Tidal forcing of seismicity and slip, ice streams B & C

Tides Project Team

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Tides, Slip, Seismics (prior work)

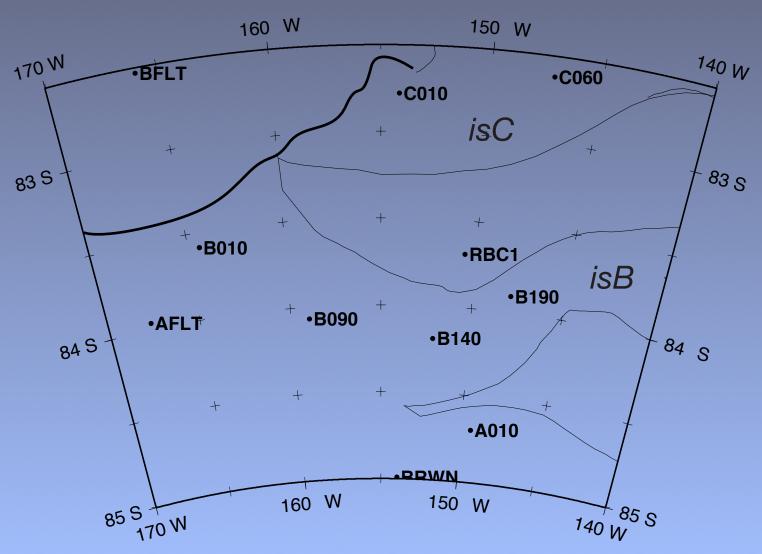
- Tides in the Ross Ice Shelf control flow of the ice plains of ice streams B, D, and E.
 - Flow is stick-slip on B, smoothly modulated on D/E
- These same tides control basal seismicity of ice stream C.
 - The seismicity was assumed a proxy for slip, lacking the ability to measure flow variations is C.
- Tide phase crucial (high vs. low)
- Tide amplitude effects unknown (spring vs. neap)

Seismicity and Slip relationship hypothesized.

- However, this relationship was not confirmed.
- Prior work on isC recorded seismicity but could not record slip (flow speeds too low)
- Prior work on D recorded slip but not seismicity (seismometers too far away, at Siple Dome and Ridge DE, not on D itself)
- Prior work on B was GPS only, so no seismometers used

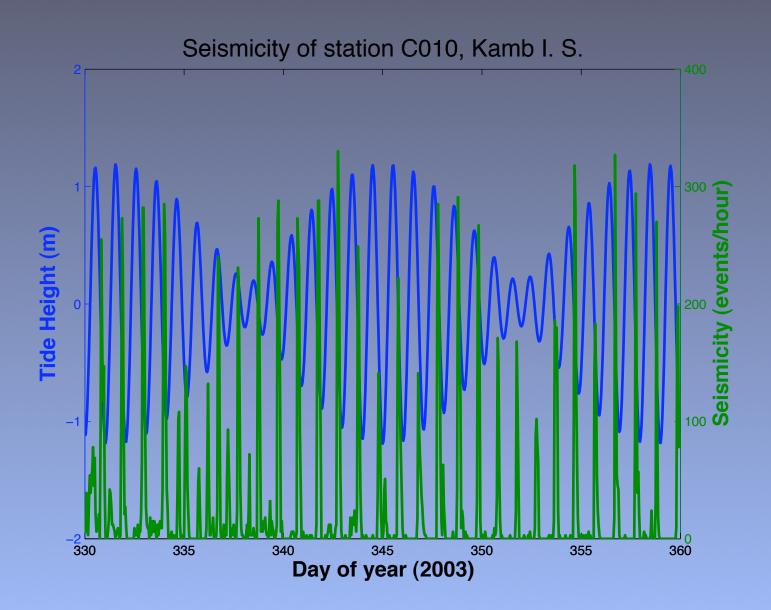
^{» (}aka, "aseismic Bob"), inside joke!

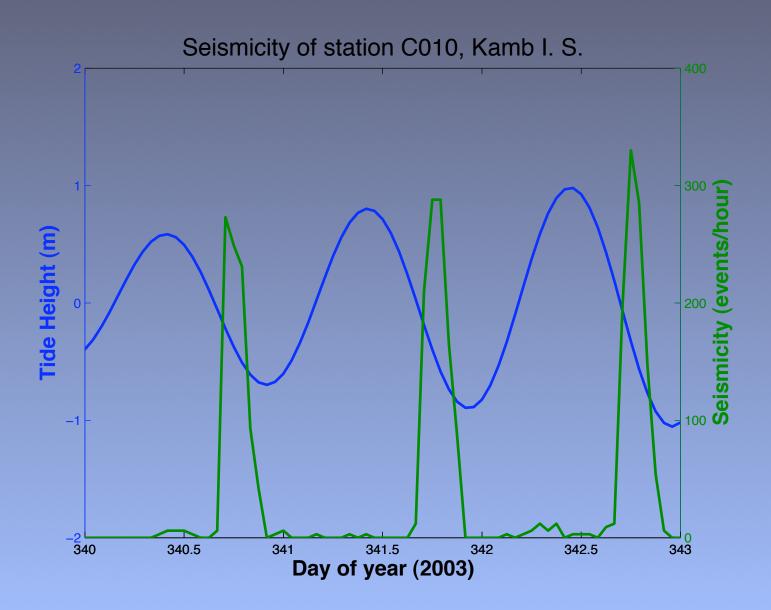
Map of study region



Combined Seismic/Slip experiment

- Stations C10, B10, B90, B140 had both seismometers and GPS receivers.
 - GPS receivers measure position every 5 min.
 - Seismometers trigger on seismic activity.
- Station BFLT measures the tide directly.





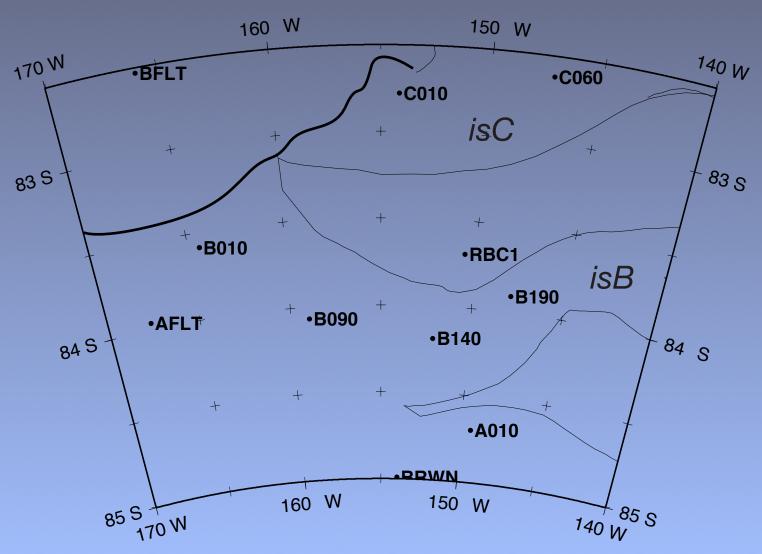
Ice stream C results compared to prior results

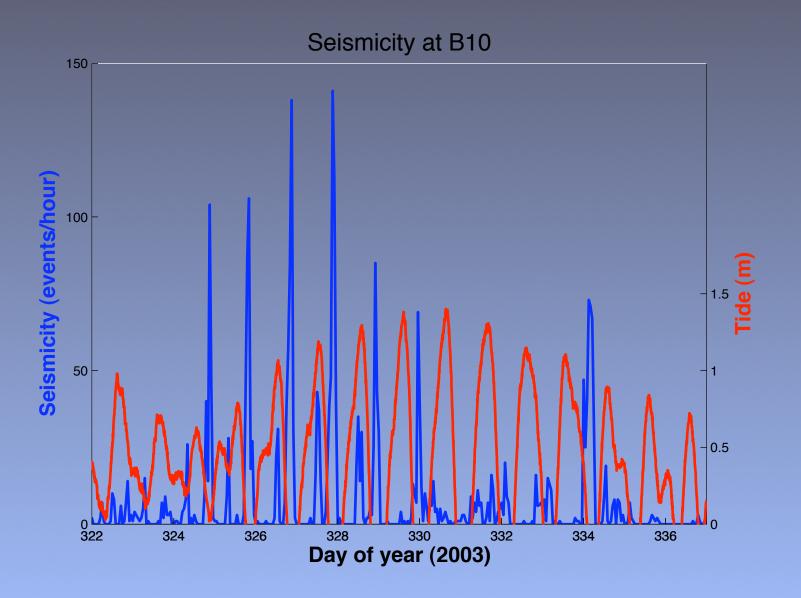
- Prior work used modeled tide and tide extrapolated from measurements taken decades earlier.
- New results agree with earlier conclusions
 - seismicity is at falling or low tide (phase matters)
 - Tide amplitude effect inconclusive (spring/neap)
 - Slip vs. seismicity still inconclusive (isC too slow)

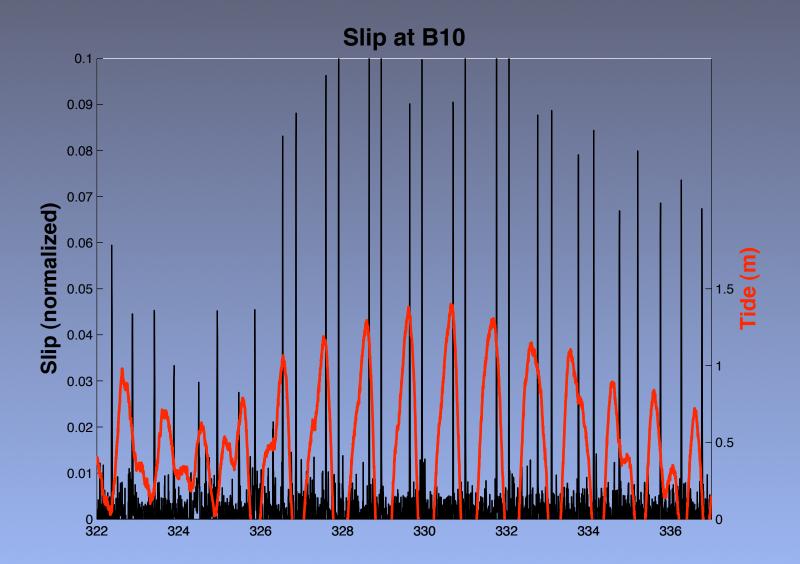
New results from isB (Whillans)

- Comparison of seismicity and slip.
- Comparison of relative importance of tide phase (high vs. low) and tide amplitude (spring vs. neap).

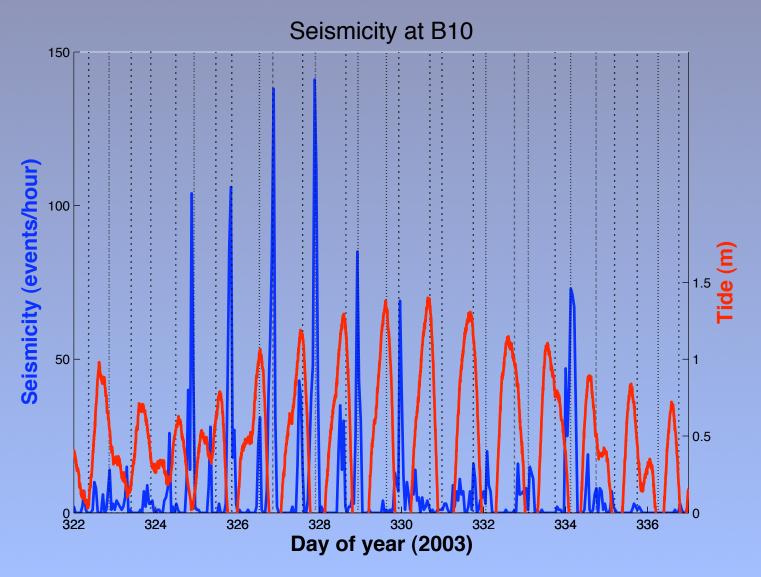
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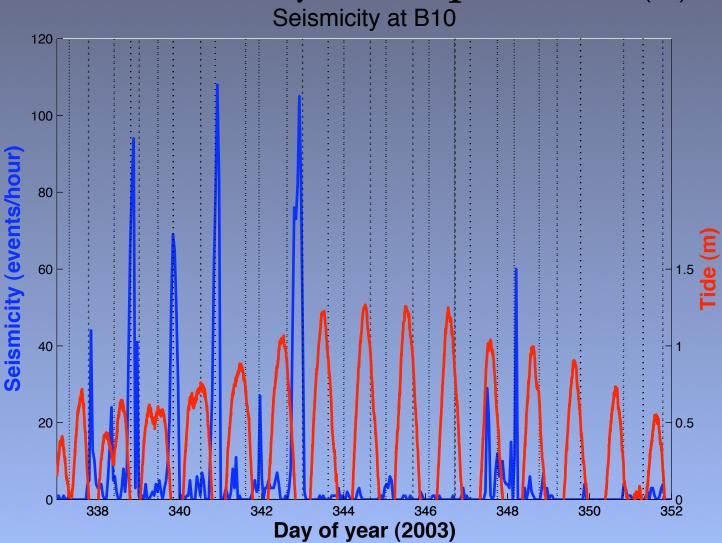




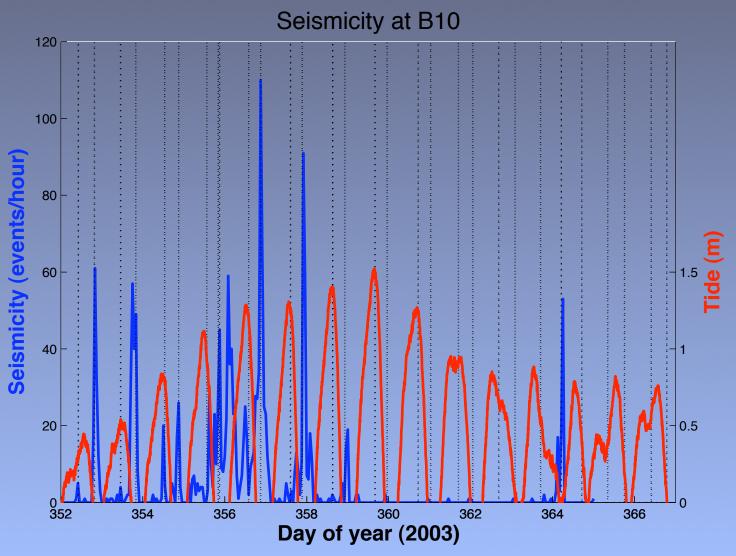
Seismicity + Slip-times (1)



Seismicity + Slip-times (2)



Seismicity + Slip-times (3)



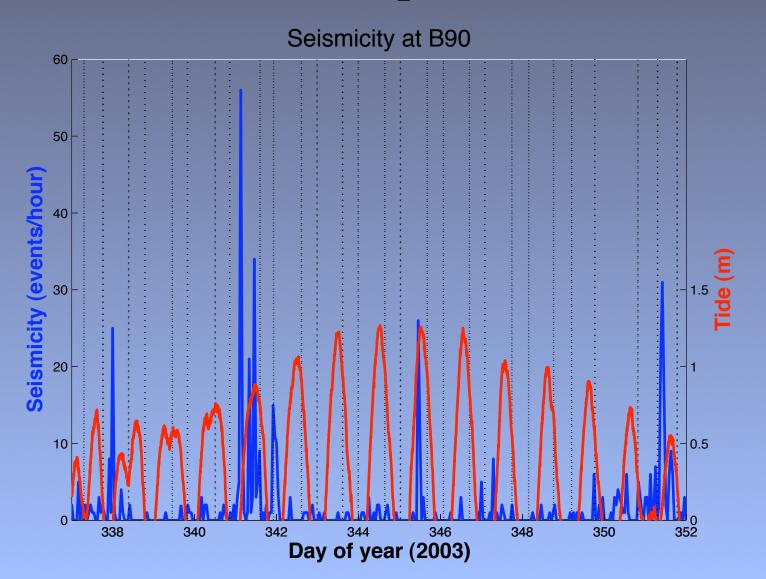
Observations

- Two slip events per tidal cycle
 - One at high tide (or soon after), one at low tide (or just before)
 - Neap tide has less-regular slip
- Seismicity is almost always associated with the 2nd slip event.
- Low (non-zero) level of seismicity throughout the day

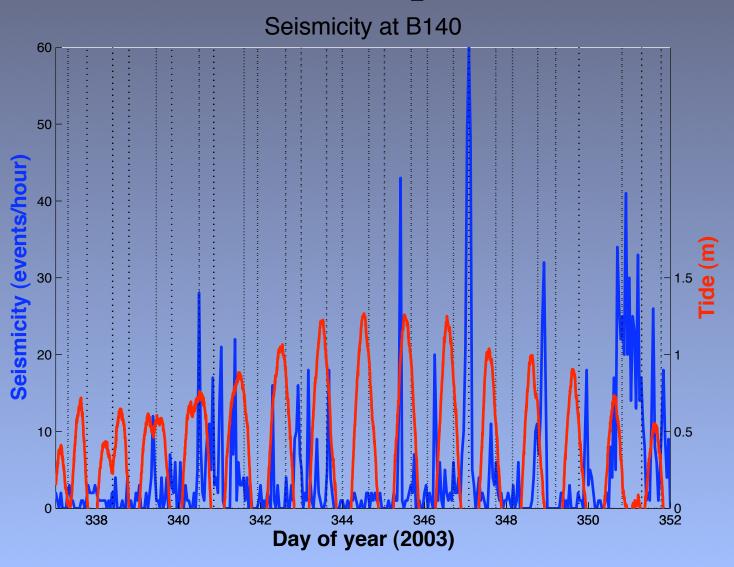
Observations (2)

- At B10 (near the grounding line of ice stream B)
 - Seismicity peaks occur during the neap to spring tide cycle
 - Seismicity is suppressed during the spring to neap tide cycle.

B90 - 90 km up from G. L.



B140 - 140 km up from G. L.



Observations (3)

- Upstream from the grounding line, the seismicity is "peaked", generally once per day, but sometimes more
 - BUT, not tied to the slip event times...
 - Not inconsistent with a model of a propagating wave from the grounding line
 - Difficult to quantify propagation speed because of difficulty correlating from station to station

Hypotheses

- The bed at the grounding line is very sensitive to the ocean tide.
 - The bed at the grounding line is weak. The strain from the slip accumulates until the bed fails.
- There is a "hysteresis" associated with the highest high tide (spring tide).
 - The bed is weakened further so that the slip of the ice stream does not cause seismic slip/failure
 - Or, water infiltrates up from the grounding line,
 allowing asiesmics sliding.

Hypotheses

- The bed is more heterogeneous upstream of the grounding line
 - More seismicity from various hypocenters
 - Continuous, low-level seismicity

Conclusions

- Basal seismicity is associated with slip at the grounding line (though not all slip is associated with seismicity)
- The tide can profoundly affect the properties of the bed.
 - Continued work will help distinguish between different hypotheses.

