



Single-seismometer Focal Mechanism and Uncertainty Estimation from Body Waves



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Introduction & Background

In a study done by Sita and van der Lee (2022), waveform data from NASA's InSight mission is used to determine the focal mechanisms of marsquakes. P- and S-wave displacement amplitudes are collected into a 3D vector

$$\mathbf{A} = (A^P, A^{SV}, A^{SH})^t$$

with corresponding noise levels estimated from seismograms recorded by a single station.

Assuming a double-couple source, the relationship between focal mechanism and observed amplitudes can be modeled using far-field body wave expressions. However, in data-sparse environments—such as Mars, where we often have only one seismometer and limited subsurface information—**relative amplitudes** become more practical to use than absolute ones.

Building on this foundation, our research introduces a mathematically robust and computationally efficient method for estimating source parameters (strike, dip, and rake). Our approach accounts for **asymmetric noise levels** across components and identifies all mechanisms that fit the data well.

Mathematical Model

Given a ray path's azimuth ϕ and take-off angles i, j of P and S waves respectively, we model the amplitudes as functions of the source mechanism's strike ψ , dip δ and rake λ , further accounting for the P and S velocities α_h and β_h at source depth h . The expressions are

$$\begin{aligned} A^P &\sim \frac{s_R(3\cos(i) - 1) - q_R \sin(2i) - p_R \sin^2(i)}{\alpha_h^3} \\ A^{SV} &\sim \frac{1.5s_R \sin(2j) + q_R \cos(2j) + 0.5p_R \sin(2j)}{\beta_h^3} \\ A^{SH} &\sim \frac{q_L \cos(j) - p_L \sin(j)}{\beta_h^3} \end{aligned}$$

$$s_R = 0.5 \sin(\lambda) \sin(2\delta)$$

$$q_R = \sin(\lambda) \cos(2\delta) \sin(\psi_r) + \cos(\lambda) \cos(\delta) \cos(\psi_r)$$

$$p_R = \cos(\lambda) \sin(\delta) \sin(2\psi_r) - 0.5 \sin(\lambda) \sin(2\delta) \cos(2\psi_r)$$

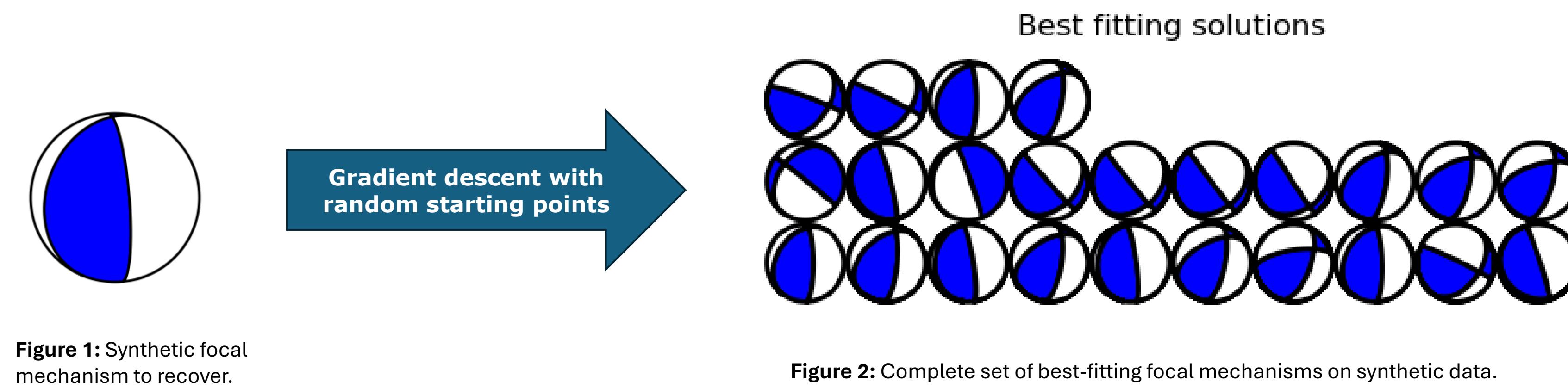
$$p_L = 0.5 \sin(\lambda) \sin(2\delta) \sin(2\psi_r) + \cos(\lambda) \sin(\delta) \cos(2\psi_r)$$

$$q_L = -\cos(\lambda) \cos(\delta) \sin(\psi_r) + \sin(\lambda) \cos(2\delta) \cos(\psi_r)$$

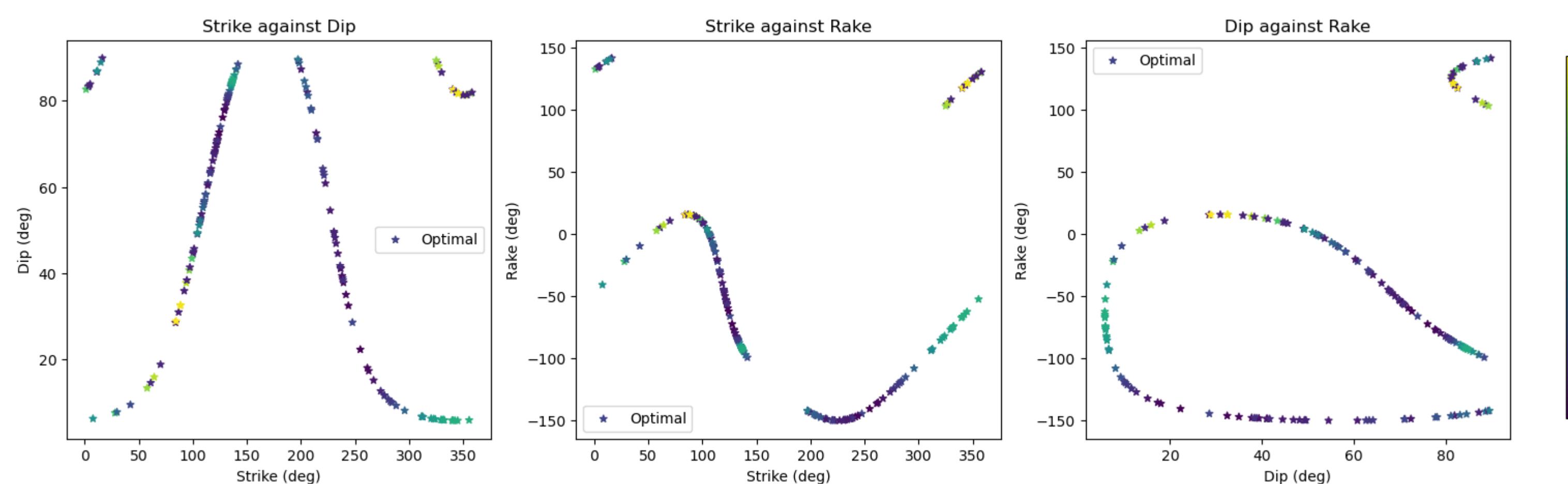
We use **cosine similarity** between observed and synthetic amplitude vectors as the misfit function.

$$\Phi(\mathbf{m}) = \frac{-\mathbf{A}_o^t \mathbf{A}_s}{\|\mathbf{A}_o\| \|\mathbf{A}_s\|}$$

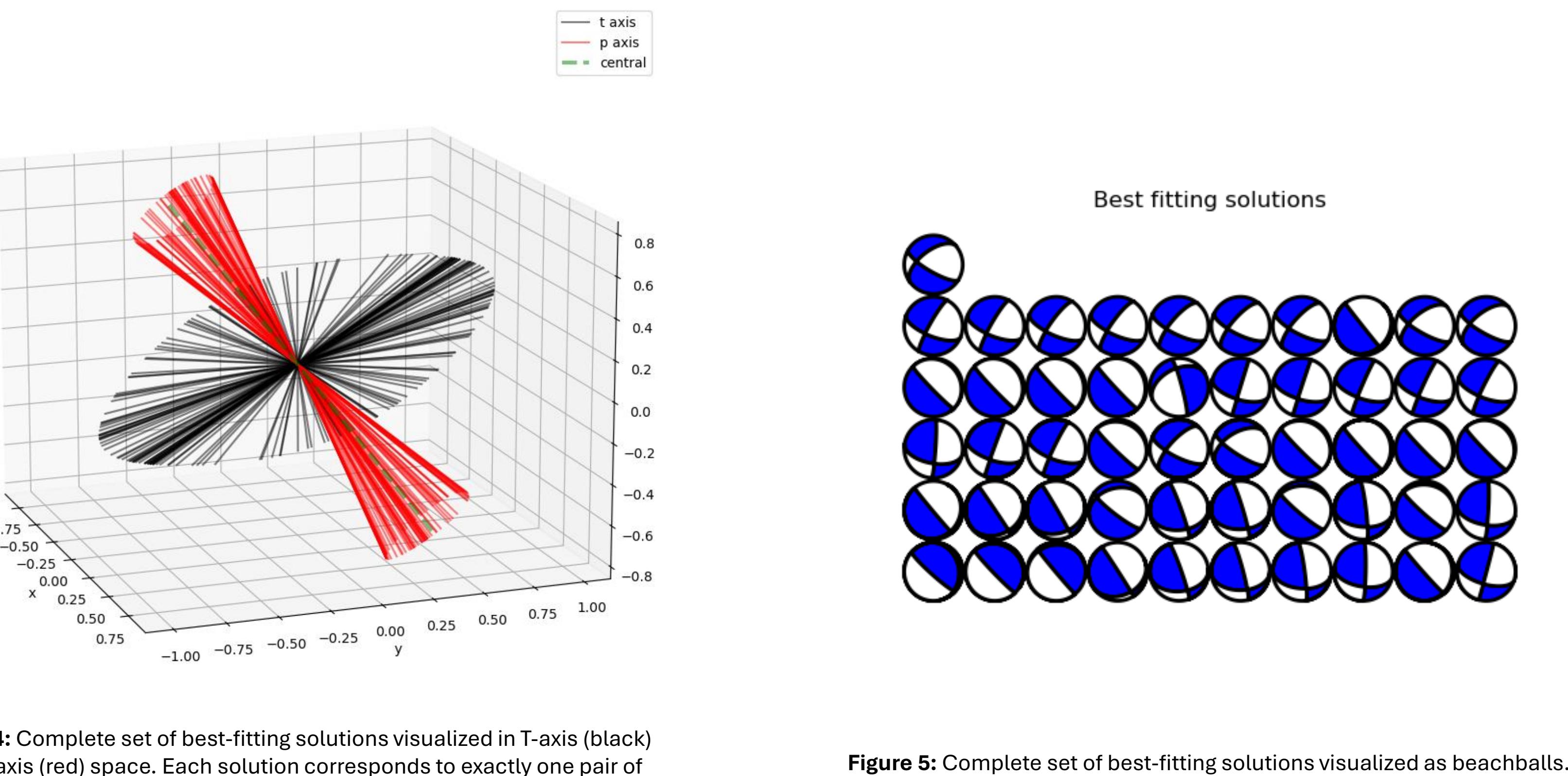
Visualizations



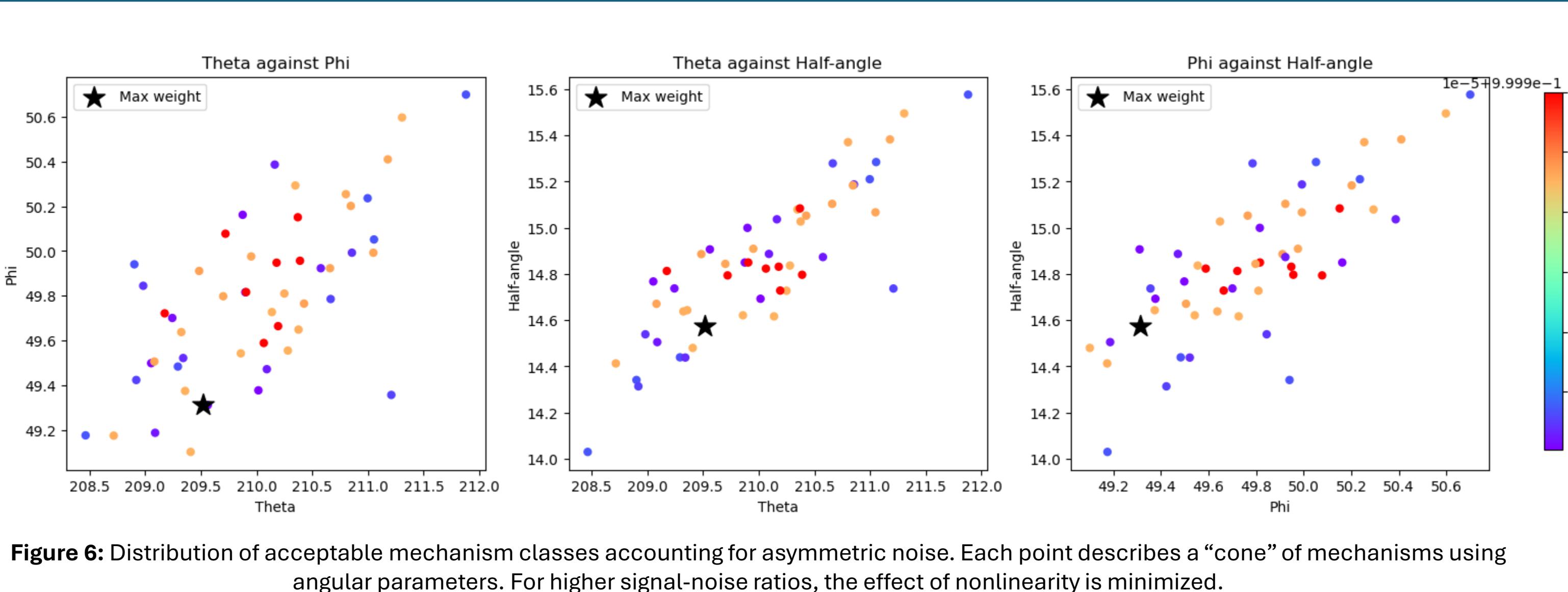
Collection 1: Using algorithmic implementation on synthetic data.



Collection 2: Using algorithmic approach on earthquake data from MBAR, Uganda.



Collection 3: Using Monte Carlo experiment to observe distribution of all classes of mechanisms that produce acceptable amplitude vectors.



Algorithmic Approach

For a highly nonlinear multivariable misfit function whose landscape is difficult to visualize directly in 4D space, we use an algorithmic approach to parameter estimation.

Using a computational graph, we can calculate gradients by backpropagation. This makes **gradient descent** effective for locating local minima. By **random sampling** of starting points, we can trace out classes of best fitting solutions.

We found that the best-fitting mechanisms usually have a stress axis sweeping out a narrow cone. We can classify mechanisms using a three angles to describe the cone, an indicator of the axis.

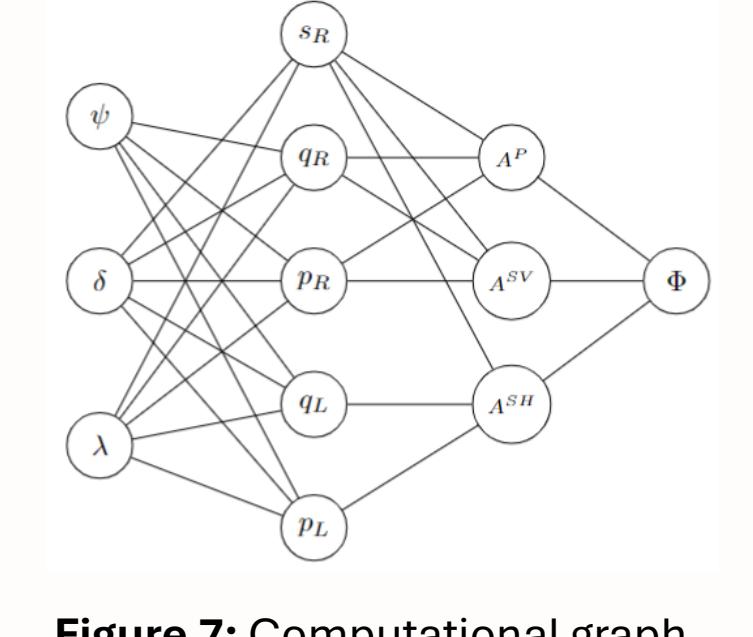
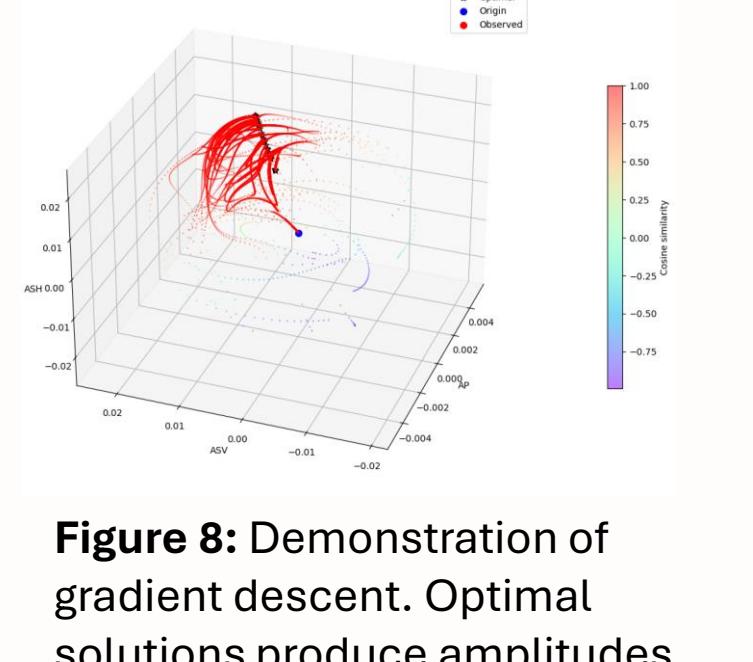


Figure 7: Computational graph for gradient calculations.



Uncertainty Measurement

Using **Monte Carlo methods**, we can investigate the distribution of mechanism classes which produce acceptable amplitudes.

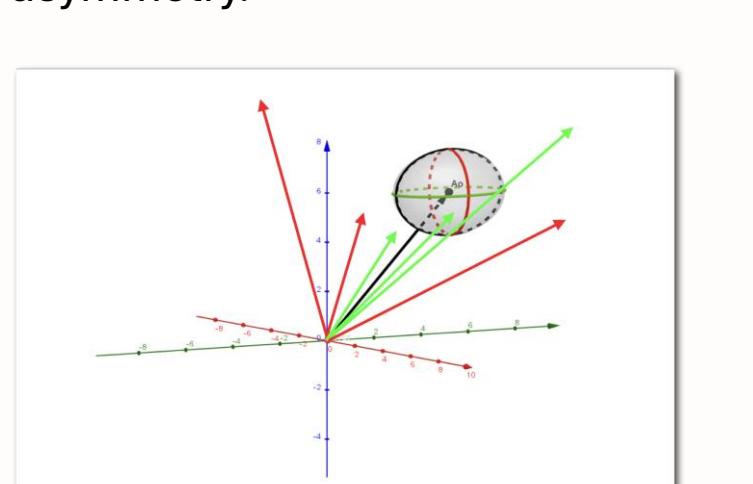
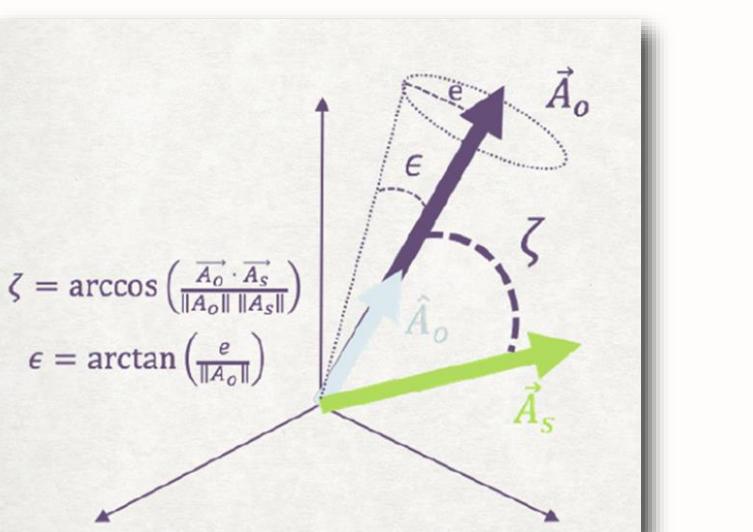
An amplitude vector \mathbf{A}_s is acceptable if it is parallel to a vector in the error ellipsoid surrounding the observed vector \mathbf{A}_o .

Given noise matrix Σ , the vector \mathbf{A}_s must lie on the line segment between boundary points \mathbf{A}_b defined by

$$\mathbf{A}_b = \left(1 - \frac{1}{C_{oo}}\right) \mathbf{A}_o \pm \left(\frac{1}{C_{oo}} \sqrt{\frac{C_{oo} - 1}{C_{oo} C_{ss} - C_{os}^2}}\right) (C_{oo} \mathbf{A}_s - C_{os} \mathbf{A}_o)$$

Constants of the form C_{xy} are matrix inner products given by

$$C_{xy} = \langle \mathbf{A}_x, \mathbf{A}_y \rangle_{\Sigma^{-1}} = \mathbf{A}_x^t \Sigma^{-1} \mathbf{A}_y$$



Conclusion

When using a single seismometer and in data-sparse settings, our methodology is more reliable in:

- Accuracy:** It can characterize entire classes of best-fitting mechanisms, which a grid search cannot do.
- Efficiency:** Gradient descent is much faster than brute force inverse methods, and it makes use of the gradient information available.
- Robustness:** It respects any asymmetries in measured noise levels for P and S waves without any major tradeoffs.

However, we have highly nonlinear transformations which make the method less reliable at low signal-noise ratios.