Martin Truffer David Maxwell

Geophysical Institute and Dept. of Math. and Stat. University of Alaska Fairbanks

7 Sep 2007, WAIS

Outline

Introduction

Kazlov-Maz'ya (KM) iteration

Some examples

Summary and Conclusions

Introduction

Kazlov-Maz'ya (KM) iteration

Some examples

Summary and Conclusions

Transition from ice sheet to ice shelf flow

- Transition from ice sheet to ice shelf flow
- Yet another free boundary problem in glaciology (-> Schoof)

Modeling challenges

 Multimodal models (such as UAF's PISM) switch from Shallow Ice Approximation to Shallow Shelf Approximation

Modeling challenges

- Multimodal models (such as UAF's PISM) switch from Shallow Ice Approximation to Shallow Shelf Approximation
- A hard switch does not capture the thermo-mechanical peculiarities of the onset area

Modeling challenges

- Multimodal models (such as UAF's PISM) switch from Shallow Ice Approximation to Shallow Shelf Approximation
- A hard switch does not capture the thermo-mechanical peculiarities of the onset area
- This might be important when modeling the evolution of onset area

Basal velocities

 Frozen - thawed transition can possibly be derived from surface measurements (radar)

Basal velocities

- Frozen thawed transition can possibly be derived from surface measurements (radar)
- This is not directly possible for the distribution of basal velocities

Basal velocities

- Frozen thawed transition can possibly be derived from surface measurements (radar)
- This is not directly possible for the distribution of basal velocities
- Hence the need for inverse methods: derive basal velocities from observations at the surface (or from remote sensing)

Outline

Introduction

Kazlov-Maz'ya (KM) iteration

Some examples

Summary and Conclusions

The forward model

 Consider 2D flow along a longitudinal cross section (of constant thickness for a start)

- Consider 2D flow along a longitudinal cross section (of constant thickness for a start)
- Consider a first order model of flow (Blatter, 1995)

- Consider 2D flow along a longitudinal cross section (of constant thickness for a start)
- Consider a first order model of flow (Blatter, 1995)
- It can be shown (Colinge and Rappaz, 1999) that the equation for a suitably normalized longitudinal velocity reduces to $\nabla \left((\epsilon + |\nabla u|^2)^{\frac{n-1}{2}} \right) \nabla u = -\frac{1}{2}$

- Consider 2D flow along a longitudinal cross section (of constant thickness for a start)
- Consider a first order model of flow (Blatter, 1995)
- It can be shown (Colinge and Rappaz, 1999) that the equation for a suitably normalized longitudinal velocity reduces to $\nabla \left((\epsilon + |\nabla u|^2)^{\frac{n-1}{2}} \right) \nabla u = -\frac{1}{2}$
- Incidentally this is the same nonlinear Poisson equation that describes full order Stokes flow of a nonlinear fluid through a glacier's cross section with no out-of-plane gradients.

 Ice thickness, surface elevation, ice rheology, and surface velocities are assumed to be known

Inverse model

- Ice thickness, surface elevation, ice rheology, and surface velocities are assumed to be known
- Surface parallel shear stress is assumed to be zero

Inverse model

- Ice thickness, surface elevation, ice rheology, and surface velocities are assumed to be known
- Surface parallel shear stress is assumed to be zero
- This problem is ill-posed and generally no solution is guaranteed

Inverse model

- Ice thickness, surface elevation, ice rheology, and surface velocities are assumed to be known
- Surface parallel shear stress is assumed to be zero
- This problem is ill-posed and generally no solution is guaranteed
- Inverse methods seek an approximate solution

 Start with a guess for the basal velocities (e.g. zero basal motion everywhere)

- Start with a guess for the basal velocities (e.g. zero basal motion everywhere)
- calculate forward problem with that guess (Dirichlet condition) at the base and zero shear stress (Neumann condition) at the top.

- Start with a guess for the basal velocities (e.g. zero basal motion everywhere)
- calculate forward problem with that guess (Dirichlet condition) at the base and zero shear stress (Neumann condition) at the top.
- Extract basal stress from that solution

- Start with a guess for the basal velocities (e.g. zero basal motion everywhere)
- calculate forward problem with that guess (Dirichlet condition) at the base and zero shear stress (Neumann condition) at the top.
- Extract basal stress from that solution
- Solve forward problem with that basal stress and the measured surface velocities

- Start with a guess for the basal velocities (e.g. zero basal motion everywhere)
- calculate forward problem with that guess (Dirichlet condition) at the base and zero shear stress (Neumann condition) at the top.
- Extract basal stress from that solution
- Solve forward problem with that basal stress and the measured surface velocities
- Extract basal velocities from that solution

- Start with a guess for the basal velocities (e.g. zero basal motion everywhere)
- calculate forward problem with that guess (Dirichlet condition) at the base and zero shear stress (Neumann condition) at the top.
- Extract basal stress from that solution
- Solve forward problem with that basal stress and the measured surface velocities
- Extract basal velocities from that solution
- Repeat and stop when the solution has sufficiently converged.

 Advantage: Method has been proved to converge for linear problem (Kozlov and Maz'ya, 1999)

- Advantage: Method has been proved to converge for linear problem (Kozlov and Maz'ya, 1999)
- Disadvantage: It converges extremely slowly

- Advantage: Method has been proved to converge for linear problem (Kozlov and Maz'ya, 1999)
- Disadvantage: It converges extremely slowly
- We have (that is D. Maxwell has) developed an acceleration scheme, which has worked very satisfactorily for several problems.

- Advantage: Method has been proved to converge for linear problem (Kozlov and Maz'ya, 1999)
- Disadvantage: It converges extremely slowly
- We have (that is D. Maxwell has) developed an acceleration scheme, which has worked very satisfactorily for several problems.
- This has been applied for the problem of isothermal flow through a glacier cross section (Maxwell et al., submitted)

Outline

Introduction

Kazlov-Maz'ya (KM) iteration

Some examples

Summary and Conclusions

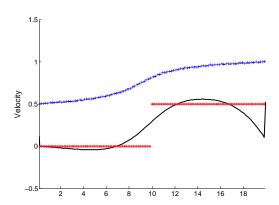
We run a forward model with a given basal velocity

- We run a forward model with a given basal velocity
- We extract the calculated surface velocity and add noise

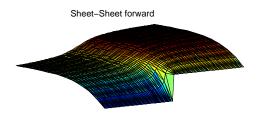
Examples with artificial data

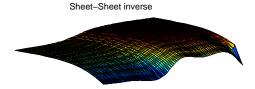
- We run a forward model with a given basal velocity
- We extract the calculated surface velocity and add noise
- We feed this data set to the inverse model and compare the resulting basal velocity to the one originally prescribed

Frozen - sliding transition

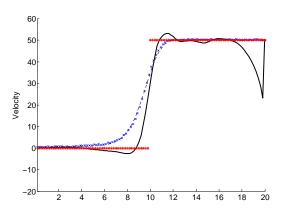


Frozen - sliding transition



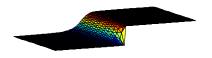


Sheet to Shelf flow transition

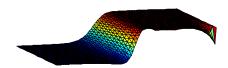


Sheet to Shelf flow transition

Sheet-Shelf forward



Sheet-Shelf inverse



Outline

Introduction

Kazlov-Maz'ya (KM) iteration

Some examples

Summary and Conclusions

Conclusions

 Accelerated KM iteration is a fast and stable method for retrieving basal velocities from surface measurements

Conclusions

- Accelerated KM iteration is a fast and stable method for retrieving basal velocities from surface measurements
- There are some fundamental limits on the spatial scale at which this can be done (one ice thickness)

Conclusions

- Accelerated KM iteration is a fast and stable method for retrieving basal velocities from surface measurements
- There are some fundamental limits on the spatial scale at which this can be done (one ice thickness)
- The method as shown assumes that rheological parameters are well known. That is a huge assumption.

Acknowledgements

NSF for funding

- NSF for funding
- You for listening