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TITANIUM-RICH CHROMITE FROM THE MOUNT AYLIFF INTRUSION, TRANSKEI: FURTHER EVIDENCE FOR HIGH-TITANIUM THOLEIITIC MAGMA

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ABSTRACT

Four profiles through the lower, picritic portion of the Mount Ayliff Intrusion from the Karoo Igneous Province of South Africa have been studied. Many analyses of chromite grains enclosed in olivine show major differences from those found in most tholeiitic rocks, in containing very high Ti and Cr, and are low in Al. Typical ranges for these elements are (as cation proportions recalculated to 32 oxygen anions) Ti: 0.2 - 1.8 (as TiO2: 0.8 - 8.8 wt%); Cr: 6.3 - 11.1 (as Cr_2O_3: 26 - 51 wt%); and Al: 2.2 - 5.0 (as Al_2O_3: 6.7 - 15.7 wt%). Some of these are the most Ti-rich chromites recorded from any tholeiite suite. The lowest Ti contents occur in the olivine gabbros immediately adjacent to the floor contact. Their compositions are typical of tholeiitic chromite and cannot be related by fractional crystallization trends to the Ti-rich chromite trends in the overlying picrite as that would involve increase in Ti and Cr.

Models involving exsolution, postcumulus modification of chromite composition by infiltrating residual magma or subsolidus re-equilibration are rejected on the following grounds. The required Cr content of original olivine would have been unrealistically high. Most analyses are of grains totally embedded in olivine and the ability for reaction with residual magma is restricted. Analyses of chromite interstitial to the olivine and enclosed in postcumulus phases, and which therefore were exposed to residual magma, do not show trends of increasing Cr and Ti. In examples where reaction with residual liquid have been documented in detail (Jimberlana and Rhum Intrusions) there are no analyses with high Ti and Cr comparable to those reported here.

Comparison of these Ti- and Cr-rich chromites with other intrusive and extrusive tholeiitic suites, and experimental studies indicates that these compositions could not form from a typical continental tholeiite, but probably crystallized from a magma with high-Ti content. The closest analogues to these chromite compositions in tholeiitic rocks are to be found in Hawaiian lava lakes, which are high-Ti tholeiites. It is suggested that the Mount Ayliff Intrusion formed from at least two injections of magma of different compositions, one being enriched in Ti.

Chromite compositions from the different profiles through the intrusion show different trends. In the Mount Evelyn profile the effects of differentiation and mixing can be identified and distinguished; while for the Siroqobeni and Ingeli profiles the chromite compositions reflect a prolonged period of slow mixing. Data from a restricted interval of the Waterfall Gorge profile at the base of the picrite indicate that the chromite formed from a typical continental tholeiite. This interpretation of multiple intrusion and mixing of geochemically distinct magmas is consistent with previously published data on Mg-rich ilmenite compositions.

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INTRODUCTION

An extensive province of intrusive and extrusive basic magmatic activity occurs in southern Africa (Cox, 1983; Bristow and Saggerson, 1983; Eales et al., 1984). The great extent of the dolerite sheets is now known to extend at least as far north as Malawi (Woolley et al., 1979; MacDonald et al., 1983). The Mount Ayliff Intrusion, which is about 600 m thick and outcrops over an area of approximately 1,000 km^2, is one of the thickest of these intrusive sheets. It has been the focus of considerable geological interest for over 100 years because of the presence of magmatic nickel-copper-platinum sulphide mineralization at the base of the intrusion (Du Toit, 1910; Scholtz, 1936; Dowset and Reid, 1967; Tischler et al., 1981; Lightfoot et al., 1984; Groves et al., 1986; Maske and Cawthorn, 1986). Based on its silicate mineralogy and differentiation sequence from picrite, through hypersthene gabbro to quartz diorite (Scholz, 1936), it has been regarded as the product of fractional crystallization of a typical continental tholeiitic magma, referred to as the Lesotho-type by Marsh and Eales (1984). However, Cawthorn et al. (1985; 1988) have documented the presence of ilmenite containing up to 10 wt% MgO in the lower, ultramafic portion of the intrusion, and suggested that the magma was possibly analogous to the incompatible-element-enriched Letaba basalt type, recognized by Bristow and Saggerson (1983) from the Lebombo and Nuanetsi regions of the Karoo Province. This interpretation has been debated by Lightfoot and Naldrett (1984) and Lightfoot et al.(1987), who suggested that the unusual ilmenite composition is the result of subsolidus re-equilibration with olivine.

In this present study the writers document the composition of Ti-rich chromite present in the lower units of this intrusion as it places further constraints on the nature of the parental magma or magmas, and crystallization history.

MOUNT AYLIFF INTRUSION

The Insizwa body is the largest of four adjacent bodies; Insizwa, Tonti, Tabankulu, and Ingeli (Fig. 1), collectively referred to as the Mount Ayliff Intrusion (Scholtz, 1936; Cawthorn et al., 1988). Typical sections are shown in Figure 2. They have a marginal medium-grained olivine gabbro, overlain by laterally variable layers of picrite and troctolite, together referred to as the basal zone. Above this occurs a central zone of hypersthene gabbro with minor olivine- and quartz-bearing horizons, and a discontinuous upper zone of quartz diorite and monzonite.

The exact age of this intrusion is unknown, but is assumed to be comparable to the geographically related dolerite sills which are about 190 ma. old (Bristow and Saggerson, 1983). Attempts to date this intrusion by Rb/Sr and Sm/Nd isotope techniques merely demonstrated that there had been multiple intrusion and/or variable contamination (Lightfoot et al., 1984).

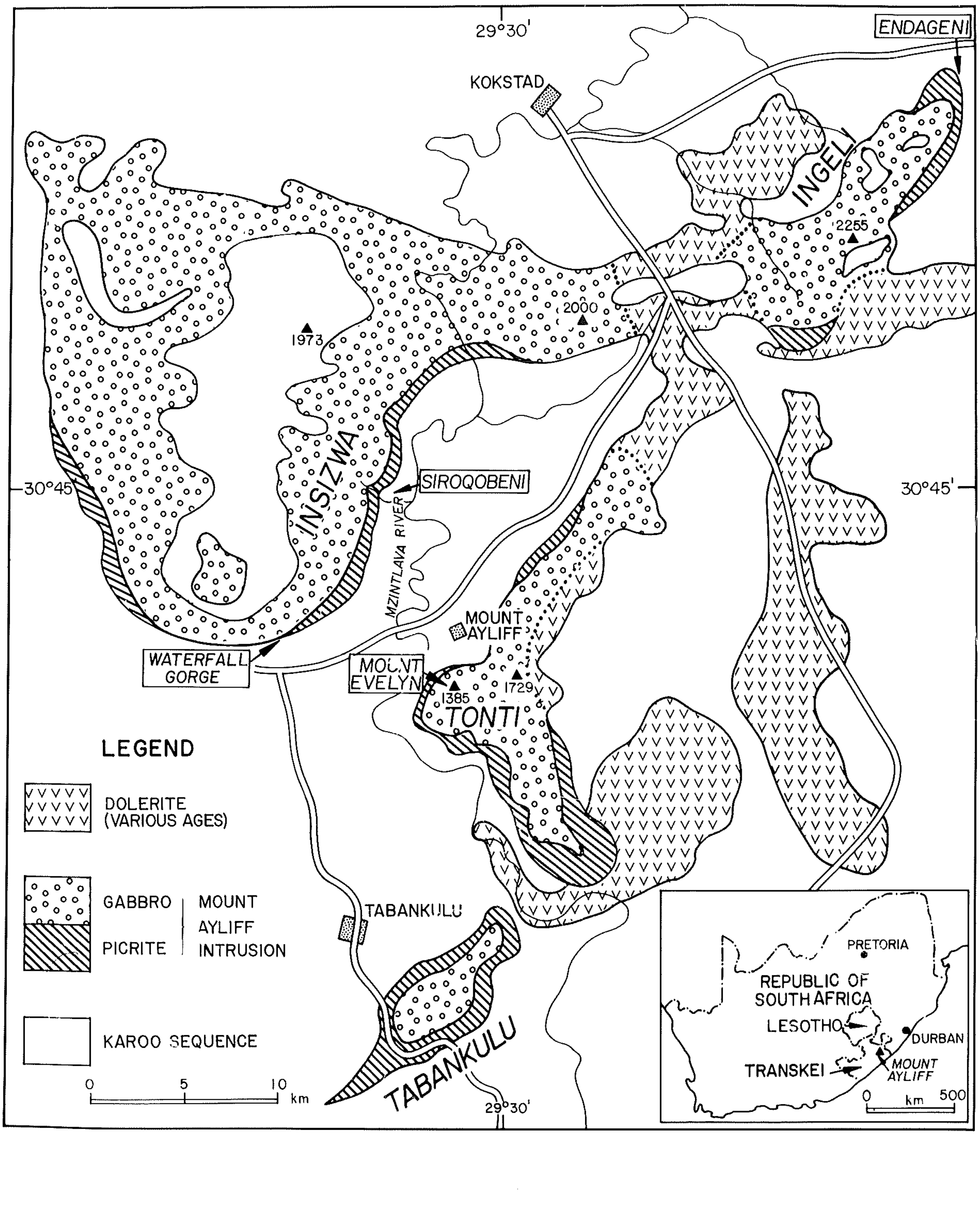

Figure 1: Simplified geological map of the Mount Ayliff Intrusion (from Maske and Cawthorn, 1986), showing localities of the profiles sampled at Mount Evelyn, Waterfall Gorge, Siroqobeni, and Endageni.

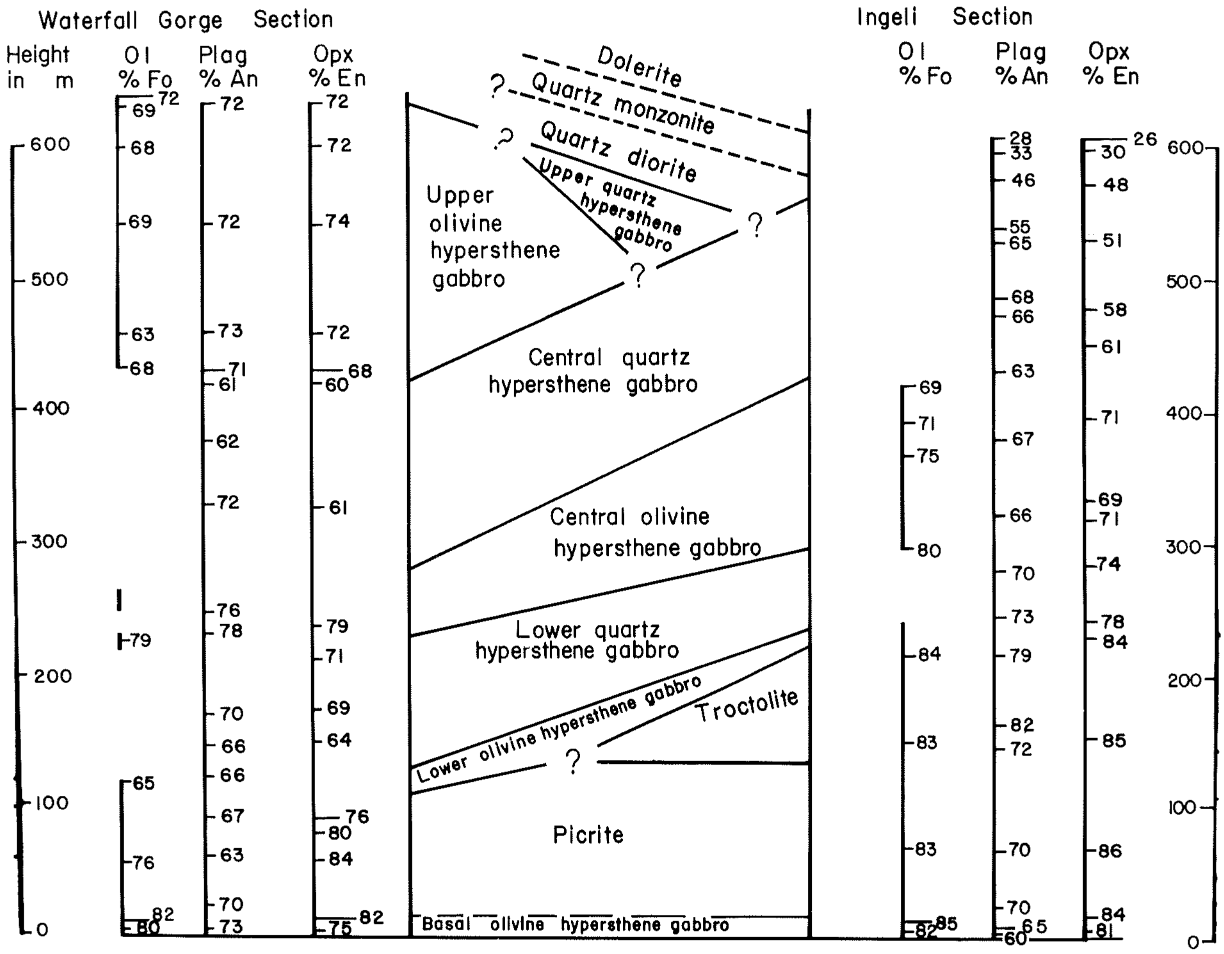

Figure 2: Typical sections through the Mount Ayliff Intrusion, including mineral compositions. Data for the lower part of Waterfall Gorge from Lightfoot et al. (1984); other data from Maske and Cawthorn (1986) and Cawthorn et al. (1986).

In the present study, chromite analyses from the picrite and under- and overlying olivine hypersthene gabbro are reported. Vertical profiles (Fig. 1) have been taken through the north end of the Tonti lobe (Mount Evelyn profile); through the Insizwa lobe, 8km northeast of the mineralized Waterfall Gorge area (Siroqobeni profile), together with some samples from the mine workings at Waterfall Gorge covering a vertical section of only 10m from the mineralized horizon at the base of the picrite; and through the Ingeli lobe (Endageni profile). These profiles offer the most nearly continuous sections through the picrite. Some additional data have been included from the Tabunkulu lobe (Lightfoot and Naldrett, 1983).

The lower contact is poorly exposed, and can usually only be located to within a few metres in the field. However, at Waterfall Gorge and at Mount Evelyn the basal contact is exposed and a thin zone of olivine gabbro is developed, which is 10 m thick at the first locality and only 2 m at the second. In some sections the lower contact of the picrite with this olivine gabbro is sharp, but in both instances examined here it is gradational. The upper contact of the picrite is not exposed on any of the specific profiles traversed; but an extremely sharp contact was located about 2 km southwest of the Mount Evelyn profile. At this boundary the modal content of olivine drops abruptly from 65 to 24 volume %. This differs from previous inferences regarding this contact (Scholtz, 1936; Lightfoot et al., 1984), where plots of modal proportions suggest a thick, gradational contact.

PETROGRAPHY

Nearly all of the samples studied are picrites. The only exceptions are the lowest (21/1) and three uppermost (21/9 to 21/11) samples from the Mount Evelyn profile. Though sample 21/1 comes from within 1m of the floor contact, it is medium grained. Olivine shows extensive corrosion features even when in contact with, or enclosed by plagioclase, although enclosure by clinopyroxene and orthopyroxene is more common (Tischler et al., 1981; Lightfoot and Naldrett, 1984). Orthopyroxene, clinopyroxene, and plagioclase are generally anhedral to subhedral, although there are a few euhedral plagioclase laths which develop an ophitic texture with the pyroxenes. The grain size of these matrix grains is typically from 0.1 to 0.2 cm. Chromite occurs either as small, euhedral grains embedded in olivine and all of the intercumulus minerals. A few show minor exsolution of ilmenite. Biotite, ilmenite, and sulphides are present as very minor constituents.

Picrite samples typically contain between 60 and 70 modal % olivine. With increasing height in the sections, the olivine grains take on an equant shape. Plagioclase, orthopyroxene, and clinopyroxene are present as poikilitic grains up to 0.5 cm in size, the pyroxenes showing slight exsolution. Ilmenite is a minor phase, euhedral when enclosed in plagioclase and pyroxene, whereas biotite is considerably more common but anhedral.

Towards the top of the picrite, plagioclase becomes more abundant, but remains poikilitic and intercumulus. The upper contact of the picrite is taken as the point where the plagioclase becomes a cumulus phase. Where this contact is exposed the modal plagioclase content increases abruptly from 26 to 46 volume %. The olivine content decreases to 20 % at this level in the Mount Evelyn profile, and then gradually disappears over a height of 50 m. Subhedral orthopyroxene and clinopyroxene are present in approximately equal proportions throughout the olivine hypersthene gabbro facies above the picrite. In the Siroqobeni profile olivine grains devoid of chromite are present a short distance above the top of the picrite and so there are no analyses from this horizon. In contrast, in the Mount Evelyn profile chromite is present in the olivine grains thoughout the entire 50 m-thick olivine-bearing horizon, from which samples 21/9 - 21/11 were taken.

The picrite contains approximately 1 to 1.5 wt% of chromite (although point counting is statistically imprecise at such low abundances), the proportion of chromite inside olivine grains being less than in the interstitial phases. The chromite grains tend to be clustered both within the olivine grains and interstitial to them. In the gabbroic rocks, the abundance decreases both in the olivine and interstitial to it. The chromite is always subhedral to euhedral, and appears to be the same size whether inside or outside the olivine grains.

TABLE 1. Analyses of chromite in olivine from the Mount Ayliff Intrusion.

Profile Mount Evelyn

Sample	21/1	21/1	21/2	21/2	21/3	21/3	21/4	21/4	21/5	21/5
Height(m)	1	1	3	3	7	7	12	12	17	17
TiO_2	0.7	3.8	4.7	7.0	5.2	6.0	3.0	4.5	3.3	5.1
Cr_2O_3	43.0	41.4	47.1	44.4	46.8	45.5	48.5	46.2	51.6	43.9
Al_2O_3	15.7	12.8	9.1	7.7	10.6	8.8	11.7	11.3	8.5	9.0
Fe_2O_3	6.5	6.6	3!2	3.5	2.0	3.3	2.2	2.7	2.6	5.9
FeO	22.8	29.0	28.9	30.1	25.8	29.8	26.1	26.3	27.7	28.5
MnO	0.3	0.3	0.3	0.4	0.3	0.4	0.3	0.3	0.3	0.4
MgO	7.1	5.0	5.2	5.6	7.9	5.4	6.2	6.9	5.3	5.6
Total	96.1	98.9	98.5	98.7	98.6	99.2	98.0	98.2	99.3	98.4
Ti	0.14	0.78	0.99	1.47	1.06	1.26	0.61	0.92	0.69	1.07
Cr	9.32	8.98	10.40	9.85	10.04	10.01	10.56	10.00	11.33	9.71
Al	5.08	4.15	3.00	2.54	3.41	2.86	3.81	3.66	2.78	2.96
Fe^{3+}	1.32	1.29	0.61	0.67	0.42	0.61	0.41	0.50	0.50	1.17
Fe^{2+}	5.25	6.73	6.82	7.13	5.86	7.03	6.06	6.08	6.50	6.72
Mn	0.07	0.07	0.07	0.09	0.07	0.09	0.07	0.07	0.07	0.09
Mg	2.89	2.05	2.17	2.33	3.20	2.22	2.56	2.83	2.19	2.35

Profile Mount Evelyn

Sample	21/6	21/6	21/7	21/7	21/8	21/8	21/9	21/9	21/11	21/11
Height(m)	23	23	39	39	58	58	62	62	93	93
TiO_2	3.2	6.3	3.4	3.4	4.8	8.8	3.6	5.2	5.1	4.3
Cr_2O_3	47.9	43.8	46.6	45.8	44.4	38.2	41.0	37.4	33.8	27.9
Al_2O_3	10.5	9.4	11.5	12.5	6.9	7.3	10.5	9.2	9.8	9.7
Fe_2O_3	3.9	3.9	3.9	4.0	8.4	8.7	9.1	11.0	14.6	20.9
FeO	26.1	29.2	26.6	25.8	30.3	28.5	29.7	31.1	31.9	32.1
FeO	26.1	29.2	26.6	25.8	30.3	28.5	29.7	31.1	31.9	32.1
MnO	0.3	0.4	0.3	0.3	0.4	0.3	0.3	0.3	0.4	0.4
MgO	6.3	6.1	6.1	6.9	4.4	7.9	4.2	3.9	3.5	2.9
Total	98.2	99.1	98.4	98.7	99.6	99.7	98.4	98.1	99.1	98.2
Ti	0.67	1.30	0.70	0.69	1.02	1.80	0.76	1.11	1.09	0.93
Cr	10.46	9.58	10.14	9.83	9.91	8.28	9.10	8.39	7.55	6.34
Al	3.43	3.06	3.72	4.02	2.28	2.37	3.47	3.09	3.27	3.28
Fe^{3+}	0.75	0.75	0.73	0.77	1.76	1.74	1.89	2.30	.2.99	4.51
Fe^{2+}	6.10	6.80	6.20	5.91	7.18	6.57	7.02	7.45	7.64	7.70
Mn	0.07	0.09	0.07	0.07	0.09	0.07	0.07	0.07	0.09	0.09
Mg	2.58	2.50	2.50	2.78	1.84	3.23	1.75	1.66	1.45	1.24

Profile Siroqobeni									Ingeli	
Sample	16/3	16/3	16/4	16/4	16/5	16/5	16/7	16/7	NGL/9	NGL/9
Height(m)	23	23	36	36	59	59	105	105	14	14
TiO_2	2.5	4.7	4.0	6.4	2.4	3.6	4.3	3.3	2.7	4.1
Cr_2O_3	40.3	33.1	38.3	31.2	40.0	35.7	37.4	25.3	41.7	38.5
Al_2O_3	7.1	7.3	6.3	6.4	7.0	6.8	9.1	12.7	12.8	11.1
Fe_2O_3	16.9	20.6	15.7	18.5	16.6	20.2	14.4	23.5	9.8	11.1
FeO	28.3	28.3	29.4	32.2	27.6	29.1	27.9	28.1	23.9	27.2
MnO	0.4	0.4	0.4	0.4	0.4	0.5	0.4	0.4	0.4	0.4
MgO	4.2	5.5	4.0	3.8	4.3	4.3	5.7	4.9	7.7	6.3
Total	99.7	99.9	98.1	98.9	98.3	100.2	99.2	98.2	99.0	98.7
Ti	0.53	0.98	0.87	1.39	0.52	0.77	0.90	0.70	0.55	0.85
Cr	9.03	7.34	8.74	7.09	9.07	7.98	8.21	5.56	8.86	8.40
Al	2.38	2.40	2.15	2.17	2.38	2.26	2.98	4.16	4.05	3.61
Fe^{3+}	3.53	4.29	3.36	3.95	3.50	4.23	2.97	4.84	1.98	2.30
Fe^{2+}	6.77	6.70	7.16	7.77	6.70	6.94	6.51	6.63	5.37	6.27
Mn	0.09	0.09	0.09	0.09	0.09	0.11	0.09	0.09	0.08	0.09
Mg	1.75	2.28	1.71	1.61	1.82	1.83	2.36	2.04	3.10	2.58

Profile Ingeli					Waterfall Gorge					
Sample	NGL/11	NGL/11	NGL/54	NGL/54	WG/1	WG/1	GC692	GC692	GC697	GC697
Height(m)	28	28	110	110						
TiO_2	2.3	1.4	1.8	3.8	2.6	2.8	0.6	0.9	2.7	4.4
Cr_2O_3	44.2	43.0	43.7	38.5	45.0	39.8	42.3	39.6	38.7	29.8
Al_2O_3	12.9	15.7	13.7	10.8	13.4	16.8	12.7	15.2	8.9	6.6
Fe_2O_3	7.1	6.9	8.8	12.2	4.7	6.0	10.8	10.3	15.5	23.7
FeO	24.0	24.0	23.4	27.3	28.9	26.9	25.7	26.3	31.2	33.0
MnO	0.4	0.4	0.4	0.5	0.4	0.4	0.4	0.4	0.5	0.4
MgO	7.3	7.2	7.7	5.7	4.7	6.2	5.6	5.7	3.5	3.0
Total	98.2	98.6	99.5	98.8	99.7	98.9	98.1	98.4	101.0	100.9
Ti	0.47	0.29	0.36	0.78	0.54	0.55	0.11	0.18	0.56	0.94
Cr	9.49	9.08	9.22	8.40	9.67	8.41	9.20	8.50	8.49	6.65
Al	4.13	4.95	4.29	3.51	4.28	5.29	4.12	4.87	2.90	2.20
Fe^{3+}	1.44	1.39	1.76	2.53	0.97	1.20	2.44	2.27	3.49	5.27
Fe^{2+}	5.45	5.34	5.22	6.31	6.56	6.01	5.70	5.79	7.00	7.57
Mn	0.09	0.08	0.08	0.11	0.09	0.08	0.10	0.09	0.11	0.10
Mg	2.93	2.86	3.06	2.36	1.88	2.46	2.30	2.29	1.46	1.25

For each sample the chromite analysis with the highest and lowest Cr_2O_3 content is presented.

Ferrous and ferric contents are calculated assuming stoichiometry, and

cation proportions are calculated for 32 oxygen anions per unit formula.

Details of analytical techniques are presented in Cawthorn et al. (1988).

CHROMITE CHEMISTRY

Only chromite grains embedded in olivine are reported here, and analyses with the highest and lowest Cr content for each sample are given in Table 1. At least eight grains per sample have been analysed. Multiple analyses of single grains indicate little variation; typically less than 0.6 wt% for Al_2O_3, Cr_2O_3 and FeO (total iron). Different grains of chromite within a single olivine grain may show slight differences. Such variations have been examined to see if chromite grains closer to the edge of olivine grains are systematically different from those nearer the centre, analogous to the findings of Evans and Moore (1968), or to see if there is any variation as a function of size, as reported by Wilson (1982). Neither of those patterns could be established in the present data. In some of the samples large differences exist between chromite analyses from different olivine grains.

Chromite compositions are plotted in Figure 3, which shows the cation proportions of Cr - Al - (Fe^{3+}+2Ti). The three complete profiles through the picrite occupy partially overlapping, but distinct fields.

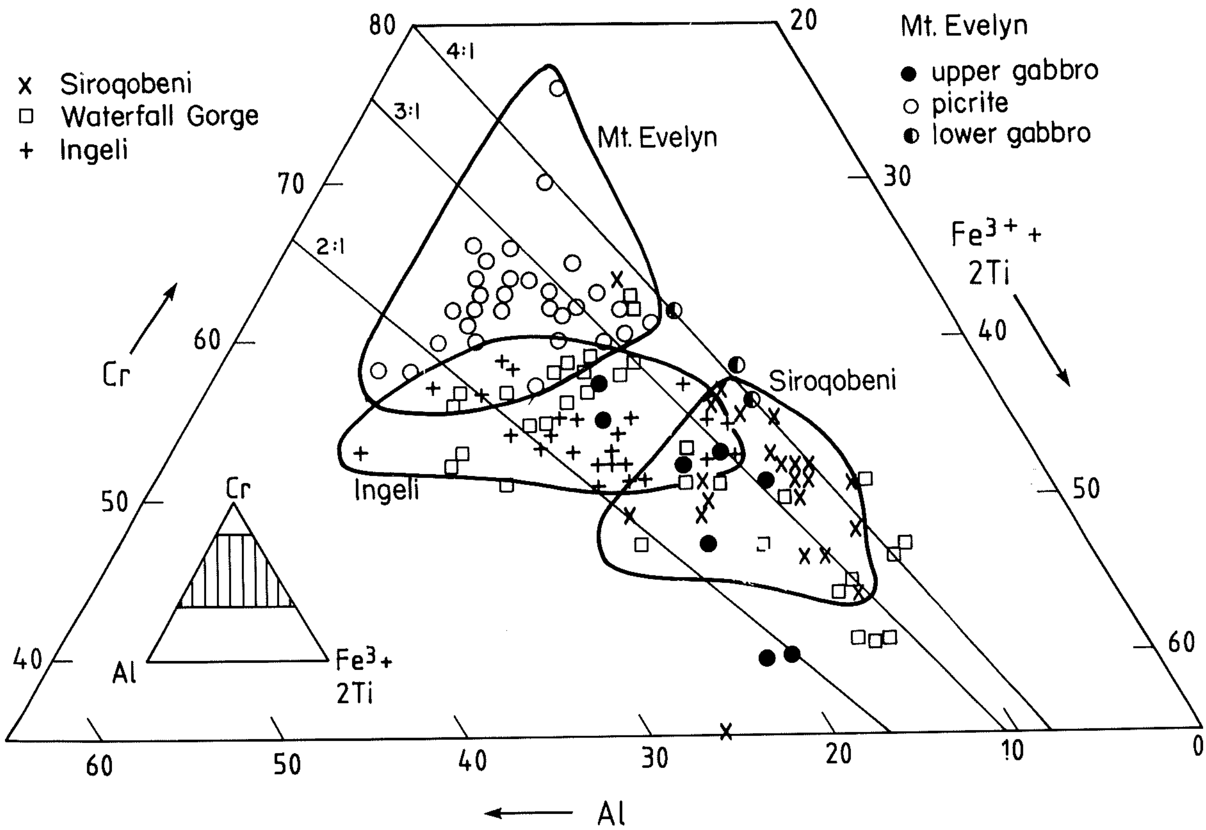

Figure 3: Plot of Cr-Al-Fe^{3+}+2Ti for compositions of chromite in olivine. Lines of constant Cr/Al ratio are shown. Samples from Mount Evelyn are shown as circles, basal olivine gabbro - half solid circles, picrite - open circles, olivine hypersthene gabbro above picrite - solid circles; Ingeli - plus signs; Siroqobeni - crosses; and Waterfall Gorge - open squares. The encircled field for Mount Evelyn refers to samples from the picrite only. In this and subsequent plots many data points are excluded because of overlap.

Those from the picrite from Mount Evelyn have the highest combined Cr+Al contents; those from Siroqobeni have lower Cr+Al contents. The samples from the Ingeli profile tend to have lower Cr:Al ratios than those from the Mount Evelyn profile. Chromite from Waterfall Gorge shows the widest range of compositions even though samples come from a restricted vertical section.

Further differences between the profiles are apparent in the cation plot of Cr+Al versus Ti (Fig. 4). (These and all subsequent cation abundances quoted and presented in figures have been calculated to 32 oxygen anions per unit formula). This diagram is designed to illustrate the relationship between elements in chromite which tend to concentrate in the early stages of differentiation (Cr and Al) and Ti which usually increases in more evolved compositions. Chromite in picrite from Mount Evelyn, Endageni, and Siroqobeni occupy distinct, subparallel trends. In contrast, the data for Waterfall Gorge show a more rapid decrease in Cr+Al in the range 0.2-0.7 cations of Ti. Chromite analyses from the upper olivine hypersthene gabbro of the Mount Evelyn profile have values straddling the gap between the Mount Evelyn picrite and Waterfall Gorge trends.

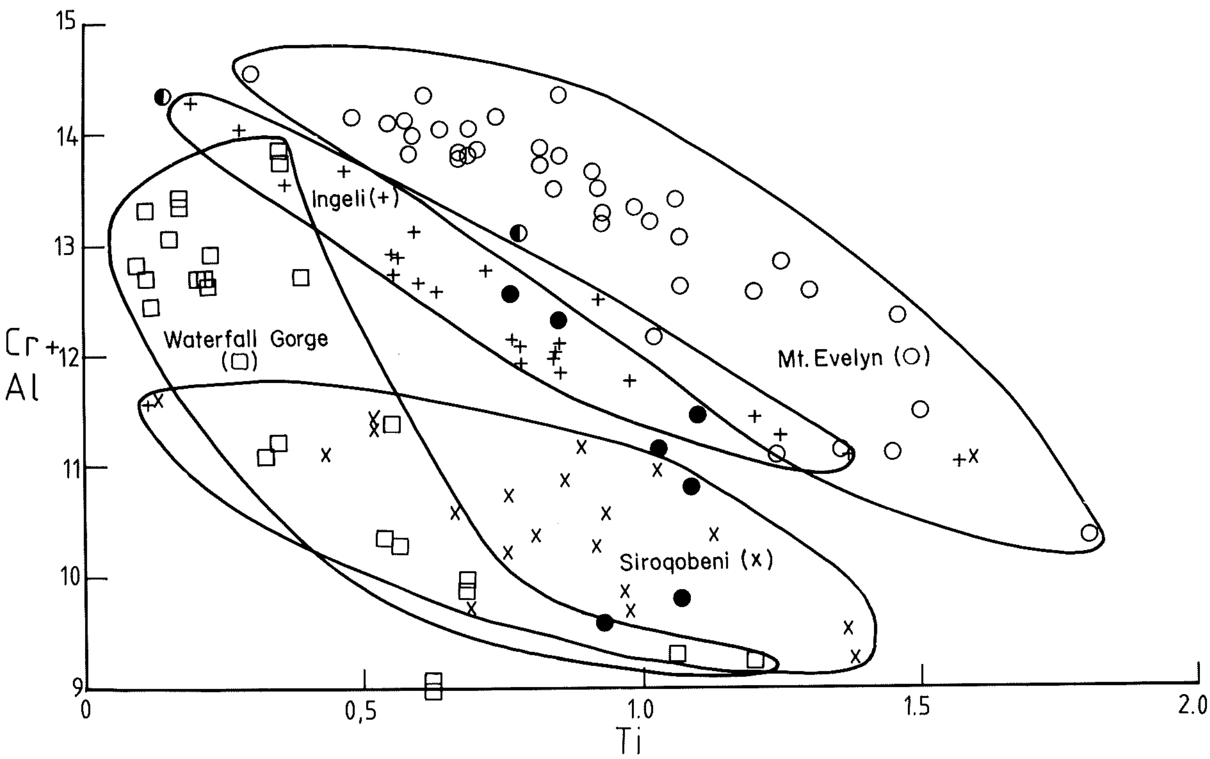

Figure 4: Plot of cation Cr+Al versus Ti for chromite compositions, based on 32 oxygen anions. Symbols are the same as for Figure 3.

DISCUSSION

Dick and Bullen (1984, p. 54) regarded chromite as "petrological litmus paper, being extremely sensitive to host rock petrogenesis". The writers concur with this view and suggest that the composition of the chromite provides further insights into the origin and evolution of the Mount Ayliff Intrusion. The most notable feature of some of these chromite compositions is their high Ti content, the highest value recorded being 1.8 cations of Ti per 32 oxygen anions for a grain with 8.1 cations of Cr. Combined high Ti and Cr is anomalous for chromite formed from tholeiitic magmas. For example, chromite in continental layered intrusions, such as the Bushveld Complex, the Great Dyke, and the Stillwater Complex, has Ti values generally less than 0.2 cations (Jackson, 1969; De Waal, 1975; Cameron, 1977: Wilson, 1982; Eales and Reynolds, 1983). In extrusive continental tholeiites chromite is very scarce, but occurs in picritic rocks such as those at Baffin Island (Clarke, 1970) and Deccan Traps (Krishnamurthy and Cox, 1977), where the chromite contains less than 0.4 cations Ti. Similarly low Ti values are reported in chromite in MORB (Dick and Bullen, 1984; Allan et al., 1988) and ophiolites (Malpas and Strong, 1975).

In order to assess the significance of these chromite compositions, it is necessary to consider what processes may have affected their chemistry. Their composition may reflect the effects of :

1. exsolution;
2. subsolidus re-equilibration with host olivine;
3. reaction with interstitial liquid; and
4. the parental magma composition(s).

Exsolution

Barnes et al. (1988) suggested that chromite inclusions in olivine in very thick komatiitic flows were the result of exsolution from the olivine host. In the present study the percent of chromite included in olivine is approximately 1.5 wt %. The average chromite analysis contains 45 wt % Cr_2O_3. Hence, to produce 1.5 % of chromite by exsolution the original olivine must have contained at least 0.68 % Cr_2O_3. Estimates of the Cr_2O_3 contents in komatiitic olivine range up to 0.3 wt% (Arndt et al., 1977). Donaldson (1982) suggested that this may be attributed to the very high temperatures of crystallization of forsteritic olivine in komatiite. This was partially confirmed by the experimental study of Murck and Campbell (1986) who showed that at 1400°C under reducing conditions olivine (Fo_{90}) could contain up to 0.3 % Cr_2O_3, falling to 0.15% (for Fo_{85}) at 1250°C. The olivine crystallizing in the picrite of the Mount Ayliff Intrusion has a composition in the range Fo_{85} - Fo_{79} (Lightfoot et al., 1984; Cawthorn et al., 1986), and hence would definitely not have contained the 0.68% Cr_2O_3 required to produce the proportion of observed chromite in olivine. Furthermore, the chromite present in the olivine is euhedral and tends to be clustered, and so is typical of magmatic processes rather than of exsolution.

Subsolidus Re-equilibration

Reaction between chromite and host silicate minerals in slowly cooled intrusions is well-documented, although extremely variable (e.g. Cameron, 1975; Hamlyn and Keays, 1979; Wilson, 1982; Hatton and Von Gruenewaldt, 1985). Lightfoot and Naldrett (1983) demonstrated this effect for chromite from the Tabunkulu lobe of the Mount Ayliff Intrusion. In order to simplify this discussion, only chromite hosted in olivine is considered here. This does not imply that there is no reaction between chromite and olivine; indeed, a geothermometer has been determined for these phases (Jackson, 1969; Evans and Frost, 1975; Fabries, 1979; Roeder et al., 1979; Sack, 1982; O'Neill and Wall, 1987). However, the major exchange will between Fe and Mg, not Ti. If Ti were to increase during slow cooling then there should be a systematic increase in Ti as a function of height in Figure 5, but the variation in Ti within one sample is greater than that between the bottom and top of the sections.

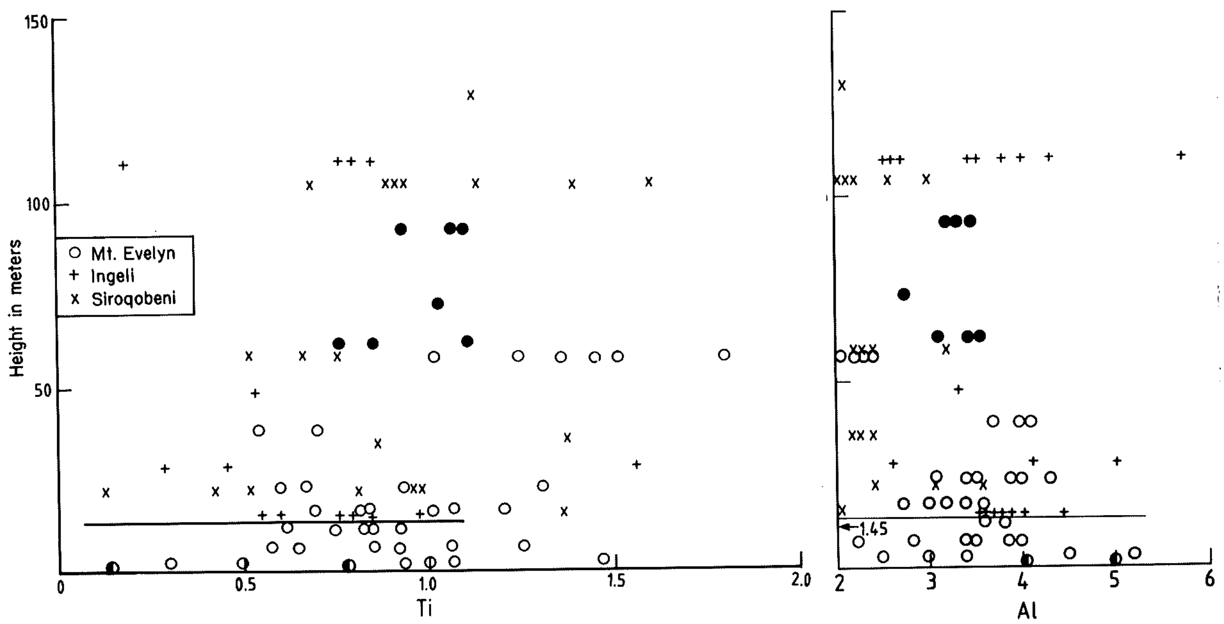

Figure 5: Plot of Ti and Al cations, based on 32 oxygen anions, versus stratigraphic height for all four profiles. Symbols the same as for Figure 3. The horizontal line indicates the range of compositions (56 analyses) found in the Waterfall Gorge section from a vertical interval of less than 10m.

Analyses of chromite enclosed in olivine from the Bushveld Complex (Hulbert and Von Gruenewaldt, 1985) and the Great Dyke (Wilson, 1982) have extremely low Ti contents. These are probably the most slowly cooled ultramafic intrusions and as these show no evidence of increased Ti in chromite then it can be inferred that high Ti contents in the chromites of the Mount Ayliff Intrusion cannot be attributed to subsolidus re-equilibration with the host olivine.

Reaction with Interstitial Liquid

Several studies have shown that chromite will react rapidly with interstitial liquid (Henderson, 1975; Ridley, 1977; Henderson and Wood, 1981). Roeder and Campbell (1985) further suggested that interstitial liquid could penetrate into olivine grains along fractures to promote such reaction. In the latter study it was suggested that the reaction involved an increase in Fe and Ti and a decrease in the Mg and Cr of the chromite. The present data are compared with those reported by Roeder and Campbell (1985) in Figure 6. No individual analyses were presented, but the average for all of the chromite grains embedded in olivine was quoted and is plotted on Figure 6. As the Jimberlana Intrusion is a large, slowly cooled, body there should have been a longer period of time for reaction to occur between chromite enclosed in olivine and interstitial liquid than in samples taken a few tens of meters above the base of the much thinner Mount Ayliff Intrusion. The low Ti content of the average analysis from Jimberlana suggests that Ti enrichment of chromite enclosed in olivine due to infiltrating magma is minimal even in that intrusion, and can be expected to be even less in the Mount Ayliff Intrusion.

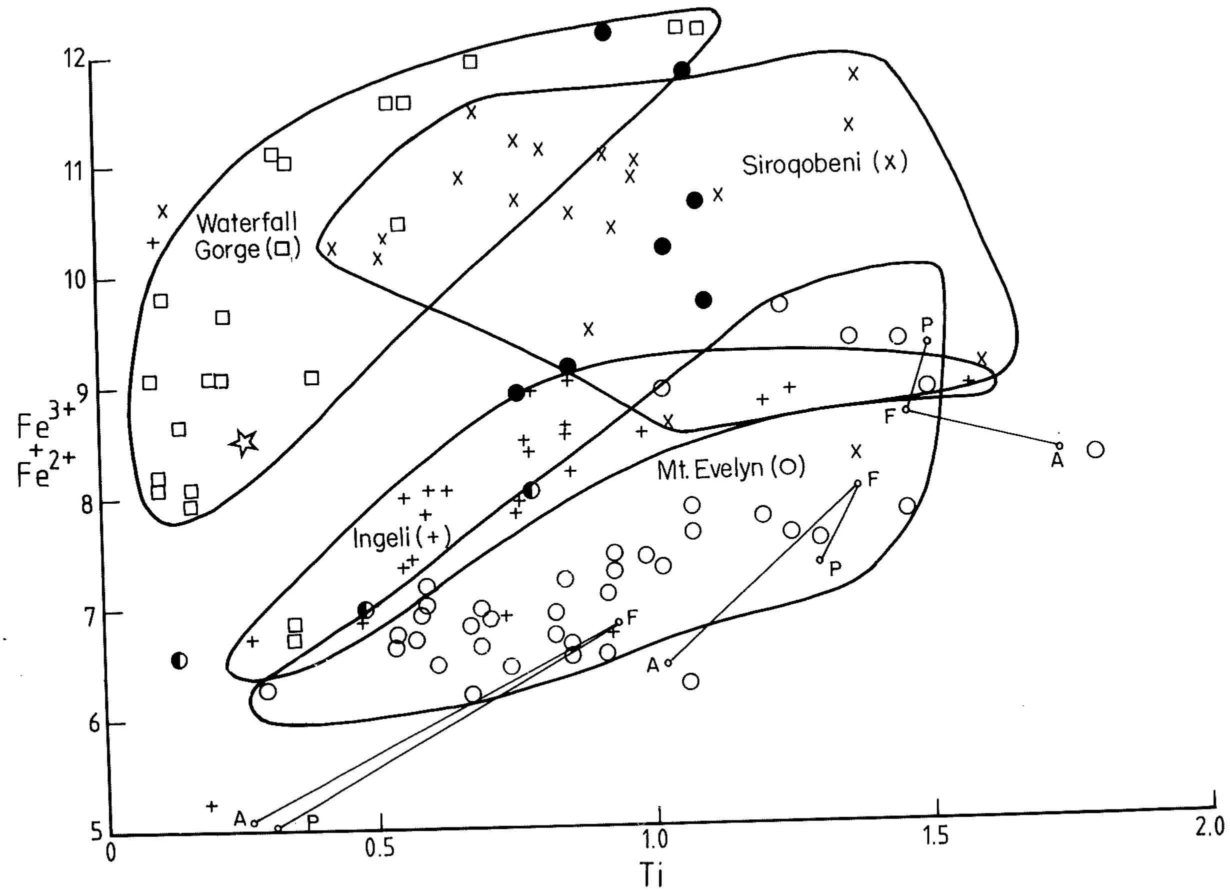

Figure 6: Plot of cations $Fe^{3+}+Fe^{2+}$ versus Ti, based on 32 oxygen anions, with the same symbols as in Figure 3. The star indicates average composition of chromite enclosed in olivine from the Jimberlana Intrusion (Roeder and Campbell, 1985). The three sets of tie-lines joining points labelled FPA refer to analyses of chromite enclosed in olivine, pyroxene, and plagioclase, respectively, in single samples from the Tabankulu lobe (Lightfoot and Naldrett, 1983).

A further test for reaction with interstitial liquid may be made by comparing compositions of chromite enclosed in olivine with those enclosed in the interstitial phases. This is demonstrated for the Tabankulu lobe in Figure 6; where chromite from olivine, plagioclase, and pyroxene in three samples are presented. One sample shows chromite to be slightly enriched in Ti when enclosed in plagioclase and pyroxene relative to that in olivine. However, in the other two cases the chromite in olivine actually contains significantly more Ti than the chromite in the other two phases, suggesting that if there has been any reaction with interstitial liquid then it has decreased the Ti content of chromite unprotected by enclosure in cumulus olivine. Thus the data for these samples argue against an increase in Ti due to reaction with interstitial liquid.

The reaction of chromite with interstitial liquid and postcumulus minerals in the Rhum Intrusion has been documented by Henderson (1975) and Henderson and Wood (1981); in the Bushveld Complex by Hulbert and Von Gruenewaldt (1985) and Hatton and Von Gruenewaldt (1985); in the Great Dyke by Wilson (1982) and in the Panton Sill by Hamlyn and Keays (1979). In none of these studies was Ti enrichment due to such processes reported. These intrusions are much thicker than the Mount Ayliff Intrusion and so time for, and extent of reaction should have been greater. There appear to be no reported analogues to the Ti-rich chromites documented here in any of the olivine-rich layered complexes even when the reaction with interstitial magma has been specifically investigated. These arguments do not disprove any contribution from interstitial magma to changes in the composition of chromite embedded in olivine, but that such reaction does not explain the distinctive compositions observed.

Primary Chromite Compositions

The distinctive features of the chromite compositions are the association of high Cr and Ti, and relatively low Al, which are very unusual for tholeiitic rocks, although they are reported in lunar spinels (Reid, 1971) and kimberlites (Haggerty, 1976; Pasteris, 1982).

The low Al contents of some of the chromite analyses is anomalous for chromite of tholeiitic origin. Experimental studies on tholeiitic compositions produce chromite with at least 3 cations Al per unit formula, and usually much higher (Hill and Roeder, 1974; Barnes, 1986; Murck and Campbell, 1986). Similarly, naturally occurring chromite in intrusive and extrusive tholeiitic rocks range from 3 - 7 cations Al (Haggerty, 1976). Some of the chromite compositions shown in Figure 5 indicate that values as low as 2 Al cations per unit formula are found, and that in the picritic rocks, generally the Al content does not excede 4 cations. Allan et al. (1988) suggested that Ti and Al demonstrate a strong negative correlation in spinels, and this is well demonstrated in the present data.

Eales and Snowden (1979) suggested that high Ti in spinels could be attributed to rapid quenching, and preservation of the true chromite composition, while during slow cooling, exsolution of ilmenite could significantly change such compositions. In none of the samples studied does the grain size or texture suggest rapid cooling or disequilibrium crystallization of chromite. Furthermore, there is no increase in Ti content close to the base of the profiles (Fig. 5), as would be expected if this process had occurred. In fact, the highest Ti contents recorded for the Mount Evelyn profile are found 58 m above the base, and close to the upper picrite - olivine gabbro contact.

The Ti content of magmatic magnetite-rich spinels is influenced by fO_2 (Buddington and Lindsley, 1964). In order to test whether anomalously reducing conditions could have caused the crystallization of high-Ti chromite, the ratio of $100*Fe^{3+}$/total Fe is plotted against Ti content of chromite in Figure 7. There is a considerable range in 100*Fe3+/total Fe ratios, but there is no negative correlation as predicted if f02 had influenced the chromite composition. In fact, for each profile there appears to be a different range of $100*Fe^{3+}$/total Fe ratios for similar Ti contents. Furthermore, from the experimental studies of Murck and Campbell (1986) and Barnes (1986), no correlation between the Ti content of the chromite and f02 was indicated.

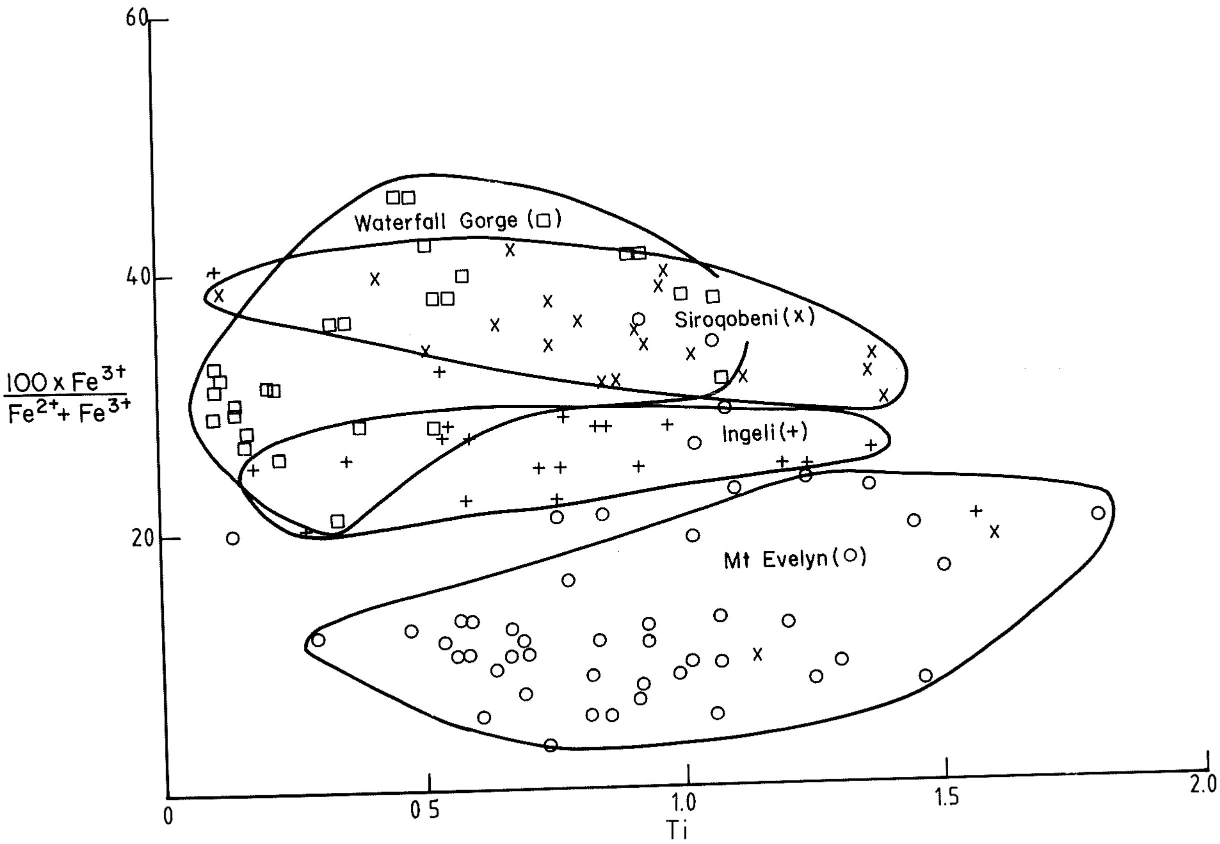

Figure 7: Plot of $100*Fe^{3+}$/8total Fe) against Ti, based on 32 oxygen anions for different profiles through Mount Ayliff Intrusion. Symbols as for Figure 3. All profiles show a similar range of Ti values but $100*Fe^{3+}$/(total Fe) varies markedly.

The extent of magmatic differentiation in these vertical sections may be evaluated by reference to the 100*Mg/(Mg+Fe) ratio of both the olivine and chromite. The former has been presented by Cawthorn <u>et al</u>. (1986), who showed that the maximum range was from Fo_{85} to Fo_{79} in the picrite, but was more iron-rich in the gabbro. The 100*Mg/(Mg+Fe) ratio of the chromite is shown in Figure 8. The general trends for the Ingeli and Siroqobeni profiles in which the samples are all picrites show little

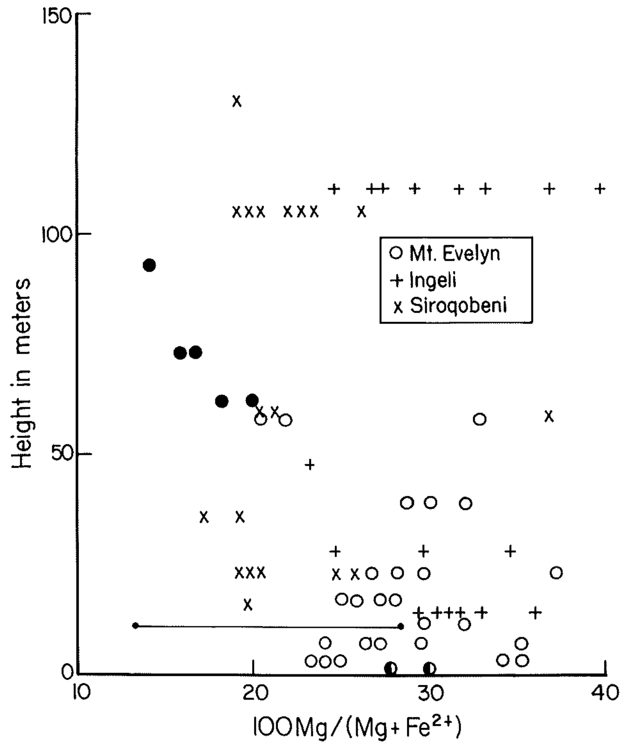

*Figure 8: Plot of 100*Mg(Mg+Fe^{2+}) versus height for profiles. The data from Waterfall Gorge are shown at constant height as the horizontal line.*

change in 100*Mg/(Mg+Fe) with height. However, for each individual sample the range of values for different chromite grains is comparable to that of the total vertical range. Hence there are no obvious differentiation trends for these profiles. Similarly, the trend for the picrites from Mount Evelyn, shows little variation in Mg/(Mg+Fe) ratio with height, but samples from the overlying olivine gabbro display decreasing values. However, as the chromite from the gabbro from the Mount Evelyn profile is not enriched in Ti compared to chromite from the picrite (Fig. 5) the high Ti contents observed for chromite in the picrite are not related to differentiation. The high Ti content of chromite associated with relatively forsteritic olivine indicates that the former is related more to the composition of the parental magma than the effects of differentiation.

Ti-Rich Chromite in other Tholeiitic Suites

No other examples of chromite with such high Cr and Ti contents in tholeiitic suites are known to the authors. The closest analogues are some spinels from Hawaii (Evans and Moore, 1968; Helz, 1987, Nicholls and Stout, 1988; Wilkinson and Hensel, 1988). They occur in lavas and lava lakes which contain relatively high TiO_2 (1.6 wt % TiO2 at 16 wt % MgO; Wilkinson and Hensel, 1988). From a single sample, Nicholls and Stout (1988) reported two chromite analyses with a range of composition from 0.38 to 0.92 cations Ti. This is not necessarily the total range in Ti for this sample, but even so illustrates that the range in Ti contents for individual samples shown in Figure 5 could be largely the result of magmatic processes.

Spinels from the Snake River Plain Basalts provide another possible analogue (Thompson, 1973), although none are as chrome-rich as those reported here. Again, the magmas contain a high TiO_2 content (3 - 4 wt % TiO_2 at 5 - 7wt % MgO).

A Model for the Crystallization of the Mount Ayliff Intrusion

Lightfoot et al. (1984, 1987) argued that the intrusion had formed by olivine fractionation and accumulation from a typical continental basalt. This was debated by Cawthorn et al. (1988), based on the fact that the coexistence of olivine and highly magnesian ilmenite is unlikely in such a magma. They suggested that the variable Mg content of the ilmenite resulted from the mixing of a low-Ti and a high-Ti magma, the latter being analogous to the Letaba Basalts documented from several localities in the Karoo Volcanic Province (Bristow and Saggerson, 1983). It can also be argued that the occurrence of coexisting chromite and ilmenite is unknown in low Ti tholeiitic magmas, but was reported by Evans and Moore (1968) from the high-Ti lavas in Hawaii. The presence of magnesian ilmenite and the high Ti content of chromite both suggest that the magma from which they formed was Ti-rich. These distinctive features are here integrated into a model for the crystallization of the Mount Ayliff Intrusion, with special emphasis on the origin of the chromite compositional relationships. The geochemical constraints imposed by the chromite compositions are shown in Figure 9 and the proposed processes operating within the magma chamber are illustrated in Figure 10.

In Figure 9 chromite from each sample is designated differently so that trends as a function of height are seen. For reference, data from other Karoo intrusions (Eales, 1979, Eales and Snowden, 1979, Eales et al., 1980), Rhum Intrusion (Henderson and Wood, 1981) and Jimberlana Intrusion (Roeder and Campbell, 1985) are included. Collectively, these define the trend of typical low-Ti tholeiite, which is designated the Karoo trend, A, in Figure 9D. Also shown in Figure 9D is the Hawaii trend, C, based on data from Wilkinson and Hensel (1988), which is produced by crystallization of a Ti-rich tholeiite.

The data from the limited section at Waterfall Gorge (Fig. 9A) plot close to the Karoo trend, and are interpreted to have crystallized from a low-Ti magma.

Analyses from the Mount Evelyn profile at 1 and 3 m scatter, some overlapping the Karoo trend while others are distinctly enriched in Ti and Cr relative to this trend. A model is suggested in Figures 10A, B, and C, whereby the injection of a low-Ti tholeiite is followed by a high-Ti tholeiite closer in composition to that which produced the Ti-rich chromite from Hawaii. If it is related to the Letaba-type of basalt it would be enriched in incompatible elements including alkalis and water, and hence less dense than the low-Ti magma. Coprecipitation of chromite and olivine from these two magmas leads to the range of chromite compositions observed (dashed line 1 in Fig. 9D). Samples from 7-39 m contain chromite with very high Ti and Cr (Fig. 9A) and formed from the high-Ti magma. These samples show no systematic change as a function of height. With differentiation the chromite composition evolves to lower Cr and higher Ti values as shown by the sample at 58 m (Fig. 9A). The general Mount Evelyn trend, B, is shown in Figure 9D for samples from 7 to 58 m. From 62 to 73 m there is a change in trend with compositions becoming depleted in Ti at constant Cr. These compositions plot between analyses from the sample at 58 m and the

Figure 9: *Plot of Cr versus Ti for chromite from different profiles from the Mount Ayliff Intrusion compared with analyses from elsewhere in the Karoo, from other intrusions and from Hawaii. A. Data for Waterfall Gorge (crosses) and Mount Evelyn, each sample from Mount Evelyn being differently designated, and stratigraphically grouped. For reference, the trend for chromite from certain Hawaiian lavas (Wilkinson and Hensel, 1988) is shown as a heavy line. B. Data for Siroqobeni profile; Hawaiian trend also included. C. Chromite from other Karoo intrusive rocks are shown by open circles (Eales and Snowden, 1979; Eales et al., 1980). Analyses from Rhum Intrusion are shown as crosses, those designated I are inclusions in olivine, E external to olivine (Henderson and Suddaby, 1971; Henderson and Wood, 1981). Analysis of chromite in olivine from the Jimberlana Intrusion denoted by star (Roeder and Campbell, 1985). Samples joined and denoted F, A, and P are chromites enclosed in forsteritic olivine, plagioclase, and pyroxene, respectively, from individual samples from the Tabankulu lobe (Lightfoot and Naldrett, 1983). Hawaiian trend included for reference. D. Schematic and simplified trends for different suites of tholeiitic rocks. Karoo, Hawaii, and Mount Evelyn refer to typical trends displayed by chromites from these localities. Dashed trends 1, 2, and 3 refer to trends in the Mount Evelyn and Siroqobeni profiles (see text).*

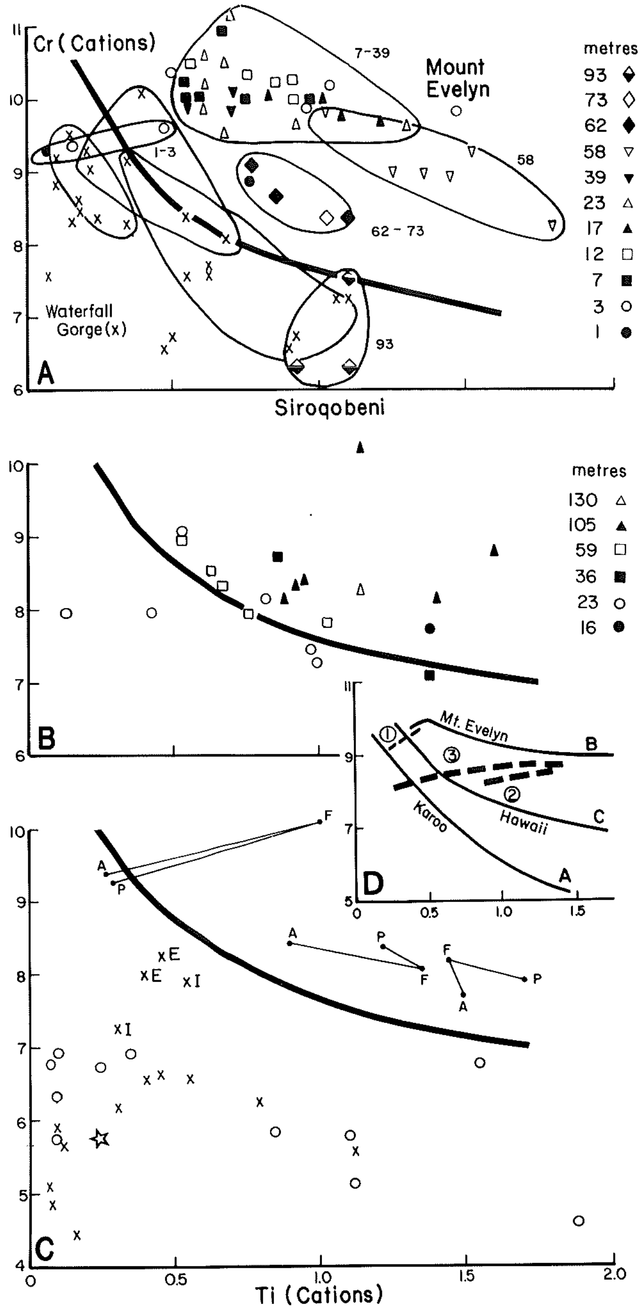

Figure 9: *See facing page*

middle of the Karoo trend. This is attributed to mixing between high Ti and low Ti magmas (dashed line 2 in Fig. 9D). The two magmas in the intrusion may mix as a result of changing densities (Figs 10D and E), or there may be injection of more low Ti magma (Fig. 10F). Further differentiation of this mixed magma produced the chromite compositions observed for the sample at 93 m, which is relatively low in Cr and high in Ti (Fig. 9A). The data for the Siroqobeni profile are shown in Figure 9B.

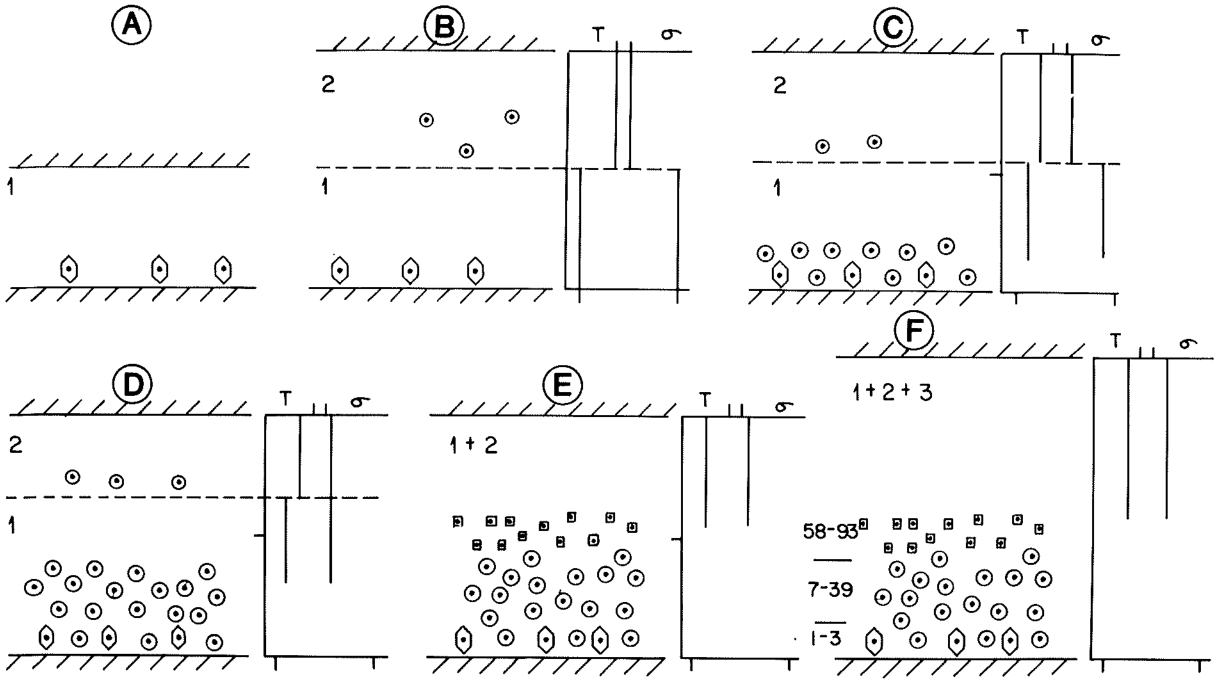

Figure 10:

Schematic sections through the intrusion depicting its evolution, by crystal accumulation, magma addition and fractionation. In stage A, magma type 1 (low in Ti) is intruded and produces low-Ti chromite in olivine (hexagons with a black dot). This produces a significantly thick layer at Waterfall Gorge, but is thin at Mount Evelyn. Stage B results when type 2 magma (high Ti) is added. The model requires it to be less dense than type 1 possibly because it is enriched in incompatible elements and volatiles. Inputs of less dense magma follow the fluid dynamic constraints discussed by Huppert et al. (1986). This magma crystallizes high-Ti chromite embedded in olivine (circles). They may remain suspended in the upper layer by turbulent convection (Tait, 1985) and eventually dump a thick succession of relatively homogeneous chromite and olivine (stage C, circles with dots). In the Mount Evelyn profile the nearly constant composition of the chromite from 7 to 39m (Fig. 9A) may result from this process. The density (σ) and temperature relationships of the two magmas are shown qualitatively on the right of each figure. The upper magma is required to be hotter than the lower so that heat will diffuse downwards and so terminate crystallization of low-Ti chromite, at least while olivine is accumulating. Stage C shows the temperature in the upper layer decreasing and the bulk density increasing to approach that of the lower layer. Crystallization of olivine will decrease the residual liquid density. Hopwever, if the olivine remains suspended in the upper magma the bulk density will increase. Alternatively, olivine and plagioclase may be coprecipitating causing the density of the residual liquid to increase but the plagioclase remains suspended in the upper liquid layer while olivine sinks, as suggested by Wilson et al. (1989) in the Great Dyke. At stage D the densities of the two layers are now equal and so mixing is initiated. The chromite and olivine from this mixing process is denoted in stage E by the square symbol enclosing a dot, and these will be intermediate between the high- and low-Ti compositions, as shown by the trends in Figure 9D, dashed lines 2 and 3, and corresponding to the compositions at 58 to 73 m in the Mount Evelyn profile. Stage F represents an alternative scenario to stages D and E, where further adddition of and entrainment by more magma similar to type 1 composition (here denoted 3) occurs. The numbers (1-3); 7-39; 58-93) in stage F correspond to the heights in metres in the Mount Evelyn profile where such chromite inclusions in olivine occur.

These display a great variation in Ti with little range in Cr. Furthermore, there is no general trend of changing composition with height; the lowest sample at 16 m having high Ti and the next sample at 23 m having some of the lowest Ti contents. The data are shown in Figure 9D as the trend denoted 3 and define a trend similar to, but more extensive than that seen for the samples from 62 to 73 m in the Mount Evelyn profile. These data are therefore interpreted as reflecting prolonged and incomplete mixing between a differentiated high-Ti magma (equivalent to the magma that produced the sample at 58 meters in the Mount Evelyn profile) and a low-Ti magma. It is therefore possible that the lateral equivalent of the lowest part of the Mount Evelyn profile (below 58 m) is either missing or very compressed in the Siroqobeni section.

The general trend for samples from Ingeli is exactly the same as for Siroqobeni, and so is not plotted in Figure 9. The slow mixing of magmas may also have taken place in this lobe of the intrusion.

The proposed intrusion relationships in Figure 10 are supported by evidence from the compositions of chromite interstitial to the olivine. The data from the Tabankulu lobe (Fig. 9C) show that chromite enclosed in olivine in two samples contains higher Ti than spinel in the interstitial phases. If the chromite and its enclosing olivine formed from the high-Ti magma and sank into the basal layer of low-Ti magma (Fig. 10C) the intercumulus minerals would form from this latter magma. Specifically the chromite interstitial to the olivine would be predicted to have a composition along the Karoo trend in Figure 9D, which is exactly what is observed. The same also applies to the pyroxenes forming the interstitial component in the picrite. They also formed from the low-Ti magma, and hence have low-Ti contents.

CONCLUSIONS

The coexistence of chromite and ilmenite in picrite in the Mount Ayliff Intrusion is unique in tholeiitic intrusive suites. These are also the most Ti-rich chromites and Mg-rich ilmenites recorded in tholeiitic rocks. These features suggest that they formed from a Ti-rich magma. The chromites are also depleted in Al which is consistent with the known incompatibility betwen Al and Ti in spinels.

Comparison with chromites in other intrusions which show evidence of reaction with interstitial liquid or of subsolidus re-equilibration indicates that these unusual compositions cannot be attributed to such processes.

Low-Ti chromites with compositions typical of tholeiitic rocks occur as inclusions in olivine in gabbroic rocks above and below the picrite, suggesting that these formed from a low-Ti magma. It is therefore suggested that the intrusion is the product of multiple injection of contrasting magma types; with different Ti contents. In one profile (Mount Evelyn) the chromite in the picrite shows evidence for restricted differentiation. In contrast, at the contact with gabbro and in two other profiles (Siroqobeni and Ingeli) there is a wide range in Ti content with little change in Cr content of the chromite which is thought to relate to mixing between two magmas, rather than differentiation.

In another profile (Tabankulu), the chromite interstitial to the olivine is low in Ti, while the chromite included in the olivine is enriched in Ti. This is interpreted as the result of sinking of olivine with its included Ti-rich chromite from an upper stratified high-Ti magma into a low-Ti magma which then crystallized low-Ti chromite among the intercumulus phases.

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