\mathcal{V}(t)} (\rho u_j) d^3r &= \int_{\mathcal{V}(t)} \left[\frac{\partial(\rho u_j)}{\partial t} + \frac{\partial}{\partial x_i}(\rho u_j) u_i \right] d^3r \\ &= \int_{\mathcal{V}(t)} \left[\rho \frac{\partial u_j}{\partial t} + u_j \frac{\partial \rho}{\partial t} + \rho u_j \frac{\partial u_i}{\partial x_i} + u_i \rho \frac{\partial u_j}{\partial x_i} + u_i u_j \frac{\partial \rho}{\partial x_i} \right] d^3r \\ &= \int_{\mathcal{V}(t)} \left[u_j \left(\frac{\partial \rho}{\partial t} + \rho \nabla \cdot \mathbf{u} + \mathbf{u} \cdot \nabla \rho \right) + \rho \left(\frac{\partial u_j}{\partial t} + \mathbf{u} \cdot \nabla u_j \right) \right] d^3r \\ &= \int_{\mathcal{V}(t)} \rho \frac{Du_j}{Dt} d^3r. \end{aligned}$$

Change of momentum can happen through either *volume forces* \mathbf{f} or *surface forces* \mathbf{t} :

$$\int_{\mathcal{V}} \rho \frac{Du_j}{Dt} d^3r = \int_{\mathcal{V}} \mathbf{f} d^3r + \oint_{\partial\mathcal{V}} \mathbf{t} ds \quad (2.5)$$

If a frame of reference is rotating - as in this case - it is no longer an inertial system and therefore pseudo forces have to be expected. Under the assumption of a constant rotation with the angular velocity Ω and a fixed rotation axis parallel to the Cartesian z-axis, the change of a quantity \mathbf{P} in the rotating frame of reference and its change in the non rotating inertial frame relate as

$$\left(\frac{d\mathbf{P}}{dt} \right)_F = \left(\frac{d\mathbf{P}}{dt} \right)_R + \boldsymbol{\Omega} \times \mathbf{P}, \quad (2.6)$$

where the subscripts F and R indicate the *fixed* and the *rotating* frame. Applying rule (2.6) twice to a position vector \mathbf{r} yields

$$\mathbf{a}_F = \mathbf{a}_R + 2\boldsymbol{\Omega} \times \mathbf{u}_R + \boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r}) \quad (2.7)$$

for accelerations \mathbf{a} . $2\boldsymbol{\Omega} \times \mathbf{u}_R$ is the Coriolis force and $\boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r})$ the centripetal force.

With the help of (2.7), $\frac{D\mathbf{u}}{Dt}$ can be transferred from the inertial frame to the frame of reference:

$$\frac{D\mathbf{u}_F}{Dt} = \frac{D\mathbf{u}_R}{Dt} + 2\boldsymbol{\Omega} \times \mathbf{u}_R + \boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r}) \quad (2.8)$$

From now on, the subscripts F and R will be cut and all quantities will be measured in the rotating frame. With (2.8), (2.5) changes to

$$\int_V \rho \frac{D\mathbf{u}}{Dt} d^3r = \int_V \mathbf{f} d^3r + \oint_{\partial V} \mathbf{t} ds - \int_V \rho (2\boldsymbol{\Omega} \times \mathbf{u} + \boldsymbol{\Omega} \times (\boldsymbol{\Omega} \times \mathbf{r})) d^3r. \quad (2.9)$$

Another pseudo force that generally appears in rotating systems, but that is neglected here, is the *Poincaré force*. Precession driven flows are the most prominent example from geophysics where it plays a dominant role (Tilgner, 2007).

The *Cauchy Theorem* relates the surface forces \mathbf{t} to the stress tensor $\underline{\tau}$ linearly via $\mathbf{t} = \underline{\tau} \cdot \mathbf{n}$, where \mathbf{n} is the normal vector. The surface term in (2.9) can thus be transformed to

$$\oint_{\partial V} \mathbf{t} ds = \int_V \nabla \cdot \underline{\tau} d^3r, \quad (2.10)$$

where $\nabla \cdot \underline{\tau} = -\nabla p + \mu \nabla^2 \mathbf{u}$ will be used as a reasonable simplification in the context of the Boussinesq approximation (see section 2.6). This expression of $\nabla \cdot \underline{\tau}$ is valid for Newtonian fluids and it is based on the assumption of a solenoidal velocity field ($\nabla \cdot \mathbf{u} = 0$) and a homogeneous dynamic viscosity μ throughout the fluid.

The body forces \mathbf{f} are the *buoyancy force* $\mathbf{f}_g = \mathbf{g}\rho$ and the *Lorentz force* $\mathbf{f}_l = \mathbf{j} \times \mathbf{B}$. The latter will be discussed in section 2.4.

In the following, the gravitational field \mathbf{g} will be discussed in more detail. As mentioned before, the system underlies an asymmetric heat flux at the CMB. This results in an asymmetric temperature field for which the penetration depth of the temperature perturbation to a spherical symmetric solution depends on the amplitude of the heat flux heterogeneity. Following the linear equation of state from the Boussinesq approximation (see section 2.6), this results in a density variation proportional to the temperature variation: $\delta\rho \sim \delta T$.

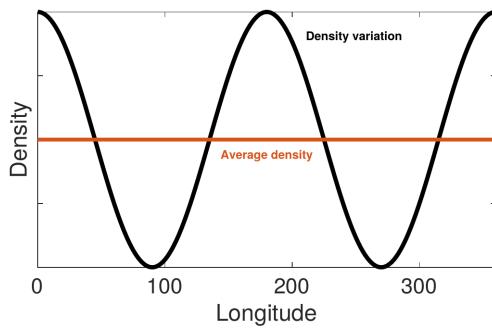


Figure 2.2: Sketch of the density along the equator, near the CMB. The deviation (black) from the average density (red) results from the asymmetric heat flux at the CMB which has an asymmetric temperature field as a consequence.

The total density can be split into one spherical symmetric part $\mathring{\rho}(r)$ that only depends on the radial level r and the density variation due to the heat flux pattern $\delta\mathring{\rho}(r, \vartheta, \phi)$:

$$\rho(r, \vartheta, \phi) = \mathring{\rho}(r) + \delta\mathring{\rho}(r, \vartheta, \phi). \quad (2.11)$$

In case of a heat flux pattern proportional to the spherical harmonic \mathcal{Y}_2^2 , the dependence of $\delta\mathring{\rho}$ on ϕ is π -periodic (see Figure 2.2) so that it can be expressed by the product $\delta\mathring{\rho}(r, \vartheta, \phi) = \mathcal{C}(r, \vartheta) \cdot \cos(2\phi)$. \mathcal{C} is only a function of r and ϑ .

Gauss's gravity law (Blakely, 1996) states

$$\nabla \cdot \mathbf{g} = -4\pi G \rho(r, \vartheta, \phi). \quad (2.12)$$

Integration over a sphere of radius r and application of Gauss's theorem yields

$$\begin{aligned} \int_{\mathcal{V}(r)} \nabla \cdot \mathbf{g} dV &= \int_0^\pi \int_0^{2\pi} r^2 \sin(\vartheta) d\vartheta d\phi \mathbf{g} \cdot \mathbf{e}_r = -4\pi G \int_0^r \int_0^\pi \int_0^{2\pi} \sin(\vartheta) r'^2 \rho(r, \vartheta, \phi) dr' d\vartheta d\phi \\ &= -4\pi G \int_0^r \int_0^\pi \int_0^{2\pi} \sin(\vartheta) r'^2 [\mathring{\rho}(r') + \delta\mathring{\rho}(r', \vartheta, \phi)] dr' d\vartheta d\phi \\ &= -8\pi^2 G \int_0^r \mathring{\rho}(r') dr' - 4\pi G \int_0^r \int_0^\pi \sin(\vartheta) r'^2 \mathcal{C}(r', \vartheta) \left(\int_0^{2\pi} \cos(2\phi) d\phi \right) dr' d\vartheta \\ &= -8\pi^2 G \int_0^r \mathring{\rho}(r') dr' \end{aligned}$$

Because the integration of $\delta\mathring{\rho}(r', \vartheta, \phi)$ over ϕ drops out for every value of r and ϑ , \mathbf{g} can be expressed by

$$g(\mathbf{r}) = -\frac{4\pi G}{r^2} \int_0^r \mathring{\rho}(r') r'^2 dr \quad (2.13)$$

and it is worth noticing that \mathbf{g} has the same form as in the spherical symmetric case (Blakely, 1996).

The buoyancy force term $\mathbf{g}\rho$ will be further discussed in section 2.6.

2.4 Maxwell equations, Lorentz force and induction equation

The liquid outer core consists of a metallic and therefore conducting fluid. Electric currents may evolve, create magnetic fields and these again may generate currents and influence the flow field. The induction equation is a transport equation for a magnetic fields \mathbf{B} 'hosted' by a fluid moving with a velocity \mathbf{u} . The Lorentz force characterizes the influence of the magnetic field on the flow field, whereas the Maxwell equations describe how electric and magnetic fields interact through charges

and currents. They form a basis for the 'magnetic part' of Magnetohydrodynamics (MHD).

In the scope of core convection, a reduced form of the Maxwell equations (Pre-Maxwell equations) suffices (Davidson, 2001):

$$\nabla \times \mathbf{B} = \mu_0 \mathbf{j} \quad (\text{Ampère's law}) \quad (2.14a)$$

$$\nabla \cdot \mathbf{j} = 0 \quad (\text{Charge conservation}) \quad (2.14b)$$

$$\nabla \times \mathbf{E} = -\frac{\partial \mathbf{B}}{\partial t} \quad (\text{Faraday's law}) \quad (2.14c)$$

$$\nabla \cdot \mathbf{B} = 0 \quad (\text{No magnetic monopoles}) \quad (2.14d)$$

Additionally, an extended version of *Ohm's law* for moving conductors

$$\mathbf{j} = \sigma(\mathbf{E} + \mathbf{u} \times \mathbf{B}) \quad (2.15)$$

and an expression for the *Lorentz force*

$$\mathbf{f}_l = \mathbf{j} \times \mathbf{B} = \frac{1}{\mu_0}(\nabla \times \mathbf{B} \times \mathbf{B}) \quad (2.16)$$

are needed. \mathbf{E} describes the electric field, \mathbf{j} the current density, μ_0 the vacuum permeability and σ the conductivity of the fluid.

The *induction equation*, a transport equation for the magnetic field, can be derived using (2.14c), (2.15) and the solenoidal character of \mathbf{B} (2.14d) and \mathbf{u} :

$$\frac{\partial \mathbf{B}}{\partial t} = \nabla(\mathbf{u} \times \mathbf{B}) + \eta \nabla^2 \mathbf{B}, \quad (2.17)$$

where $\eta = \frac{1}{\sigma \mu_0}$ is the magnetic diffusivity.

2.5 Conservation of energy and the light component

The conservation of internal energy in a fluid reads

$$\rho \frac{De}{Dt} = -\nabla \cdot \mathbf{q}_{\text{T}} - p(\nabla \cdot \mathbf{u}) + \Phi, \quad (2.18)$$

where the internal energy per unit mass is described by e . It can be changed by either *volume compression* $-p(\nabla \cdot \mathbf{u})$, *viscous dissipation* Φ or a *heat flux* \mathbf{q}_{T} through the fluid surface. In the context of the Boussinesq approximation (section 2.6), viscous dissipation Φ is negligible and $\nabla \cdot \mathbf{u} = 0$. Furthermore, using the *Fourier law* $\mathbf{q}_{\text{T}} = -k_{\text{T}} \nabla T$ and the *perfect gas* approximation $e = c_{\text{P}}T$, (2.18) can be transformed to

$$\frac{DT}{Dt} = \kappa_{\text{T}} \nabla^2 T. \quad (2.19)$$

Here, $\kappa_{\text{T}} = \frac{k}{c_{\text{P}}\rho}$ is the thermal diffusivity and T the fluid temperature. Internal sources of e in the liquid outer core are completely omitted in this study.

The derivation of equation (2.19) was adopted from Kundu and Cohen (2008).

The light component is released at the ICB as the inner core slowly crystallizes. It serves as an additional source of buoyancy. In this model, the light component per unit mass, C , can only change by a flux \mathbf{q}_c through the surface of the fluid. According to the equation for the conservation of the internal energy above,

$$\rho \frac{DC}{Dt} = -\nabla \cdot \mathbf{q}_c$$

can be transformed to

$$\frac{DC}{Dt} = \kappa_c \nabla^2 C, \quad (2.20)$$

using $\mathbf{q}_c = -k_c \nabla C$ and introducing $\kappa_c = \frac{k_c}{\rho}$.

2.6 The Boussinesq approximation

In most geophysical applications of fluid dynamics, the Boussinesq approximation is a reasonable simplification of the full equations.

Starting from an adiabatic reference state, density, pressure, temperature and composition can be separated into a reference state value (denoted by an overbar) that is only dependent on the radial level r and a fluctuating part (denoted by a prime):

$$T = \bar{T}(r) + T', \quad C = \bar{C}(r) + C', \quad \rho = \bar{\rho}(r) + \rho', \quad p = \bar{p}(r) + p' \quad (2.21)$$

Further on, it is assumed that the typical length scale of the system (here: the shell thickness d) is small compared to the scale heights in the reference state. Perturbations to that state shall be small compared to the adiabat. As a result, the reference state becomes independent of position.

Whether the assumption of a negligible small superadiabaticity is valid in the context of core convection is still debated (Anufriev et al., 2005). The alternative is to model either the fully compressible equations or to use the anelastic approximation that, in contrast to the Boussinesq model, allows density stratification. Jones (2007) extensively discusses the implications of the choice of one of these approaches.

From the Boussinesq approximation it follows for the system of equations that

- the equation of state takes a linear form:

$$\rho' = -\bar{\rho}(\alpha_T T' + \alpha_C C'), \quad (2.22)$$

where α_T , and α_C signify the thermal and compositional expansion coefficient.

- the equation of continuity transforms to

$$\nabla \cdot \mathbf{u} = 0 \quad (\text{solenoidal character of } \mathbf{u}). \quad (2.23)$$

- the dissipative heating term can be neglected in the equation of momentum and internal energy (see section 2.3 and 2.5).
- the material properties such as α_T , α_C , μ_0 , c_P , η , κ_T and κ_C stay constant throughout the fluid.

An important advantage of the modified equations is that sound waves are *filtered out*. This reduces the numerical costs without curtailing any relevant physical processes, since the short time scales of sound waves does not have to be resolved numerically. In the differential form, the equation of momentum (2.9) now reads

$$\frac{D\mathbf{u}}{Dt} = -2\boldsymbol{\Omega} \times \mathbf{u} - \nabla \left(\frac{\pi'}{\bar{\rho}} \right) + \nu \nabla^2 \mathbf{u} + \frac{1}{\mu_0 \bar{\rho}} (\nabla \times \mathbf{B}) \times \mathbf{B} + \frac{\mathbf{g}}{\bar{\rho}} \rho', \quad (2.24)$$

where $\nu = \mu/\bar{\rho}$ is the new kinematic viscosity. The centrifugal potential is incorporated into the pressure fluctuation term $\pi' = p' - \frac{\bar{\rho}\Omega^2 s^2}{2}$, where s is the distance from the z-axis in cylindrical coordinates. The hydrostatic pressure gradient of the reference state $\nabla \bar{p}$ is balanced by the gravity force $\mathbf{g} \bar{\rho}$ so that only ρ' appears equation (2.24).

2.7 Boundary conditions

This section is dedicated to the boundary conditions.

For the velocity field they are chosen to be *no-slip*, e.g. $\mathbf{u} = 0$ on the ICB ($r = R_i$) and the CMB ($r = R_0$), in order to be comparable to other studies. Several authors state that these *rigid* BC reflect what is realistic for the interaction between the inner and outer core and the outer core and the mantle, respectively (Glatzmaier and Roberts, 1995; Christensen and Aubert, 2006; Trümper et al., 2012). On the other hand, Zhang and Busse (1987) or Kuang and Bloxham (1997) argue that the Ekman layers which result from rigid boundaries are negligibly thin in the Earth's core due to its extremely small viscosity. Because today's numerical models are forced to use a far to high viscosity, they over emphasize the effect of Ekman layers and therefore *stress-free* BC are more appropriate. For a better comparison, this work follows the approach that was used by related studies (Olson and Christensen, 2002; Aubert et al., 2008a; Hori et al., 2014).

According to Trümper et al. (2012), Derichlet type BC are chosen for the chemical component. A compositional gradient, $\Delta C = C'_{\text{ICB}} - C'_{\text{CMB}}$, is imposed across the shell. In a more earth-like scenario, one would apply Neumann type conditions with fixed compositional influx at the ICB and zero outflux at the CMB (Braginsky and Roberts, 1995). This would further increase the complexity of the model and is therefore adjourned in this case.

The mantle is assumed to be an insulator due to its rocky content (Dormy et al., 1998). The question whether the inner core should be treated as an insulator or not was discussed by Wicht (2002). He found that the effect of a conducting inner

core on the flow field and the magnetic field is rather small. It is hence reasonable to choose $\sigma = 0$ in the inner core in order not to be obliged to solve the induction equation in the full sphere. In the mantle and the core, the magnetic field \mathbf{B} is the solution to

$$\mathbf{B} = -\nabla\Phi \quad (2.25)$$

with Φ being a scalar potential that follows the Laplace equation

$$\nabla^2\Phi = 0. \quad (2.26)$$

The BC are expressed through the fact that \mathbf{B} has to match the solution of (2.25) at the ICB and the CMB.

The thermal BC is of the Neumann type and will be discussed in more detail in the following section.

2.8 CMB - heat flux pattern

A heat flux balance between the inner and the outer core boundary is assumed. The total influx at the ICB Q_i equals the total outflux Q_o at the CMB:

$$-Q_i = Q_o \quad (2.27)$$

$$\Leftrightarrow \int_{S_{\text{ICB}}} \kappa_T \nabla T \Big|_{\text{ICB}} \cdot \mathbf{e}_r dS = \int_{S_{\text{CMB}}} \kappa_T \nabla T \Big|_{\text{CMB}} \cdot \mathbf{e}_r dS \quad (2.28)$$

$$\Leftrightarrow \int_{S_{\text{ICB}}} \frac{\partial T}{\partial r} \Big|_{\text{ICB}} dS = \int_{S_{\text{CMB}}} \frac{\partial T}{\partial r} \Big|_{\text{CMB}} dS \quad (2.29)$$

For the spectral decomposition it is important to notice that only the 0th order spherical harmonic \mathcal{Y}_0^0 yields values $\neq 0$ when being integrated over a closed surface S :

$$\int_S \frac{\partial T}{\partial r} dS = \int_S \left(\frac{\partial T}{\partial r} \right)_0^0 \mathcal{Y}_0^0 dS,$$

with $\left(\frac{\partial T}{\partial r} \right)_0^0$ being the spectral coefficient of degree and order 0. Thus, (2.29) can be transformed into the spectral domain via

$$\begin{aligned} \int_{S_{\text{ICB}}} \left(\frac{\partial T}{\partial r} \right)_0^0 \Big|_{\text{ICB}} \mathcal{Y}_0^0 dS &= \int_{S_{\text{CMB}}} \left(\frac{\partial T}{\partial r} \right)_0^0 \Big|_{\text{CMB}} \mathcal{Y}_0^0 dS \\ \Leftrightarrow \left(\frac{\partial T}{\partial r} \right)_0^0 \Big|_{\text{ICB}} R_i^2 &= \left(\frac{\partial T}{\partial r} \right)_0^0 \Big|_{\text{CMB}} R_o^2. \end{aligned}$$

This allows to formulate a simple relation between the mean radial temperature gradient at the inner and outer boundary:

$$\left(\frac{\partial T}{\partial r}\right)_0^0 \Big|_{\text{ICB}} = \left(\frac{\partial T}{\partial r}\right)_0^0 \Big|_{\text{CMB}} \frac{R_o^2}{R_i^2} = -\beta \frac{R_o^2}{R_i^2} = -\beta \frac{1}{a^2} \quad (2.30)$$

with $\beta := -\left(\frac{\partial T}{\partial r}\right)_0^0 \Big|_{\text{CMB}}$ as prescribed temperature gradient of degree and order 0 at the CMB.

Since no internal sources of heat are assumed, the stationary form of (2.19) takes the form of a Laplace equation. It reads

$$\nabla^2 T = 0 \quad (2.31)$$

and with (2.30) the appropriate BC are

$$\text{at the ICB: } \left.\frac{\partial T}{\partial r}\right|_{\text{ICB}} = -\beta \frac{1}{a^2} \mathcal{Y}_0^0 \quad (2.32)$$

$$\text{at the CMB: } \left.\frac{\partial T}{\partial r}\right|_{\text{CMB}} = -\beta \mathcal{Y}_0^0 + \text{Amp}_l^m \mathcal{Y}_l^m. \quad (2.33)$$

Amp_l^m is the amplitude of the heat flux heterogeneity for a heat flux pattern \mathcal{Y}_l^m . Its definition follows the work of Hori et al. (2014):

$$\text{Amp} = \frac{q_{\max} - q_{\min}}{2q_{\text{mean}}} \quad (2.34)$$

The mean heat flux at the CMB q_{mean} results from the temperature gradient β at the CMB. q_{\max} and q_{\min} are the extrema of the lateral heat flux variation.

In order to implement the BC into the numerical model, a (conductive) solution to (2.31) has to be found. In the most general form, it reads

$$T = \sum_{l,m} [a_l^m r^l + b_l^m r^{-l-1}] \mathcal{Y}_l^m(\vartheta, \phi) \quad (2.35)$$

and its radial derivative

$$\frac{\partial T}{\partial r} = \sum_{l,m} [l a_l^m r^{l-1} - (l+1) b_l^m r^{-l-2}] \mathcal{Y}_l^m(\vartheta, \phi). \quad (2.36)$$

At the ICB, (2.32) and (2.36) yield

$$-b_0^0 R_i^{-2} = -\beta \frac{1}{a^2} \Rightarrow b_0^0 = \beta R_o^2 \quad (2.37)$$

for $l = m = 0$ and the algebraic equation

$$la_l^m R_i^{l-1} - (l+1)b_l^m R_i^{-l-2} = 0 \quad (2.38)$$

follows for all $l > 0$ and $m > 0$.

For the CMB, (2.33) and (2.36) yield

$$la_l^m R_o^{l-1} - (l+1)b_l^m R_o^{-l-2} = Amp_l^m. \quad (2.39)$$

Since only a heat flux pattern with $l = m = 2$ will be used here, (2.38) and (2.39) simplify to a linear system of equations

$$\begin{pmatrix} 2R_o & -3R_o^{-4} \\ 2R_i & -3R_i^{-4} \end{pmatrix} \begin{pmatrix} a_2^2 \\ b_2^2 \end{pmatrix} = \begin{pmatrix} Amp_2^2 \\ 0 \end{pmatrix}. \quad (2.40)$$

Inserting a_2^2 , b_2^2 and b_0^0 into (2.35) gives a solution to (2.31) that respects the boundary conditions (2.32) and (2.33):

$$T_{\text{cond}}(r, \vartheta, \phi) = \left(a_0^0 + \frac{R_o^2 \beta}{r} \right) \mathcal{Y}_0^0 + Amp_2^2 \cdot \frac{R_o^2 2a^5 r^3 - 3R_o^2 (a-1)^5 r^2}{6D^4(a-1)^2(1-a^5)} \mathcal{Y}_2^2(\vartheta, \phi) \quad (2.41)$$

a_0^0 is an arbitrary integration constant that is chosen to be 0 in the following. This conductive temperature profile with its spherically asymmetric part will be introduced into the equations in the density fluctuation part ρ' in section 2.10.

The effect of the CMB - heat flux pattern on the conductive temperature field is shown in Figure 2.3(a) and (b). Regions, where the heat flux from the core to the mantle is increased (heat flux maxima are located at 90° and 270° east), are characterized by relatively low temperatures because they are cooled more effectively. Vice versa, regions of reduced heat flux (heat flux minima are located at 0° and 180°) show higher temperatures than average.

The amplitude of heterogeneity, Amp_2^2 , from now on referred to as q^* , plays a central role for the heat flux pattern. Figure 2.4 shows the core-mantle heat flux along the equator for two values of q^* . If the amplitude is chosen to be greater than 1, negative (inward) heat flux occurs in distinct regions around the flux minima. Heat flow from the mantle to the core produces gravitationally stable regions with respect to temperature. Together with an always destabilizing compositional gradient, this could yield non linear double-diffusive effects as fingering or layering. A scenario in which gravitationally stable regions play a role in core convection was recently discussed for mercury (Tian et al., 2015). Yet for the Earth, where the heterogeneities in core-mantle heat flux are considered to be rather mild (Aubert et al., 2008a; Olson and Christensen, 2002; Glatzmaier et al., 1999), Nakagawa and Tackley (2008) expect that heat flows from the mantle to the core where compositionally dense piles settle at the base of the mantle. Because of that and due to the fact that this study

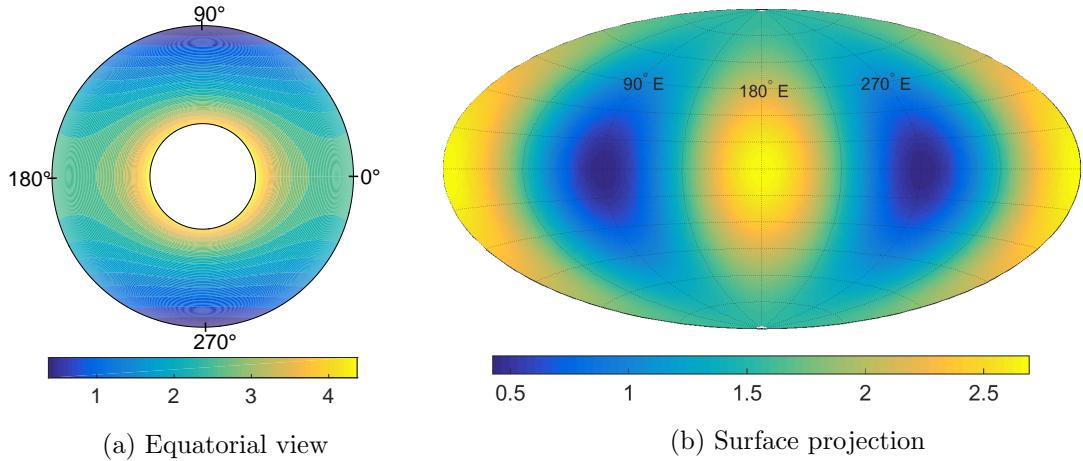


Figure 2.3: Conductive temperature field T_{cond} in the equatorial plain (left) and as a *Hammer-Aitov-projection* at the CMB (right) in a non dimensional form (see section 2.9). The amplitude of the heat flux heterogeneity is $\text{Amp}_2^2 = 1$, heat flux maxima are located at 90° and 270° , respectively. See Figure 2.5 for a schematic overview.

aims to explore the effect of a heat flux pattern on core convection from a general perspective, values of $q^* > 1$ are likewise regarded.

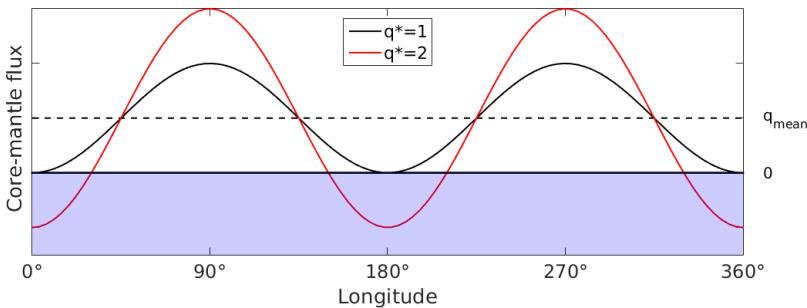


Figure 2.4: Plot of the CMB - heat flux along the equator for amplitudes of heterogeneity $q^* = 1$ and $q^* = 2$. For $q^* > 1$, heat partially flows from the mantle to the core (blue region).

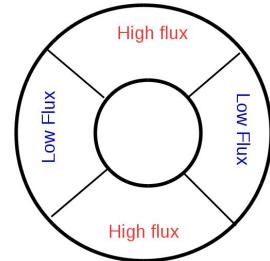


Figure 2.5: Sketch of the equatorial plain and the distribution of high and low heat flux at the CMB.

2.9 Nondimensional equations

A common procedure in fluid dynamics is to *rescale* the equations introduced so far in order to reduce the number of parameters and to make the relevant physical processes intuitively more accessible. The new scales are adopted from Trümper et al. (2012), they are summarized in table 1.

Variable	Symbol	Scale
Length	\mathbf{r}	$D = R_o - R_i$
Time	t	D^2/ν
Velocity	\mathbf{u}	ν/D
Temperature	T	βD
Composition	C	ΔC
Pressure	π	$\bar{\rho}\nu^2/D^2$
Magnetic Field	\mathbf{B}	$\sqrt{\eta\Omega\mu_0\bar{\rho}}$

Table 1: Overview of the new scales that are introduced to non dimensionalize the magneto-hydrodynamic system of equations that was introduced in the sections 2.2 to 2.6.

The application of these scales to (2.23), (2.24), (2.19), (2.20) and (2.17) yields a system of non dimensional equations:

$$\hat{\nabla} \cdot \hat{\mathbf{u}} = 0, \quad (2.42a)$$

$$\begin{aligned} \frac{D\hat{\mathbf{u}}}{Dt} = & -\hat{\nabla}\hat{\pi}' - \frac{2}{Ek} \mathbf{e}_z \times \hat{\mathbf{u}} + \hat{\nabla}^2 \hat{\mathbf{u}} + \frac{1}{EkPr_m} (\hat{\nabla} \times \hat{\mathbf{B}}) \times \hat{\mathbf{B}} \\ & + (Ra_T \hat{T}' + Ra_C \hat{C}')(1-a)\hat{r} \mathbf{e}_r, \end{aligned} \quad (2.42b)$$

$$\frac{D\hat{T}'}{Dt} = \frac{1}{Pr_T} \hat{\nabla}^2 \hat{T}', \quad (2.42c)$$

$$\frac{D\hat{C}'}{Dt} = \frac{1}{Pr_C} \hat{\nabla}^2 \hat{C}' \text{ and} \quad (2.42d)$$

$$\frac{\partial \hat{\mathbf{B}}}{\partial \hat{t}} = \hat{\nabla} \times (\hat{\mathbf{u}} \times \hat{\mathbf{B}}) + \frac{1}{Pr_m} \hat{\nabla}^2 \hat{\mathbf{B}}. \quad (2.42e)$$

Nondimensional quantities are denoted by $\hat{\cdot}$ and they are related to their dimensional counterparts via $(\hat{\cdot}) = scale(\cdot)$.

The following parameters of similarity appear in the equations:

Thermal Rayleigh number :	$Ra_T = \frac{\alpha_T g \beta D^4}{\nu^2}$
Compositional Rayleigh number :	$Ra_C = \frac{\alpha_C g \Delta C D^3}{\nu^2}$
Ekman number :	$Ek = \frac{\nu}{\Omega D^2}$
Thermal Prandtl number :	$Pr_T = \frac{\nu}{\kappa_T}$
Compositional Prandtl number :	$Pr_C = \frac{\nu}{\kappa_C}$
Magnetic Prandtl number :	$Pr_m = \frac{\nu}{\eta}$

The thermal and compositional Rayleigh numbers are measures for the vigor of convection due to thermal and compositional buoyancy sources, respectively. The definitions of Ra_T and Ra_C slightly differ because of the different scales chosen for temperature and composition (see table 1). ' $\beta D'$ and ' $\Delta C'$ ' refer to either Neumann or Derichlet type boundary conditions for T and C , respectively.

Whether a fluid is constrained rather by rotation or viscosity is described via the Ekman number.

The distinction between a thermal and a compositional Prandtl number is a key feature of this study. It allows for differences in the dynamical response of the system to either thermally or compositionally dominated convective forcing. Roughly spoken, a large Prandtl number promotes viscous effects in a fluid, whereas a small one promotes inertia effects. The magnetic Prandtl number relates viscous and magnetic diffusion and therefore decides how much kinetic energy is needed to sustain a magnetic field.

In the following, all $(\hat{\cdot})$ will be omitted, since only non dimensional quantities are mentioned.

The rescaled inner and outer radii of the spherical shell are $R_i = 0.539$ and $R_o = 1.539$, respectively.

2.10 Conductive and convective Fluctuations

The temperature field T' and the chemical field C' , each reduced by the reference state fields \bar{T} and \bar{C} , can be decomposed into a *conductive part* which is the full solution if the system is subcritical and its *convective perturbations* θ and ζ :

$$T'(\mathbf{r}) = T_{\text{cond}}(\mathbf{r}) + \theta(\mathbf{r}) \quad C'(\mathbf{r}) = C_{\text{cond}}(\mathbf{r}) + \zeta(\mathbf{r})$$

The stationary solutions T_{cond} and C_{cond} can be obtained analytically (see section 2.8 for the temperature field), so that *PARODY* has to solve only for θ and ζ . The conductive solution C_{cond} follows from the stationary form of (2.20) and the boundary conditions described in section 2.7. The temperature and chemical equation transform to

$$\frac{D\theta}{Dt} = \frac{1}{\text{Pr}_T} \nabla^2 \theta - (\mathbf{u} \cdot \nabla) T_{\text{cond}} \quad (2.43a)$$

$$\frac{D\zeta}{Dt} = \frac{1}{\text{Pr}_C} \nabla^2 \zeta - (\mathbf{u} \cdot \nabla) C_{\text{cond}}, \quad (2.43b)$$

when the decomposition is inserted.

Furtheron, the equation of momentum slightly changes, when temperature and composition are split into two parts.

T_{cond} has a spherically symmetric and a spherically asymmetric part

$$T_{\text{cond}}(r, \vartheta, \phi) = t_0^0(r) + t_2^2(r, \vartheta, \phi),$$

where $t_0^0(r)$ is proportional to the constant \mathcal{Y}_0^0 and $t_2^2(r)$ to $\mathcal{Y}_2^2(\vartheta, \phi)$ that is a function of ϑ and ϕ .

$t_0^0(r)$ and the spherical symmetric solution $C_{\text{cond}}(r)$ can be incorporated into the reduced pressure fluctuation term π' via

$$\pi = \pi' - (1-a) \left(\text{Ra}_{\text{T}} \int r t_0^0(r) dr + \text{Ra}_{\text{C}} \int r C_{\text{cond}}(r) dr \right).$$

Yet, it is not possible to find a potential τ for the asymmetric part that satisfies

$$\nabla \tau = \text{Ra}_{\text{T}} t_2^2(r) (1-a) r \mathbf{e}_r,$$

since this would demand $\partial \tau / \partial \vartheta = \partial \tau / \partial \phi = 0$, a condition that cannot be fulfilled by t_2^2 being a function of ϑ and ϕ . Thus, t_2^2 appears in the buoyancy term of the momentum equation:

$$\begin{aligned} \frac{D\mathbf{u}}{Dt} = & -\nabla \pi - \frac{2}{\text{Ek}} \mathbf{e}_z \times \mathbf{u} + \nabla^2 \mathbf{u} + \frac{1}{\text{EkPr}_m} (\nabla \times \mathbf{B}) \times \mathbf{B} \\ & + (\text{Ra}_{\text{T}}[\theta + t_2^2] + \text{Ra}_{\text{C}}\zeta)(1-a)r \mathbf{e}_r \end{aligned} \quad (2.44)$$

2.11 Diagnostics

This chapter briefly introduces the diagnostic quantities and techniques that are used in the course of this work:

- Vorticity

The vorticity $\boldsymbol{\omega}$ describes the local rotation of a fluid and can be computed via the curl of the velocity:

$$\boldsymbol{\omega} = \nabla \times \mathbf{u} \quad (2.45)$$

- Helicity

Helicity is generated where a flow \mathbf{u} is oriented parallel to its vorticity vector $\boldsymbol{\omega}$. It can be computed by the dot product

$$\mathcal{H} = \mathbf{u} \cdot \boldsymbol{\omega} \quad (2.46)$$

- Magnetic Reynolds number

- Rossby number

- Elsasser number

- Temporal average

3 Numerical Method

4 State of research

On influence of Prandtl number:

- Zhang Spiralling columnar convection in rapidly rotating spherical fluid shells.

Ont he collection of magnetic field lines by cyclones:

- Jones: Planetary magnetic fields and fluid dynamos

The presence of a magnetic field has an effect on the flow field in a self-sustained dynamo via the Lorentz force term in equation 2.44. Nevertheless, it is widely assumed that convection columns elongated parallel to the rotation axis persist, even if the influence of the Lorentz force term in Equation 2.44 is comparable to or higher than the Coriolis force term (Christensen, 2011).

5 Outline of the study

Describe the parameters used and show earth values

- The Prandtl numbers will be held fixed at $\text{Pr}_T = 3$ and $\text{Pr}_C = 0.3$ during this
- uniform refers to $q=0$

6 Thermo-chemical core convection under the influence of the CMB heat flux pattern

In this section, some typical features of non magnetic convection in a rotating spherical shell that is heterogeneously cooled from above are described based on a few showcases. Therefore, the focus lies on the heat flux pattern at the CMB and thus on the thermal driving, not on the thermo-*chemical* nature of the model. Nevertheless, a thermo-chemical case will play a role in section 6.5. The Prandtl numbers will be held fixed at $\text{Pr}_T = 3$ and $\text{Pr}_C = 0.3$, the Ekman number as well at $\text{Ek} = 10^{-4}$, if not marked differently.

The influence of the new boundary conditions can best be observed in the long-term averaged flow (Olson and Christensen, 2002). Consequently, all data that is shown in this section refers to such averages, if not marked differently. The length of the data sets used is at least three viscous diffusion times.

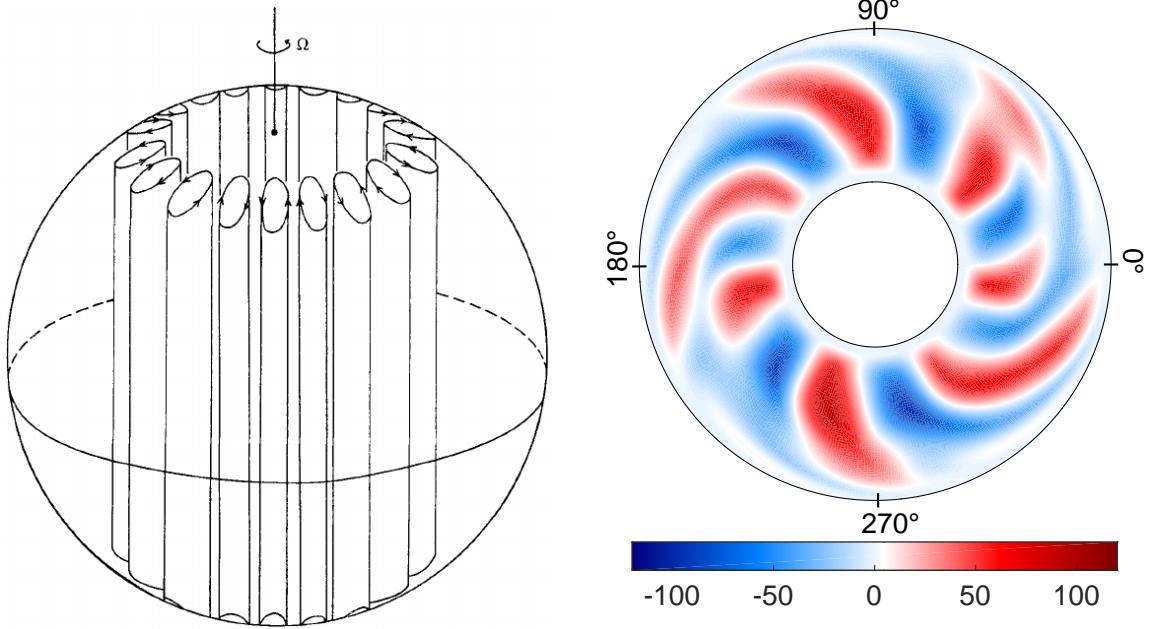


Figure 6.1: Sketch of the onset of convection for a rotating spherical shell. Convection columns form adjacent to the tangential cylinder (Busse, 1970). The columns can be visualized by different quantities. Here, they are sketched as contours of constant vorticity.

Figure 6.2: Snapshot of the radial velocity in the equatorial plain (view from the north pole, like all other equatorial views in the following) for a uniform case with $\text{Ra}_T = 4.4 \cdot 10^6$, $\text{Pr}_T = 0.3$, $\text{Ek} = 10^{-4}$ and $q^* = 0$. The convection columns are tilted and slightly deformed compared to Figure 6.1.

6.1 The onset of convection in the uniform case

In a rotating system that is homogeneously forced, convection sets in in the form of a thermal Rossby wave (Busse, 1970, 2002). The *Taylor-Proudman theorem* (see section 6.2) imposes a column like structure to the system that is aligned with the axis of rotation (z -axis, see Figure 6.1) and drifts in azimuthal direction. Busse (2002) describes a mean flow instability that leads to differential rotation (the variation of azimuthal velocity with radius), evolving under the influence of a spherical geometry. Differential rotation presumably plays a central role in the magnetic field generation as it is an ingredient for the α -effect (Roberts, 2007).

According to the analytical theories of Jones et al. (2000) and Dormy et al. (2004), the azimuthal length scale of each convection column would be extremely small in case of the Earth's rotation rate. Therefore, this scenario is assumed to be unrealistic and the dominance of the relative influence of rotation compared to the one of buoyancy is questioned. The alternative is a rather three-dimensional, turbulent regime (King et al., 2009), for which, on the other hand, the mechanism of the magnetic field production is an open question.

6.2 The geostrophic and the thermo-chemical wind balance

The rotation rate of the LOC, measured in terms of the inverse of the non dimensional Ekman number Ek^{-1} , is very high (see section 5). The *Rossby number* $Ro = \frac{u}{\Omega L}$ quantifies the relative importance of the inertia force compared to the Coriolis force in the equation of momentum. If it is small, a system is predominantly rotationally constrained and this suggests simplifications, accordingly. The inertia term drops out and viscosity can be neglected as well due to $\text{Ek} = \frac{\nu}{\Omega D^2} \ll 1$. In the absence of gravitational instabilities, this yields the *geostrophic balance*

$$-\frac{2}{\text{Ek}} \mathbf{e}_z \times \mathbf{u} = \nabla \pi, \quad (6.1)$$

where the rotational force is balanced by the pressure gradient (Vallis, 2006; Jones, 2007). The *Taylor-Proudman theorem* follows from (6.1), if the curl is applied:

$$(\mathbf{e}_z \cdot \nabla) \mathbf{u} = 0. \quad (6.2)$$

It states that variations in the velocity field parallel to the axis of rotation (z -axis) are suppressed. The system tends to take on two dimensional solution, although this is never realizable in a sphere which naturally has sloping boundaries (Busse, 2002). Relation (6.1) slightly changes, if buoyancy forces due to temperature or compositional variations cannot be neglected. The *thermo-chemical wind balance* then describes, how lateral gradients in T and C cause flow gradients parallel to the axis of rotation (Jones, 2007):

$$-2 \frac{\partial \mathbf{u}}{\partial z} = (1 - a) \text{Ek} \nabla \times (\text{Ra}_T(\theta + t_2^2) + \text{Ra}_C \zeta) \mathbf{r}. \quad (6.3)$$

It can be obtained, when adding the buoyancy term from equation (2.44) to the force balance (6.1) and taking the curl again. Equation (6.3) will be important in the following, because the conductive response to the heat flux pattern, t_2^2 , prescribes lateral temperature gradients to the system. Additionally, dynamical variations θ and ζ , especially along the polar coordinate ϑ , are assumed to drive thermo-chemical winds in the sphere (Trümper et al., 2012).

6.3 The formation of radial flux patches by azimuthal temperature gradients in the equatorial plain

The non axisymmetric distribution of up- and downwellings in the time-averaged flow is a feature that is distinct to non homogeneous forcing provided by the heterogeneous core-mantle heat flux. Zhang and Gubbins (1992) and Dietrich et al. (2016) analytically show that the azimuthal position of in- and outward radial motion in the equatorial plain is determined by the azimuthal temperature gradient $\partial T / \partial \phi$. Equation (6.3) is used as a starting point, whereupon only thermal forcing is regarded ($\text{Ra}_C = 0$). Since the z -axis plays a central role in the geostrophic context, cylindrical

coordinates are used with s as a radial distance from the axis of rotation. According to Dietrich et al. (2016), the velocity field is split into a geostrophic part \mathbf{u}^g and an ageostrophic part \mathbf{u}^a with $\mathbf{u} = \mathbf{u}^g(s, \phi) + \mathbf{u}^a(s, z, \phi)$, where the geostrophic part per definition is independent of the z -position. To compute the geostrophic flow, the z -average of the z -component of (6.3) is needed,

$$-2 \left\langle \frac{\partial u_z}{\partial z} \right\rangle_z = (1 - a) \text{EkRa}_T \langle \mathbf{e}_z \cdot \nabla \times (\theta + t_2^2) \mathbf{r} \rangle_z, \quad (6.4)$$

where the average of a function f can be computed by

$$\langle f \rangle_z(s, \phi) = \frac{1}{2H} \int_{-H}^{+H} f(s, z, \phi) dz,$$

with $H = \sqrt{R_o^2 + s^2}$ as the half height of a cylinder at the radial position s (Gillet and Jones, 2006). The right-hand-side (rhs) of (6.4) can be transformed to (now in spherical coordinates)

$$\text{rhs} = (1 - a) \text{EkRa}_T \left\langle \mathbf{e}_z \cdot \left(\frac{1}{\sin \vartheta} \frac{\partial}{\partial \phi} (\theta + t_2^2) \mathbf{e}_\vartheta - \frac{\partial}{\partial \vartheta} (\theta + t_2^2) \mathbf{e}_\phi \right) \right\rangle_z \quad (6.5)$$

The conductive temperature profile t_2^2 is symmetric with respect to the equator, hence its lateral gradient antisymmetric and it drops out when averaged over the z -axis. The temperature variation θ is likely to be distributed equatorial symmetric as well and therefore it is neglected, likewise. In cylindrical coordinates, (6.5) now reads

$$\text{rhs} = (1 - a) \text{EkRa}_T \frac{1}{s} \left\langle \frac{\partial}{\partial \phi} (\theta + t_2^2) \mathbf{e}_z \cdot \mathbf{e}_\vartheta \right\rangle_z. \quad (6.6)$$

Following Gillet and Jones (2006), the left-hand-side of (6.4) transforms to

$$-2 \left\langle \frac{\partial u_z}{\partial z} \right\rangle_z = -\frac{1}{H} \int_{-H}^{+H} \frac{\partial u_z}{\partial z} dz = - \left(-u_s^g \frac{2s}{H^2} \right). \quad (6.7)$$

The slope of the spherical boundary is responsible for the fact, that a radial velocity component enters the z -averaged equation in order to balance the boundary condition.

In the equatorial plain, $\mathbf{e}_z \cdot \mathbf{e}_\vartheta = -1$, $r = s$ and the radial velocity u_s^g in cylindrical coordinates equals the radial velocity u_r^g in spherical coordinates. It can now be expressed in terms of the azimuthal temperature gradient:

$$u_r^g = -\frac{H^2}{2r^2} (1 - a) \text{EkRa}_T \left\langle \frac{\partial}{\partial \phi} (\theta + t_2^2) \right\rangle_z. \quad (6.8)$$

The position of the maximum (minimum) azimuthal temperature gradient should thus coincide with radial downwellings (upwellings) (see relation (6.8) that contains a minus sign) in the equatorial plain. Since only the z -averaged / geostrophic part

of the velocity field \mathbf{u} contains radial components ((Dietrich et al., 2016), u_r^g is the full radial flow in this approximation.

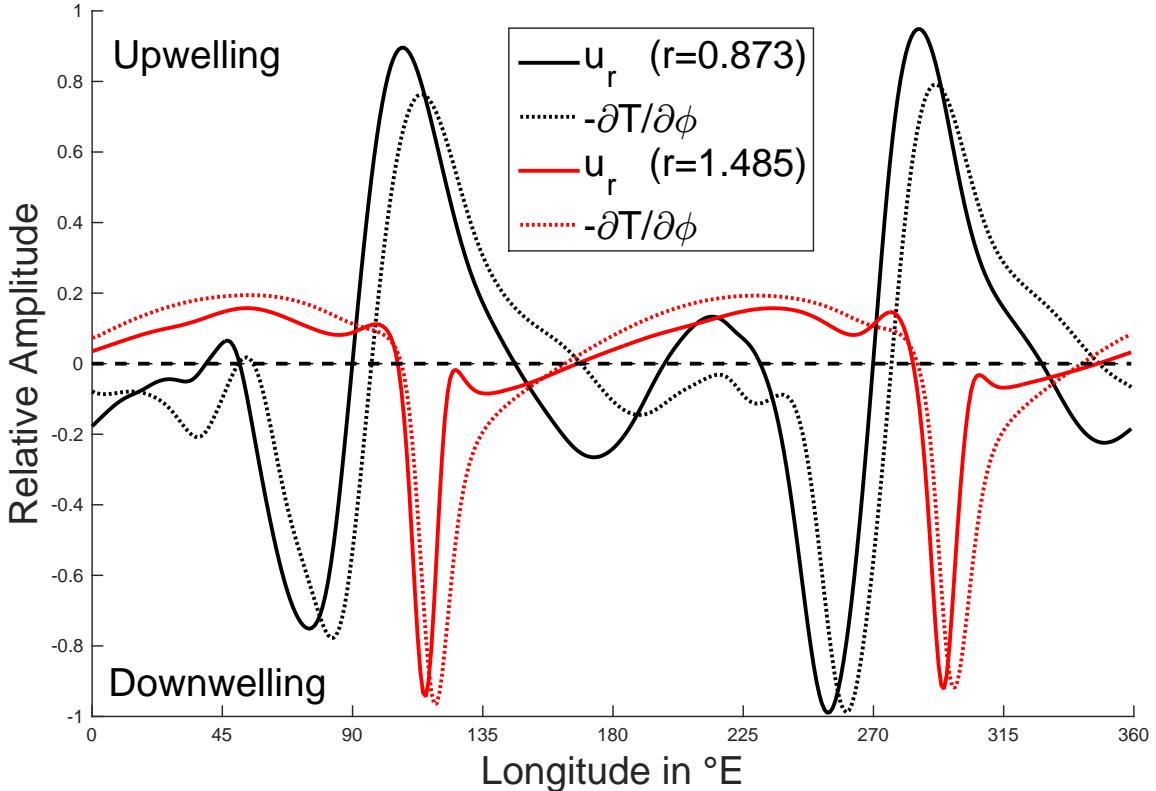


Figure 6.3: Azimuthal profiles of the radial velocity u_r (solid lines) and the negative azimuthal temperature gradient $-\partial(\theta + t_2^2)/\partial\phi$, here $-\partial T/\partial\phi$ (dotted lines), at mid depth (black, $r = 0.873$) and at the top of the free stream near the CMB (red, $r = 1.485$), normalized with respect to their individual maximum value. All values are measured in the equatorial plain. As parameters, $\text{Ra}_T = 2 \cdot 10^6$ and $q^* = 2$ are used besides the Prandtl and Ekman number as mentioned above.

Figure 6.3 shows two azimuthal profiles of the the radial velocity u_r (solid lines) and the negative azimuthal temperature gradient $\partial(\theta + t_2^2)/\partial\phi$ (dotted lines) in the equatorial plain. One is taken at mid depth (black lines), the other one on top of the free stream near the CMB (red lines). All values are normalized with respect to their maxima in order to make them comparable in one plot.

The forcing in this case is purely thermal with a Rayleigh number $\text{Ra}_T = 2 \cdot 10^6$ that is slightly overcritical: $\text{Ra}_T \approx 3 \cdot \text{Ra}_{\text{crit}}$. The critical Rayleigh number was taken from Hori et al. (2012) who computed Ra_{crit} for both, Neumann and and Derichlet type boundary conditions for the forcing component.

It becomes clear that the ϕ -gradient of the temperature serves well as an explanation for the location of the two radial downwellings near the CMB. The two negative extrema of the radial velocity coincide nearly perfectly with the maxima of

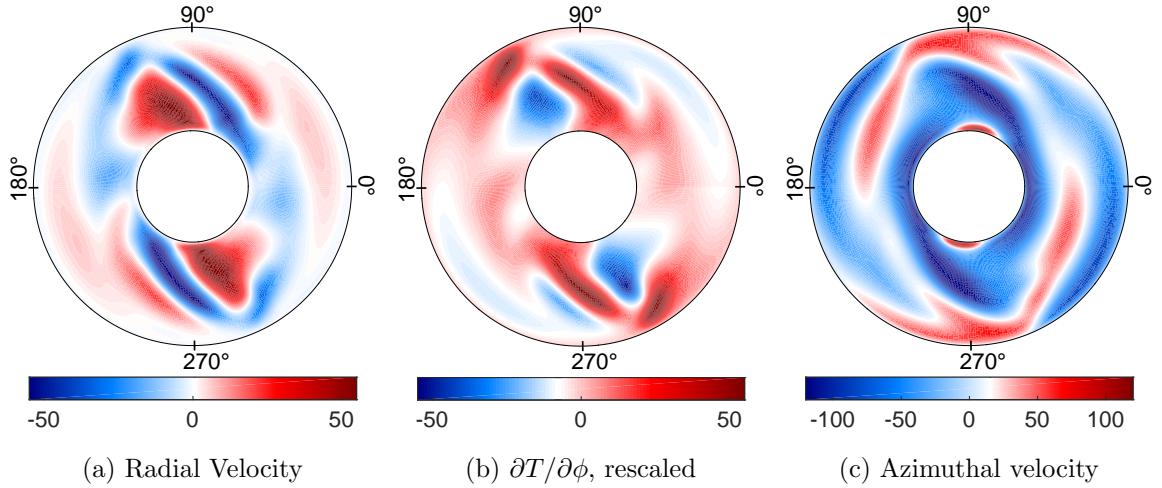


Figure 6.4: Equatorial contours of (a) the radial velocity, (b) the azimuthal temperature gradient $\partial(\theta + t_2^2)/\partial\phi$ and (c) the azimuthal velocity. The values in (b) are rescaled according to Equation (6.8) (without the minus sign) in order to be comparable to (a). The case is the same as in Figure 6.3.

the azimuthal temperature gradient. Only the small kinks in front of and behind the extrema of u_r cannot be explained by $\partial(\theta + t_2^2)/\partial\phi$ in this first order approach.

In a purely conductive regime with $\text{Ra}_T < \text{Ra}_{\text{crit}}$, the maxima of $\partial(\theta + t_2^2)/\partial\phi = \partial t_2^2/\partial\phi$ (because $\theta = 0$ in the subcritical case) would be located 45° E of the heat flux maxima (located at 90° E and 270° E, see Figure 2.3 for the conductive temperature field) (Zhang and Gubbins, 1992). The fact that the phase shift is only 30° here (see Figure 6.3) can be attributed to the supercriticality of this case. More specifically, the azimuthal velocity field is responsible for the westward shift of the downwelling because it inhibits strong retrograde flow near the CMB around 135° E (see Figure 6.4(c)).

In the upper region of the core, the two downwellings (around 120° and 300° E) are more pronounced than the complementary upwellings (around 45° and 225° E, see Figures 6.3 and 6.4(a)). Dietrich et al. (2016) explain this asymmetry with the azimuthal part of the non linear temperature advection term $-\frac{\partial\theta}{\partial\phi}u_\phi$ (see Equation (2.43a)) that promotes positive azimuthal temperature gradients and therefore downwellings at the corresponding longitudes.

The radial flow in the deeper regions of the core (e.g. the solid black line Figure 6.3) has a more complex form and is thus not as directly as in the upper layers related to the \mathcal{Y}_2^2 heat flux pattern. Broad and pronounced upwellings develop right under the CMB downwellings at $\sim 110^\circ\text{E}$ and $\sim 290^\circ\text{E}$. They are also visible in the temperature field in Figure 6.5 in form of hot plumes.

The downwellings observed at the CMB bend in jet like structures in retrograde direction towards the ICB (see Figure 6.4(a)). The bending happens due to strong zonal westward flows at mid depth at $\sim 90^\circ\text{E}$ and $\sim 270^\circ\text{E}$, respectively (see Figure

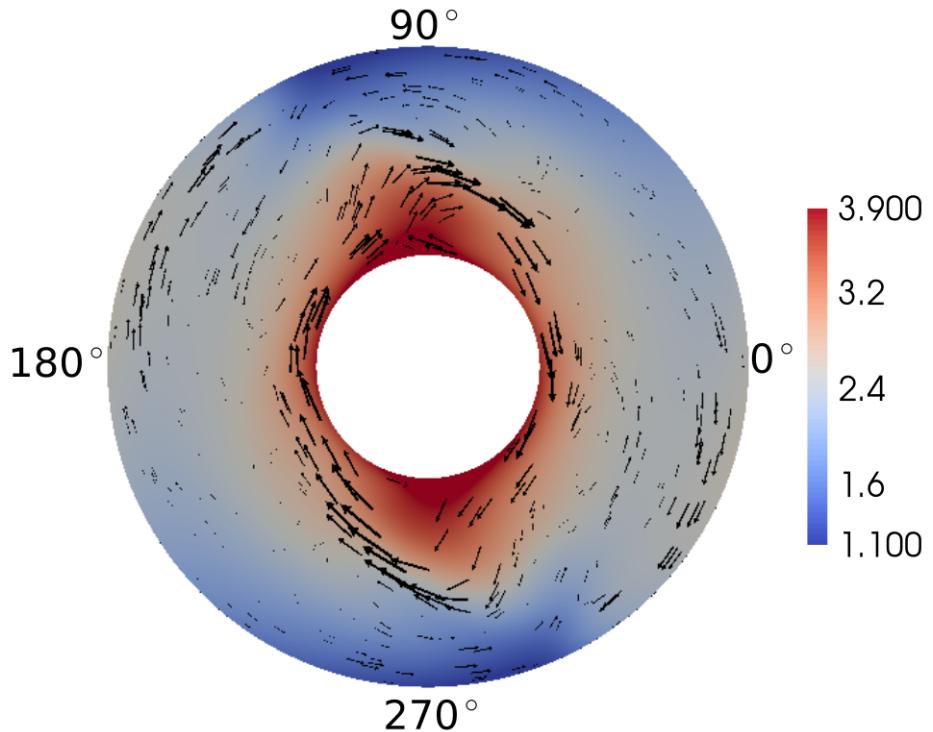


Figure 6.5: Horizontal velocity vectors in the equatorial plain, scaled by the corresponding flow magnitude. The background contour shows the full non dimensional temperature field $T = \theta + T_{\text{cond}}$. The two strong upwelling plumes at the ICB (at $\sim 110^\circ\text{E}$ and $\sim 290^\circ\text{E}$) and the bending jet like downwellings right above are clearly visible.

6.4(c)).

Although the flow structure in the deeper core regions is more complex than at the CMB, equation (6.8) and therefore the thermo-chemical wind relation provide a good explanation for the radial flow field (compare Figure 6.4(a) and (b)).

The small phase shift between $\partial T / \partial \phi$ and u_r , observable in Figure 6.3 (dotted and solid lines), illustrates that the assumption of negligible viscosity (see section 6.2) is not fully justified in case of $\text{Ek} = 10^{-4}$. Zhang and Gubbins (1992) state that this shift vanishes if the rotation rate is further increased and therefore the influence of the viscosity decreased.

The azimuthal flow in the equatorial plain is, in reaction to the radial flow, convergent (divergent) in the neighborhood of downwellings (upwellings). This can be observed either in the full horizontal flow field (Figure 6.5) or in the radial and azimuthal flow fields (Figure 6.4(a) and (c)). The downwellings are each 'fed' from the lateral direction and vice versa.

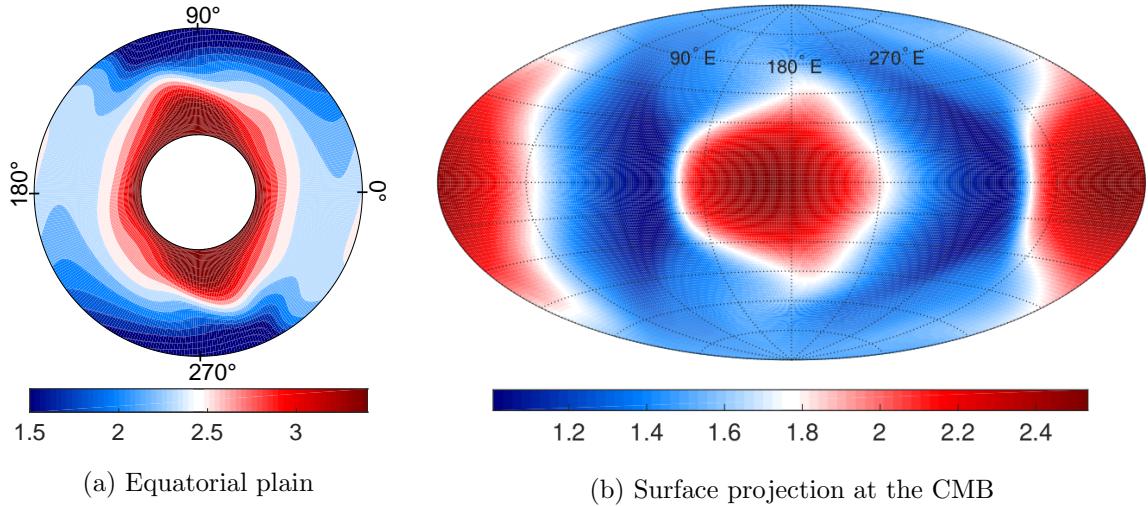


Figure 6.6: Full temperature field $T = \theta + T_{\text{cond}}$ for the same case as above in (a) the equatorial plain and (b) as a surface projection right under the CMB. (a) is the same graph as in Figure 6.5, but the color map is 'shortened' in order to pronounce the phase shift of the temperature field.

6.4 The effect of the flux pattern on the temperature field

Figure 6.6(a) shows the same equatorial view on the temperature field as Figure 6.5 in the previous section. The color map was slightly rescaled to make the phase shift between the temperature maxima at the CMB at $\sim 170^\circ$ and $\sim 350^\circ$ and the heat flux maxima (at $\sim 90^\circ$ and $\sim 270^\circ$) visible. The reason for the phase shift lies in the flow field and can be found in the previous chapter: Strong radial motion occurs predominantly in the region under the heat flux maxima. The up- and downwellings visible in Figure 6.5 intensify the cooling mechanism, thus go along with a colder core region. The hot regions coincide with strong azimuthal motion along the CMB that inhibit the transport of heat out of the core. (EKMAN dependence ???) The surface projection of the temperature field at the CMB (Figure 6.6(b)) confirms that the temperature field in the slightly super critical case nearly is a phase shifted version of the conductive temperature field (Figure 2.3(b)) (Zhang and Gubbins, 1992).

6.5 Stationary vortex columns attached to mantle heterogeneities

As mentioned above, convection in rapidly rotating systems is organized in a column like structure at onset (Busse, 1970). The question whether this structure is maintained, even in the highly super critical LOC, has important implications for the understanding of the geodynamo. The so called *vortex columns*, visualized in Figures 6.1 and 6.8, play a key role in the magnetic field production.

Although motion parallel to the z -axis is strongly constrained by rotation (see Equation (6.2)), the spherical form of the boundaries necessarily relaxes this constraint

locally and vertical flow is unavoidable (Busse, 2002; Gillet and Jones, 2006). An important example for the relaxation of the Taylor-Proudman theorem in this context is the *secondary circulation* inside of vortex columns (Olson et al., 1999; Aubert et al., 2008b; Christensen, 2011), sketched in Figure 6.7.

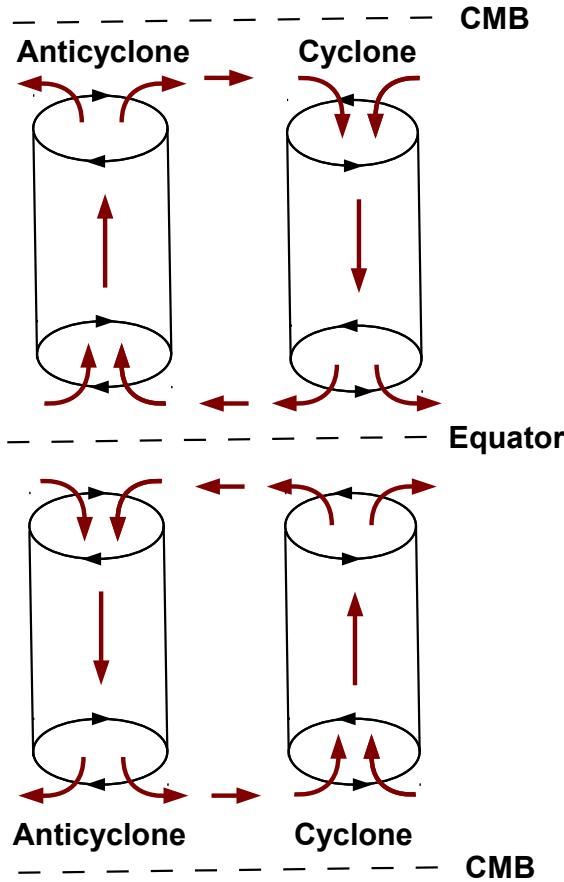


Figure 6.7: Sketch of the secondary circulation induced by a pair of a cyclonic and an anticyclonic vortex. The two convection columns can be imagined as part of the sketch in Figure 6.1, a situation typical for the onset of convection.

Figure 6.7 shows how a cyclone causes a divergent flow at the equator, whereas an anticyclone acts converging at the equator. The opposite is true near the CMB. A cyclone and an anticyclone together, they commonly appear in pairs, form a *circular* motion that superimposes the primary vortial motion. The interaction of a vorticity carrying fluid column and a vertical flow produces helicity \mathcal{H} . Besides the differential rotation that is necessary for the ω -effect, the presence of helicity is a condition for the occurrence of the (macroscopic) α -effect (Roberts, 2007; Christensen, 2011).

Ekman pumping and *suction* are boundary layer effects that are distinct to systems with no-slip conditions (as in this study, see section 2.7). They evolve in the interplay between the rotation dominated bulk and the viscosity dominated Ekman layer. In *cyclones*, i.e. vortex columns with positive vorticity $\omega > 0$, Ekman pumping induces a vertical flux away from the CMB towards the equator. Ekman suction causes motion towards the CMB in anticyclones, where $\omega < 0$. Although both mechanisms take place in the boundary layers, they potentially induce vertical motion in the whole column. A more detailed description can be found in Pedlosky (2013).

The variation of buoyancy along the z -axis inside of a vortex column acts as a second source of vertical motion (Olson et al., 1999). As in the case of Ekman suction and pumping, this motion is oriented equatorwards in cyclones and away from the equator in anticyclones. Yet, it is independent of the viscous boundary condition and therefore it makes this discussion applicable to a wider range of model configurations.

The interplay of the azimuthal and the radial velocity field in the equatorial plain (Figure 6.4(a) and (b)) causes a vorticity distribution that is quite different from the one in the uniform case, depicted in Figure 6.8 in form of a snapshot.

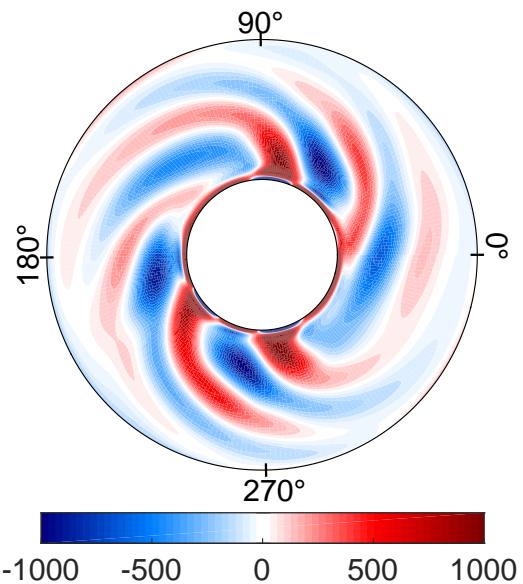


Figure 6.8: Snapshot of the z -component of vorticity in the equatorial plain for a slightly super critical, uniform case: $\text{Ra}_T = 10^6$.

6.6 kinetic energies, wave numbers..

kinetic energy increases with increasing q^*

7 Thermo-chemical dynamo action

7.1 The influence of chemical forcing on magnetic field properties

8 Summary

9 Conclusion and Outlook

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