

Surface and basal crevassing of the Larsen C Ice Shelf: Implications for ice shelf stability

Daniel McGrath¹, Konrad Steffen¹, Ted Scambos², Eric Rignot^{3,4}, Jose Luis Rodriguez⁵, Waleed Abdalati^{1,6}

¹ CIRE, University of Colorado at Boulder, Boulder, Colorado

² NSIDC, University of Colorado at Boulder, Boulder, Colorado

³ University of California, Earth System Science, Irvine, California

⁴ Jet Propulsion Laboratory, Pasadena, California

⁵ CECS, Valdivia, Chile

⁶ NASA Headquarters

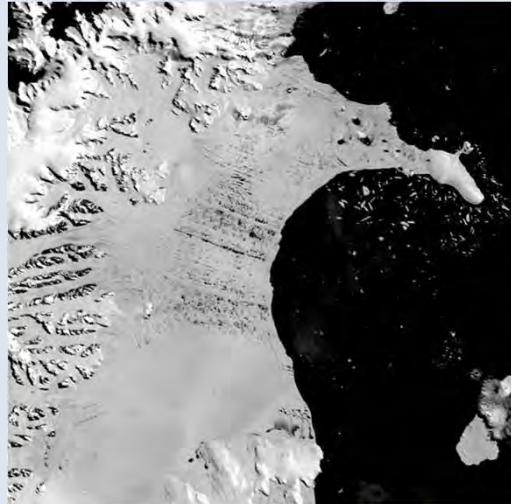
WAIS Workshop
September 21-23, 2011



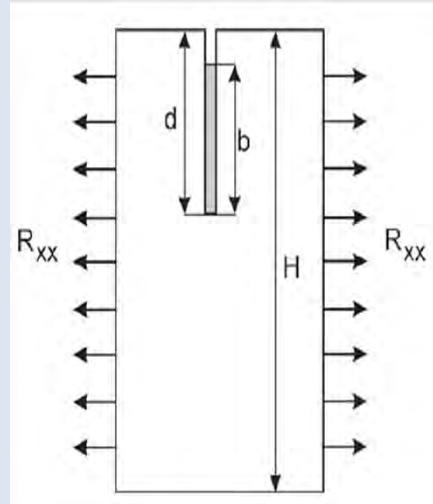
**British
Antarctic Survey**

NATIONAL ENVIRONMENT RESEARCH COUNCIL

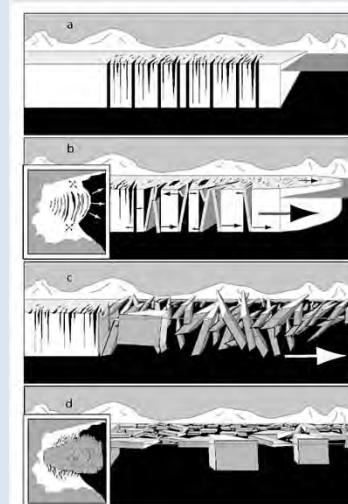
What have recent collapses told us about ice shelf processes? “climate induced mechanical instability”



Larsen B; pre-collapse



Van der Veen, 2007



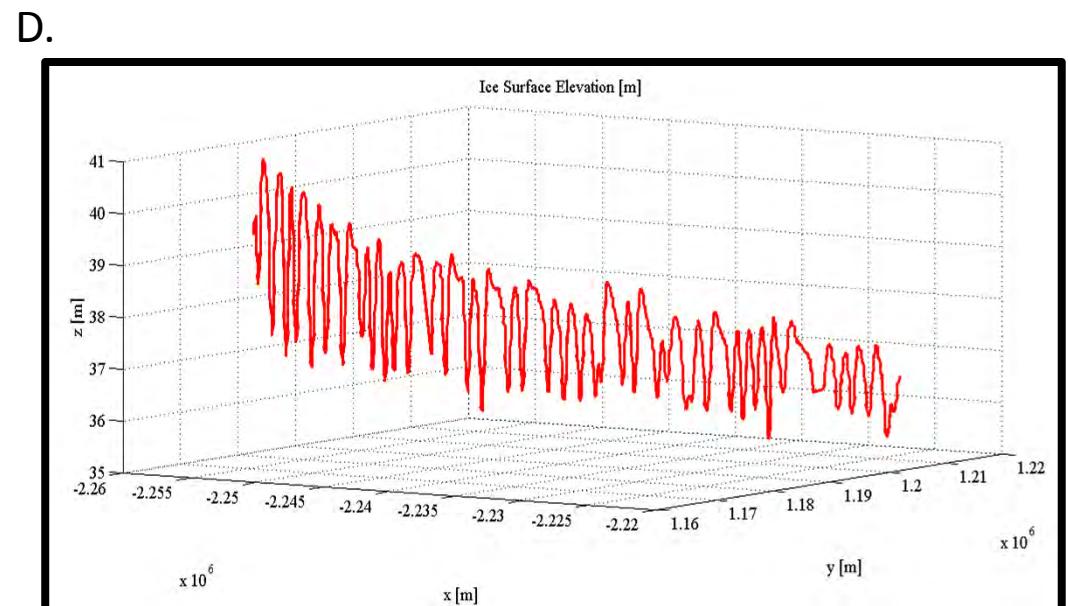
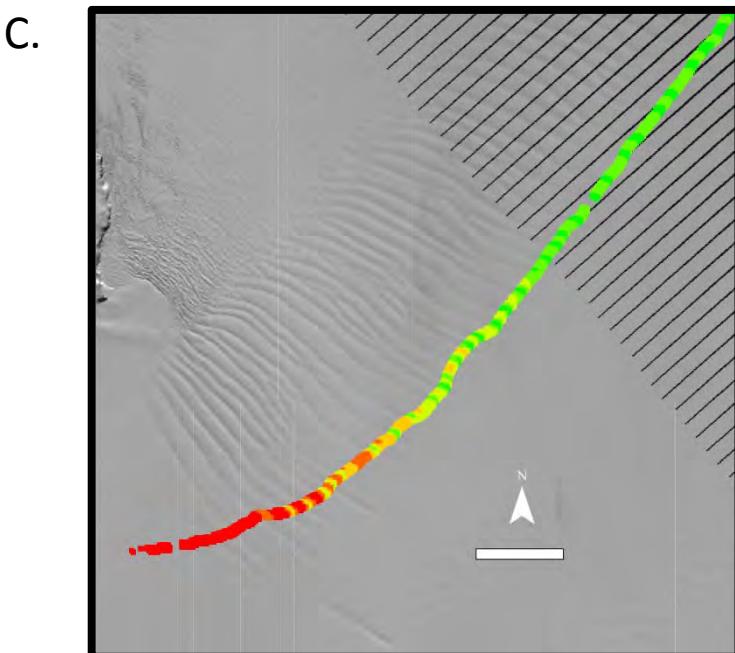
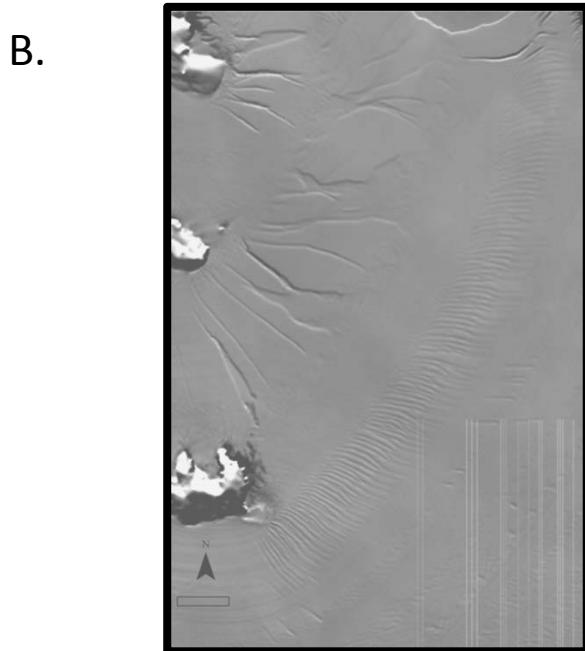
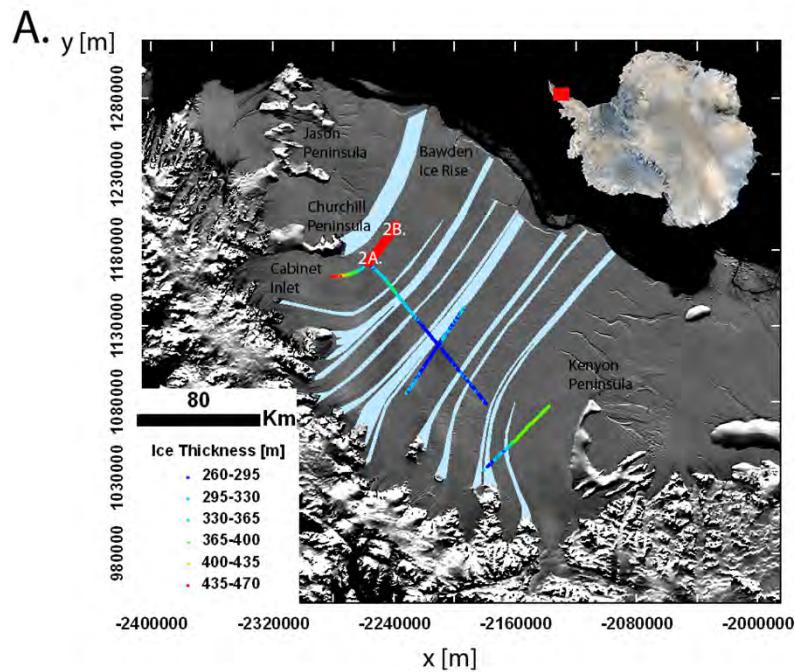
MacAyeal and others, 2003



Larsen B; post-collapse

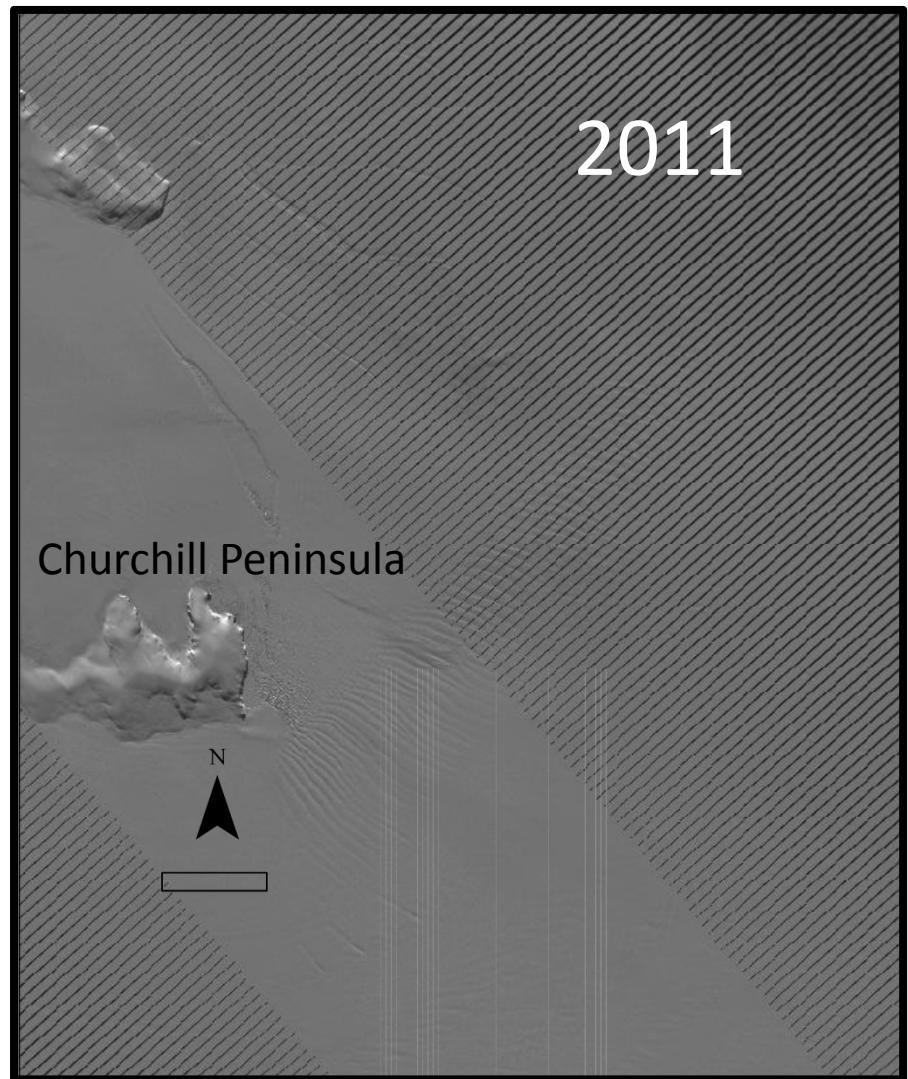
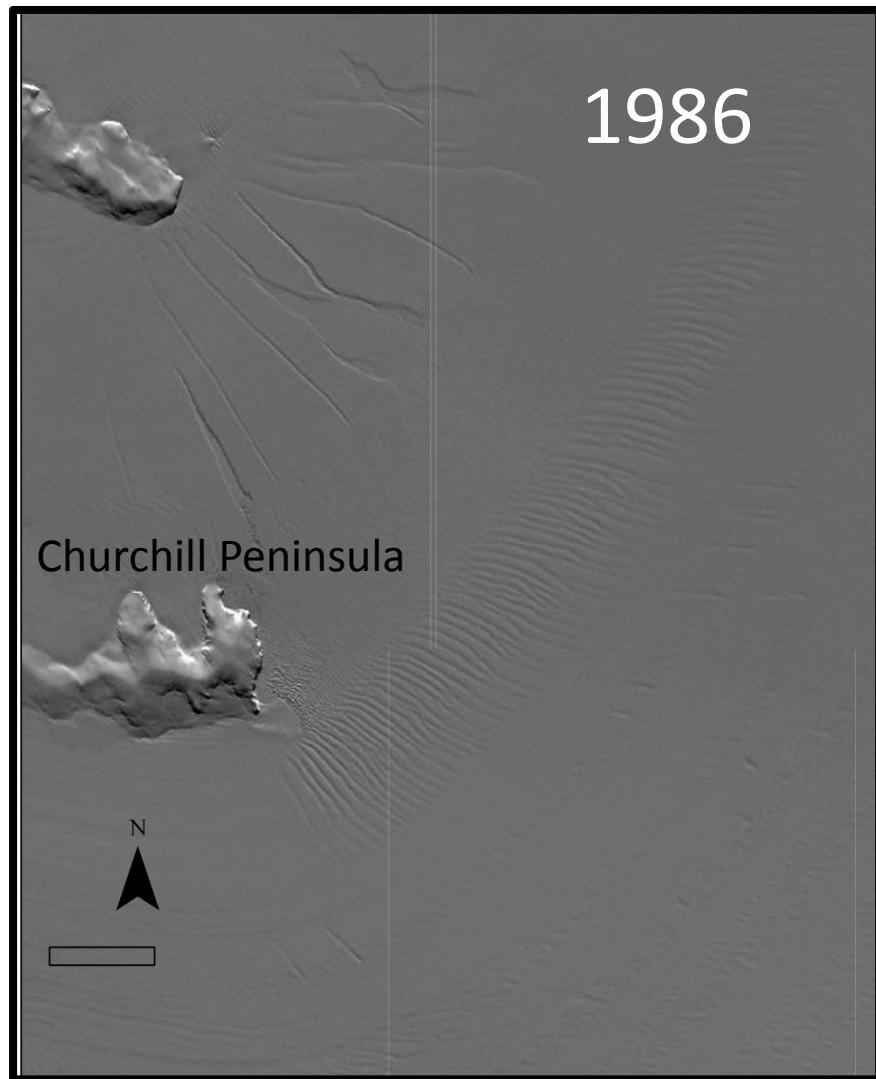
Pre-conditioning of the ice shelf:

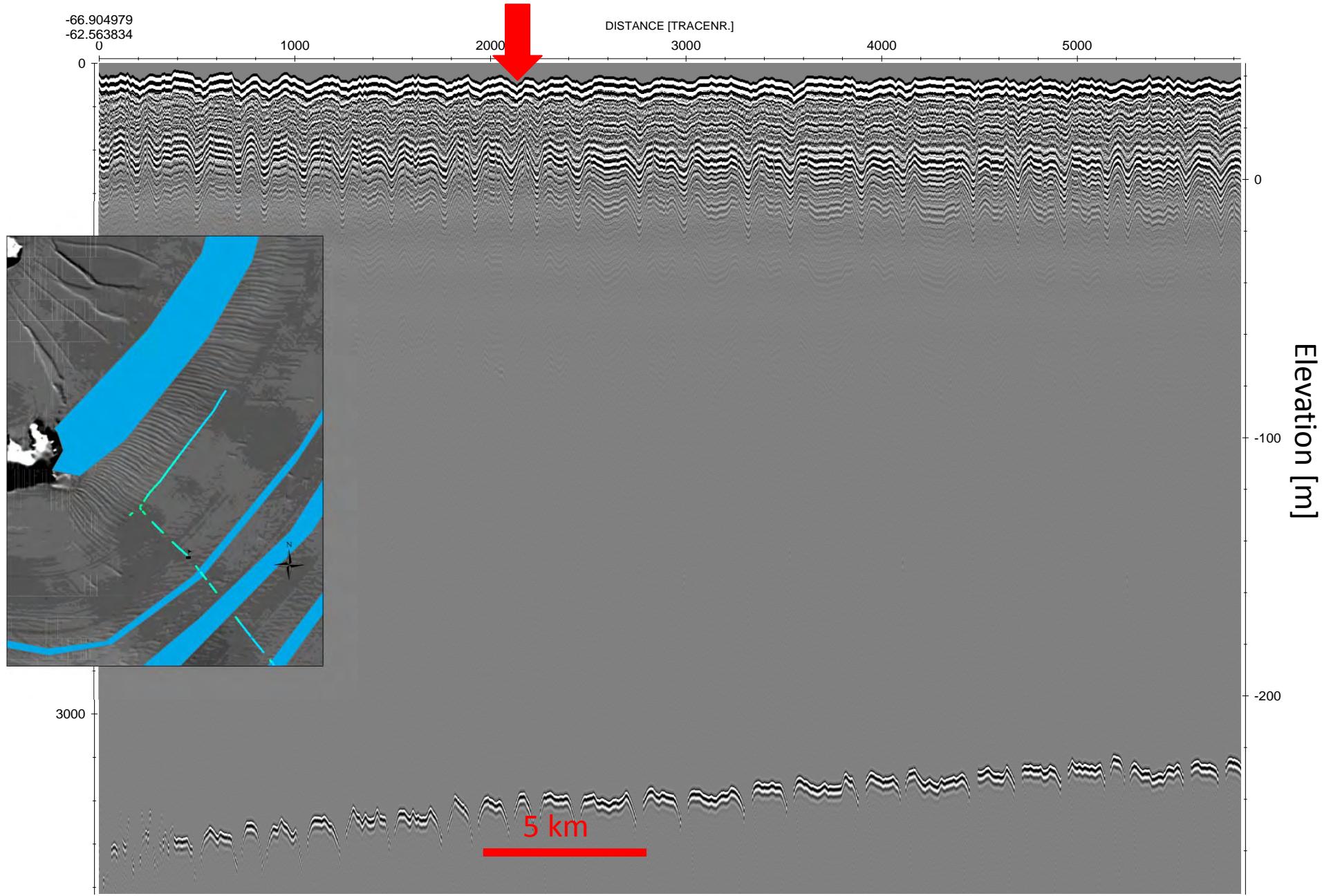
- a) Thinning (firn densification, basal melt)
- b) Ability to support surface melt ponds
- c) Reduced backstress from shear margins
- d) Surface and/or basal crevassing and rifting
- e) Reduced marine ice accretion/weakened suture zones

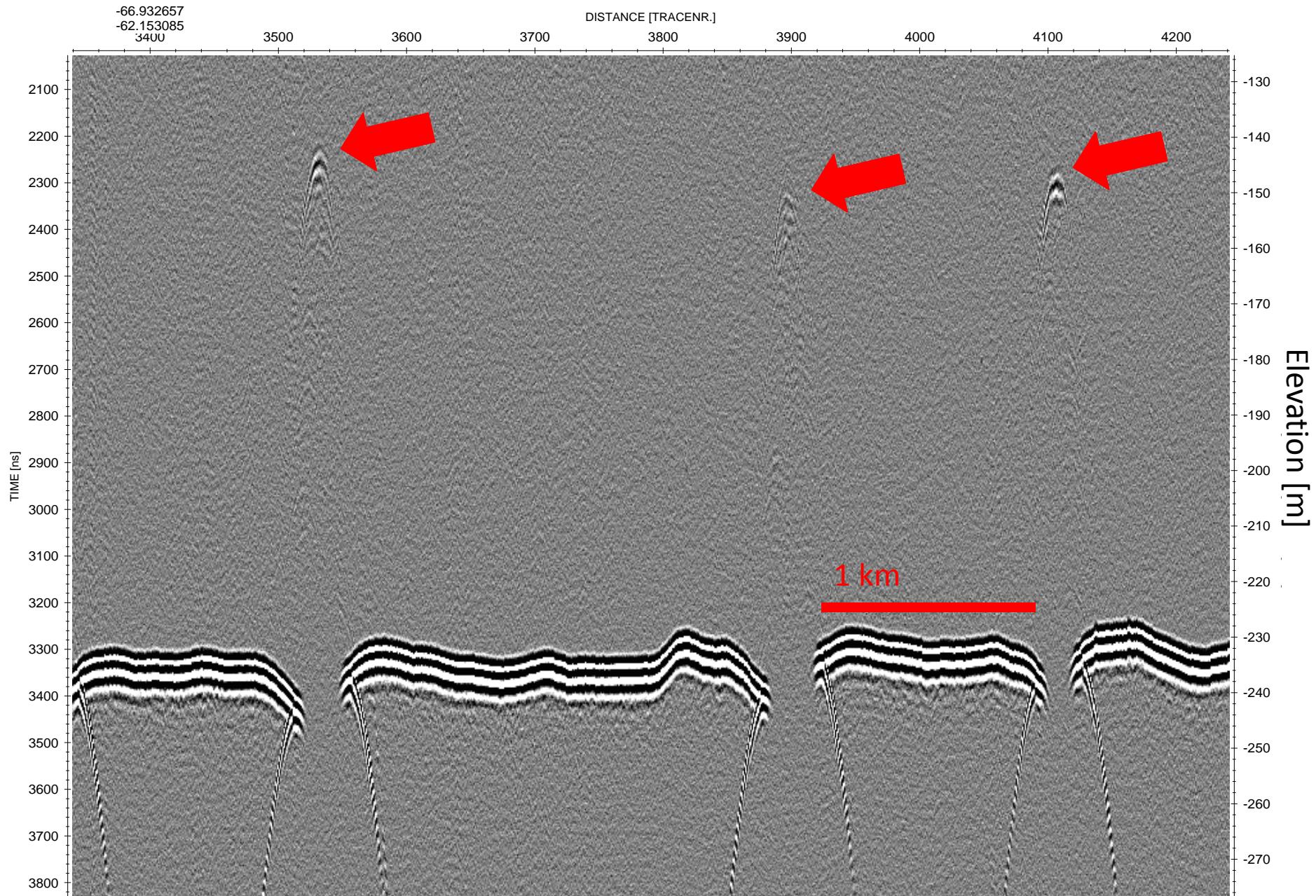


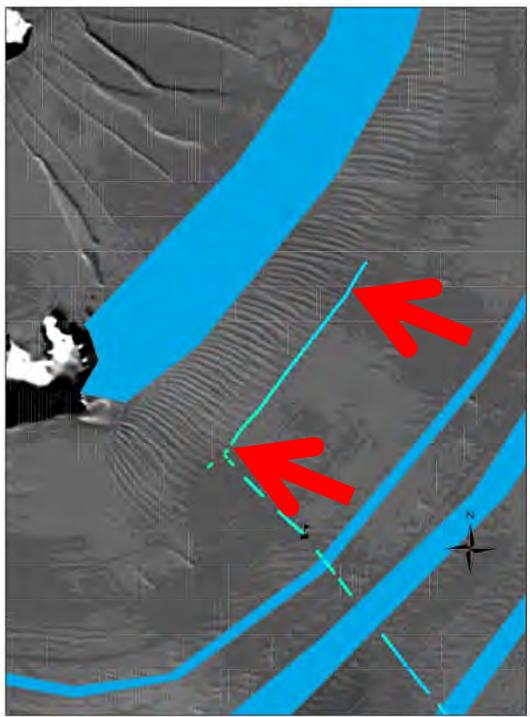
NASA IceBridge ATM surface elevation (2009)

Stable surface features of Larsen C





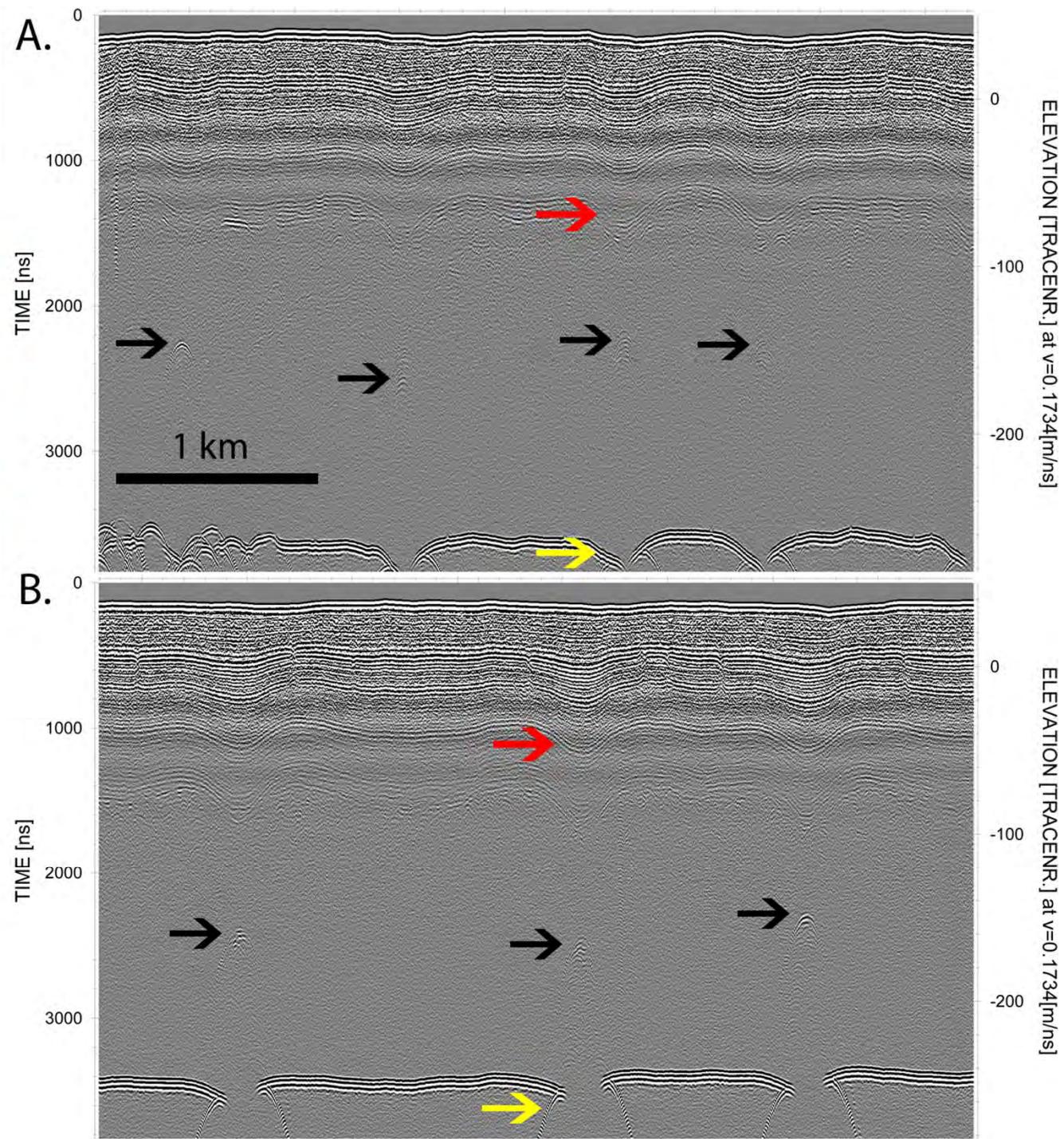




Spacing: 0.54 to 2.04 km, mean of 1.2 km

Height: 69 to 130 m, 24 to 43% of ice thickness

Width: ~120 m at end of transect



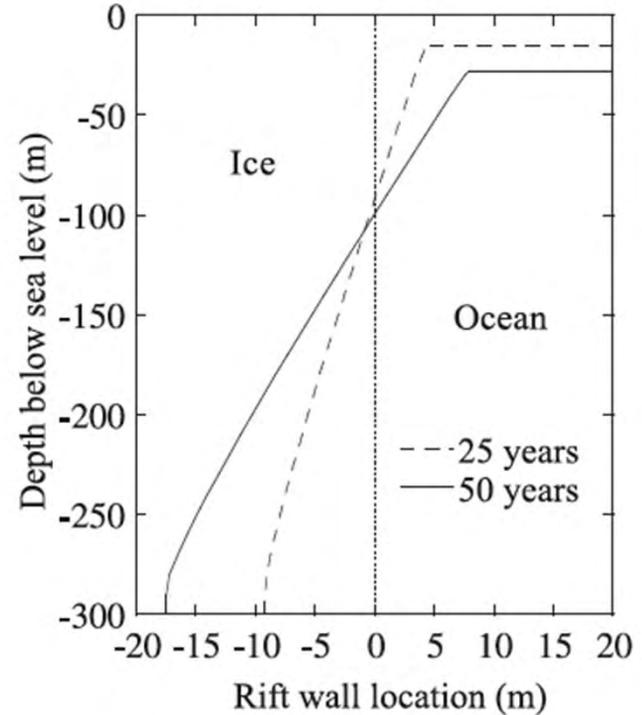
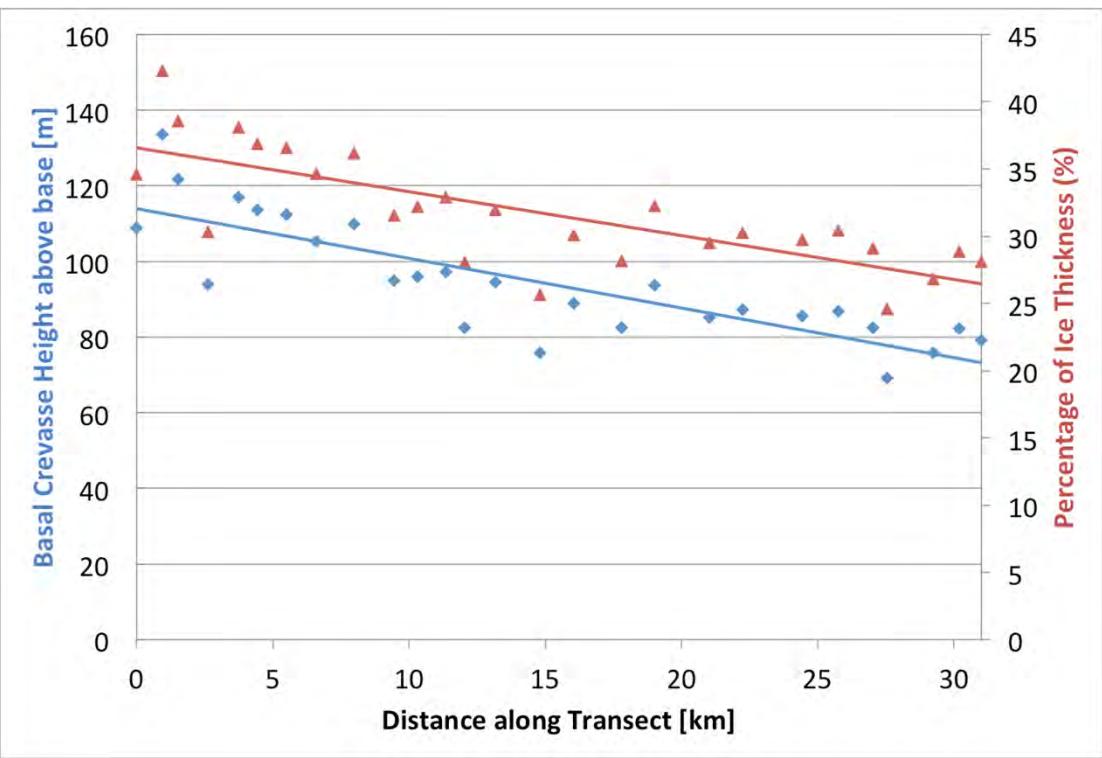
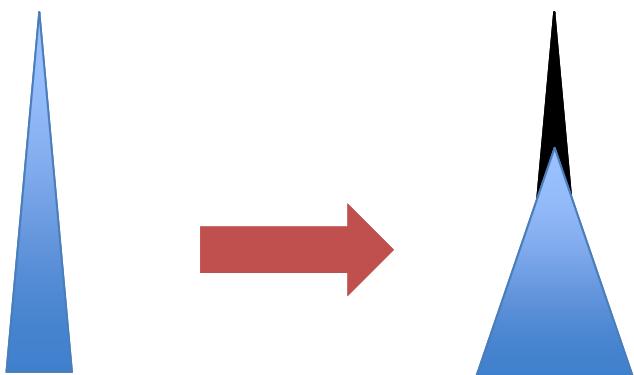
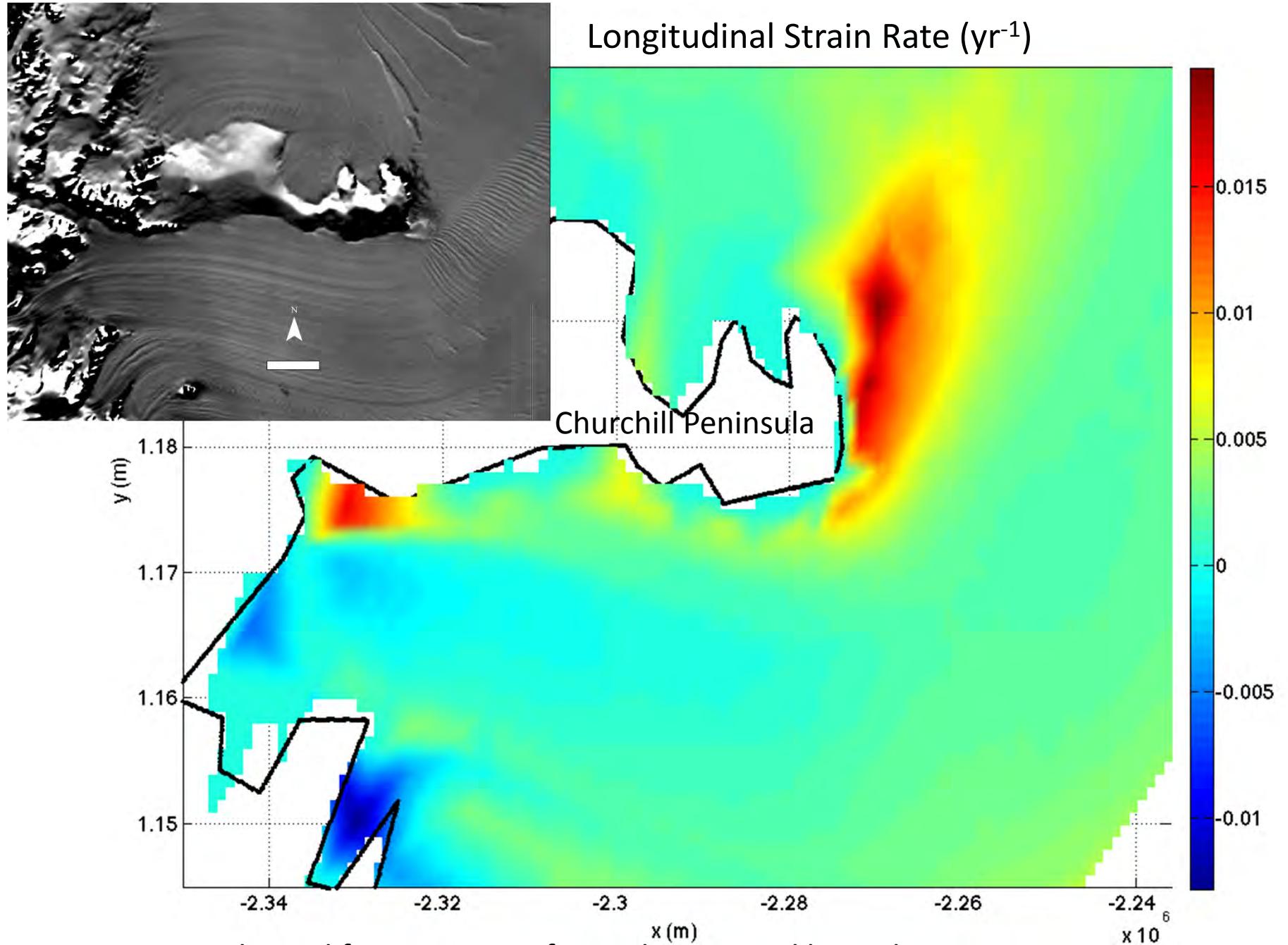


Figure 5. The profile of the rift's wall after a 25-year integration and at the end of a 50-year standard experiment compared with the initial wall location (which is represented by the vertical dotted line at the middle of the horizontal axis). The horizontal extensions of the two curves at the top indicate the interface between the accumulated ice and the water at the corresponding times. Notice that the horizontal distance scale is exaggerated relative to the vertical one.

Khazendar and Jenkins, 2003



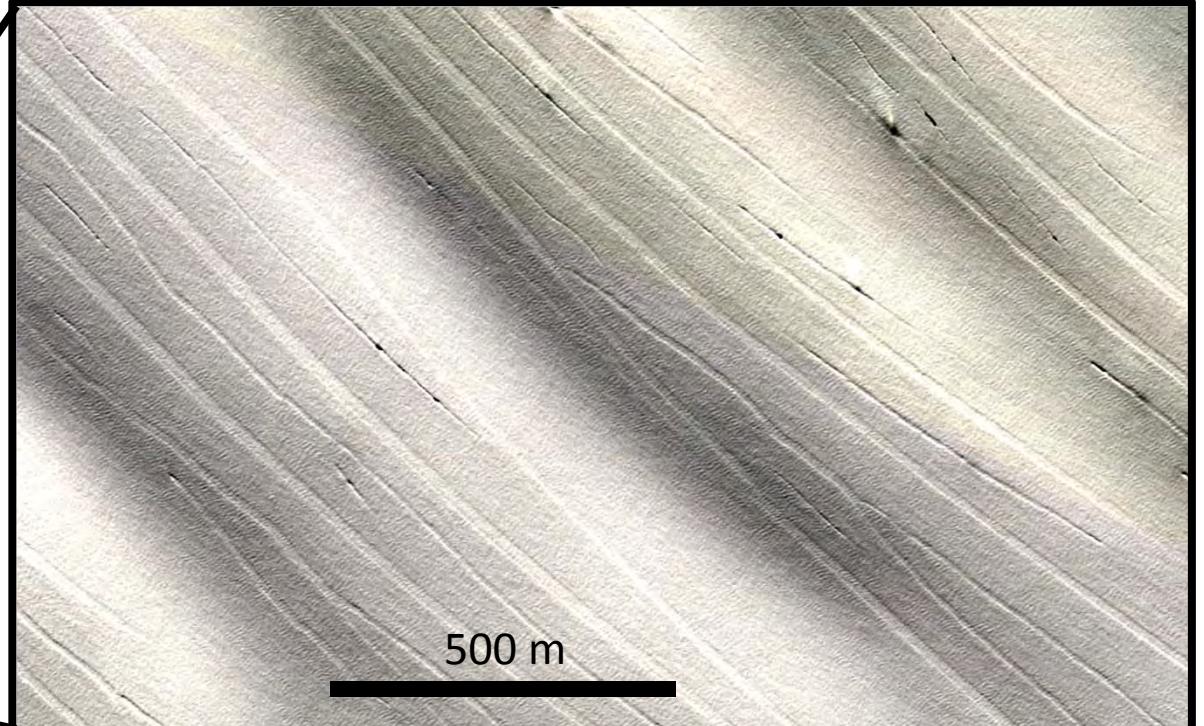
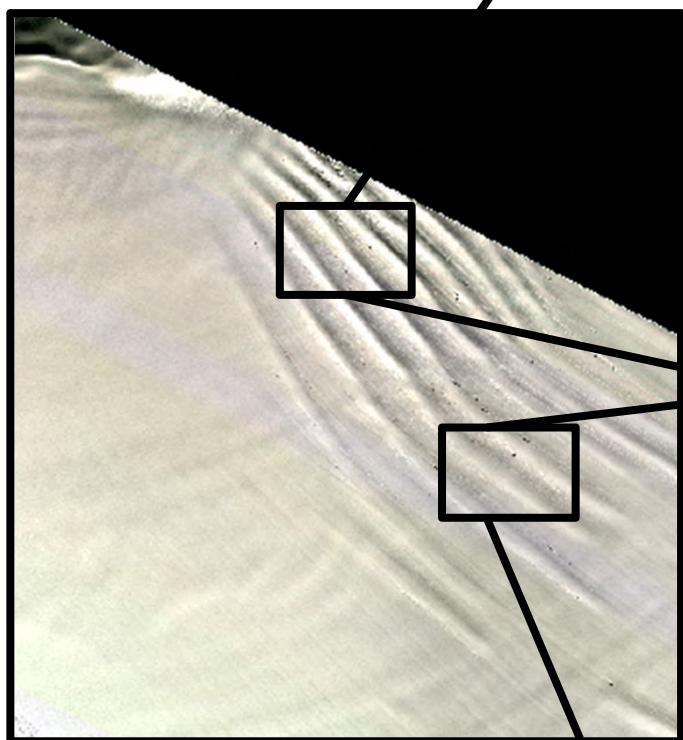


Strain rates derived from InSAR surface velocity, speckle tracking on ALOS PALSAR 2008 data. E. Rignot, 2011

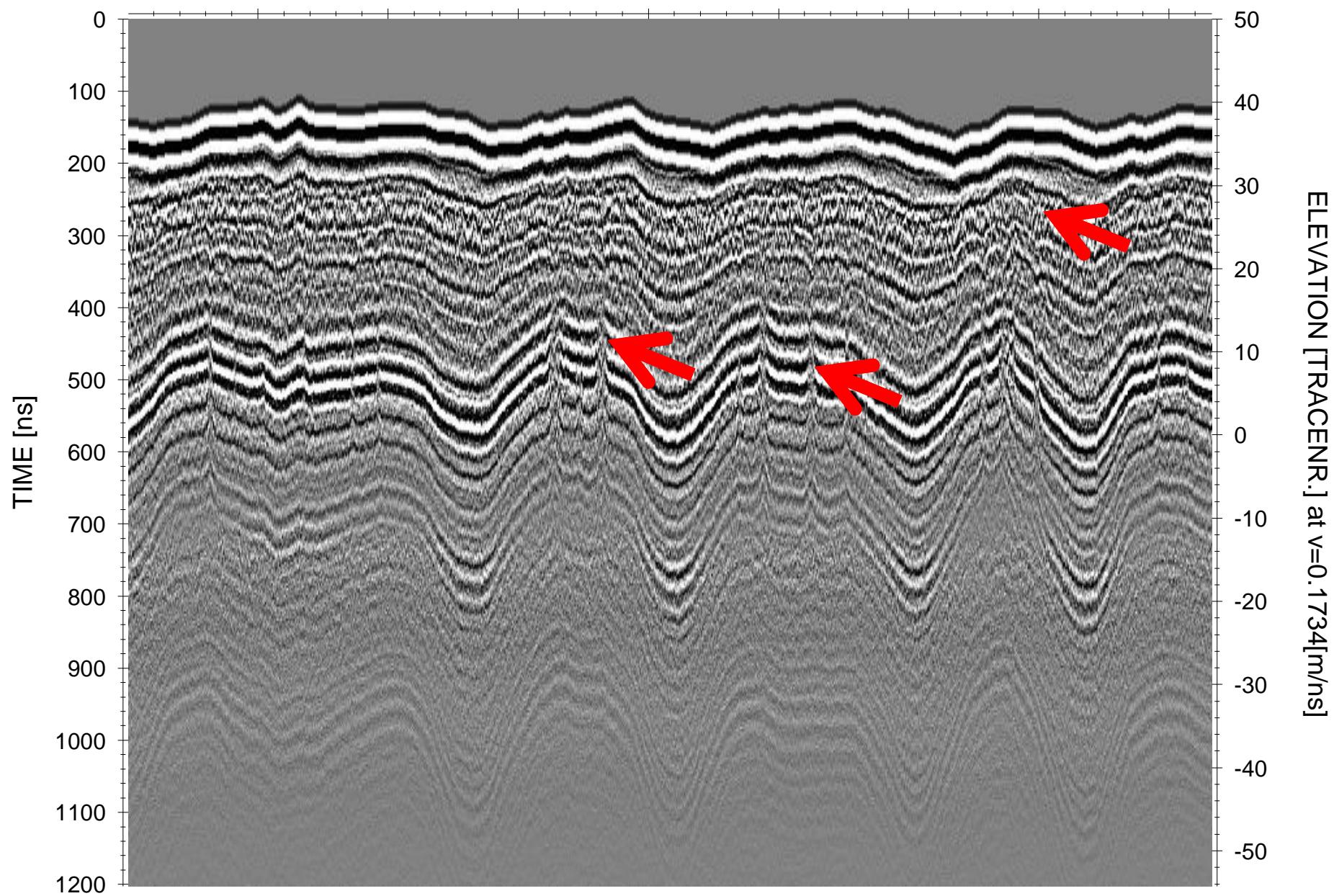
Why basal crevasses but no surface crevasses?

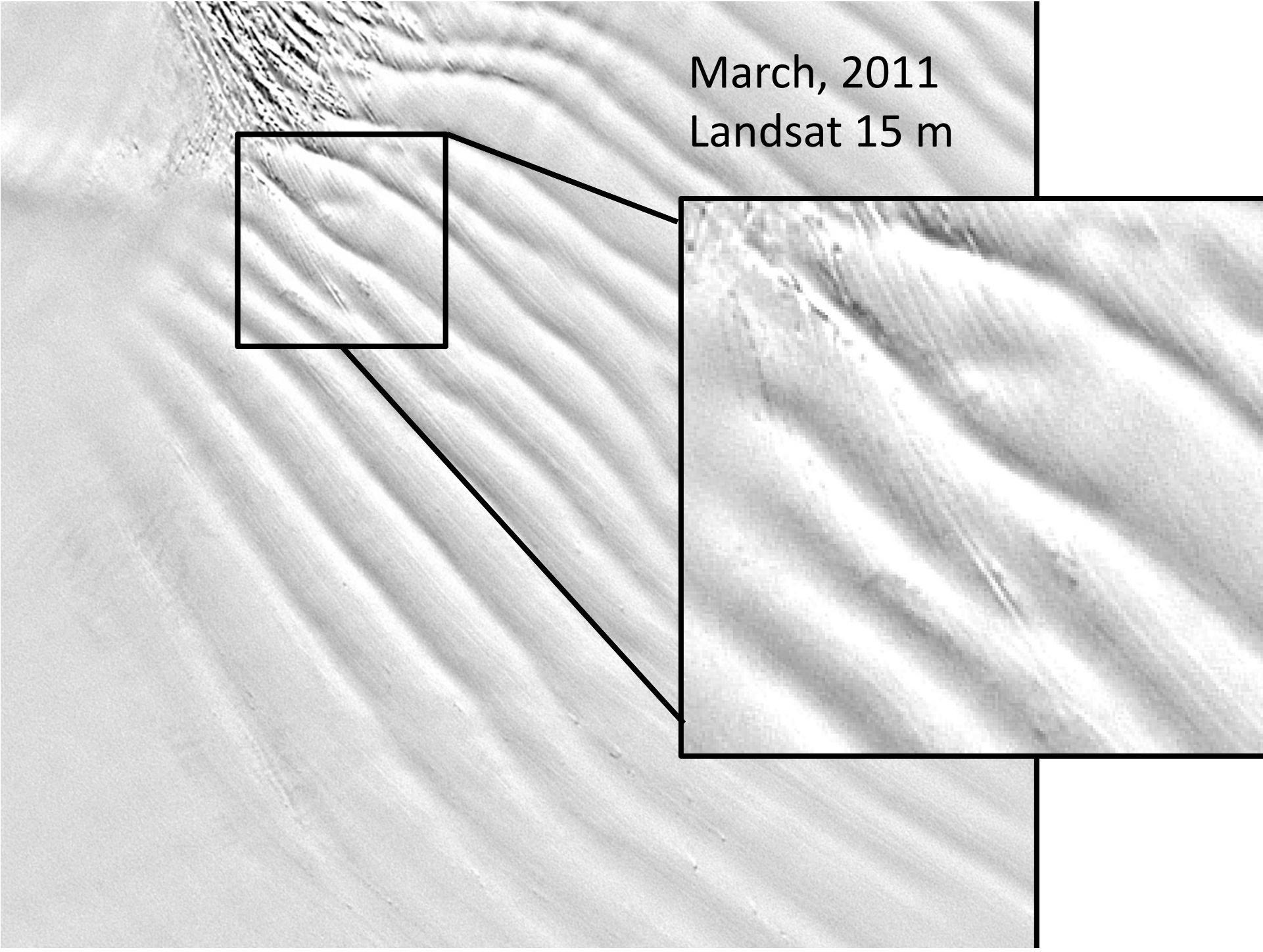


**Geoeye ~2 m;
2010**

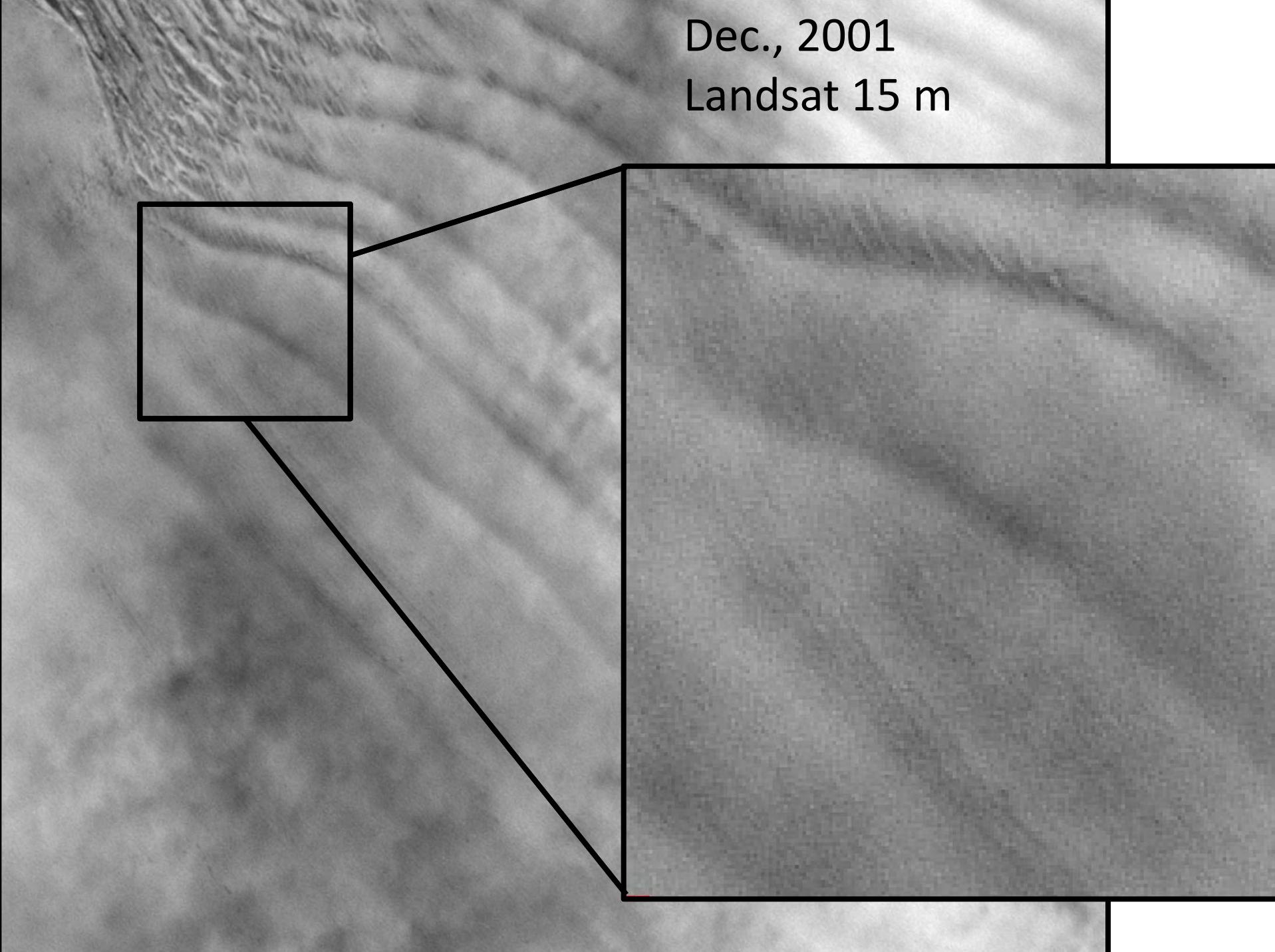


Surface crevassing at crests of surface undulations





March, 2011
Landsat 15 m

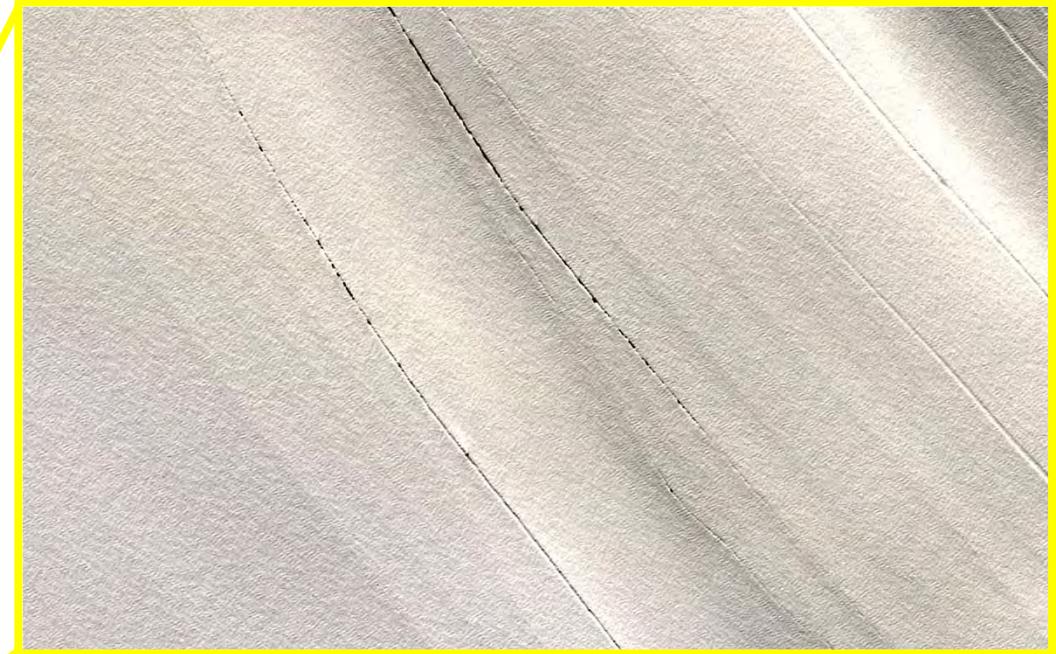
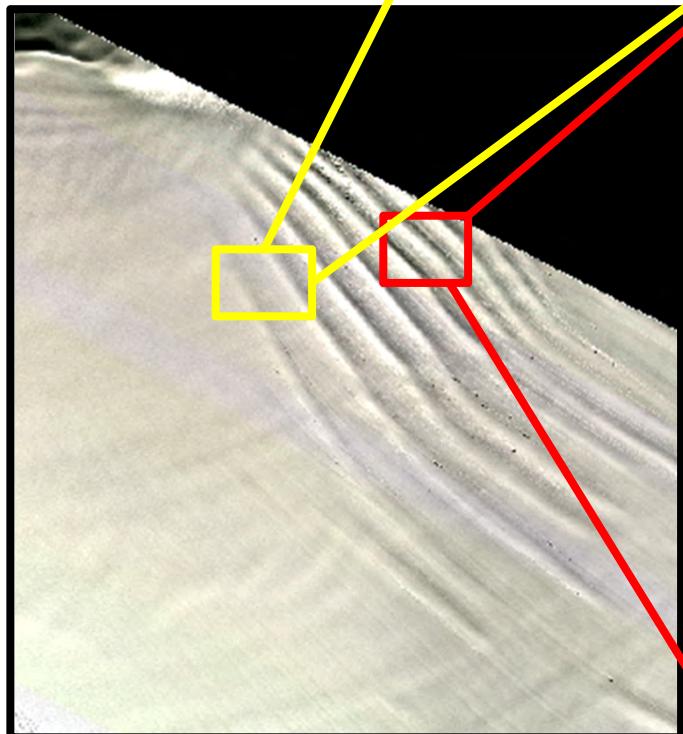


Dec., 2001
Landsat 15 m

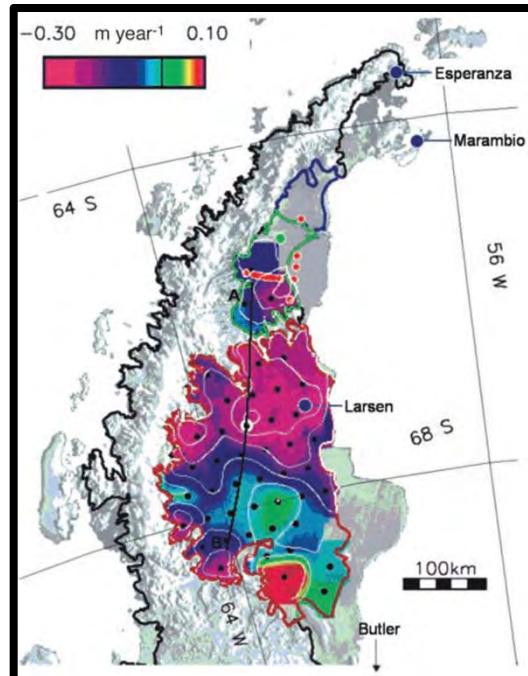
Timing/Mechanism?

-same time/stress as for basal crevasses

-bending stress induced by troughs → hydrostatic compensation of basal crevasses

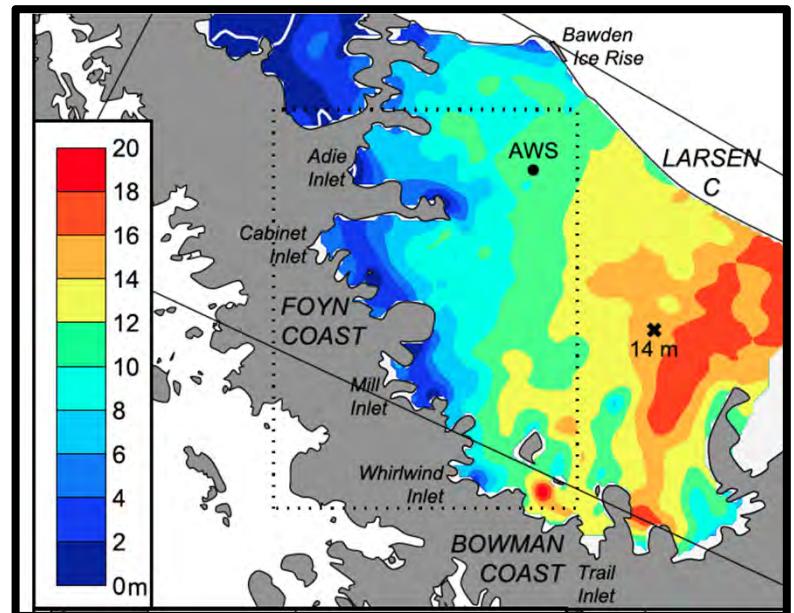


A.



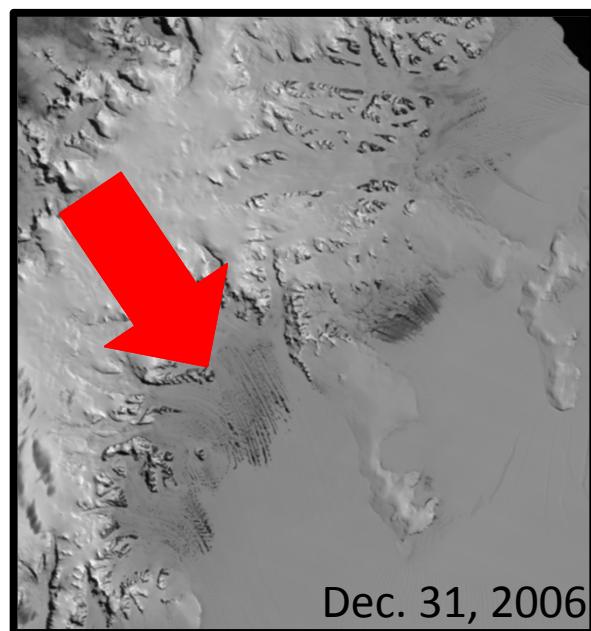
Shepherd and others, 2003

B.



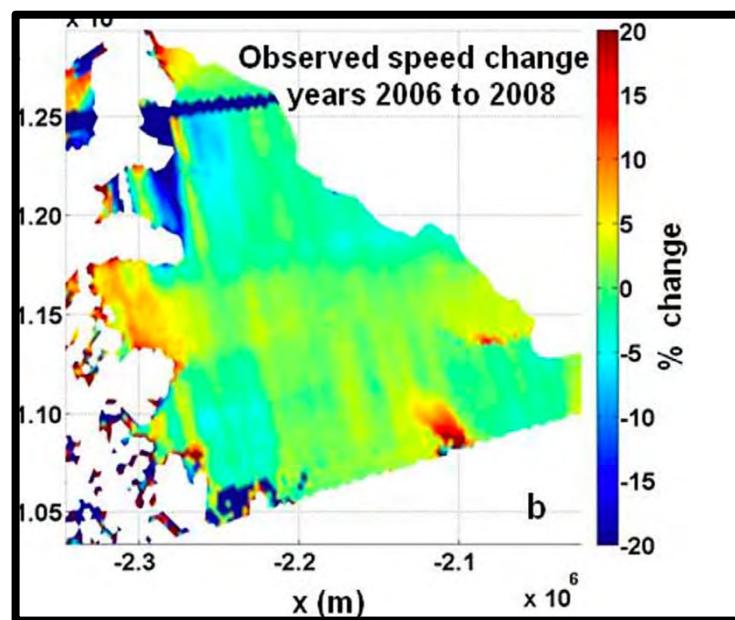
Holland and others, 2011

C.

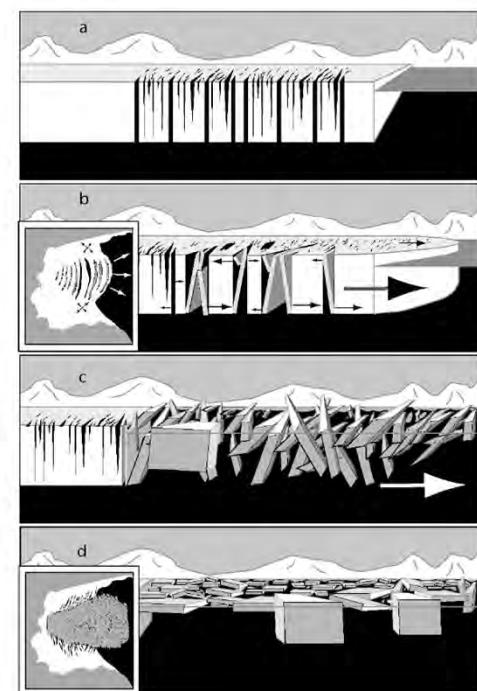
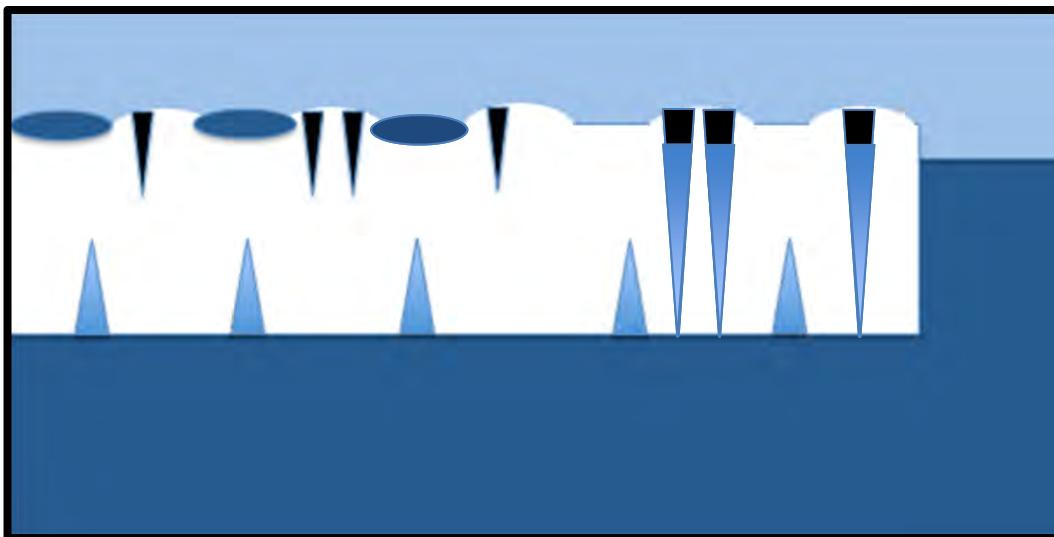
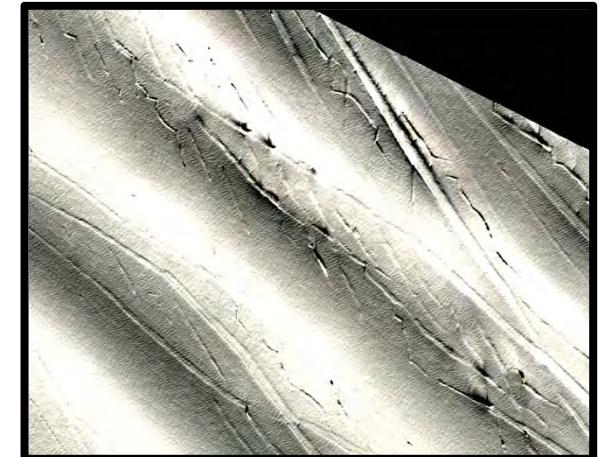
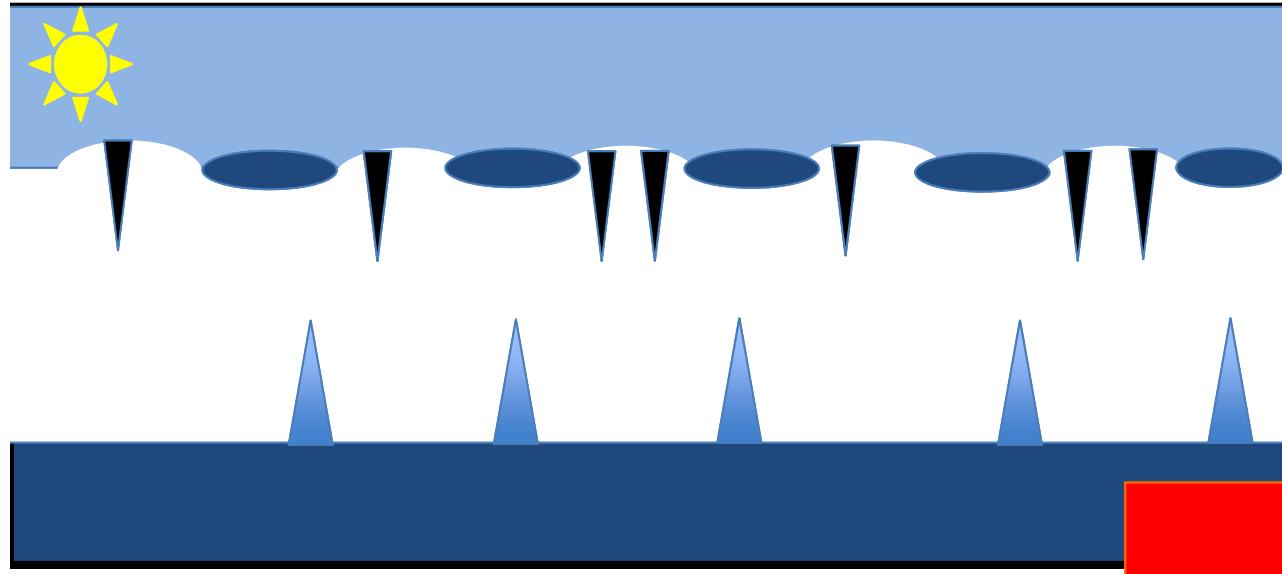


Dec. 31, 2006

D.



Khazendar and others, 2011



MacAyeal and others, 2003



Crevassing of the Larsen C ice shelf: Implications for ice shelf stability

Thanks to British Antarctic Survey for excellent field support; P. Morin and PGC for high resolution imagery; L. Padman for providing CATS2008a tide model; P. Holland for providing marine ice bands; W. Krabill and IceBridge/NSIDC for ATM data



Questions? Comments.

