

# **Validation of fault displacements from dynamic rupture simulations against the observations from the 1992 Landers earthquake**

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## **ABSTRACT**

Coseismic fault displacements in large earthquakes have caused significant damage to structures and lifelines on and near fault lines. Fault displacements, therefore, represent a real threat, especially to distributed infrastructure systems. For the infrastructures that are unable to fully avoid near active faults, displacement demands used in engineering design are defined using probabilistic fault displacement hazard analyses (PFDHA). But PFDHA models are sparse and poorly constrained partly due to the scarcity of detailed fault displacement observations. Advancements in dynamic rupture simulation methods make them an attractive approach to address this important issue. Because fault displacements can be simulated for various geologic conditions as constrained by current knowledge about earthquake processes, they can be used to supplement the observation datasets, providing on-fault displacements, and when used with appropriate constitutive models for the bulk medium, they can capture off-fault inelastic deformations as well. For viable extrapolation, simulations must first be validated against data. In this paper, we summarize the calibration and validation of the dynamic rupture model against the observations of the well-documented 1992 Landers earthquake. We defined a preferred model that reproduces several first-order fault displacement metrics (such as the on-fault partition of the total displacement, the mean fault zone width and the location of the peak displacement). Simulated ground motions consistent with the observations ensure all physics important to modeling have been properly parameterized. For the extrapolation, we generated a suite of dynamic rupture models to quantify expected fault-displacement metrics, their intercorrelations and magnitude dependencies, which are in part supported by the Landers and other recent earthquakes. This validation/extrapolation exercise paves the way for using dynamic rupture modeling to quantitatively address fault displacement hazard on a broader scale. The results are promising and are expected to be useful to inform PFDHA model development.

## **INTRODUCTION**

Coseismic fault displacements in large earthquakes have caused significant damage to structures and lifelines located on or near fault lines. Examples include severely damaged bridges, dams, and tunnels in the 1999 Chi-Chi earthquake, pancaked buildings in the 1999 Kocaeli earthquake, distorted railroads in the 1975 Guatemala earthquake, and ruptured water lines during the 2019 Ridgecrest Earthquake Sequence. Fault displacements, therefore, represent a real threat, especially to distributed infrastructure systems that can be intersected by faults in more than one location. This seismic hazard has been addressed by recommendations and

regulations restricting construction near active fault zones. An example is the Alquist-Priolo Earthquake Fault Zoning Act (AP Act, Appendix A of CGS 2018) in the state of California, a law enacted following the 1971 San Fernando earthquake. The AP Act requires, for example, that a new construction used for human occupancy be more than 30 m away from a mapped fault trace or if within that zone, requires a comprehensive site-specific investigation be performed. Similar legislations have been implemented in Japan, New Zealand and China as well. However, there are situations for which the avoidance of the fault zone is not possible or not practical (e.g., gas and water supply and distribution networks, roads and certain dams). Thus, guidance on the potential fault displacement hazard is needed to inform engineering design.

For ground motions hazard, the design is informed by a combination of deterministic and probabilistic seismic hazard analyses (DSHA, PSHA). SHA is well established for ground motion hazard, but its counterpart for fault displacement probabilistic hazard analysis (PFDHA) is not as mature. The PFDHA approach gained broader interest during the assessment of the Yucca Mountain nuclear waste repository site (Stepp *et al.*, 2001; Youngs *et al.*, 2003). It has since been extended to include more fault displacement data from various fault mechanisms and revised uncertainty quantification approaches (Moss and Ross, 2011; Petersen *et al.*, 2011; Chen and Petersen, 2019; Nurminen *et al.*, 2020). However, compared to the well-developed empirical ground-motion models (GMMs, a.k.a. GMPEs), empirical fault displacement models are sparse and poorly constrained partly due to the paucity of detailed high-resolution fault-displacement observations, despite recent large database building efforts (e.g., Sarmiento *et al.*, 2019; Baize *et al.*, 2020). The reasons for this paucity of data include (1) a low probability of earthquake ruptures that reach the surface (tend to be from a subset of large magnitude events, and large magnitude events are rare), (2) a long recurrence period between co-located surface ruptured earthquakes (~ hundreds of years) that doesn't allow the comparison of potentially similar events, (3) a technical difficulty of exhaustively collecting perishable fault displacements over very large areas. These data limitations make the use of earthquake rupture simulations an attractive solution to support PFDHA modeling as part of the Fault Displacement Hazard Initiative (FDHI, Bozorgnia *et al.*, 2021).

Before simulation models can be used, they must first be validated against observations for which detailed datasets exist. The validated models can then be used to generate scenarios and simulated datasets for events we haven't yet observed. A similar approach was used to support the development of GMMs. For example, Day *et al.* (2008) simulated the effect of sedimentary basin on long-period response spectra, Donahue and Abrahamson (2014) developed a hanging-wall effect model based on simulations performed with the Southern California Earthquake Center (SCEC) Broadband Platform (BBP, Maechling *et al.*, 2015), and Goulet *et al.* (2021) used BBP simulations to model the magnitude scaling for the central and eastern U.S. Results from these simulations contributed to the Next Generation Attenuation (NGA) Project GMMs.

In this study, the physics-based model used is referred to as the dynamic rupture model; it constructs spontaneously evolving earthquake ruptures under mechanical causative conditions (e.g., fault geometry, friction law, stress condition and surrounding rock properties). It makes it ideal for the simulation of displacements, which vary along the fault plane due to the dynamic response to imposed stresses. Our project involves the application of the dynamic rupture model to California (CA) events, its validation against well-documented case histories, and its use to

predict displacements for scenarios and events we have not yet experienced. Those synthetic displacement datasets can in turn be used alone or combined with empirical observations to support PFDHA model developments.

In this paper, we present the validation and calibration of our method against the 1992 **M** 7.3 Landers earthquake. We selected this event for the first validation exercise due to the high quality and relative completeness of its displacement and ground motion data. The Landers earthquake is the third-largest California event of the twentieth century. It occurred within a roughly 70 km long zone of NNW-trending dextral faults accommodated by the Eastern California Shear Zone (Figure 1). The Landers earthquake was initiated near the southern end and propagated unilaterally toward the northern end. Field mapping by geologists revealed particularly complex patterns of surface rupture (Figure 1, Liu *et al.*, 2003). Comparison of the surface rupture with previously mapped fault traces (pre-rupture mapping, Figure 1), shows that the Landers event spanned two (previously) disconnected groups of faults with the previously unknown Kickapoo Fault (indicated by an arrow in Figure 1). This pre/post-event comparison of surface expression highlights the complexity and the difficulty in predicting future displacements from pre-rupture fault mapping (in this case, as defined in the AP Act at the time). It also highlights the motivation behind our choice of geometry for the validation, which is a simplified model, similar to what is documented in the AP Act. In that sense, the validation is performed with the end goal in mind of not reproducing every single displacement observed in Landers, but to capture the general displacement characteristics observed, using an ensemble of simulation results. This is consistent with how PFDHA will be used, that is, based on a scenario defined by a magnitude and geometric fault information. This is also consistent with how GMMs are developed and used, based on scenario parameters.

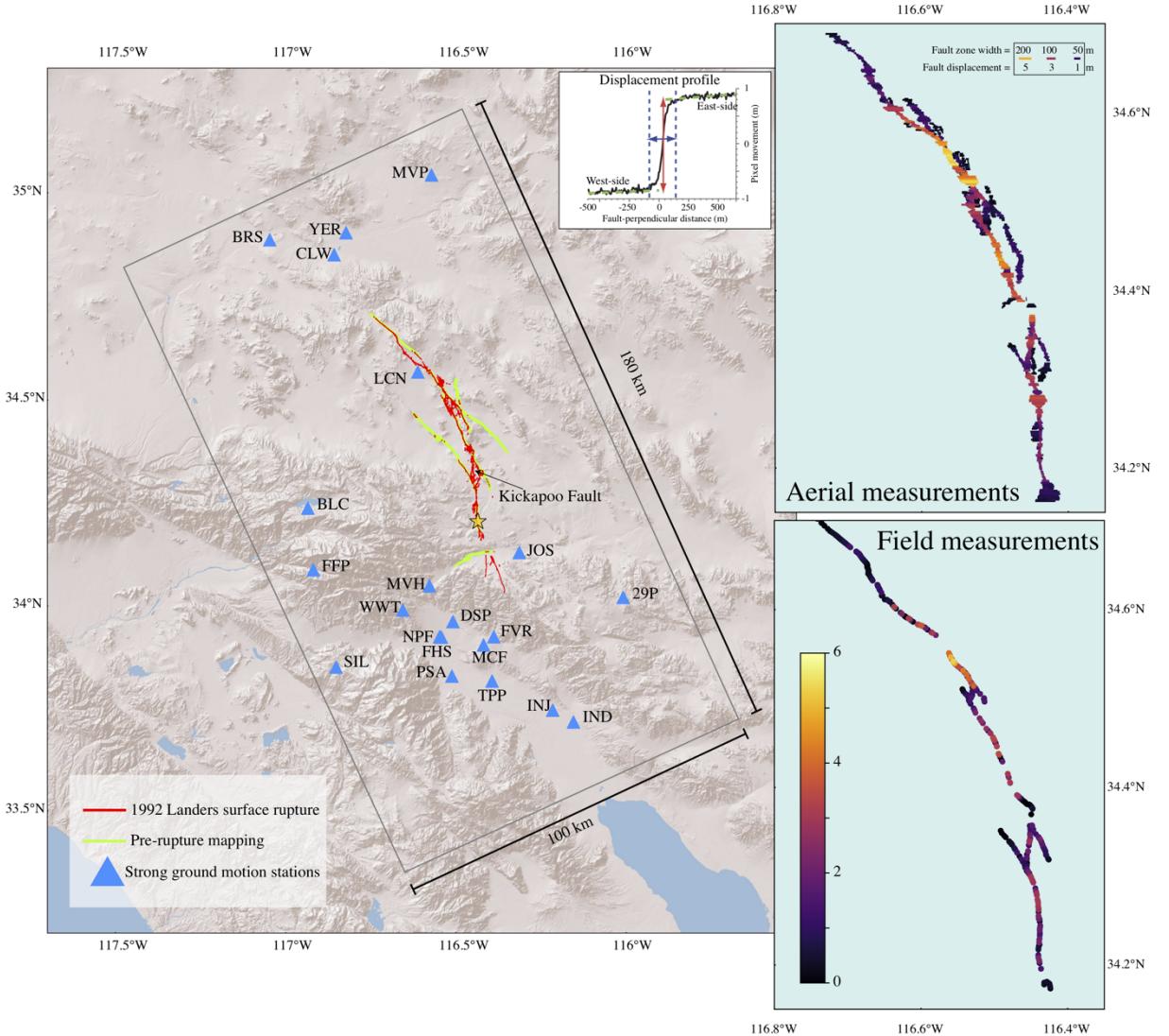


Figure 1. Map of the 1992 Landers earthquake source region. The red lines in the left geographic map indicate the surface rupture strands involved in Landers and the green ones indicate the pre-rupture mapping in this region. Both datasets are compiled into the FDHI Database (Sarmiento *et al.*, 2019). The black arrow points out the location of the Kickapoo Fault. Blue triangles are seismic stations at which ground motions were recorded. The upper small panel shows an example of fault displacements interpreted from aerial measurements (Milliner *et al.*, 2015). The two right panels illustrate the fault displacements from aerial (top) and field (bottom) measurement (Sarmiento *et al.*, 2019). The fault zone width (FZW) is also shown in the right upper panel.

In addition to the detailed field geological mappings (e.g., Liu *et al.*, 2003), a high-resolution fault displacement profile was provided through the Co-registration of Optically Sensed Images and Correlation (COSI-Corr) technique, which performs aerial measurements of surface deformation over a wider aperture than that of the field mapping (Milliner *et al.*, 2015; Milliner *et al.*, 2016). Such a technique overcomes the limitation that traditional field mapping is unable to capture the complete fault displacements at a high resolution along the entire fault

rupture region, including right along the fault strand and at distances away from it (referred to as "off-fault displacement", or OFD). The distributed off-fault displacements, while practically challenging to collect, can be important to engineering applications sensitive to strains such as high-speed rail, for example (Yoon *et al.*, 2020). Figure 1 illustrates the difference between aerial and field observations maps for Landers. The aerial measurements capture the presence of OFD and this aperture coverage leads to larger total displacement than those strictly observed on the fault by field measurements (Milliner *et al.*, 2015; Milliner *et al.*, 2016). The distributed OFD in the forms of warping, granular flow, rigid-block rotation, secondary faulting and microfracturing is common in single coseismic events and over repeated earthquake cycles (e.g., Rockwell *et al.*, 2002; Osokin *et al.*, 2007; Zinke *et al.*, 2014; Teran *et al.*, 2015). The COSI-Corr technique is also capable of providing the strain localization information by the interpretation and definition of a fault zone width (FZW) across which the fault displacement shows a gradual transition (Figure 1). The existence of OFD has a significant implication for the fault displacement model development and incorporation of the complete near-field deformation pattern will lead to a more reliable fault displacement hazard prediction. The advent of aerial datasets and COSI-Corr are now enabling the integration of such strain fields into PFDHA models.

In this study, we develop a dynamic rupture model for Landers and validate the fault displacements against the available empirical datasets. Although our focus is on displacement, we also perform a validation of the ground motions resulting from that model to confirm that the physics has been appropriately modeled and parameterized. The validation step involves sensitivity analyses and a calibration of the model parameters to be used to capture at least some of the modeling uncertainties. With the intent of later generating simulations for CA events not yet observed, we develop a suite of dynamic rupture models (varied magnitudes) by relaxing the fault small-scale geometry and stress condition, introducing a certain level of extrapolation of the Landers-based models. The dynamic rupture model was recently successfully applied to reproduce the observed fault-displacement profile of the 2010 M 7.0 Darfield earthquake (Dalguer *et al.*, 2019), a notably complex fault. In contrast, our simulations are designed to be general enough in their parameterization and not detailed to the point that the results apply to a single specific fault modeled in detail, so as to supplement the empirical datasets and provide guidance for PFDHA modeling. In addition, this type of simulation is computationally demanding and provided that several suites of large-scale simulations are to be performed, simplifications in modeling are necessary (as described below) to make the scope of work tractable.

## DYNAMIC RUPTURE MODEL SETUP FOR THE LANDERS SCENARIO

### Simulations in support of PFDHA modeling

The motivation for our project is to support PFDHA modeling by supplementing the limited empirical displacement dataset with simulation data. The forward simulations will need to cover, for example, different fault geometries, magnitudes and states of stress that are representative of expected future events in CA, our focus region. To be useful in informing PFDHA modeling scaling, the simulations must be specific enough to capture realistic fault ruptures while being general enough to capture the key features of the displacement field, without being fault specific. Our modeling strategy to develop a suite of models and simulations that together will capture the expected displacement of future events. Before the simulations are used by PFDHA modelers, the underlying models must be validated against empirical

observations. By performing the validation in a manner consistent with how the forward simulations will be generated (general ensemble models), we ensure a consistency between the validated models and the forward simulations. That is, the validation goal is not to replicate every displacement detail from the Landers event, but to make sure that a calibrated ensemble of simulations can capture the trends in observations. This type of validation strategy was successfully used for ground motion validation of the SCEC BBP (Goulet *et al.* 2015, Dreger *et al.* 2015). In the forward sense, the PFDHA models will require simple input parameters such as  $\mathbf{M}$ , fault dimensions and dip, fault mechanism, so our goal is to generate suites of events that span these attributes.

The faults that ruptured during the Landers earthquake are very complex: they consist of curved, branched, and segmented non-planar faults (Figure 1). While similarly complex fault systems have been modeled by dynamic rupture (Dalguer *et al.*, 2019; Aochi *et al.*, 2003; Tago *et al.*, 2012; Wollherr *et al.*, 2019), such that they capture the complexity of the real event, they remain event-specific. Our application requires modeling simplifications to meet the goals described above and also to be computationally tractable. In the following subsections, we describe our modeling choices to meet those goals by focusing on the dynamic rupture for simulating the 1992 Landers earthquake. To build a spontaneously-evolving dynamic rupture model, a number of causative physical modeling choices and assumptions are required that can be summarized in the following general categories, of so-called “ingredients” (Harris *et al.*, 2009): 1) fault geometry and initial stress condition, 2) fault failure criterion (fault constitutive law that specifies the condition under which the fault is allowed to slip) and, 3) surrounding material (rock) properties. To develop dynamic rupture models that are readily extendable to other scenarios and events, the models are based on relatively simple physics models and parameterized following general criteria loosely informed by past Landers simulation studies, as described below. The choice of modeling assumptions and the parameter values were selected following suites of sensitivity analyses that are not described in this paper. We define a preferred model, selected to be most representative of the as-observed Landers event, and a suite of ensemble models representing plausible earthquakes of varying magnitude along the same base problem setting (fault geometry and stresses),

## Fault geometry and initial stress condition

Consistent with our simulation philosophy, we first define a simplified geometry for the Landers event, using the rough pre-rupture mapped fault traces as a guide. The three-segmented connected fault was originally developed by Wald and Heaton (1994) and followed by most seismologists (e.g., Cohee and Beroza, 1994; Olsen *et al.*, 1997). Our simplified model (Figure 2a) consists of three linear segments loosely fitting the following faults: the Landers Johnson Valley Segment (LJVS) in the south where the rupture was nucleated, the Homestead Valley Segment (HVS) in the central part of the fault, and the Camp Rock Emerson Valley Segment (CREVS) in the north (thick black line in Figure 2a). Their strike azimuths are  $355^\circ$ ,  $337^\circ$  and  $315^\circ$  from the south to the north, respectively, leading to a fault length of 84 km with two kinks roughly at third-lengths. The strike-slip fault is modeled from the surface to a depth of 17.5 km (Figure 2b).

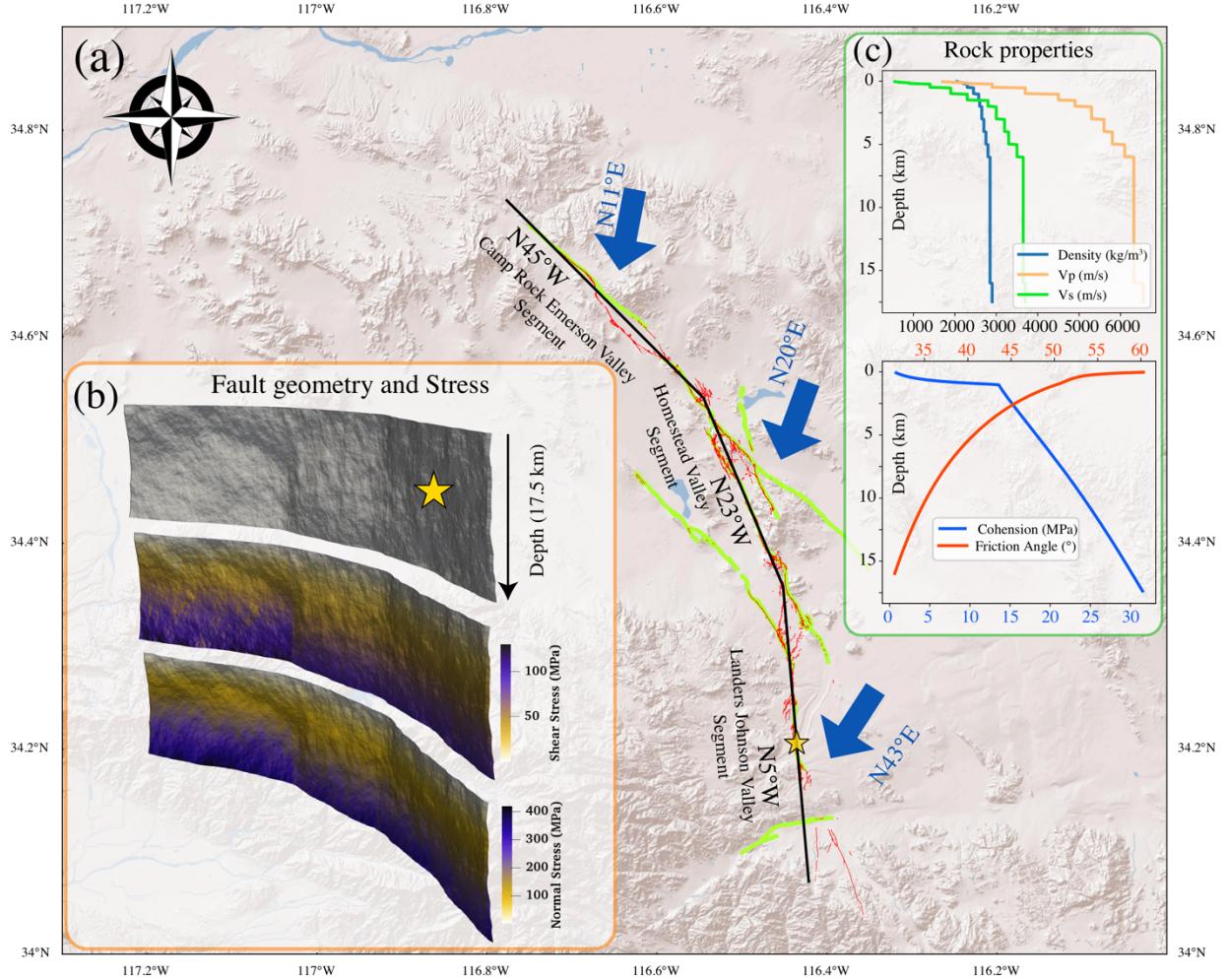


Figure 2. Simulation model setup. (a) Landers surface rupture strands (thin red lines), pre-rupture mapping (green lines) and the simplified 3-segment fault line defined for the simulation (thick black line). Blue arrows indicate the orientations of the horizontal maximum principal stresses on different segments. (b) The lower left panel shows the fault plane model which includes small-scale fault roughness, with shear and normal stress projected on the resulting complex fault surface. (c) The upper right panel shows the rock properties, including depth variable density, P and S velocities, cohesion and friction angle, which are used for the dynamic rupture simulation.

To make our fault plane more realistic, we superimposed a geometrically self-similar roughness on the original planar geometry (Figure 2b). This effectively improves the along-strike variability to be more consistent with observed data, as recommended by Bruhat *et al.* (2020). The fault roughness level is defined by the amplitude-to-wavelength ratio  $\alpha$  (Equation A1 in Shi and Day, 2013) for which a value between  $10^{-3}$  and  $10^{-2}$  is proposed to correlate with the fault maturity (Power and Tullis, 1991; Sagy and Brodsky, 2009) as wear process from repeated slip decreases  $\alpha$  in mature faults. Fault maturity can be inferred from fault structural complexities, coseismic rupture speed and the ratio of deep slip to surface slip (Dolan and Haravitch, 2014; Perrin *et al.*, 2016). It is noted that a quantitative relationship between fault roughness level and fault maturity has not been established to date due to lack of data and relies on a qualitative assessment. Based on documentation of the subshear rupture velocity, shallow slip deficit and

appreciable structural complexity from the 1992 Landers earthquake (Sieh *et al.*, 1993; Wald and Heaton, 1994; Fialko, 2004; Dolan and Haravitch, 2014), we set  $\alpha=0.05$  to represent a fault with a relatively low maturity. In contrast, a more structurally mature fault such as the single prominent fracture surface of the Punchbowl Fault was found with a value of  $\alpha=0.001$  based on laboratory tests from field samples (Chester *et al.*, 2004).

To construct the initial stress condition, we essentially follow Wollherr *et al.* (2019) and adapt their stress computation to our own geometry. The vertical principal stress is defined as an intermediate principal stress  $\sigma_2$  that is an average of the horizontal maximum and minimum principal stresses (i.e.,  $\sigma_2=(\sigma_1+\sigma_3)/2$ ). It varies linearly with depth as the difference between lithostatic pressure and hydrostatic pore fluid pressure:

$$\sigma_2(h) = \int_0^h \rho(z)gdz - \rho_w gh \quad (1)$$

with gravity  $g = 9.81 \text{ m/s}^2$ , depth-dependent rock density  $\rho(z)$ , water density  $\rho_w$  and depth  $h$ . The difference between  $\sigma_1$  and  $\sigma_2$  ( $\Delta\sigma=\sigma_1-\sigma_2$ ) corresponds to the radius of the horizontal Mohr-Coulomb stress circle, which implies:

$$\tau^s = \sin(2\theta) \tau^n = \sigma_2 - \cos(2\theta) \quad (2)$$

where  $\theta$  is the angle of the horizontal maximum principal stress with the fault plane and  $\tau^s$  and  $\tau^n$  are shear and normal stresses projected on the fault plane. A relative prestress ratio  $R$  is defined as the fault stress drop over the fault strength drop:

$$R = \frac{\tau^s - \mu_d \tau^n}{\mu_s \tau^n - \mu_d \tau^n} \quad (3)$$

where  $\mu_s$  and  $\mu_d$  are static and dynamic friction coefficient from the slip-weakening friction law we adopted in this study (next section). By combining Equations (2) and (3),  $\Delta\sigma$  is obtained as a function of  $R$  and  $\sigma_1$  and  $\sigma_3$  are accordingly derived.  $R$  can be chosen through numerical experiments with the observational constraints of seismic moment and rupture velocity. Furthermore, it is noted that the ratio between the initial shear and normal stresses are exclusively determined by  $R$ ; they are independent of the stress and fault plane orientations. For simplicity, we maintain a constant  $R=0.4$  along the full length of the fault as our preferred model. The orientation of the two horizontal principal stresses are taken directly from Wollherr *et al.* (2019), with the maximum principal stresses imposed at N11°E, N20°E and N33°E orientations for CREVS, HVS and LJV, respectively (Figure 2a). This makes the northern CREVS the least favorable segment (56° between the fault and the maximum principal stress) characterized by a combination of larger shear and normal stresses relative to the other two segments at any given depth (Figure 2a). This model setup implying a larger potential stress drop on CREVS (for frictional coefficients set equal for all faults), which is consistent with the observation of a substantial amount of slip occurring near the CREVS and HVS connection (Liu *et al.*, 2003; Milliner *et al.*, 2015), something also revealed in inversion studies (e.g., Cohee and Beroza, 1994; Wald and Heaton, 1994; Olsen *et al.*, 1997).

	Landers Johnson Valley Segment	Homestead Valley Segment	Camp Rock Emerson Valley Segment
Fault orientation (°)	5 NW	23 NW	45 NW
Maximum principal stress orientation (°)	33 NE	20 NE	11 NE
Static friction coefficient	0.5	0.5	0.5
Dynamic friction coefficient	0.2	0.2	0.2
Critical slip distance (m)	0.6	0.6	0.6
$R$ (Preferred model)	0.4	0.4	0.4
$R$ (Ensemble models)	0.45	0.45	0.35
	0.5	0.5	0.4
	0.35	0.35	0.25

Table 1. Summary of modeling parameters. The preferred model involves a single set of parameters and the ensemble models consist of 48 individual models spanning four sets of  $R$  ratios and 12 realizations of the fault roughness.

The dynamic rupture is nucleated on the LJV S using a circular zone with a 3-km radius at a depth of 7 km (yellow star in Figure 2). Within this zone, the rupture is forced radially outward with a fixed slow rupture velocity of 2 km/s. Beyond that nucleation region, the rupture is allowed to spontaneously propagate.

## Fault failure criterion

In this study, we use a linear slip-weakening friction law (e.g., Ida, 1972) where the shear stress is bounded by the fault strength  $\tau^c$  defined as:

$$\tau^s \leq \tau^c = \begin{cases} c_0 + (\mu_s - \mu_d)l/D_c\tau^n, & l \leq D_c \\ c_0 + \mu_d\tau^n, & l > D_c \end{cases} \quad (4)$$

where  $\mu_s$  and  $\mu_d$  are static and dynamic friction coefficients, respectively, and  $D_c$  is the critical slip-weakening distance,  $l$  is the slip, and  $c_0$  is the frictional cohesion. All the frictional parameters are set constant across the fault plane with values as listed in Table 1. We set the static friction coefficient  $\mu_s = 0.5$  on LJV S and HSV and to 0.6 on CREVS, consistent with the regionally-inferred effective value of 0.6 by Gross and Kisslinger (1997). The slightly increased value on CREVS is imposed so that the fault strength is over the peak shear stress which prevents the rupture from unexpectedly nucleating on this specific segment. The dynamic

friction coefficient  $\mu_d$  is set to 0.2 across the fault and the critical slip-weakening distance  $D_c$  is 0.6 m, close to that used in previous studies (Olsen *et al.*, 1997; Wollherr *et al.*, 2019). These parameters along with the  $R$  ratio (0.4) synergically facilitate a sustained rupture with an observationally-consistent seismic moment of  $6\text{-}16 \times 10^{19}$  Nm and rupture velocity of 2.5–2.9 km/s (Cohee and Beroza, 1994; Wald and Heaton, 1994; Cotton and Campillo, 1995).

The material in the shallow crust transitions from ductile in the uppermost crust (top few km) to brittle in the remaining seismogenic zone. The main purpose of the shallow ductile layer is to increase energy dissipation. This makes the friction property have a depth dependence (e.g., Brune and Anooshehpoor, 1998; Day and Ely, 2002). In the rate-and-state friction framework, this layer has been proposed to follow a velocity-strengthening mechanism (Kaneko *et al.*, 2008) that is supported by laboratory experiments (Marone, 1998) and field observations of interseismic creep (Lindsey *et al.*, 2014) and shallow afterslip following large earthquakes (Perfettini and Avouac, 2007). We set the frictional cohesion  $c_0$  (in the slip-weakening law) to linearly decay from 1 MPa at the surface to zero below 3 km to mimic the ductile-to-brittle transition. These modeling choices are consistent with our goal to keep the model as simple as possible while making sure it produces an appropriate description of aggregate displacement at the surface.

## **Rock properties**

During an earthquake, some of the energy dissipation comes from inelastic deformations. The deformations are expected to be larger in the near-fault area where the stress concentration and release are the largest. Geodetic-based studies have inferred that the slip at depth is systematically larger than at the surface, a phenomenon called shallow slip deficit (SSD). That SSD, as well as the OFD, is attributed to the shallow distributed inelastic deformation (e.g., Kaneko and Fialko, 2014; Roten *et al.*, 2017). Both SSD and OFD sensitive to the material strength beyond which rock inelastically deforms were observed in the Landers earthquake. We model the SSD and OFD using a plasticity model borrowed from continuum mechanics. Specifically, we follow Roten *et al.* (2017) and implement the generalized Hoek-Brown model (Hoek *et al.*, 2002) that accounts for pre-existing cracks in rock strength by introducing the Geological Strength Index (GSI). Detailed descriptions of the Hoek-Brown model can be found in Hoek *et al.* (2012). For our purpose of simplicity, we don't implement a low-velocity fault zone as did in Roten *et al.* (2017). We tuned the properties to match observation, selecting an average quality rock mass (defined in Roten *et al.*, 2017) with  $GSI = 50$ , intact material constant  $m_i=12$  and unconfined compressive strength  $\sigma_{ci}=80$  MPa, corresponding to disintegrated but weakly interlocked limestone. We implement depth-variable equivalent friction and inelastic cohesion (lower panel of Figure 2c), which approximate the Hoek-Brown yield surface with a commonly used Drucker-Prager criterion (Drucker and Prager, 1952) through the relations in Hoek *et al.* (2002). The plasticity model is a simplification of the more complex physics at play (involving brittle failures), but it is computationally efficient and sufficient for our specific application. The surrounding elastic medium is defined by a 1D layered velocity model as illustrated in the upper panel in Figure 2c. The velocity model was adopted from the SCEC BBP as representative of the structure of the Mojave Desert (Goulet *et al.* 2015).

## **Numerical approach and performance**

We simulate fault displacements by numerically solving the three-dimensional (3D) elastoplastic spontaneous rupture propagation problem with the Support Operator Rupture Dynamics (SORD) code (Ely *et al.*, 2009; Shi and Day, 2013; Wang and Day, 2020). This generalized finite difference method has been used in numerous studies of spontaneous dynamic ruptures and ground motions (e.g., Wang and Day, 2017; Vyas *et al.*, 2018; Wang *et al.*, 2019; Savran and Olsen, 2020). This code also has been verified through a number of benchmarks including testing with small-scale fault geometrical roughness, plasticity, and thermal pressurization as part of the SCEC-based technical activity group on dynamic code rupture verification (Harris *et al.*, 2009; Harris *et al.*, 2018). The cohesive zone requires at least  $\sim 5$  grid points for ensuring numerical convergence (Day *et al.*, 2005) and is well resolved by our discretization (Figure S1). Considering the minimum shear wave velocity of 550 m/s, we estimate that our numerical solution provides accurate waveforms for frequencies up to roughly 1Hz. The model contains 1.9 billion hexahedral elements and simulates a 60 s rupture. The computation takes  $\sim 3$  hrs using 8192 processors on Frontera at the Texas Advanced Computing Center (TACC) and on Theta at the Argonne Leadership Computing Facility (Argonne National Laboratory). Both systems were used in this project.

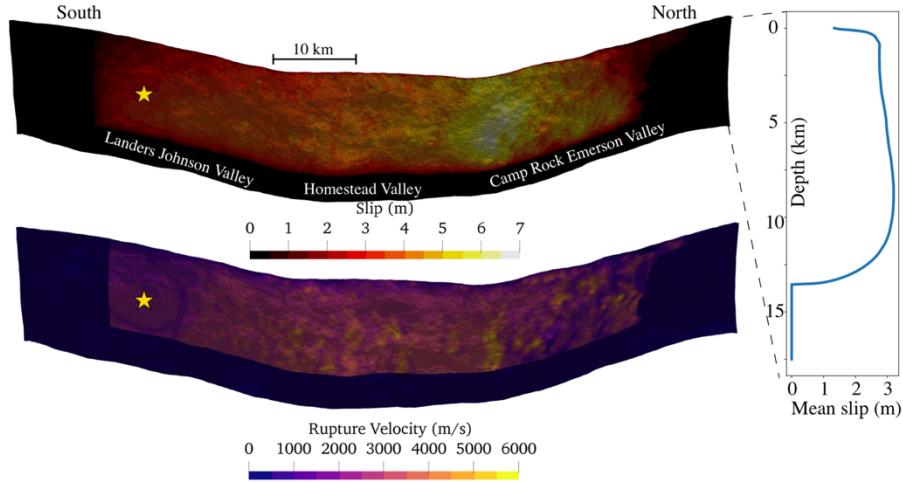


Figure 3. Snapshot of achieved slip (left top) and rupture velocity (left bottom) for our preferred model, with the location of the hypocenter indicated by a yellow star. The right plot shows the mean slip averaged over the strike as a function of depth.

## RESULTS

### Preferred dynamic model

In this section, we further describe our preferred model, intended to best represent the Landers event (given the modeling assumptions and choices described above). The seismic moment ( $10.7 \times 10^{19}$  Nm corresponding to **M7.3**) is within the range of observational studies ( $6 \times 10^{19}$  to  $16 \times 10^{19}$  Nm), the slip-weighted stress drop is 14.0 MPa, also in agreement with that used in previous dynamic modeling studies (e.g., Peyrat *et al.*, 2001; Wollherr *et al.*, 2019) and kinematic source modeling studies (Day *et al.*, 1998; Bouchon & Campillo, 1998). Figure 3 illustrates the coseismic slip and rupture velocity over the whole irregular fault plane. The peak slip exceeds 6 m near the smoothed kink between HVS and CREVS (around 40 km away from the hypocenter). This simulated result is supported by the observations (Perrin *et al.*, 2016) that

the zone of the largest fault displacement and the hypocenter location are uncorrelated: they can be close to each other (e.g., the 1999 Hector Mine and the 2010 El-Mayor Cucapah) but more often distant (e.g., the 1992 Landers and the 2002 Denali). The simulated depth-varying mean slip in Figure 3 shows that the slip is substantially tapered in the top 1 km and that we achieved a strong plastic effect, leading to a shallow slip deficit of ~40%, within the range (30-60%) estimated for the Landers earthquake (Kaneko and Fialko, 2011). In addition to those observation-consistent static deformation parameters, the rupture velocity is nearly 3 km/s at the seismogenic depth and as slow as 1-2 km/s at the shallow depth (Figure 3), which is also in agreement with previous finite-fault studies findings (Cohee and Beroza, 1994; Wald and Heaton, 1994; Cotton and Campillo, 1995). In all, results for both the simulated static and kinematic parameters are consistent with those estimated previous Landers studies, suggesting that our preferred dynamic model captures the first-order characteristics of the Landers earthquake in the aspects of overall deformation and seismic radiation.

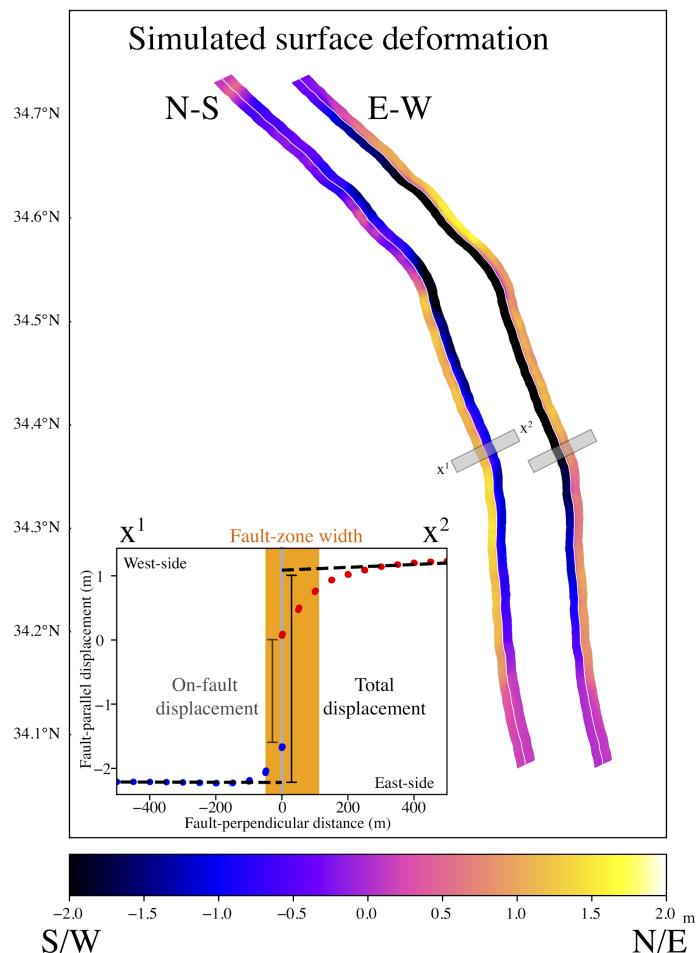


Figure 4. Simulated surface deformation and fault displacement from the preferred dynamic model. Two components of horizontal deformation within 3 km to the fault trace. A representative simulated fault displacement across the fault-perpendicular distance is shown in an inset figure where the total displacement, on-fault displacement, fault-zone width are defined. Dashed lines are the regression lines from the far-field deformations.

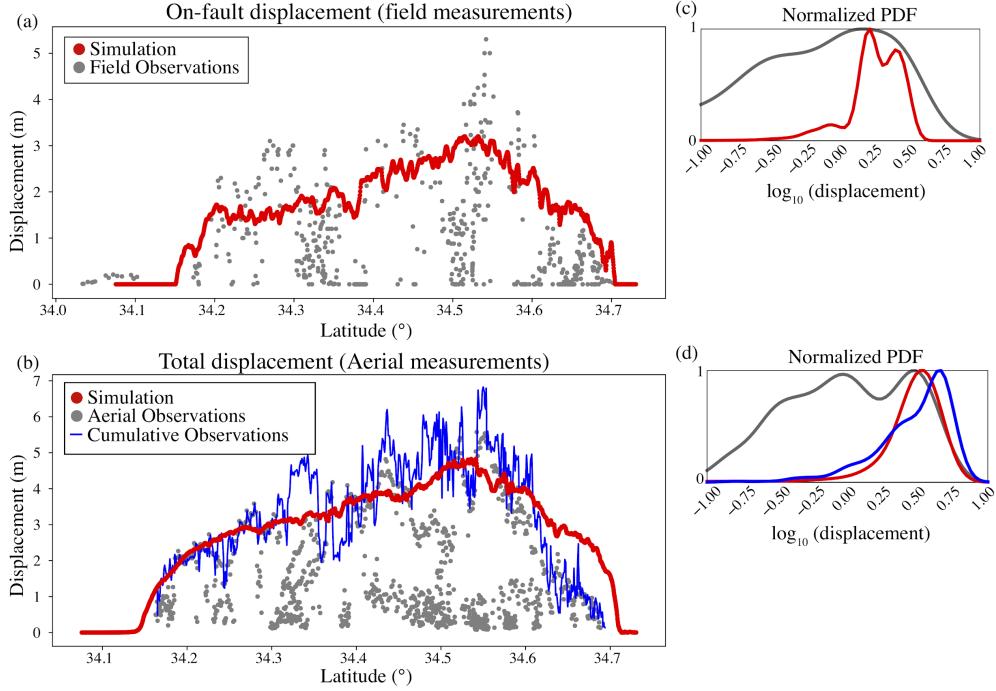


Figure 5. (a) Simulated on-fault displacement and comparison with field mapping. (b) Simulated total displacement and comparison with aerial mapping. A cumulative displacement from the aerial mapping (blue curve) is superimposed. Normalized probability density function of observations and simulations in (a) and (b) are plotted in (c) and (d), respectively. The legends in (c) and (d) are consistent with those (a) and (b).

The simulated horizontal surface deformation near the fault trace (Figure 4) shows an overall dextral shear movement (for clarity, we show the north-south and east-west components next to each other). The sinuosity on the fault trace (Figure 4) illustrates the effect of the self-similar fault roughness applied to the three-segment model. For the purpose of this project, we define two sets of fault displacements: on-fault displacement that is measured on the fault trace analogous to the field estimates and total displacement that is measured over a wider aperture analogous to the aerial estimates. To mimic the subpixel correlation process used in Milliner *et al.* (2015), we devised a sliding window (gray box in Figure 4) with a 2 km fault-perpendicular width and a 200 m fault-parallel length. Inside the window, simulated fault-parallel displacements (computed from the horizontal deformation and the averaged strike orientation of the local fault trace) are stacked (inset in Figure 4) and exhibit the characteristic pattern revealed by the subpixel correlation method (Figure 1 and Milliner *et al.*, 2015). With regards to the stacked fault-parallel displacement, we measure the on-fault displacement by computing the displacement difference on the split nodes representing the fault (thick gray line in inset). The total displacement is measured by computing the difference of displacements that are on-fault intercepts of lines best fitting the far-field deformations ( $>500$  m off the fault) on each side. The fault zone width (FZW) is defined as the fault-normal extent of observable surface shear in the form of smoothed transition of displacement (orange shaded box in the inset). The stacked fault-parallel displacement is close to an error function shape. Hence we can define the FZW with a single number (and using a systematic approach that is repeatable) by fitting the normalized spatial derivative of fault-parallel displacement of each side with the derivative of an

error function:  $\exp(- (x/a)^2)$  in which  $x$  is the distance to the fault and  $a$  is the FZW. We estimate the on-fault displacement, total displacement and FZW within each stacking window, thereafter computing the corresponding OFD as the difference between the total and on-fault displacement. Finally, we compute the OFD ratio as the ratio between OFD and the total displacement, and compile all these metrics into a single dataset for follow-up analyses.

On-fault displacements of the preferred model and field measurements (as compiled in Sarmiento *et al.*, 2019) are compared in Figure 5a. We observe a good agreement in the global spatial distribution of the peak displacements from both datasets near  $34.5^\circ$  latitude corresponding to the connection between the HVS and CREVS fault. This suggests that the orientations of the causative stress and faults have a first-order impact on the distribution of fault displacement. However, the along-strike variability of the field-derived fault displacement is much greater than that of the preferred model. The simulated fault is ruptured along a single continuous strand in contrast to a multi-segmented rupture pattern in the actual event (Figure 1 and 5a). This comes from the simplifications in our model setup, that by design is not defined to replicate the complex patterns in part controlled by structural complexities (Milliner *et al.*, 2015) and shallow fault discontinuities (Oglesby, 2020). Figure 5b compares the total displacement between the simulation and the aerial measurements (as compiled in Sarmiento *et al.*, 2019). A cumulative displacement interpreted from aerial observations (Milliner *et al.*, 2015) that sums up all single-strand measurements is also superimposed. The simulated total displacement generally fits the interpreted observation-based cumulative displacement well, but with a lower along-strike variability, as noted above. For a more quantitative comparison, we also compute normalized probability density functions (PDF) of distributions of both simulated and observed displacements using a  $\log_{10}$  scale and illustrate them in Figure 5c and 5d. These PDFs show that our preferred model simultaneously captures large-scale statistical features (i.e., shapes and mean values) of on-fault and total fault displacements but misses the small-scale characteristics.

Quantitative analysis of the along-strike variability is as follows. Along-strike strain (the change in displacement per unit of along-strike distance) reflecting the extent to which the displacement varies along the strike is up to  $10^{-3}$  in the simulation; it is similar to that of observations of  $2 \times 10^{-3}$  in the Landers (Milliner *et al.*, 2015) and  $10^{-3}$  documented for other events (e.g., Michel and Avouac, 2006; Oglesby, 2008; Elliott *et al.*, 2009). Another metric representing the along-strike variability is the Hurst exponent of the power spectral density (PSD) of the fault displacement. The Hurst exponent of the simulated on-fault displacement is 0.63 (Figure S2), which is close to that on the main rupture path of the Landers ( $0.55 \pm 0.099$ ) and a modeling prediction ( $\sim 0.65$ ) based on a comparable roughness level (Bruhat *et al.*, 2020). Figure S2 shows that the Hurst exponent of the on-fault displacement is smaller than that of the total displacement (0.88), implying that the on-fault variability is smoothed when a wider aperture in detection is applied. Nevertheless, it is noticeable that the Hurst exponents of both on-fault and total displacement are systematically larger than that of the cumulative displacement (0.44 Milliner *et al.*, 2015), suggesting that the simulated geometrical fault roughness may be insufficient to represent the variability of a structurally complex event such as the Landers and that additional complexities (e.g., surrounding material and friction heterogeneity) may be needed in modeling to capture them (Bruhat *et al.*, 2020).

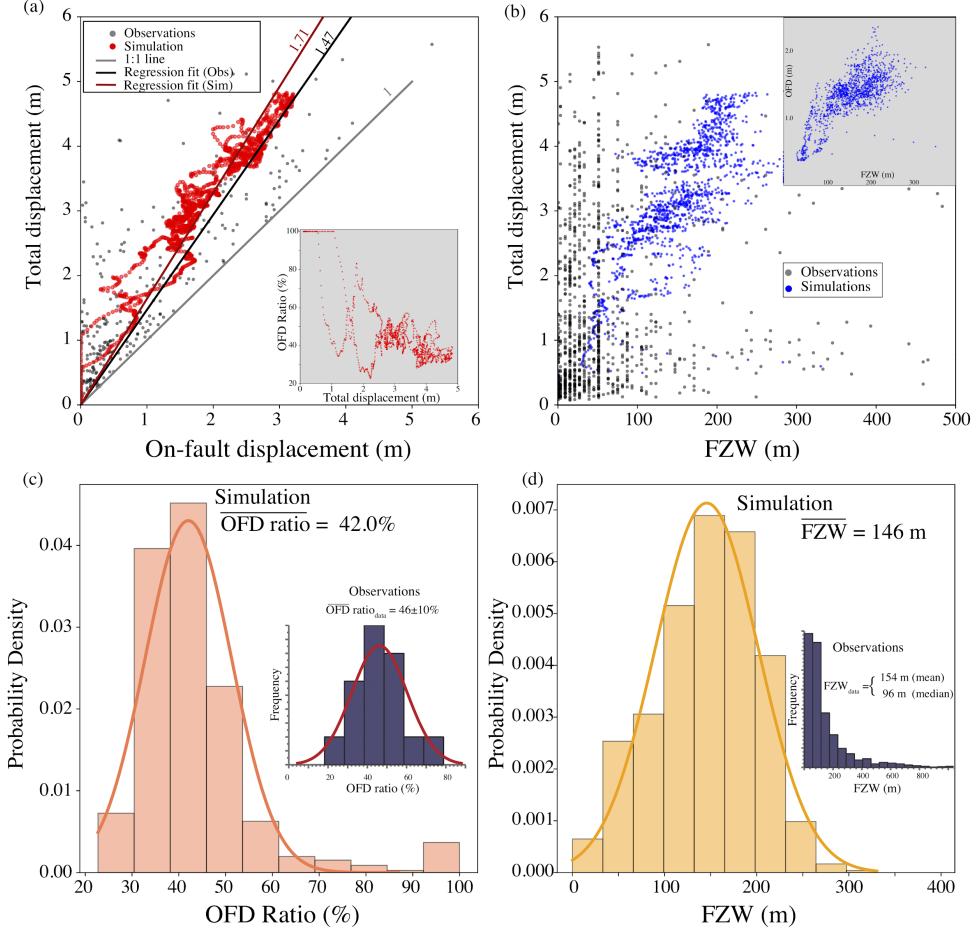


Figure 6. Comparison of simulation results (preferred model, Sim.) for different displacement metrics against aerial observations (Obs) as interpreted by Milliner et al. (2015). (a) Correlation between the total and on-fault displacement from simulations and observations. The inset shows the correlation between OFD ratio and total displacement of the preferred model. (b) Correlation between the total displacement and the FZW from simulations and observations. (c) Histograms of the simulated OFD ratio with inset showing that from observations. A fitted normal distribution (pink line) is superimposed. (d) Histogram of the simulated FZW with inset showing that from observations.

Figure 6a shows the partition of total and on-fault displacements at points along the fault. Both the simulations and observations indicate that considerable displacement occurs in the off-fault region (as indicated by the larger than 1 slope from data regression) and that the trends are quite similar for both datasets, with the simulations exhibiting a larger slope. The simulated OFD ratio (inset in Figure 6a) has a globally negative correlation with the total displacement. Such a correlation that the OFD ratio is lower in high-slip portions of ruptures has been reported for Landers (Milliner *et al.*, 2015) and for other recent earthquakes such as the 2010 El Mayor-Cucapah earthquake (Teran *et al.*, 2015) and the 2019 Ridgecrest Earthquake Sequence events (Gold *et al.*, 2021).

The shear strain field can be important for certain engineering applications and can be approximated by the halved ratio between the total displacement and the FZW. In the simulations

dataset, the total displacement is positively correlated with the fault zone width (Figure 6b) and the average ratio of around  $2.5 \times 10^{-2}$  corresponds to a shear strain of  $\sim 1.3 \times 10^{-2}$  similar to that estimated from the aerial correlation map (Milliner *et al.*, 2016). In addition, the OFD positive correlation with the FZW (inset in Figure 6b) in the simulation is also consistent with observations in the Landers (Milliner *et al.*, 2015) and the Ridgecrest (Gold *et al.*, 2021) events. However, the observed aerial dataset (gray dots in Figure 6b) show that the total displacements tend to be concentrated within a small FZW relative to the simulated dataset (blue dots in Figure 6b), implying that significant localized shear strains due to very complex fault geometry are missed by the simulation.

Statistical characteristics of the OFD ratio and FZW in the simulation dataset are plotted in Figures 6c and 6d. The simulated OFD ratio follows a typical normal distribution with a mean value of 42%, showing agreement with that of the aerial dataset. For the FZW, the simulation dataset follows a normal distribution as well with a mean value of 146 m close to the mean of the aerial dataset (154 m). However, the FZW from the aerial dataset distinctly follows an exponential distribution and contains a great number of small-scale ruptures that are missed in the dynamic model. Given our simplifying modeling assumptions, the impact of small scale fractures or strands is not captured by the simulations. This is a compromise we have to contend with. A potential solution would be to filter the data below a certain resolution, something to be investigated once many scenarios have been validated. However, the FZW of other earthquakes may follow distinct distributions due to a change of tectonic settings. For instance, the 1999 Hector Mine earthquake is instead much closer to a normal distribution (Milliner *et al.*, 2016).

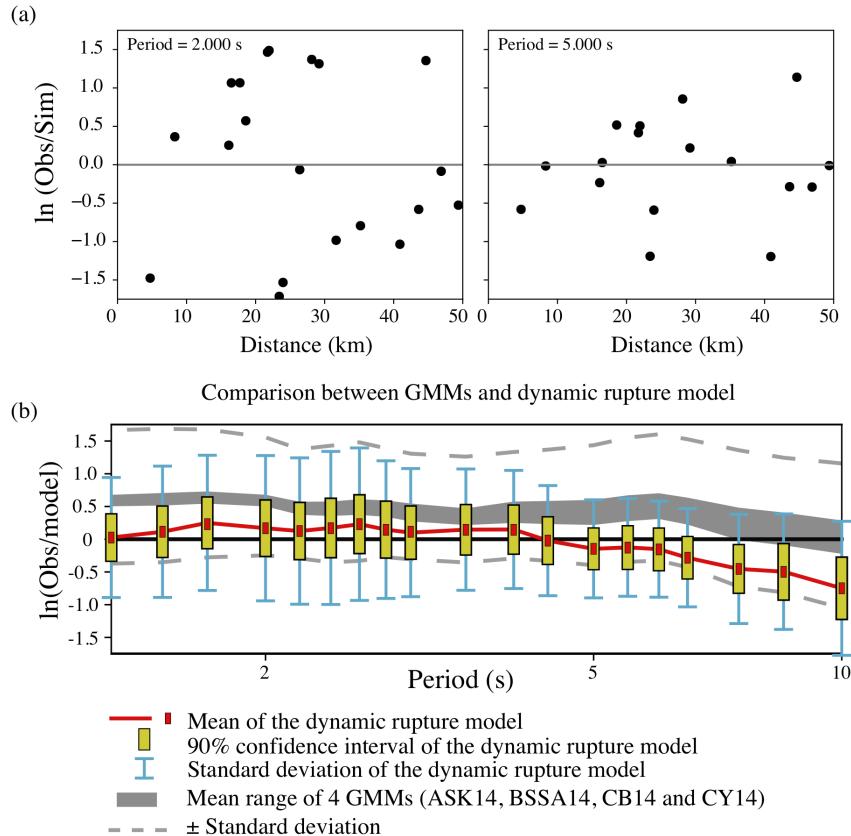


Figure 7. (a) Station-wise bias plots comparing 5%-damped RotD50 between the simulations and ground motions observations at periods of 2 and 5 s. (b) Statistics of the GMMs (Abrahamson *et al.*, 2014; Boore *et al.*, 2014; Campbell and Bozorgnia, 2014; Chiou and Youngs, 2014) and simulation are relative to the observation fit from all the stations. The red curve indicates the mean of the dynamic model while the yellow box and blue bar represent the 90% confidence interval of the mean and the standard deviation centered around the mean, respectively. The gray shade is the range of mean bias of four GMMs and dashed lines depict standard deviation of bias.

We also perform a validation of simulated ground motions against recordings from the event to ensure that all physics important to modeling have been properly parameterized. The goal of this validation is therefore more a consistency check than a thorough validation and is based on pseudo-spectral acceleration (PSA), and not waveform comparisons. The locations of seismic stations are denoted by blue triangles on the map (Figure 1) at which we simulate surface ground motions. We post-process the simulated and observed ground motions using the SCEC BBP (Maechling *et al.*, 2015) tools. The BBP computes relevant statistics from observation and simulation datasets, parameters needed for comparisons with empirical GMMs, and a goodness-of-fit (GOF) in terms of log residuals ( $\ln(\text{Obs}/\text{Sim})$ ). Figure 7a shows a station-wise comparison of 5%-damped RotD50 PSA between the data and the simulation for periods of 2 s and 5 s. Although data is limited, the mean GOF is near zero, the scatter is comparable to that from other studies (e.g., Dreger *et al.*, 2015), and there is no systematic trend with distance. Figure 7b shows the same GOF metric against spectral period aggregated across all the stations. The figure includes a comparison of the GOF from our simulation and from an aggregate of four GMMs (Abrahamson *et al.*, 2014; Boore *et al.*, 2014; Campbell and Bozorgnia, 2014; Chiou and Youngs, 2014). We find that the simulated ground motions produced by the dynamic model on average better match the observations relative to the GMMs, as the dynamic model inherently addresses the event-specific nature of Landers in terms of source and wave propagation. GMMs provide median global predictions of ground motions, with the between-event variability captured by an external event term (between-event residual). Additionally, the dynamic model produces fit variability comparable to that from GMMs for a given observed earthquake. The dynamic rupture model overpredicts ground motions periods beyond 5s (Figure 7b). This is likely due to the use of a 1D velocity model where the S wave speed at the shallow depth is as low as 550 m/s, systematically amplifying the simulated ground motions of stations located on the much stiffer structures in the north (Figure S3). Such deviations from a GOF of zero are consistent with previous results (Dreger *et al.*, 2015) for ground motions simulated in a 1D layered medium, and are not a concern. Overall, the ground motions resulting from our preferred model are reasonable and confirm that the modeling assumptions are adequate and do not lead to unrealistic ground motions.

## Suite of predictive models

The development of PFDHA models requires that several datasets be available for the definition of a stable prediction of median and standard deviation of expected future displacement metrics. Simulations can supplement the empirical datasets, which are rarely as detailed and complete as for the Landers event. Our first set of such forward simulations are built using the model setup developed through the Landers validation. While preserving most input parameters of the dynamic model (e.g., fault length, friction properties, and rock materials), we relax fault roughness and on-fault stress conditions to capture modeling uncertainties, thus generating models with a wide range of resultant magnitudes. Specifically, we create 12 distinct

self-similar fault roughness realizations with roughness levels identical to that used in the preferred model and 3 additional stress conditions specified by a segment-varying  $R$  that accounts for the uncertainties of the stress amplitudes and orientations (see details in Table 1). Consequently, a suite of 48 scenarios is simulated by iterating each roughness realization and stress condition. For analyzing the magnitude dependence of fault-displacement metrics, we divide all resulting scenarios into 5 bins based upon their magnitudes from 7.0 to 7.5.

Figure 8 illustrates on-fault and total displacements of magnitude-binned dynamic models and the Landers event. As the magnitude increases, the surface rupture length grows from about 40 km to 70 km, which roughly follows the scaling between the surface rupture length and the magnitude predicted by Wells and Coppersmith (1994) except that at the high-magnitude end, because the rupture was originally bounded for the M 7.3 Landers event. Earthquakes with lower magnitudes tend to partially rupture the fault, as expected, and they terminate near the kink structure between the HVS and CREVS faults, which supports the plausible unfavorable bend argued in Wollherr *et al.* (2019). Once the kink is overcome and the CREVS segment begins to rupture, the peak displacement is found near the HVS and CREVS connection point or on CREVS as a consequence of a combination of orientations of the fault and stress. Although the surface rupture is rather complex, the first-order spatial pattern of displacement such as the zone of peak displacement could be controlled by the large-scale tectonic settings, for example, the orientations of fault and stress (in this study) or the fault maturity (Perrin *et al.*, 2016), which may lay a physical foundation for implementing additional predictive simulations.

The simulated total displacement is again systematically larger than the on-fault displacement in each magnitude bin as shown in Figure 8. The extent to which the on-fault and total displacements are different will be discussed later. To test the predictivity of the ensemble models, we the amplitude and gradient of the binned simulated displacements with observations from the Landers event. Through this test, we find that Landers is closer to the scenarios with the magnitude between 7.2 and 7.4 (violet and pink curves in Figure 8b) enclosing the ground truth (7.3) Landers magnitude. This test implies that the physics-based dynamic rupture model is not only likely to quantitatively address observationally-consistent fault displacements, but also able to simulate events with a wide magnitude range by capturing physical plausible uncertainties, to support the PFDHA model development.

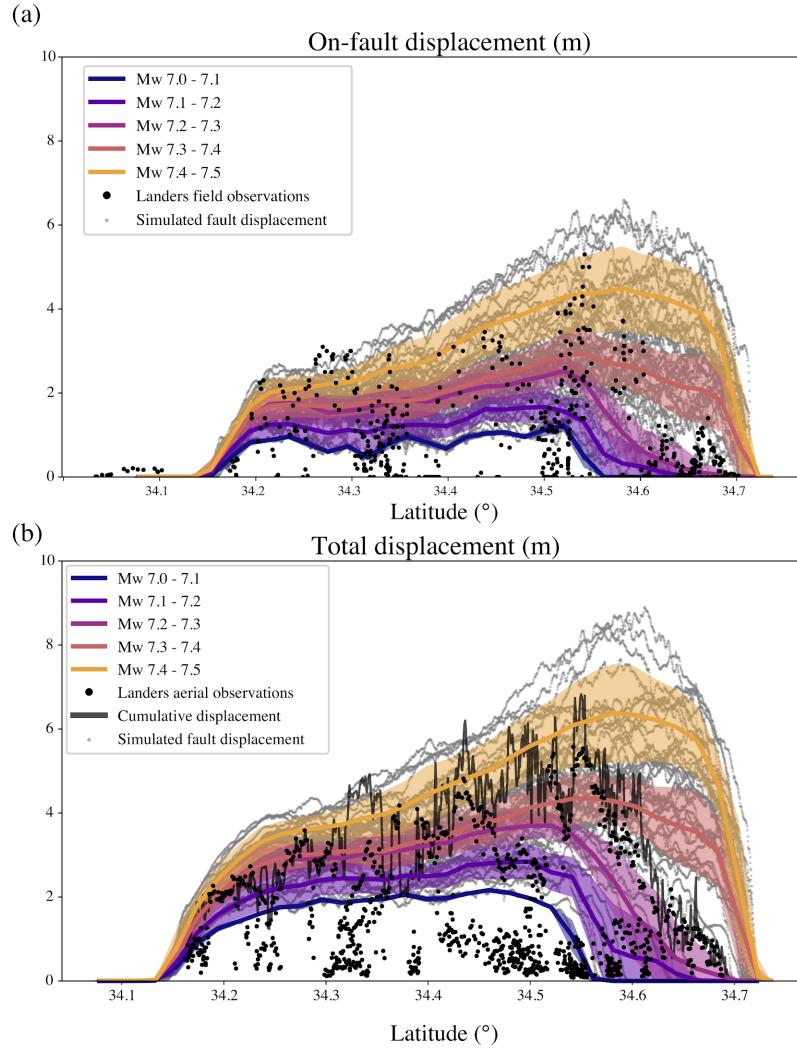


Figure 8. Simulated (a) on-fault and (b) total displacements from the suite of dynamic model simulations. The results are binned by magnitude with the mean (think solid curve) and standard deviation (shaded area) of displacements shown for each bin. Gray and black dots are simulated and observed displacements from Landers, respectively. The thin black curve in (b) is the cumulative displacement interpreted from Landers aerial observations (identical to the blue curve shown in Figure 5).

Next, we present the histograms of total displacement, OFD, OFD ratio and FZW of simulations within each magnitude bin and their mean values and standard deviations varying over the magnitude range (Figure 9). As seen with the preferred model, all these quantities approximately follow normal distributions. Although some of the histograms exhibit a certain level of some skewness, we fit a normal distribution as a way to provide general trends in mean and standard deviation values. The rightmost column in Figure 9 aggregates those mean and standard deviations as a function of magnitude. The ensemble models indicate that as the magnitude increases, total displacement, OFD and FZW all increase while the OFD ratio tends to decrease. The ensemble simulations centered on a magnitude of 7.3 slightly overpredict the total displacement and OFD relative to those (1.52 m and 0.8 m) observed in Landers. This is

consistent with the observations made on Figure 6 as these quantities in the Landers both follow the exponential distribution with an abundance of small-scale ruptures not modeled in our simulations. The decreased OFD ratio with increased magnitudes infers that surface deformation is more localized on the high-slip fault plane (large magnitude), which is widely supported by studies of the Landers (Milliner *et al.*, 2015), Hector Mine (Milliner *et al.*, 2016), El Mayor-Cucapah (Teran *et al.*, 2015) and Ridgecrest (Gold *et al.*, 2021) events. This result is also in agreement with the general geological findings that the extent of OFD is enhanced by structurally immature faults (Hollingsworth *et al.*, 2012; Dolan and Haravitch, 2014) and the notion that a mature fault is prone to bear larger magnitude earthquakes (Bletery *et al.*, 2016; Perrin *et al.*, 2016).

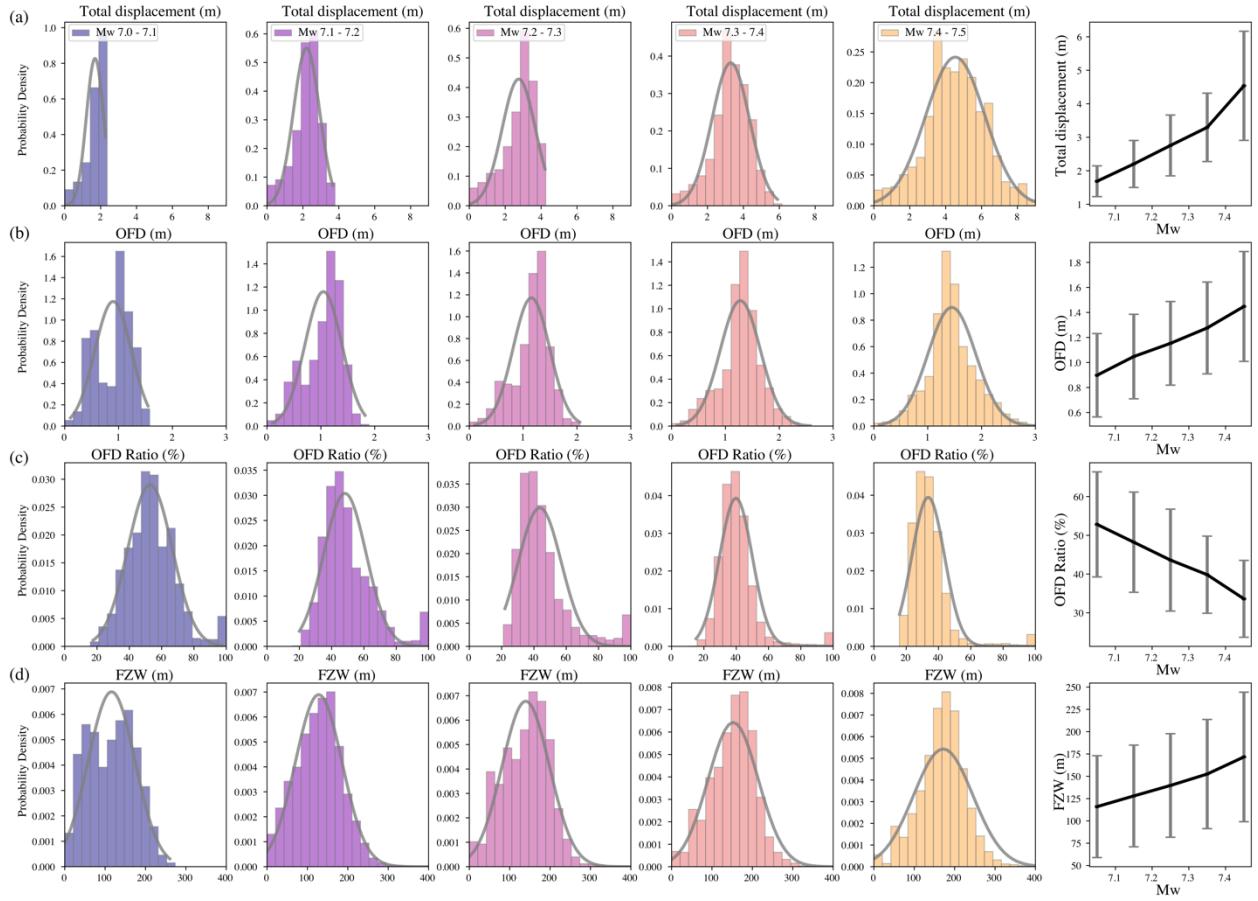


Figure 9. Histograms of (a) total displacement, (b) OFD, (c) OFD ratio and (d) FZW within each magnitude bins. The magnitude-dependent mean and standard deviation determined by the normal distribution fit (gray thin curve) are plotted in the rightmost column.

Figure 10 illustrates the correlations among the fault-displacement metrics and their dependencies on magnitudes. As seen in the observed dataset, the total displacement on average increases with the OFD (Figure 10a). The mean curves of the total displacements show a subtle magnitude dependence and stay in a relatively narrow zone (around one standard deviation, the shaded area in Figure 10a). A similar pattern is unveiled in the positive correlation between the OFD and the total displacement (Figure 10b). In the Landers dataset, there are numerous data points very near the 1:1 line, highlighting that off-fault deformations are identified but little

on-fault slip is observed (Figure 10b) at certain locations along the fault. Similar examples of little or no discreet faulting but large distributed deformation have been observed in the 2005 M6.5 Bam (Fialko *et al.*, 2005) and the 2010 Darfield (Van Dissen *et al.*, 2011) earthquakes. In contrast, correlations of the total displacement and OFD with the FZW show manifest magnitude dependencies (Figure 10c and 10d). As the FZW approaches around 100 m, the total displacement and OFD monotonically increase with the FZW (inelastic deformation is limited in this phase). Then both present an approximate saturation over the FZW from 100 to 300 m and then tend to taper to a larger FZW. The magnitude-dependent saturation is a consequence of the yielding effects caused by the implemented off-fault plasticity which is also found to saturate the near-fault ground motions and directivity amplification (Wang and Day, 2020). A similar saturation pattern of total displacement near 1 and 3 m is noted in the Landers dataset (Figure 10c). The reduced total displacement above the FZW of ~300 m is because the large FZW is mostly found at the rupture terminations or structurally complex segments where the total displacement is generally small.

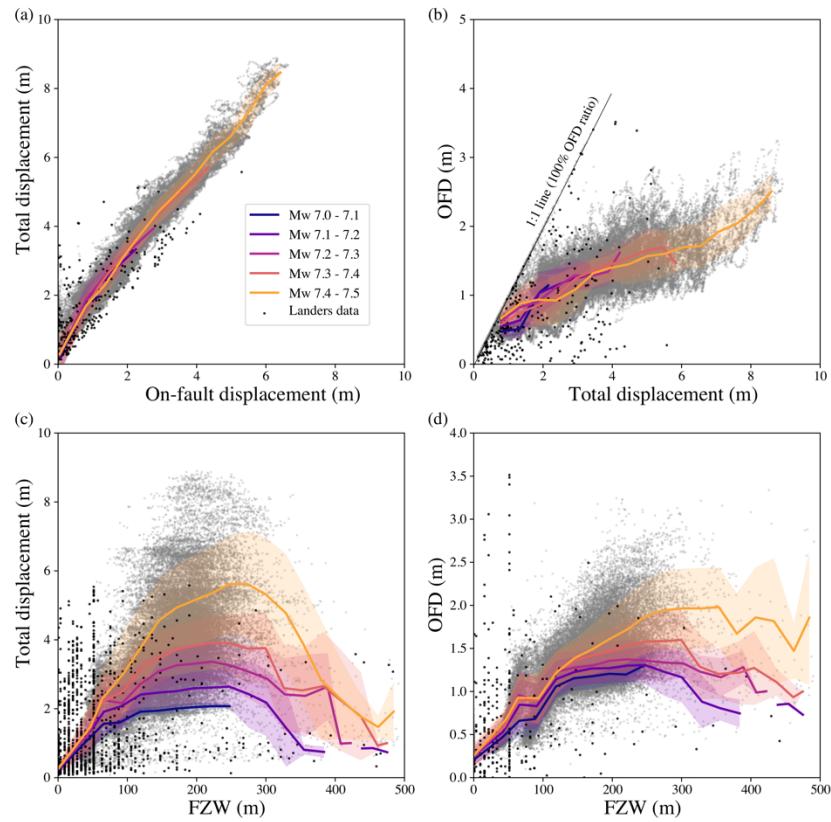


Figure 10. Correlation relationships between (a) total and on-fault displacement, (b) OFD and total displacement, (c) total displacement and FZW and, (d) OFD and FZW within each magnitude bin for the suite of models. The mean and standard deviation are represented by the colored curve and shaded area. Background gray dots and foreground black dots are the simulation dataset and the Landers observations, respectively.

Scaling laws relating mean displacement to magnitude are an essential component of PFDHA models (e.g., Petersen *et al.*, 2011; Field *et al.*, 2014). Provided that the recently observed total fault displacements from the aerial mapping are systematically greater than those from the field mapping, the pre-existing scaling relation (e.g., Wells and Coppersmith, 1994)

missing the off-fault deformation may lead to a significant bias of fault displacement hazard. As shown in Figure 11, the mean values of total displacements in the 1992 Landers (3.41 m in Milliner *et al.*, 2016), 1999 Hector Mine (2.84 m in Milliner *et al.*, 2016), 2010 Darfield ( $\sim 2.5$  m in Quigley *et al.*, 2012) and 2013 Balochistan ( $\sim 8$  m in Zinke *et al.*, 2014) earthquakes are systematically larger than those predicted by empirical models (Wells and Coppersmith, 1994; Wesnousky, 2008). However, thus far, fault displacement datasets containing aerially-derived total displacements are too sparse to be built upon. This is one area for which simulations can supplement the empirical datasets. We compute the mean and median displacement of on-fault and total displacement for our suite of models and plot them in Figure 11. The simulated on-fault displacements (mean and median) are within the empirical prediction range and the simulated total displacements are consistent with the limited COSI-Corr observations. This comparison suggests the dynamic rupture method has the capability to capture observation-consistent characteristics of both on-fault and total displacements. The steepening of the slope beyond M7.3 comes from the fault length being fixed for the original Landers scenario. Extending the model so that it can accommodate larger areas will change that relationship and this is something we will investigate in future work.

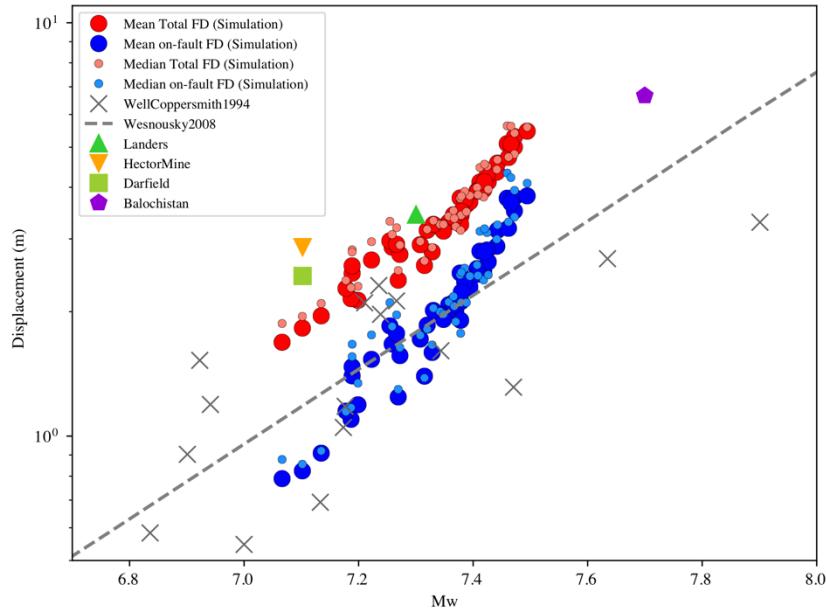


Figure 11. Scaling relationship of mean displacement with magnitude for simulated and recorded strike-slip earthquakes. Red and blue circles represent the mean values of the total and on-fault displacements from the suite of simulations, respectively, while light red and blue circles represent the corresponding median values. Crosses show the data included in Wells and Coppersmith (1994). The dashed line is the power-law regression of the mean displacement from Wesnousky (2008). Four data points of the mean displacements aerially estimated (Landers, Hector Mine, Darfield and Balochistan earthquakes) are superimposed.

## DISCUSSION

In this study, we developed a dynamic rupture model that creates observation-consistent fault displacements and relevant metrics with those from the 1992 Landers earthquake. However, when compared to the detailed fault displacement observations, two major discrepancies are

emergent. The first is that our simplified dynamic rupture simulations substantially overlook the small-scale fault displacements as represented by the exponential distributions of the total displacement, OFD and FZW. These small-scale observations may have significant implications on the mechanisms of surface rupture and off-fault deformation but are physically and computationally challenging to be modeled. The second discrepancy is that the simulated along-strike variability useful for constructing PFDHA models (Lavrentiadis and Abrahamson, 2019) is systematically lower than that of the observation dataset. As an indicator, the slopes or Hurst exponents obtained from numerical simulations are often overestimated. It is a likely consequence of the model simplification without considering other controls on displacement heterogeneity. , we prioritized the modeling of the intermediate-to-large scale features of fault displacements. We note that part of these discrepancies may also be due in part to how the observational data is interpreted. Given that datasets collected from either field or aerial mapping are sampled as a data cloud, a strike variable fault displacement used for evaluating the spectral signatures is subjectively created, for example, by selecting a major rupture path (Bruhat *et al.*, 2020), cumulating projected data back to a fault line (Milliner *et al.*, 2015) or integrating the local maximum fault displacements over distance bins (Gold *et al.*, 2021). These subjective processes may lead to distinct evaluations for the same event or dataset. For instance, Bruhat *et al.* (2020) and Milliner *et al.* (2015) use different methods to obtain distinct spectral signals (i.e., PSD) for the Lander earthquake.

The dynamic rupture model is explicitly site-specific. That is, the predicted variability in fault displacements at a given site inherently includes all the characteristics of the 3D fault and surrounding material specific and unique to that site. To broaden the applicability and enhance the predictivity of dynamic rupture models for future site-specific events, the uncertainties of the modeling approach and input parameters have to be appropriately addressed in the physics-based modeling. By generalizing our simulation set, we make it more broadly usable to inform the scaling of displacements with different ingredient characteristics, making them more broadly usable to support PFDHA development. We expect that our future work including additional scenarios will allow us to set better bounds on the parameter space to sample to capture more of the CA events. As we develop new rupture models spanning different geometries, fault mechanisms and stress fields, we will refine, and most likely expand, that parameter space so it is consistent with several past events. We are also considering expanding our modeling to include brittle damage as an improvement over to the plasticity model described here. While others have successfully used damage mechanics in the dynamic rupture model (e.g., Yamashita, 2000; Dalguer *et al.*, 2003; Bhat *et al.*, 2012; Lyakhovsky and Ben-Zion, 2014; Thomas and Bhat, 2018), we plan to evaluate an implementation to investigate its impact on displacement modeling. While both types of models limit stress at high strain, a key difference is that the brittle damage changes the elastic moduli of the host rock while plasticity does not. Thus the resultant low-velocity zone widely documented in the field may impact the distribution of peak fault displacement (Cappa *et al.*, 2014) and spatial extent of fault zone width (Roten *et al.*, 2017). Specifically the microfracture-based brittle damage model in Thomas and Bhat (2018) may support modelling of the currently missed small-scale surface ruptures.

We summarize below the need for continued work in related fields of study that will support the dynamic rupture modeling input for the PFDHA. The continued improvement in the fields will impact the selection of input parameters for key ingredients, even for simplified models such as ours

- (a) **Fault mapping.** Consistently mapping principal and secondary fault traces is a basis of PFDHA to address the potential fault displacement hazard. The post-earthquake exhaustive reconnaissance mapping of coseismically ruptured faults is valuable for supporting validation (e.g., Ponti *et al.*, 2020, Goulet *et al.*, 2021). Furthermore, the pre-rupture fault mapping from geomorphology (e.g., Arrowsmith and Zielke, 2009) and its comparison with the post-event fault rupture data may provide clues about seismic hazards and issues related to the repeatability of specific rupture geometries on the surface. Both types of fault mapping are beneficial to construct site-specific dynamic rupture models by adopting the original or simplified fault geometry.
- (b) **Fault history.** The fault history revealed from paleoseismic trenching and geomorphological surveys can also support more relevant dynamic rupture models. As stated in Milliner *et al.* (2015) and Hollingsworth *et al.* (2012), the off fault deformation and fault zone width are controlled by the fault maturity represented by the cumulative slip, which can provide insight into how much off fault deformation will be accommodated in future earthquakes. The fault maturity inferred can then provide constraints on the level of fault roughness to implement to appropriately reproduce the desired along-strike displacement variability and/or high-frequency ground motion (Shi and Day, 2013). In our dynamic rupture models, the structural controls shown in Figure S5 impacted the modeling decision, where the large OFD ratio and FZW occur in the framed segments of either a big change in strike angle or with the rupture termination. Additionally, a detailed segmented fault history is likely used to constrain fault and stress properties. For example, the cyclicity (e.g., periodic cycle or supercycle) revealed from paleoseismic studies may provide insight on the buried fault structure, the interseismic stress transfer mode and the degree of fault segment coupling (Scharer *et al.*, 2017; Philibosian and Meltzner, 2020). For a periodic paleoseismic record on a certain segment of the fault, a long recurrence period accompanied with dominantly large magnitude events would imply higher friction than if the paleoseimicity shows a more regular moderate magnitude stress release, for example (Dal Zilio *et al.*, 2019). Hence, continued work on such paleoseismic studies is important and can provide useful constraints for rupture modeling..
- (c) **Stress condition.** The stress condition is another poorly constrained parameter. It is temporally and spatially variable as it evolves with tectonic processes and multi-phase interplay (e.g., fluid transportation). Relative to its magnitude, the orientation of the stress field is less challenging to be constrained and as seen in our study, the orientations of fault and stress synergistically make a first-order impact on the fault displacement (e.g., peak displacement location). Close monitoring of regional seismicity and stress orientation offers an opportunity to address a variability caused by the in-situ stress in the physics-based modeling framework. The SCEC Community Stress Model (CSM, e.g., Bird, 2017) is one of such efforts that construct a suite of models and constraints on the stress and stress rate in the southern California lithosphere.
- (d) **Rheology of surface material.** The rheology of surface materials impacts the relative amount of OFD for a given total displacement. The OFD was found to be large in areas of younger sediments (Osokin *et al.*, 2012; Quigley *et al.*, 2012; Milliner *et al.*, 2016), but

this relationship does not always appear consistent (Milliner *et al.*, 2015; Gold *et al.*, 2021). Understanding this complex pattern deserves continued theoretical and observational studies, especially for the more complicated circumstance where the deformation occurs at sites of sediment-bedrock interfaces. In our current model, we use the 1D layered velocity structure for simplicity, but dynamic rupture models capable of enabling the heterogeneous and bi-material complexity will be developed for understanding and predicting the distributed off fault deformation in such complicated scenarios.

## CONCLUSION

Coseismic fault displacements in large earthquakes have caused significant damage to structures and lifelines near faults and impacted structural designs and legislations over the years. Unlike well-developed empirical ground-motion models, empirical fault displacement models are sparse and poorly constrained partly due to the paucity of detailed fault displacement observations. The physics-based dynamic rupture model has the potential to address this limitation. The overarching goal of this study is to develop and validate the dynamic rupture model so that it can reproduce observed fault displacements from past events and serve to generate datasets for yet unobserved events. We focused this initial study by validating and calibrating the dynamic rupture model with the M7.3 1992 Landers earthquake. Our validation and forward simulation exercise included (1) the definition of a preferred dynamic rupture model that systematically reproduces first-order metrics of the fault displacement (such as the mean ratio between the total displacement and on-fault displacement, the mean OFD ratio, the mean FZW and the location of the peak displacement) and (2) the development of a suite of dynamic rupture models quantifying expected fault-displacement metrics, their scaling and magnitude dependencies that are in part supported by Landers and other recent earthquakes.

Through these validations, we demonstrate the dynamic rupture model has a promising potential to develop physics-based fault displacement datasets that alone or combined with empirical observations support PFDHA models. Furthermore, dynamic rupture models recently make a push to extend deterministic ground-motion simulations to higher frequencies used for engineering applications (e.g., Andrews and Ma, 2016; Withers *et al.*, 2019). Our validation study highlights that the dynamic rupture model can provide defensible forward simulations for both ground motion and fault displacement hazards, serving communities of PSHA and PFDHA development. Encouraged by this validation exercise, more scenarios and events will be implemented with a similar approach to validate the dynamic rupture model, to include other strike-slip events (e.g., the M7.1 Hector Mine and the M7.1 Ridgecrest earthquakes), reverse mechanism events (e.g., Chi-Chi earthquake) and events of relatively small magnitudes (M6-7, e.g., the M6.4 Ridgecrest earthquake).

## DATA AND RESOURCES

The open-source software package SORD can be downloaded from <https://github.com/wangyf/sordw3>. We cite the FDHI dataset that is still under development (<https://www.risksciences.ucla.edu/nhr3/fdhi/databases>). Ground motion sanity check was performed using BBP (<https://github.com/SCECcode/bbp>). Figures in this article were prepared using Paraview (<https://www.paraview.org/>) and Matplotlib (<https://matplotlib.org/>). All websites were last accessed in February 2021.

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