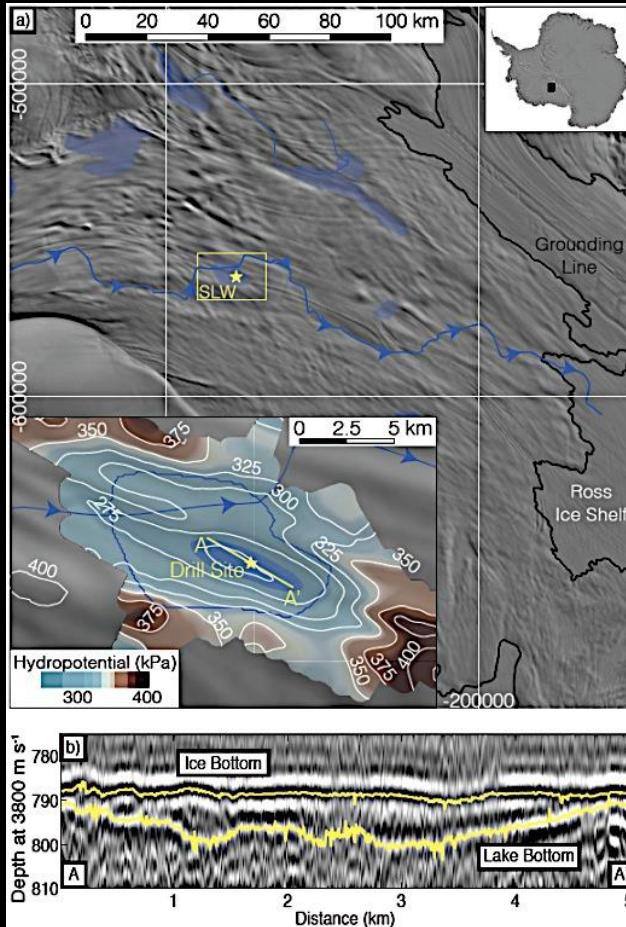


# Environment and Processes of Subglacial Lake Whillans, West Antarctica



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# Main Points

SLW sediment:

- comes from upstream and local marine sources
- shows evidence of various degrees of subglacial shear
- also evidence of recent dissolution (microbial mediation?)
- water saturated to compacted till - appears normally consolidated
- vertical clast fabric formed by decoupling during refilling
- last loading effect now over-printed

Lake discharge-recharge at low velocity

- ice recouples with lake bed at some lowstands
- till deformed into lake basin
- then ice re-floats and unloads
- no evidence of fluvial erosion or transport in subglacial flooding events
- flood velocities too low (<0.4 m/s) to entrain significant volumes of sediment
- water flow likely occurs in broad anastomosing sheets

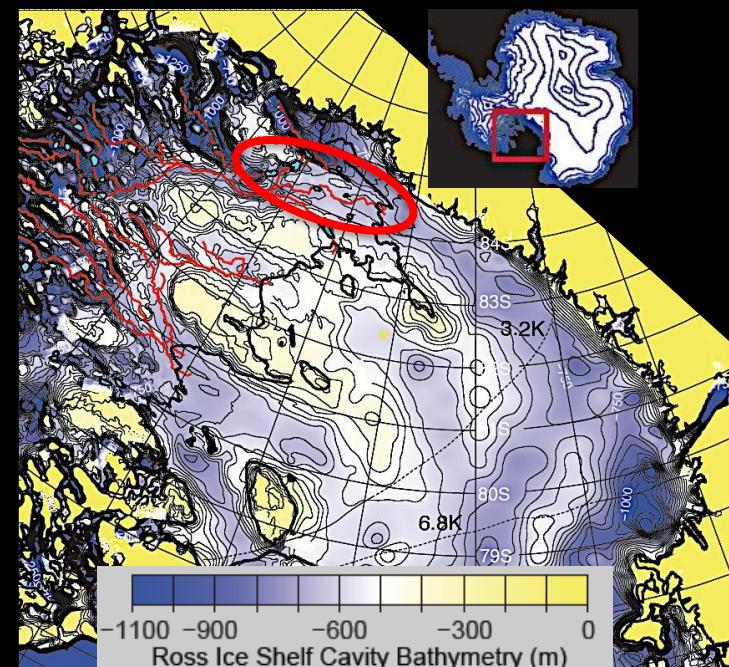
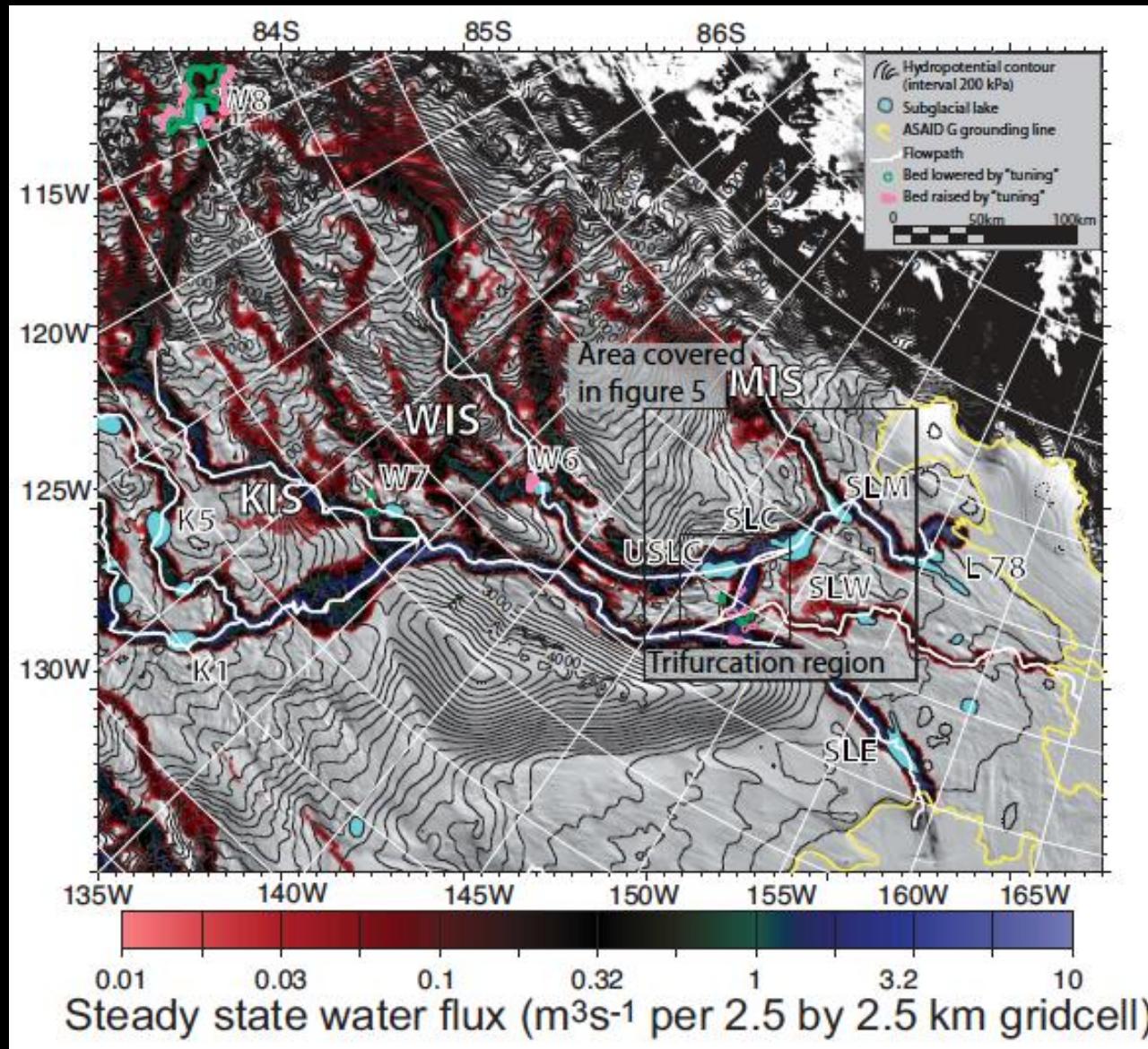


Image: Sasha Carter

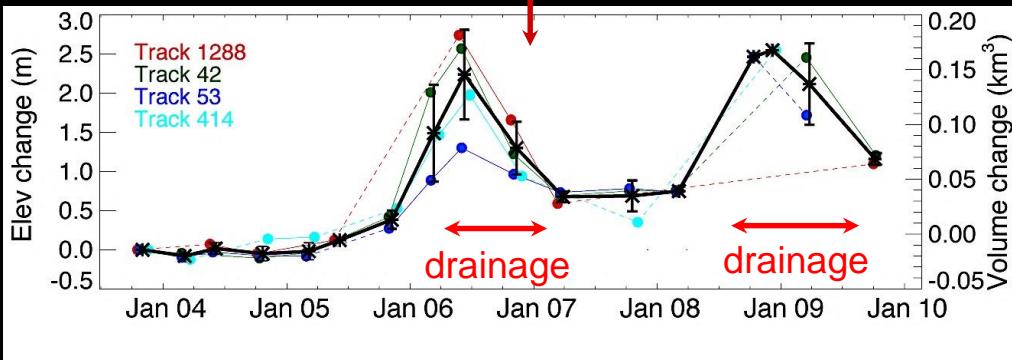
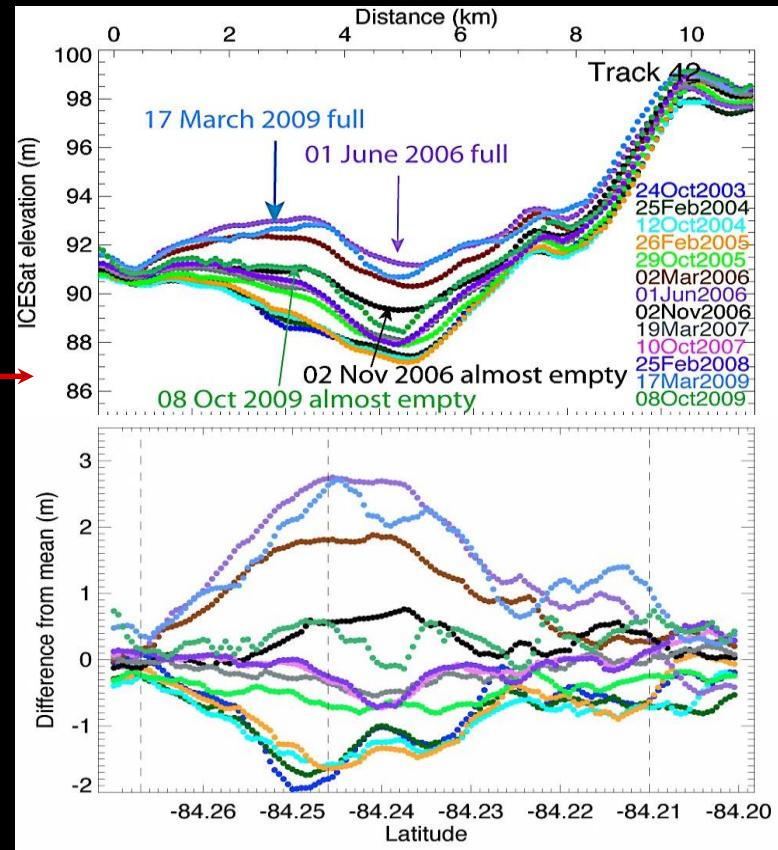
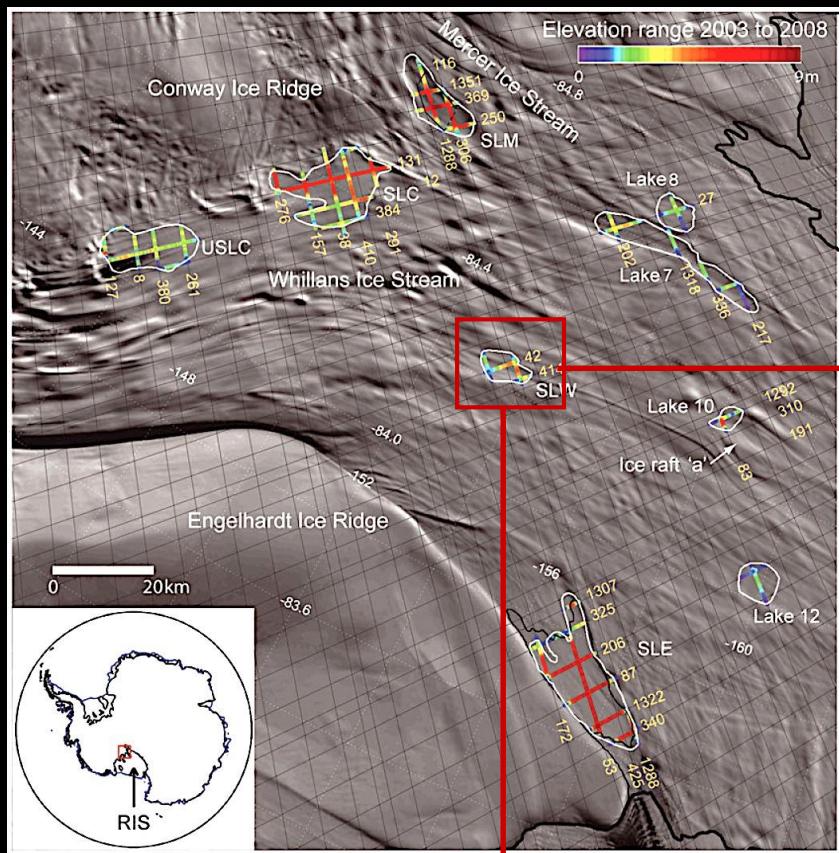
# A dynamic hydrological system

Lake is on a branch of a network of subglacial drainage



# A dynamic hydrological system

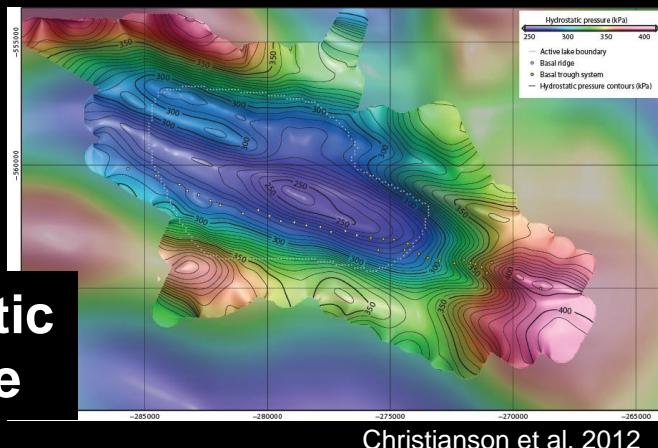
Lake discharges and refills on a period of a few years



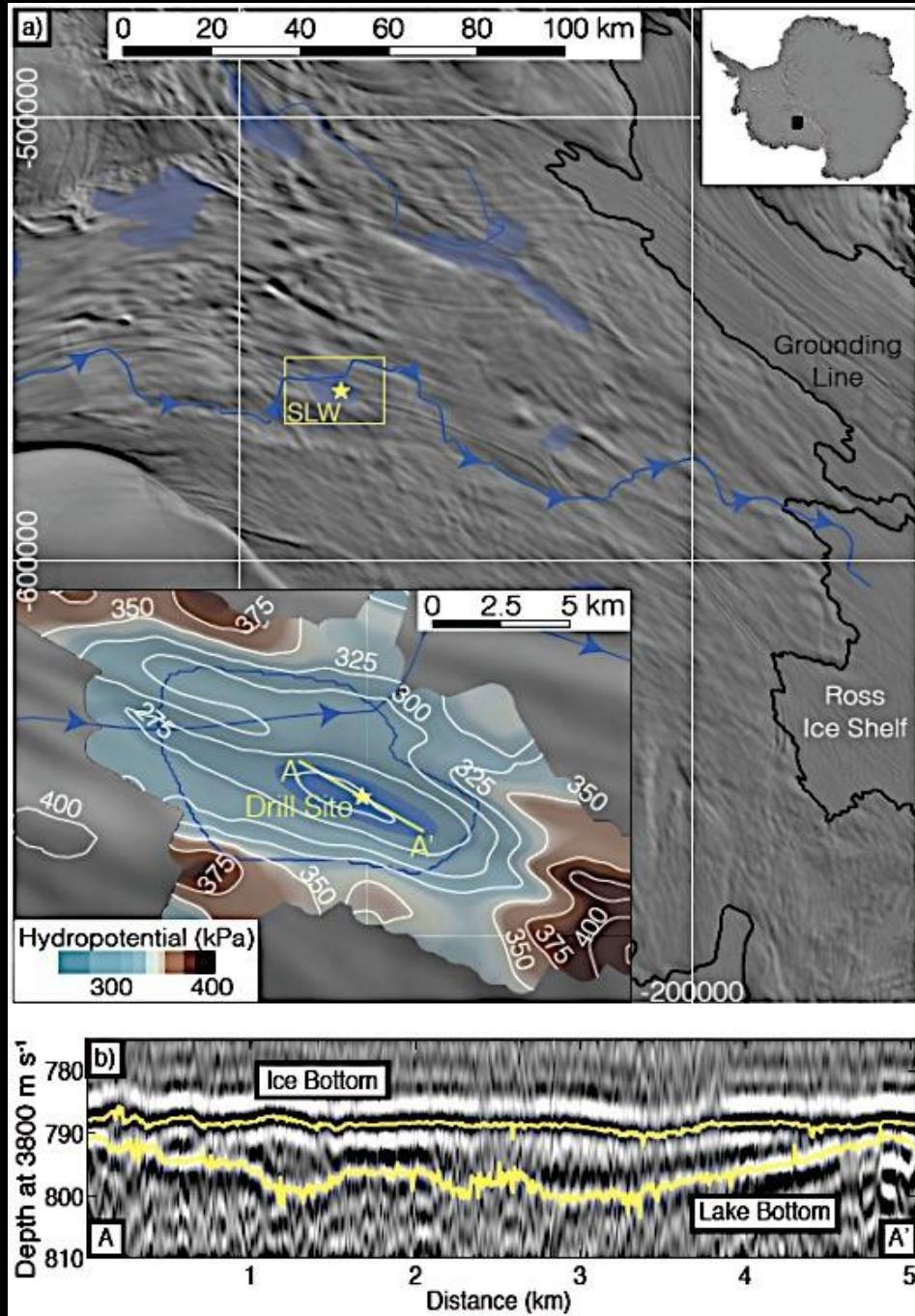
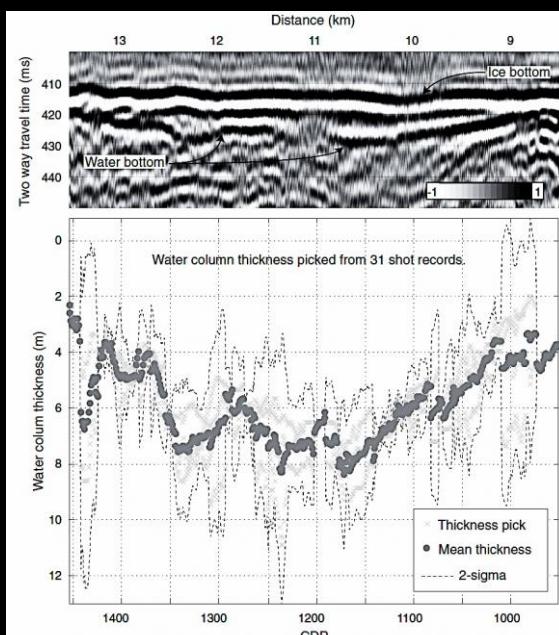
SLW has drained twice during the ICESat mission (2006 & 2009)  
The 2009 discharge was captured during the last two ICESat campaigns  
Now refilling

# Geophysical team provided definition of the lake

and best site to access



Seismic reflection





## Traversing from McMurdo



# Camp set up by SLW



**Lake bottom sediment rucked-up in front of camera**

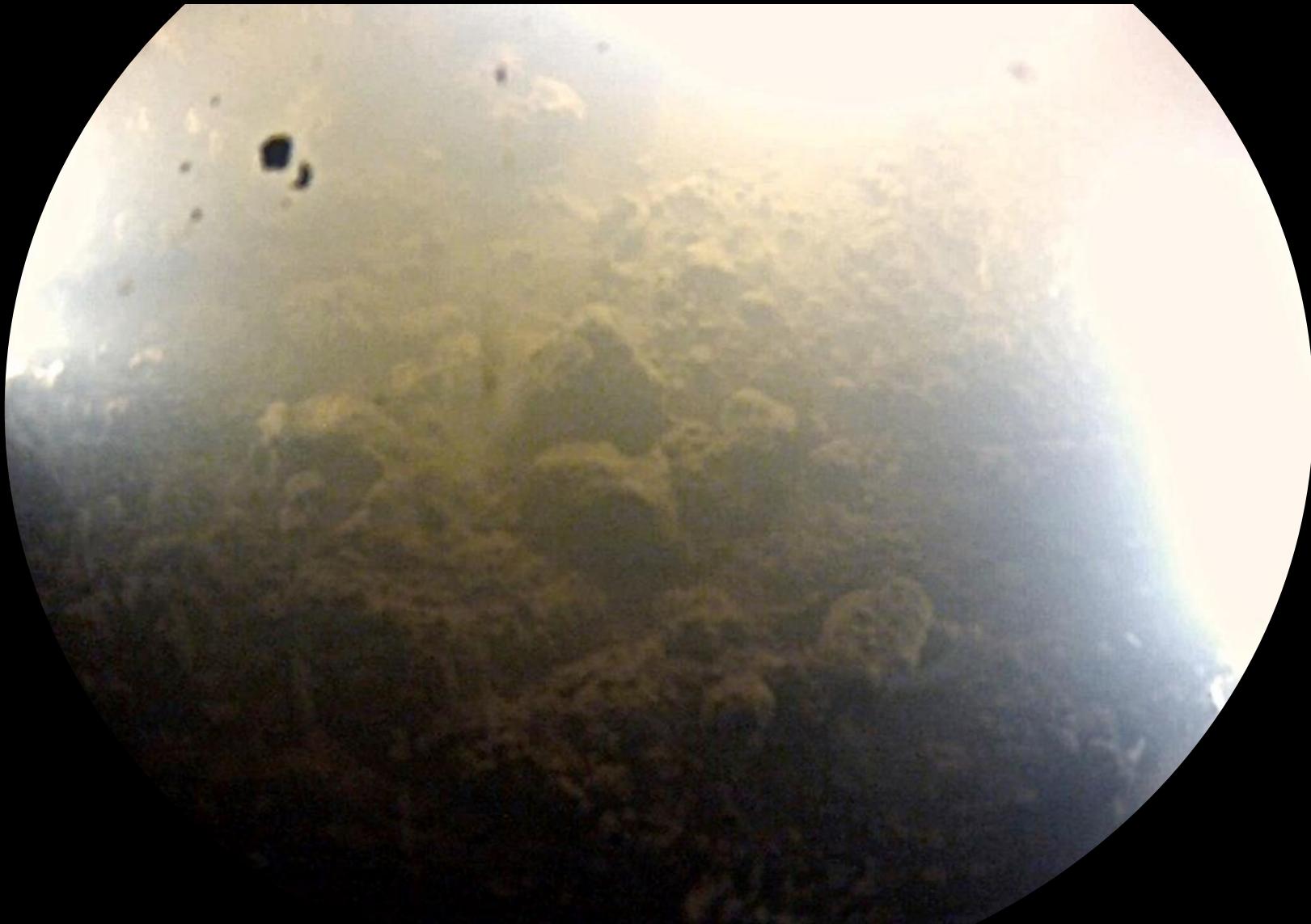
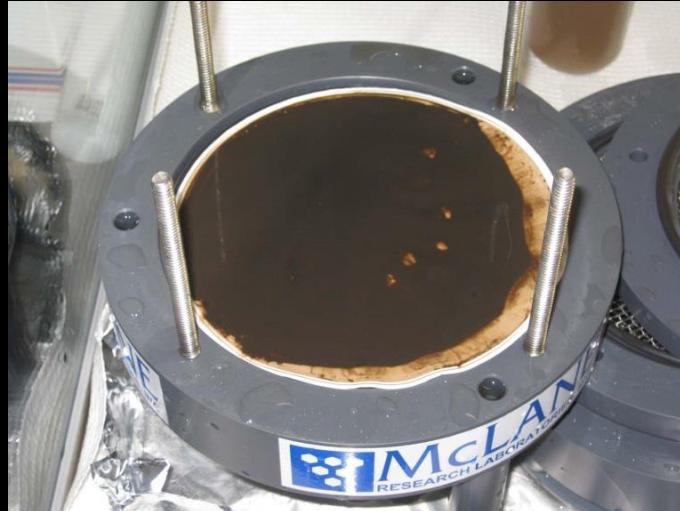


Image: Alberto Behar

# Major WISSARD borehole science goals were achieved



Three ~0.4 liter water samples



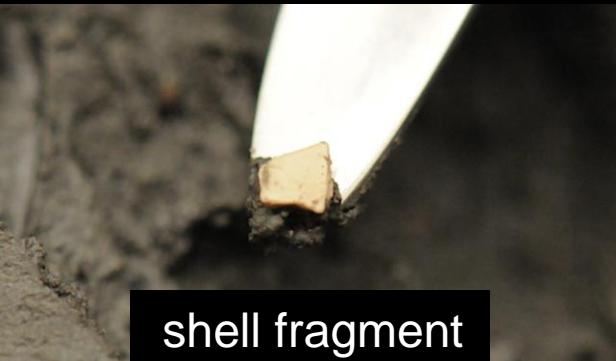
Suspended sediment filtered from lake water onto a 0.2 micron filter



Sediment from percussion, piston and multi-cores

# Sediment core analyses

ITRAX XRF scanner & XRF  
Geotek physical props scanner  
X-radiography  
Grain Mineralogy  
Clast lithology  
Biomarkers  
Particle size  
Clast fabric  
Paleomag NRM and AMS  
Grain & clast surface microtextures  
Thin section micromorphology  
Moisture content  
Strength tests



shell fragment

Core degassing



Sediment  
strength tests

# General Description

Below ~1m of turbid water

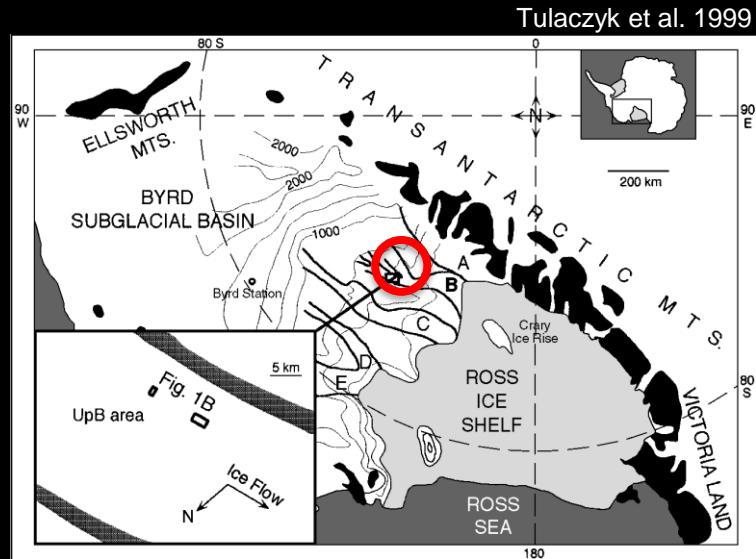
Local basal melting + flow from upstream

Structureless, clast-poor (<10%)  
muddy diamictite

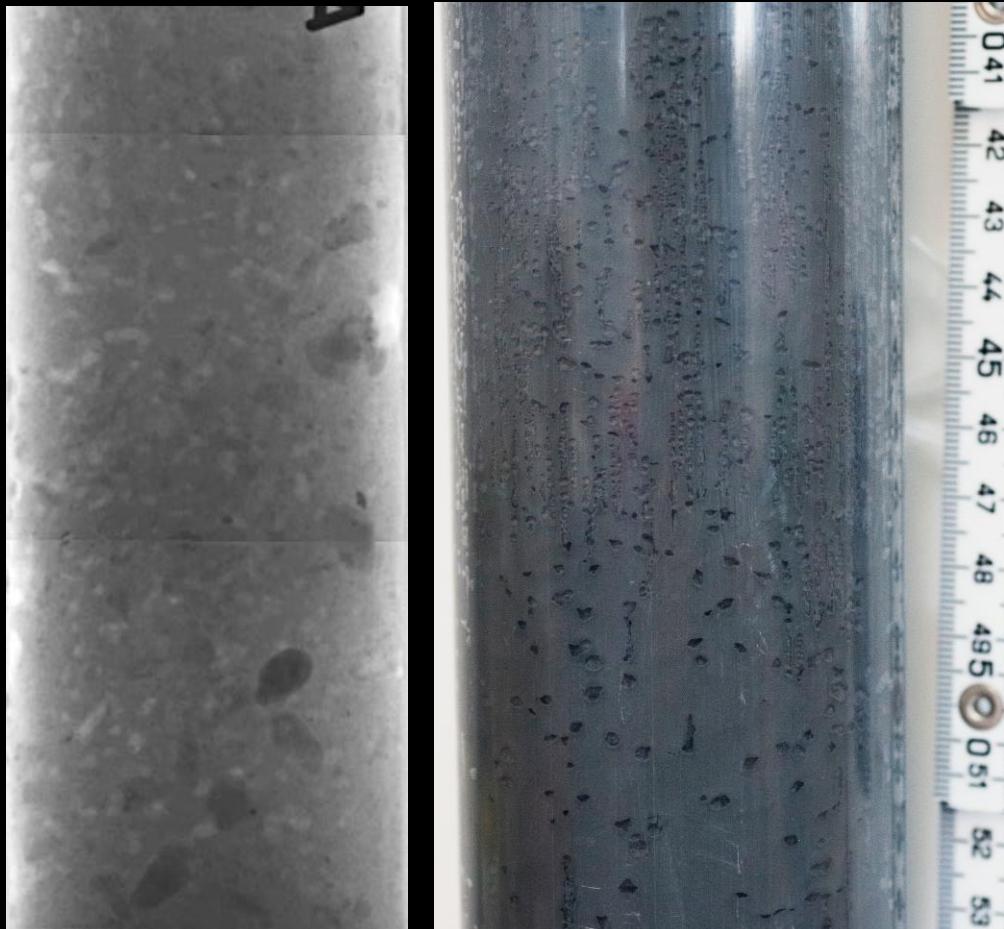
Water saturated to compacted

Appears very homogeneous

Compare with upstream ("UpB")



## Core degassing



X-radiograph

Image through clear  
core liner

# X-radiographs

- homogeneous - some weak layering?
- clast abundance - variable, < 10%
- clast orientation - locally preferred

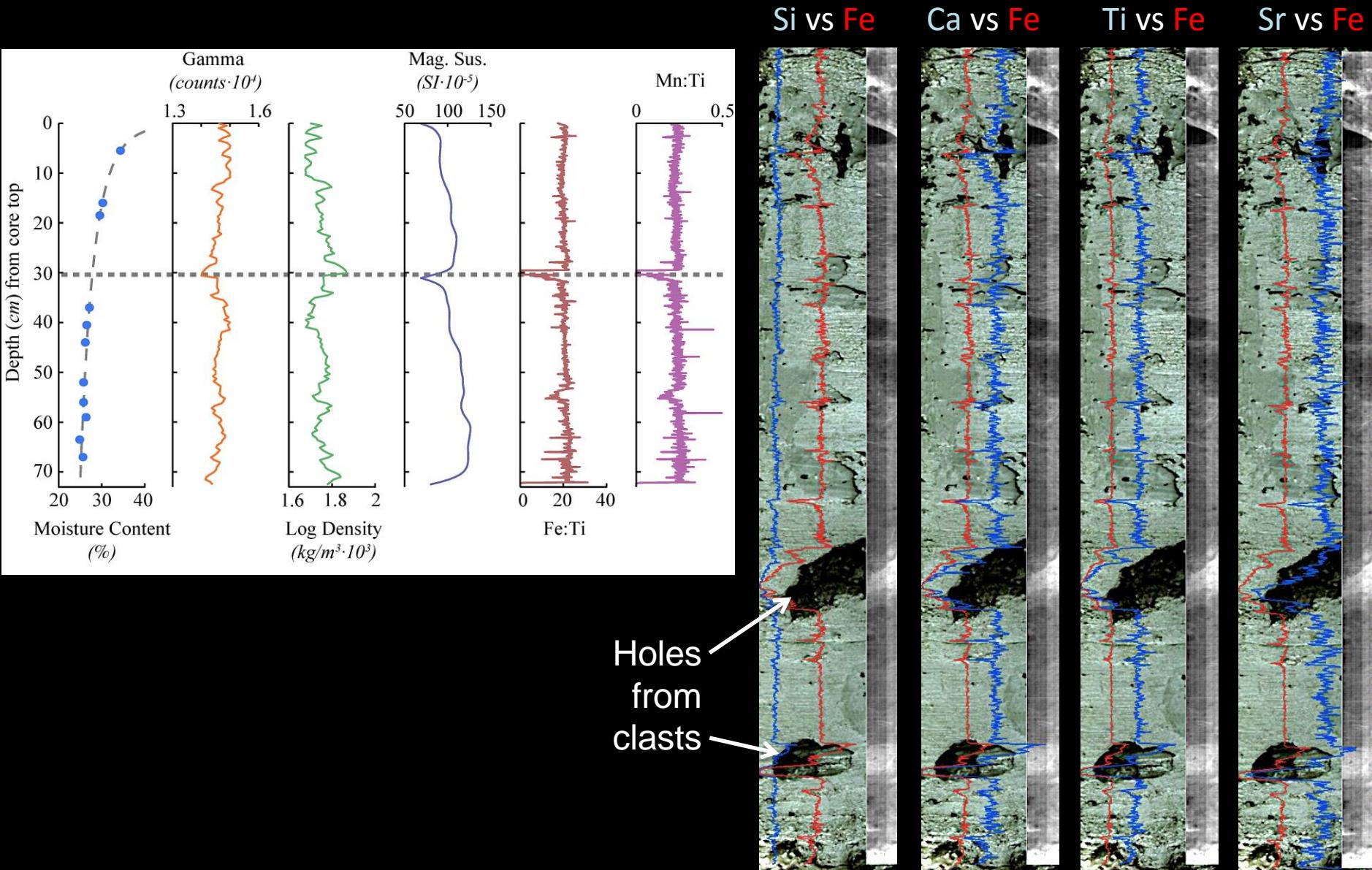
↙ *artifacts*

.84m-long core

.41m-long core

# Homogeneity shown by ITRAX XRF

(Percussion core)



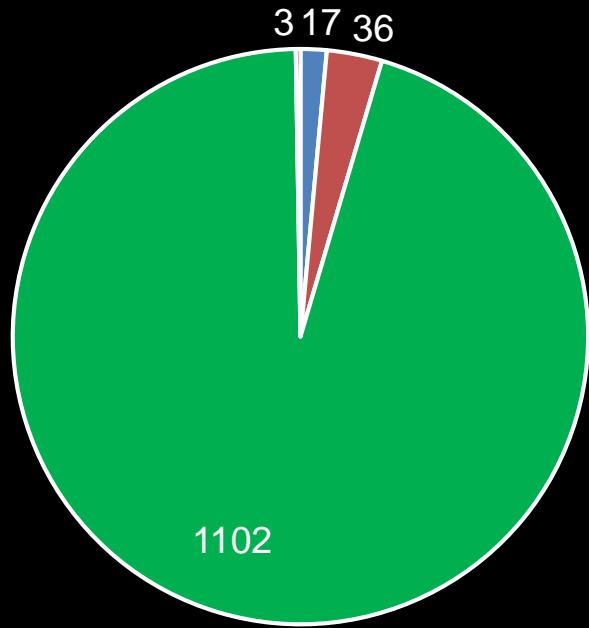
# Provenance analysis

## Methods

- Clast lithology
- Sand mineralogy
- Major/trace element chemistry
- Molecular biomarkers

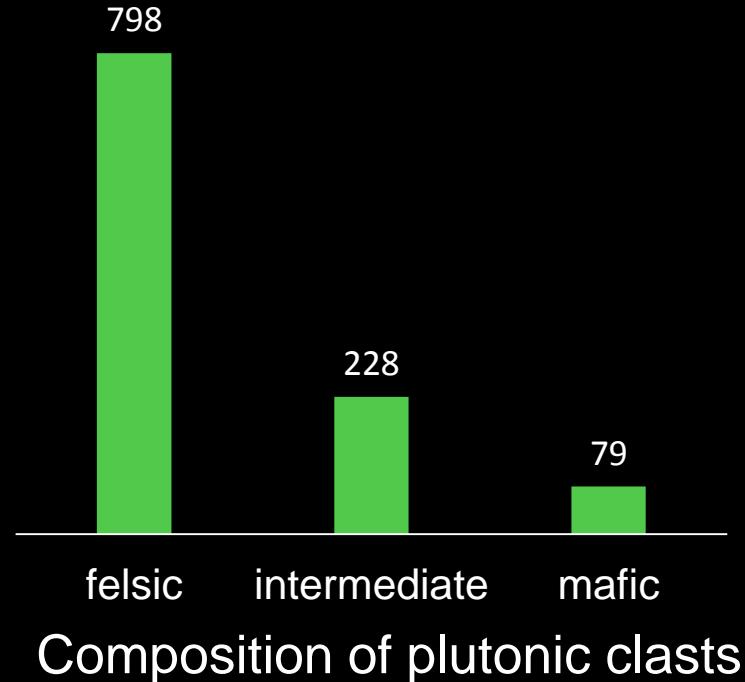
## Clast lithologies

1158 clasts - mostly from granules to fine pebbles



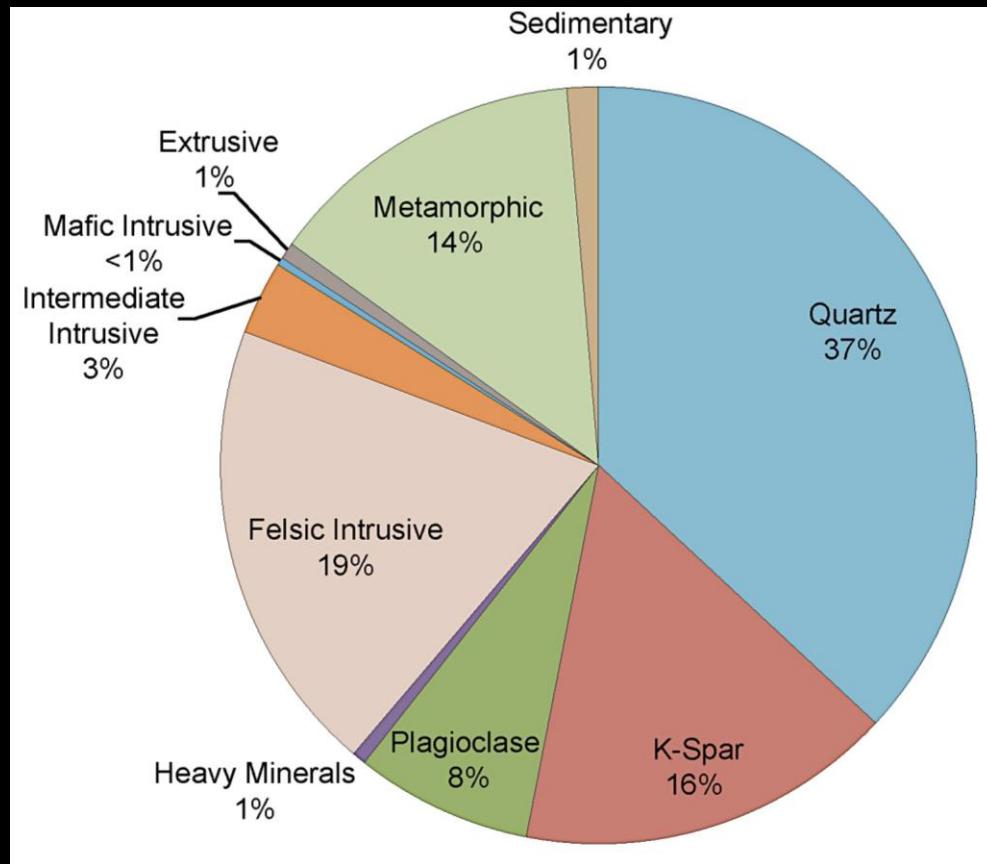
■ Sedimentary  
■ Plutonic

■ Metasedimentary  
■ Volcanic

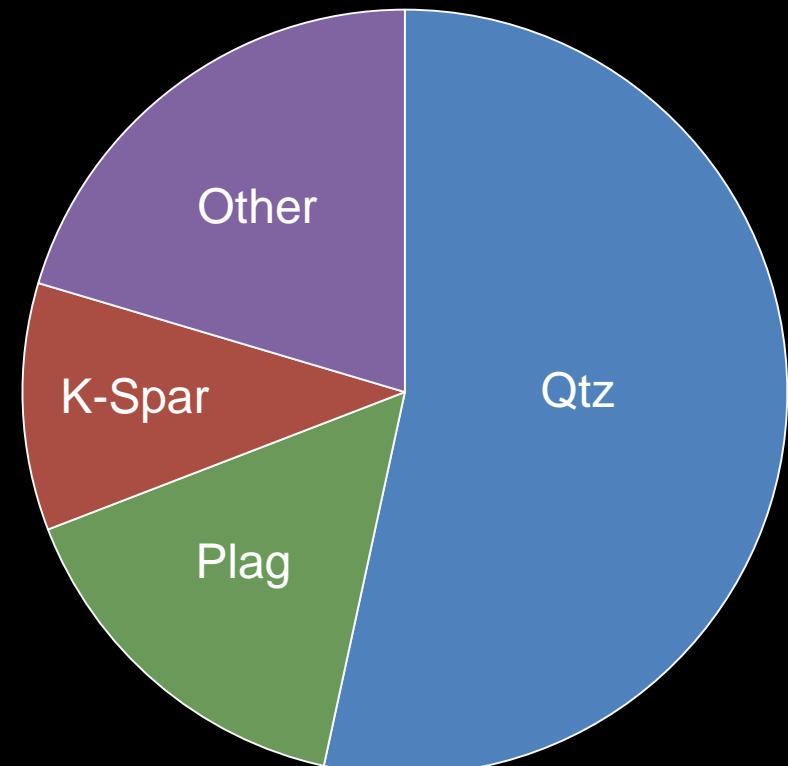


# 0.5-2mm sand mineralogy

SLW



UpB



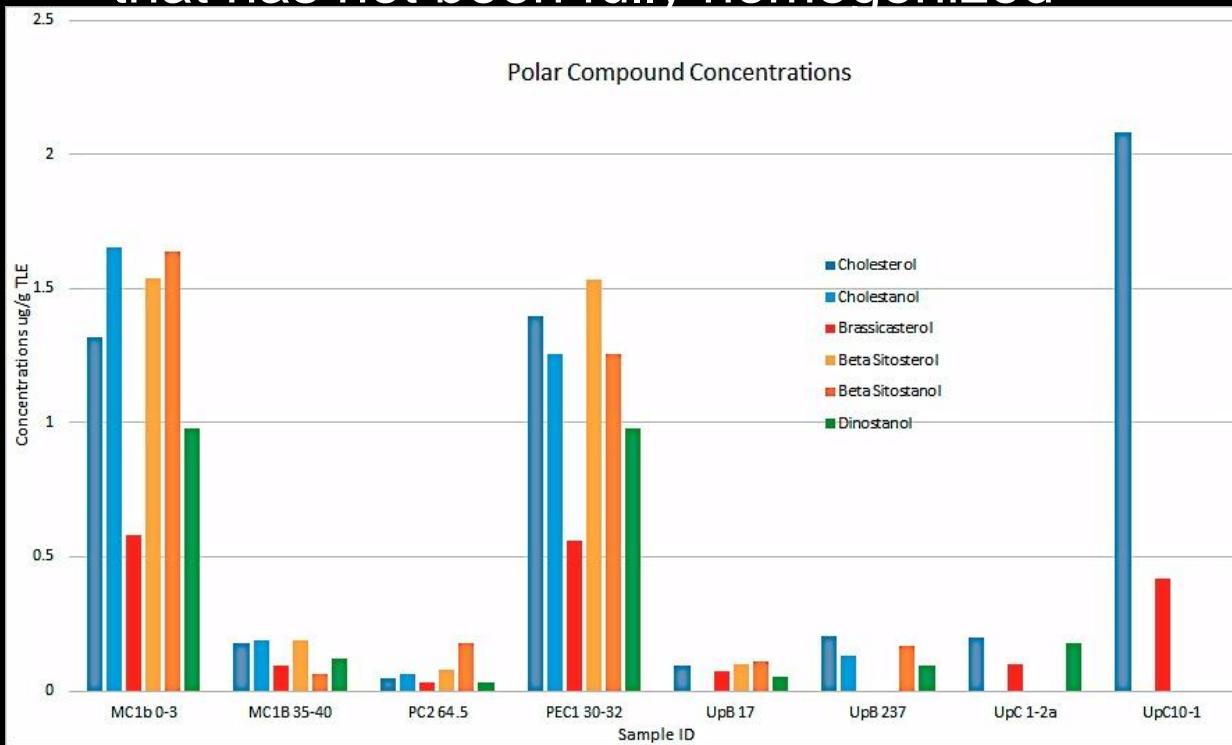
Tulaczyk et al. 1999

Clasts and sands indicate similar sources to UpB  
But there are differences

# Molecular Biomarkers

Appears to suggest

- Downstream increase in marine biomarkers
- Local source of related marine sediment that has not been fully homogenized



Jeremy Wei  
Isla Castañeda's lab

Polar compounds found in sediments:

- Brassicasterol – specific to diatoms
- Dinostanol – thought to be specific to dinoflagellates

# SLW water column geochemistry

- Consistency in chemical composition between three casts
- Nutrients N and P present in water column
- $\delta^{18}\text{O}$  indicates glacial meltwater as dominant water source
- Br/Cl indicate a diluted seawater signal

	pH	$\text{Na}^+$	$\text{K}^+$	$\text{Mg}^{2+}$	$\text{Ca}^{2+}$	$\text{F}^-$	$\text{Cl}^-$	$\text{Br}^-$	$\text{NO}_3^-$	$\text{SO}_4^{2-}$	$\text{PO}_4^{3-}$	DIC ( $\text{HCO}_3^-$ )	$\delta^{18}\text{O}$
Cast 1	8.0	5118	175	473	1034	28	3657	5.1	1.4	1228	4.5	2095	-38.0
Cast 2	8.2	5285	177	477	1020	31	3827	5.8	0.7	1255	4.3	2130	-38.0
Cast 3	8.1	5389	178	486	1023	31	3904	6.6	0.4	1272	5.5	2109	-38.1

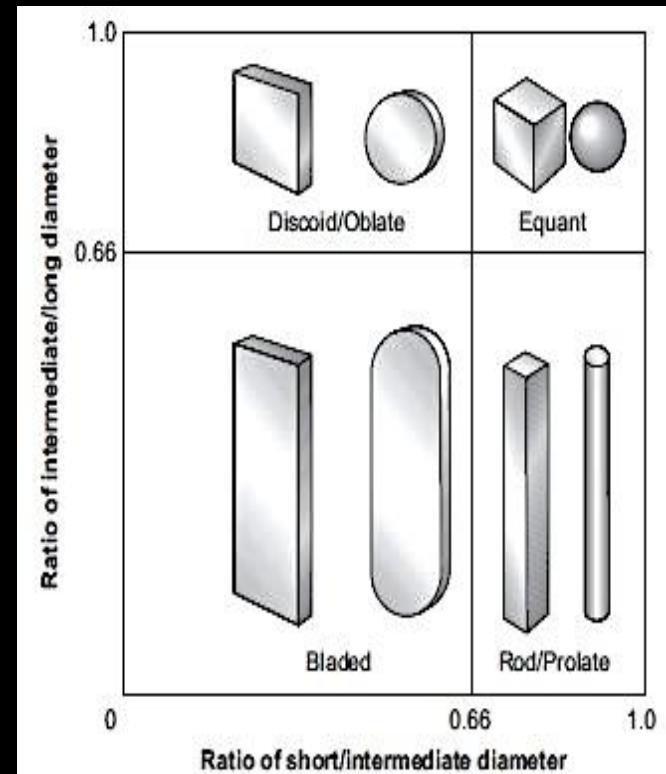
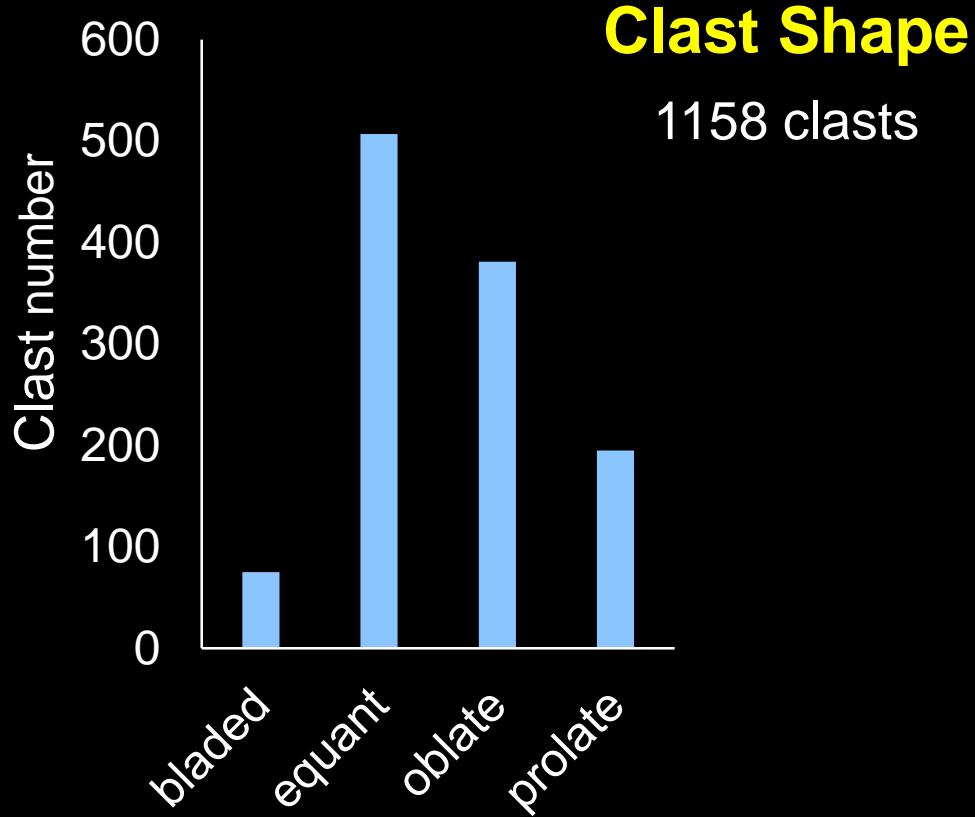
Units are ueq l<sup>-1</sup> except for  $\delta^{18}\text{O}$  and pH

Mark Skidmore

Br/Cl ratio of SLW is 0.00153, seawater is 0.00155 (Holland 1978)

3.8 mM Cl = 1/144 strength seawater

# Particle shapes and surface textures to assess subglacial processes and dynamics



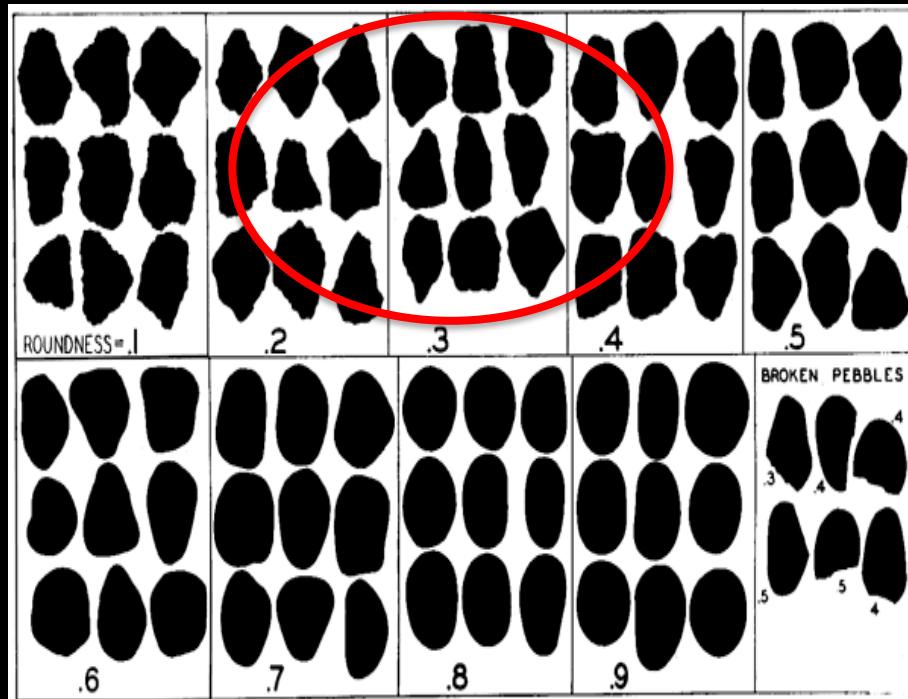
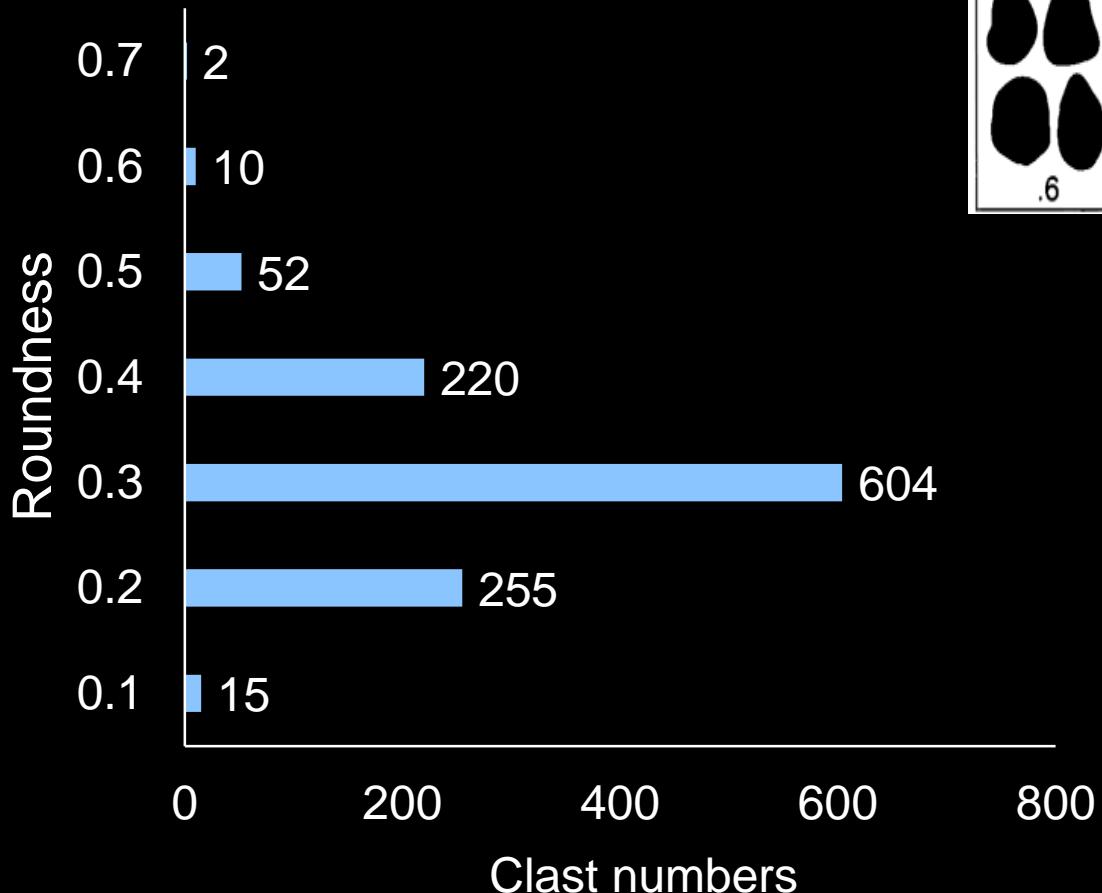
Dominant lithology – igneous & metasedimentary

Typical glacial shape for lithology

# Clasts

## angular, faceted and striated

Typical of subglacial till –  
not fluvial



Feature	Percentage of Samples
Faceted	10%
Striated	28%

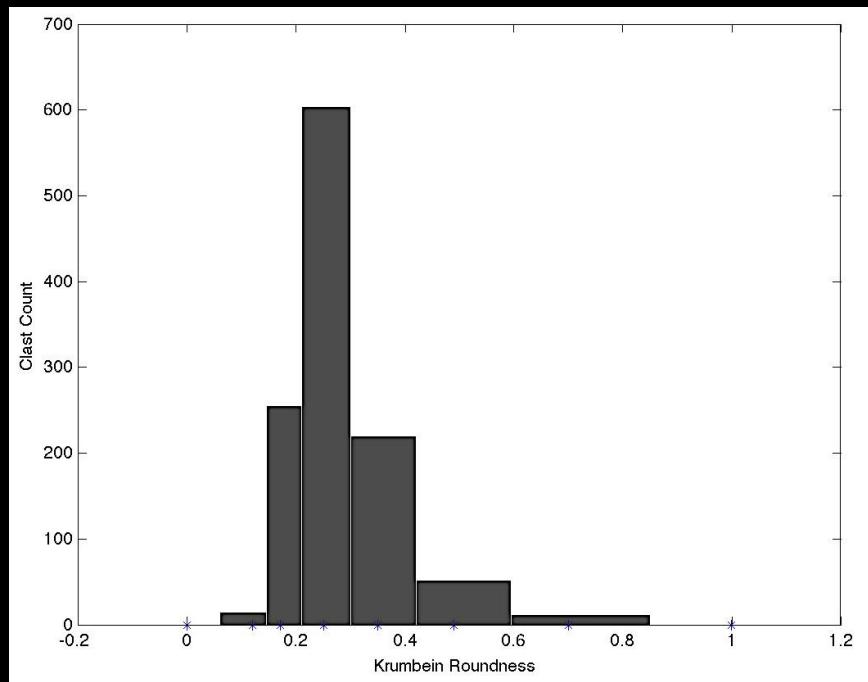
# Clast features compared

## UpB

- Striae: ~0.9% (?)
- Facets: 50%
- Roundness: subangular

## SLW

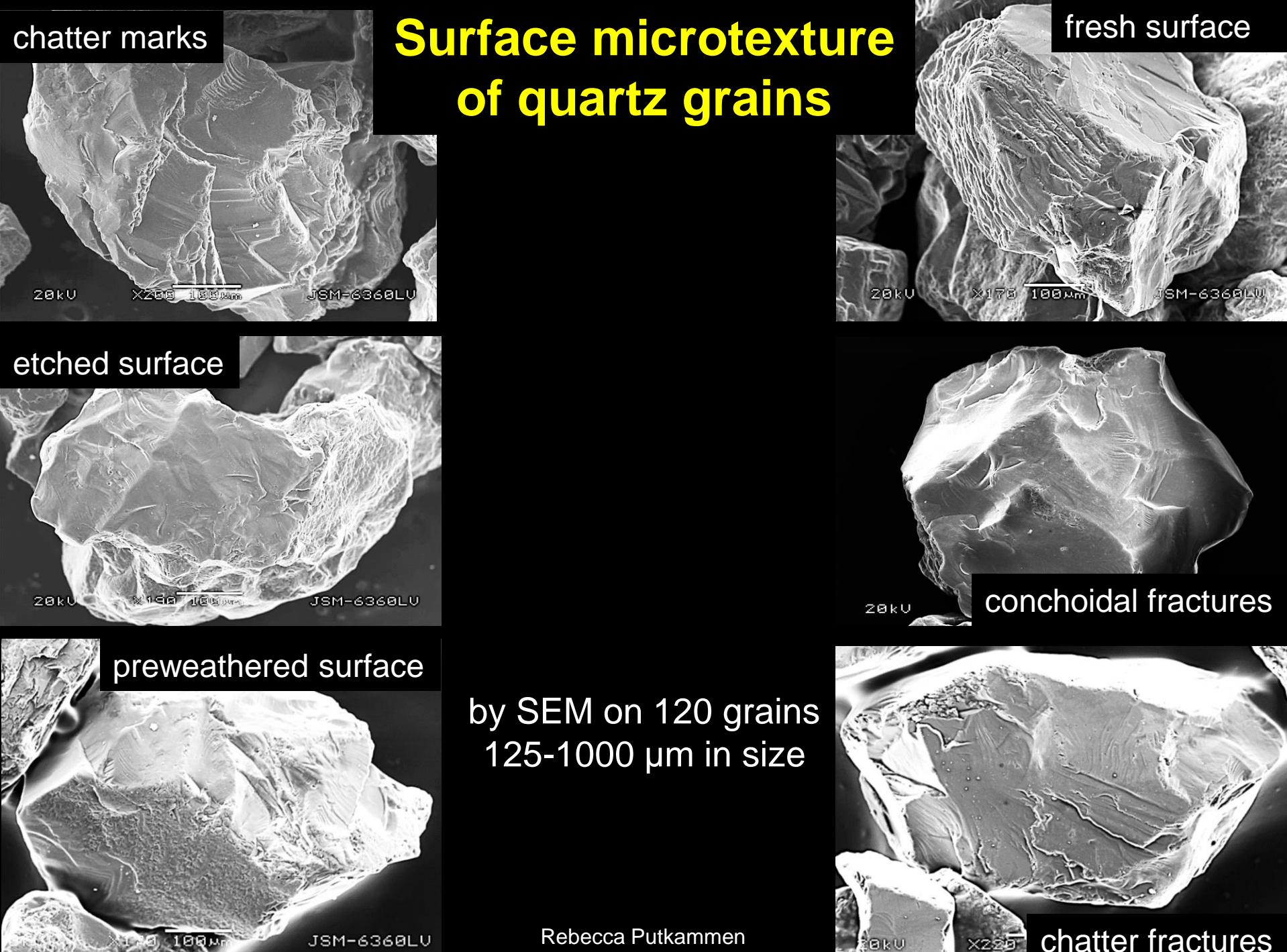
- Striae: 28%
- Facets: 10%
- Roundness: subangular



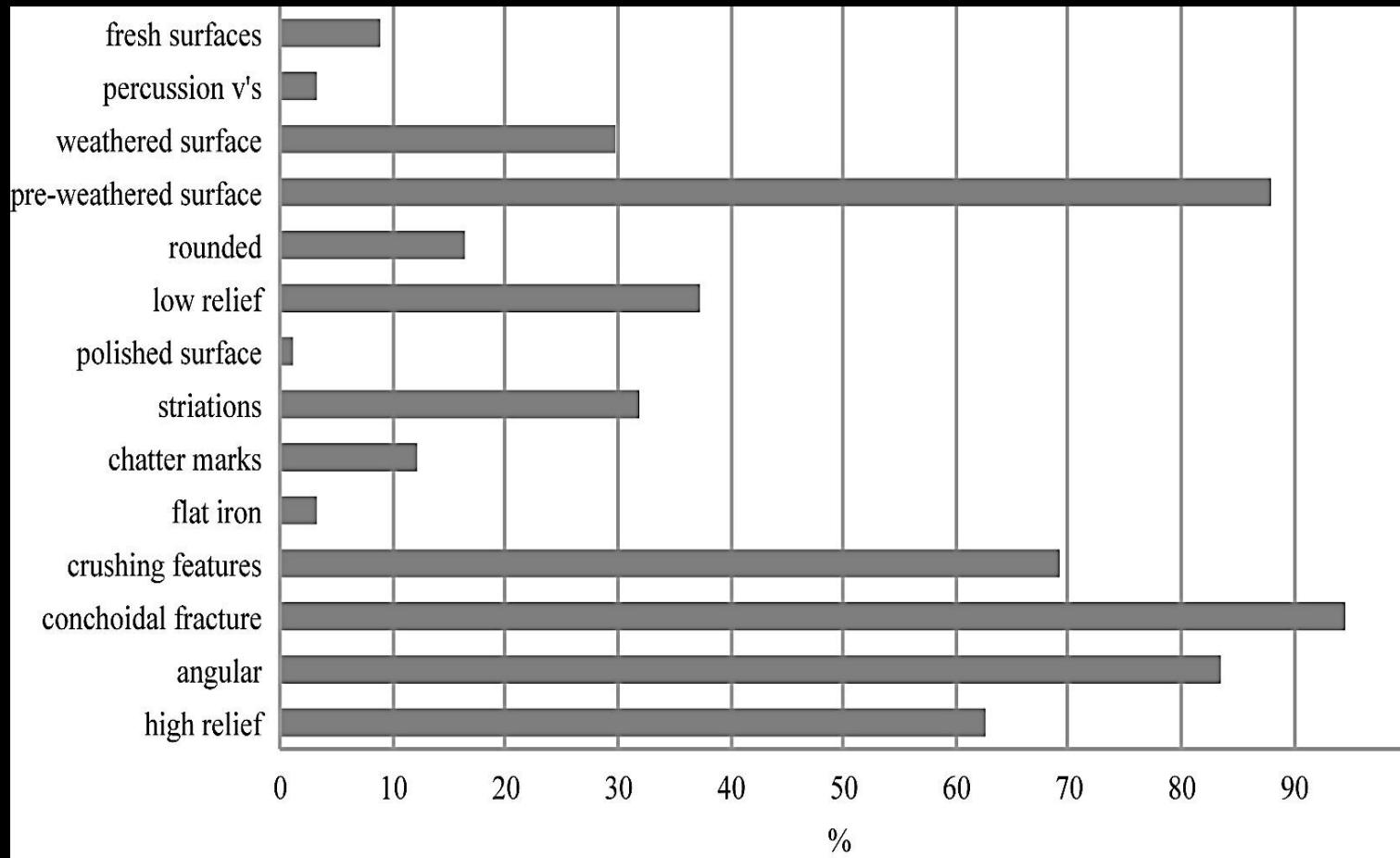
More striae but fewer facets downstream

less bedrock interaction  
more in-till contacts

# Surface microtexture of quartz grains

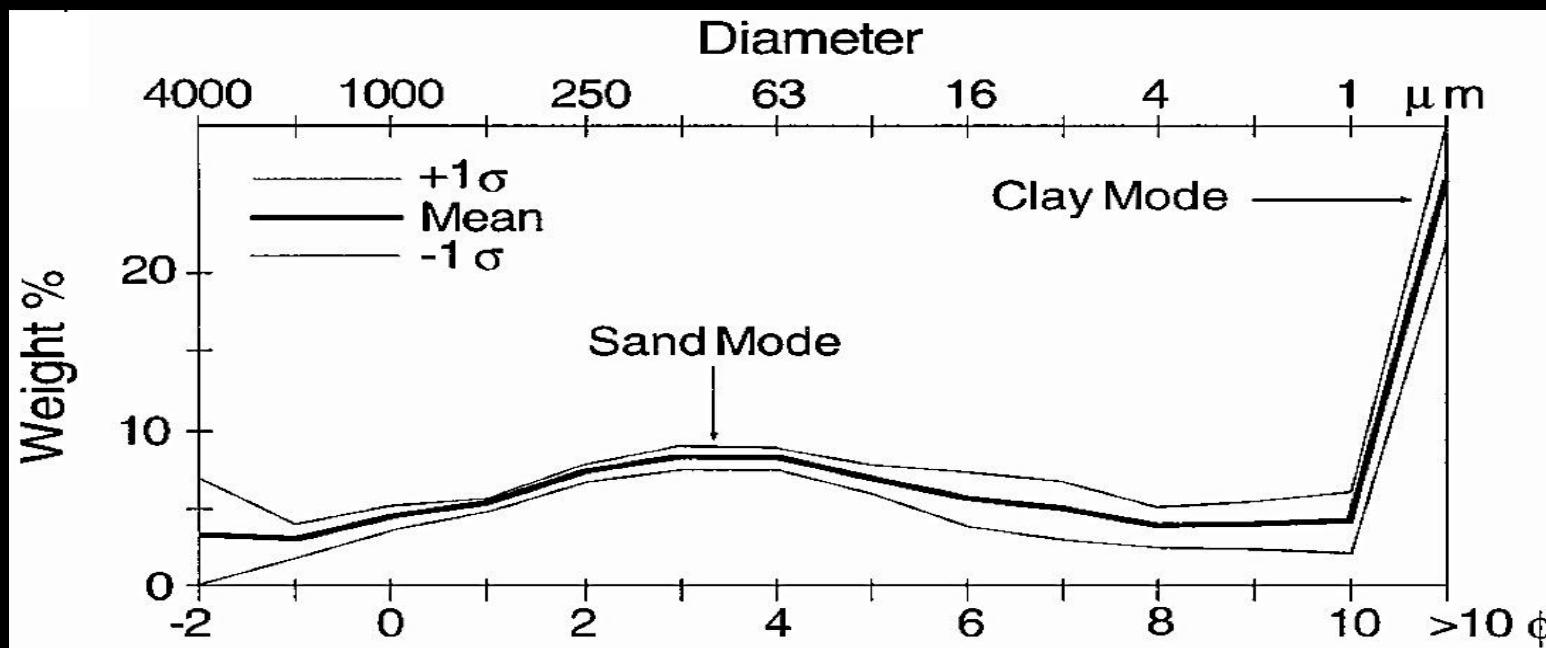
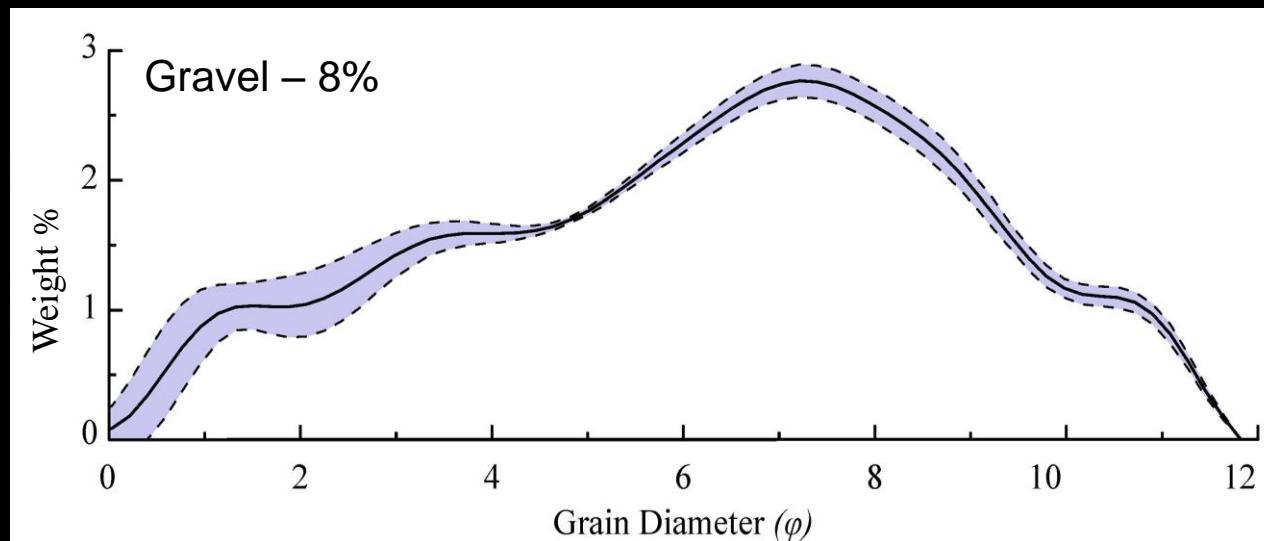


- Microtextures differ from UpB
- Infer a more complex history
  - weathering → crushing → fracture → abrasion → dissolution
  - → variable shear



# SLW till particle size distribution

Finer (silt) mode to UpB



Tulaczyk et al., 1998

# Texture of sediment infers

- no sorted sediments
- may be larger volumes of water moving around during discharge-recharge events
- but given gradients and inferred conduit size
- velocities likely to be low at  $<0.4\text{ms}^{-1}$  in 1m deep water column

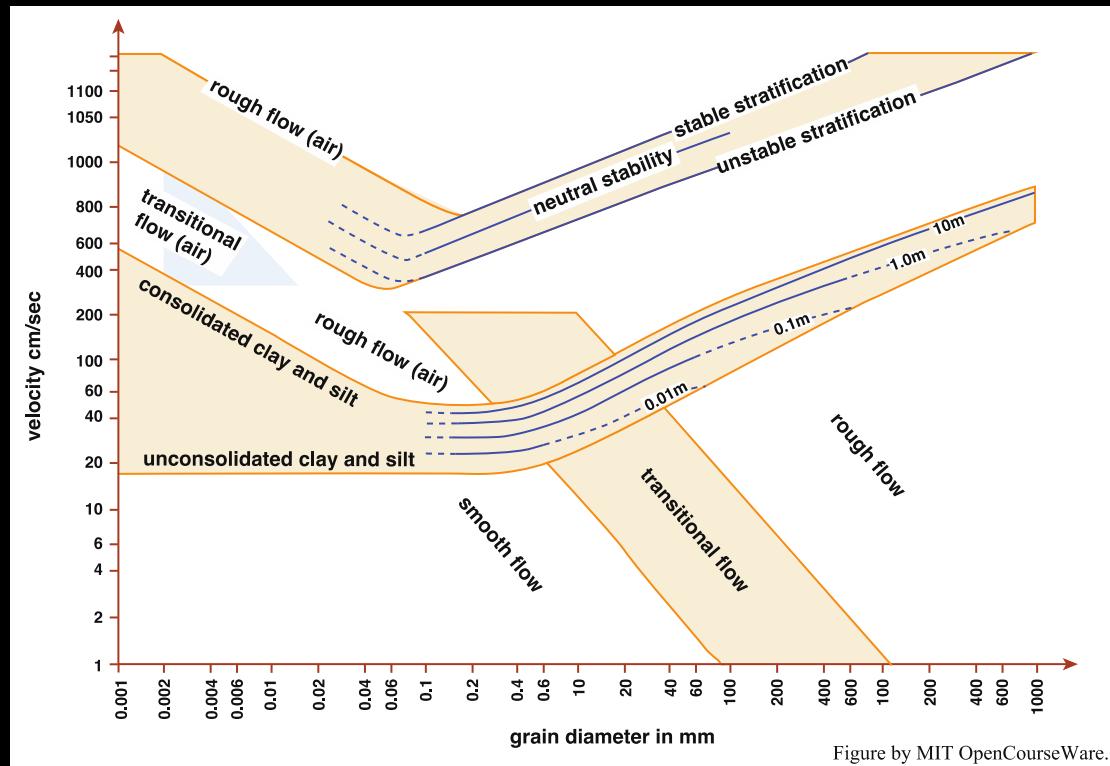
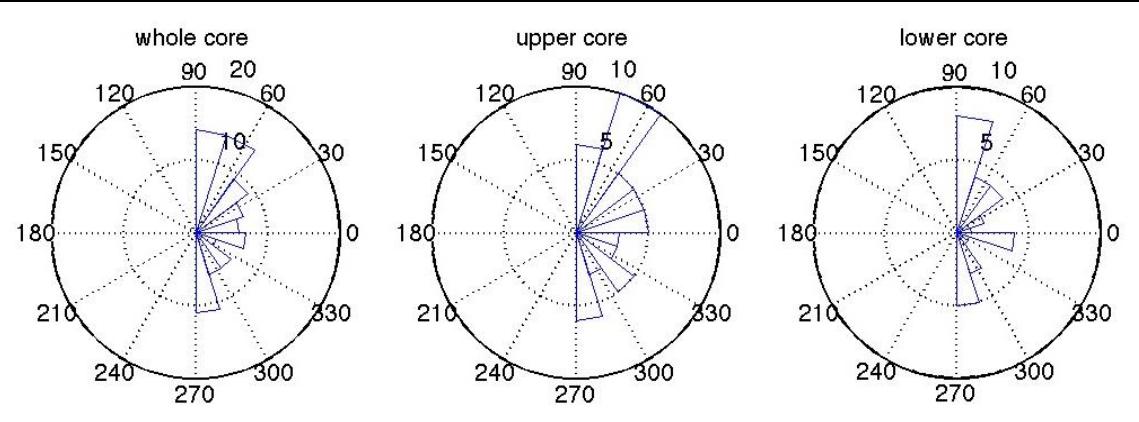
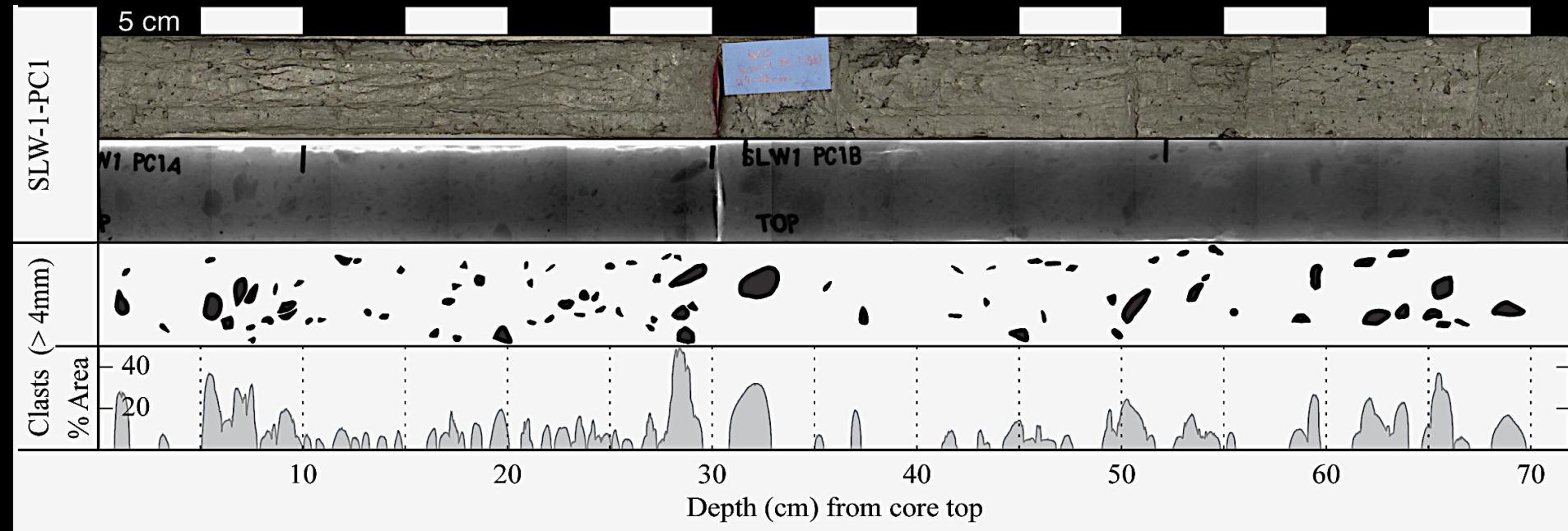


Figure by MIT OpenCourseWare.

Hjorth-Sundborg plot

# Clast fabric

Core section shows >4mm diameter clasts  
Measure apparent long-axis on 82 clasts



- weak vertical fabric
- not typical sheared till fabric
- formed with decoupling?

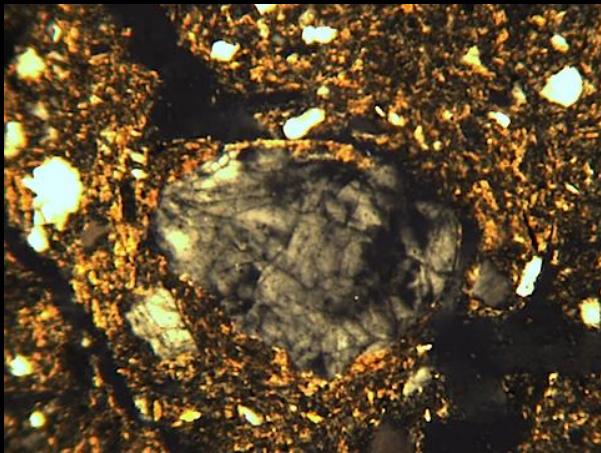
# Till microfabric

SLW1-PEC1: 23.5-31cm ↑

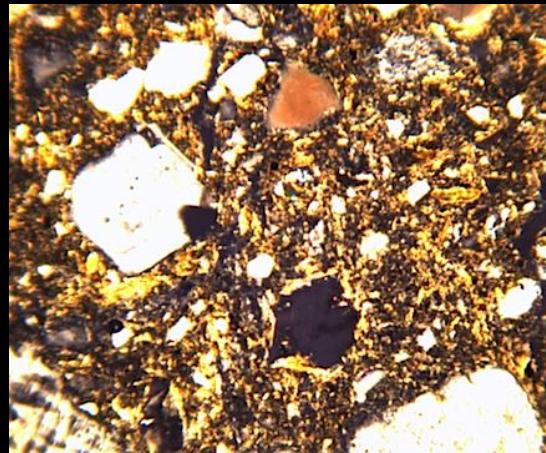
SLW1-PEC1: 23.5-31cm ↑



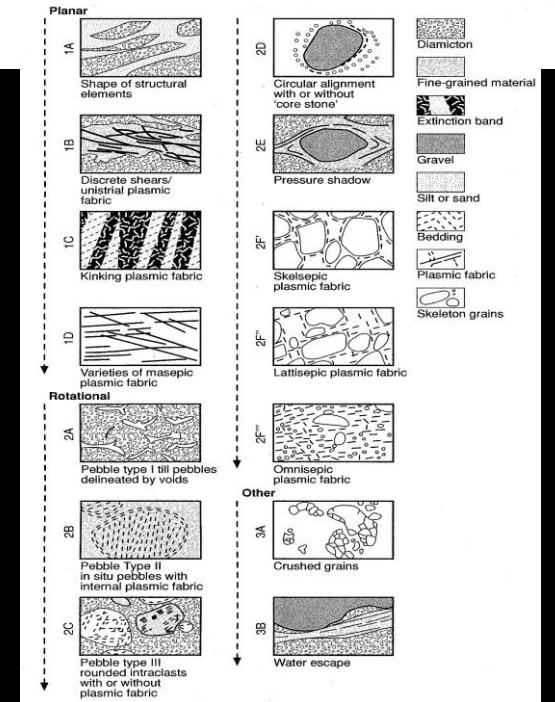
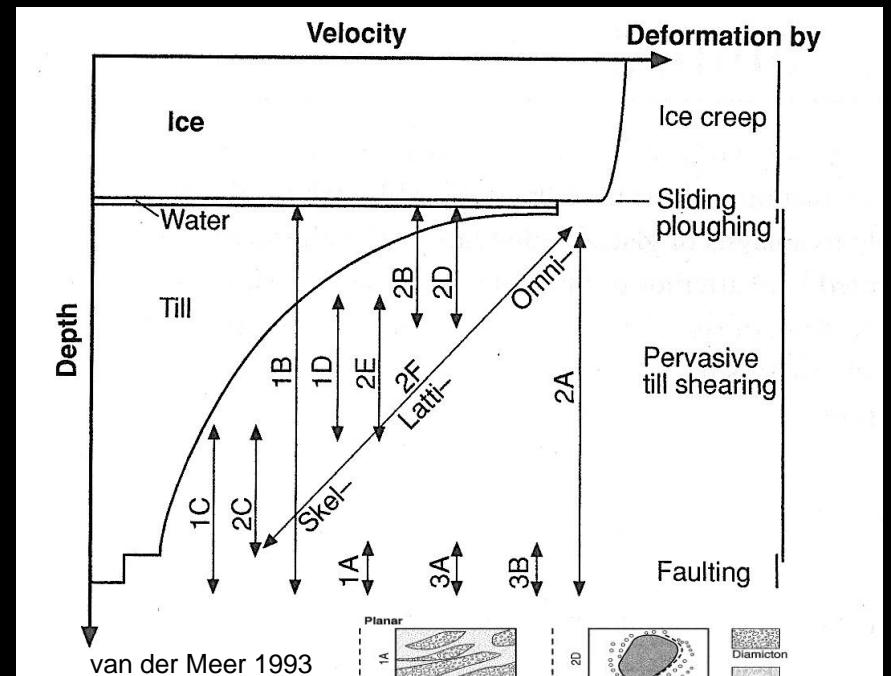
Fractures



Skelsepic

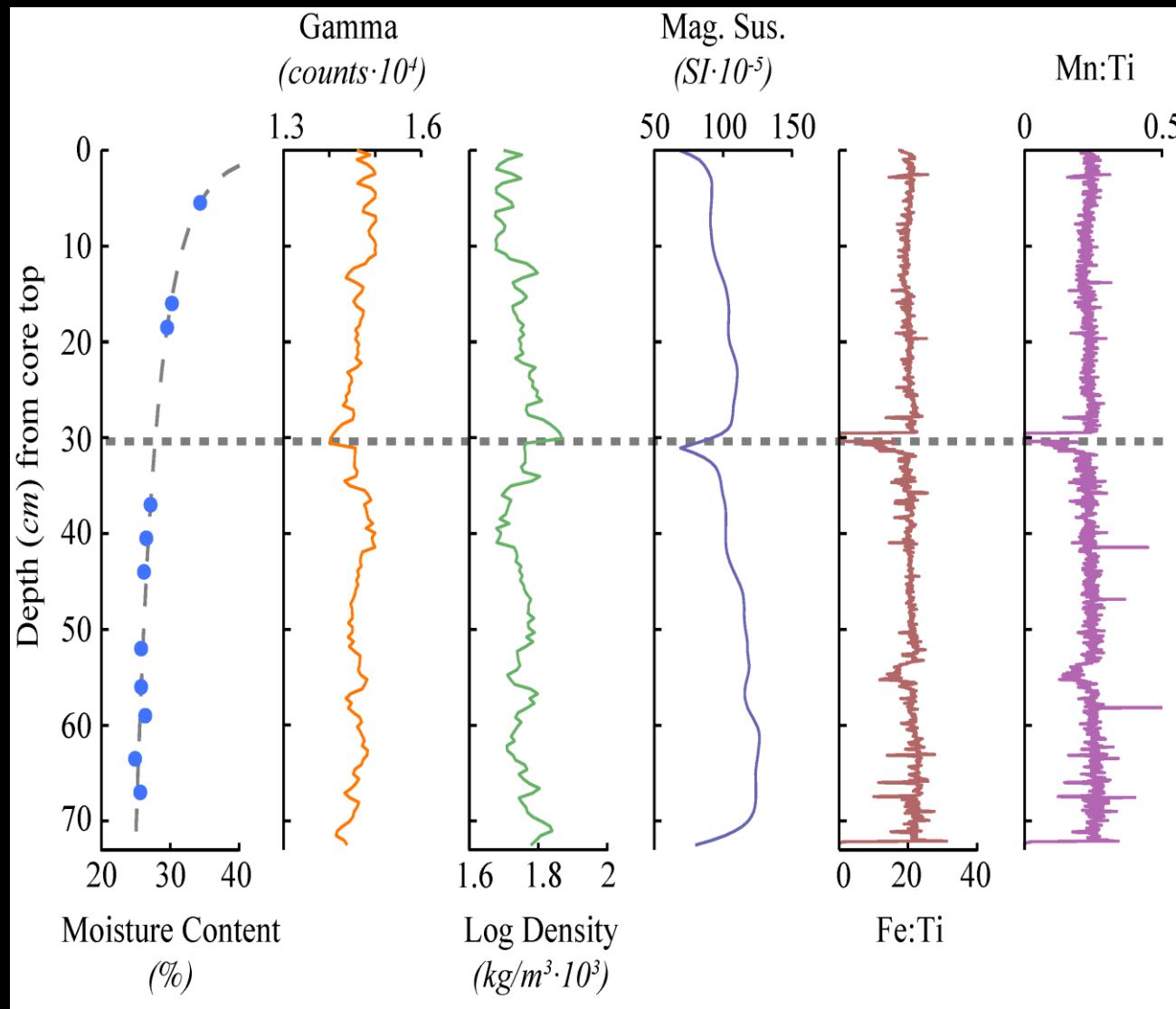


Lattispeic



# Physical properties

e.g.

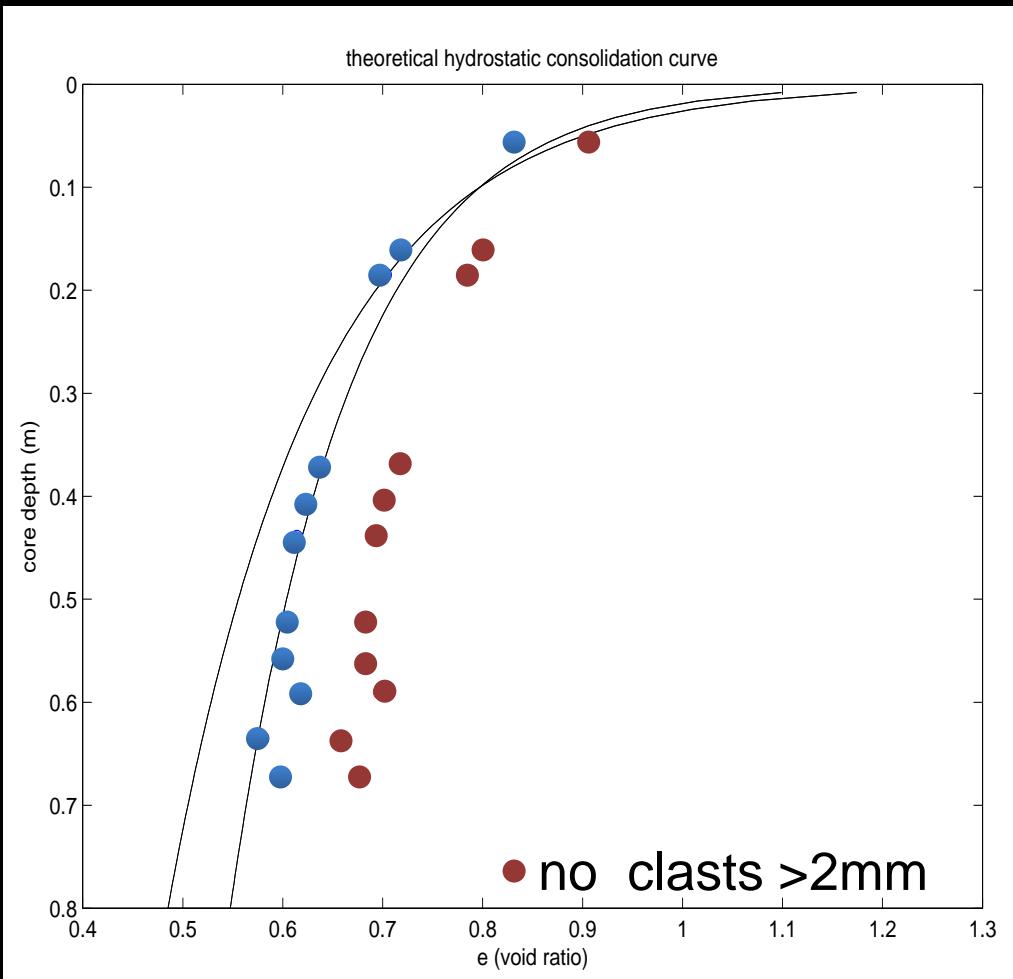


perhaps most interesting is water content  
and inferred consolidation

# Consolidation vs depth

SLW till normally consolidated

appears any loading effect from last ice touch-down during lake drainage has been compensated



For hydrostatically consolidated sediment:

$$e = e_o - C_p \log \frac{N_o}{N_{eo}}$$

e: void ratio

$e_o$ : ref. void ratio

C: compressibility

$N_o$ : effective normal pressure

$N_{eo}$ : ref. effective normal pressure

$$N_o = N_{eb} + (\rho_t - \rho_w)gz$$

$$\frac{dN_e}{dz} = (\rho_t - \rho_w)g \approx 10 \text{ kPa m}^{-1}$$

converted water content to void ratio

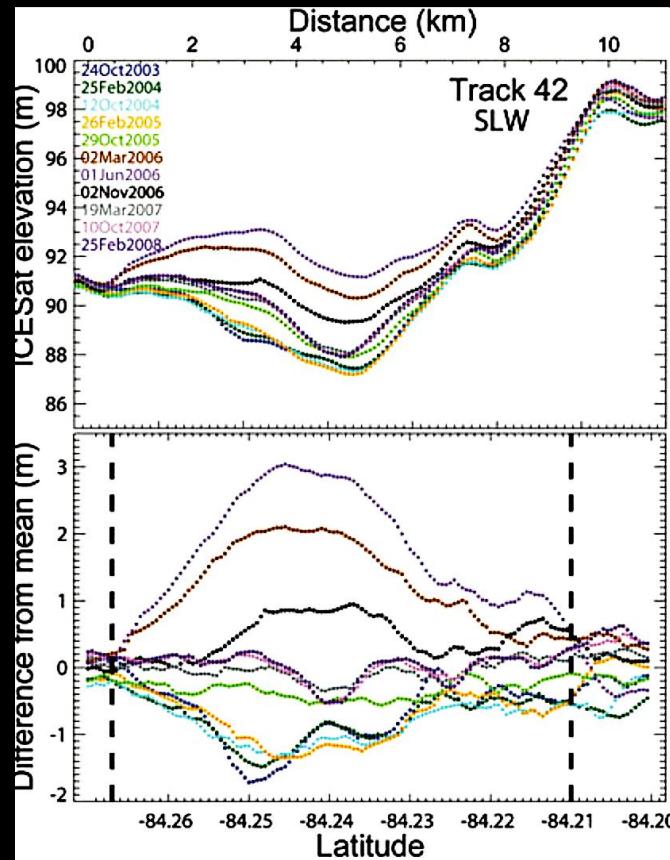
# Environmental Interpretation

No signs of sorting or lag surfaces

- quiescent conditions in the lake
- negligible deposition or erosion of lake sediment
- “floods” not typical floods – flow only at cm/s due to low surface gradients and wide conduits

But during ‘lowstands’ ice recouples with bed

- deforming new till into the lake
- mixes any older thin lake sediment into till



# Conclusions

## SLW sediment

- homogeneous, structureless, clast poor, muddy till
- water saturated to compacted

## Sediment sources

- minor differences with those of UpB
- at least a local marine component

## Clasts, grains and microfabric

- a more complex transport history than UpB
- strong (but variable) glacial shear and recent dissolution

## Lake discharge-recharge at low velocity ( $<0.4\text{ms}^{-1}$ )

- floods move across WIP but no sign of fluvial erosion or transport
- flow unlikely via persistent conduits, most likely in braided sheets

## At some lowstands ice recouples with lake bed (last 2004?)

- till deformed into lake basin
- then re-floats and unloads
- common for lakes on ice plains but unlikely in deeper interior lakes

## Till appears normally consolidated

- loading effect from last lake drainage now compensated

## Clasts have weak vertical fabric

- formed with decoupling during lake refilling

