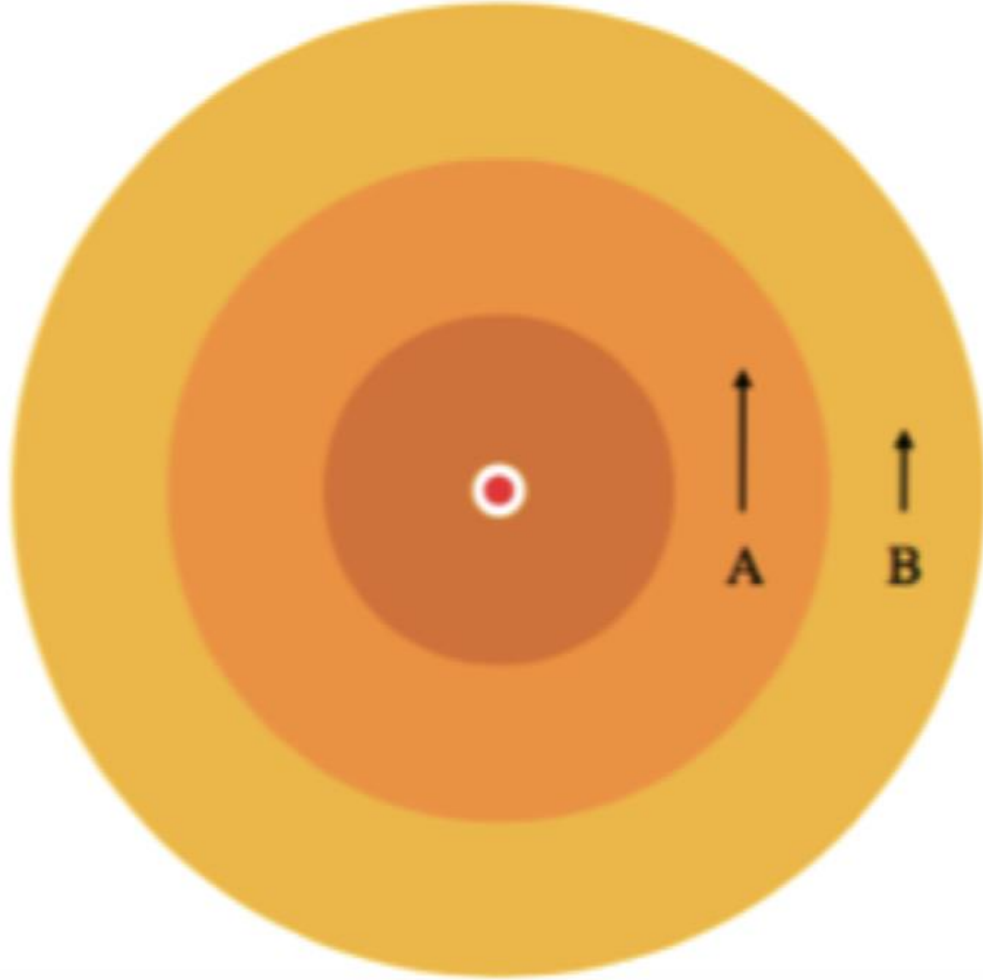


## Shear + Viscosity = Heating



If there is shear between the fluid elements A and B, there will be exchange of bulk momentum between the fluids.

Along with it, the work of the viscous force leads to conversion of kinetic energy into heat.

In astrophysical contexts, the viscosity is not molecular but turbulent (more about it next module of the class), but the result is similar: massive viscous heating.

## Viscous heating powers accretion disks



## Reynolds number

Given the Navier-Stokes equation

$$\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} = -\frac{1}{\rho} \nabla p + \mathbf{f} + \nu \nabla^2 \mathbf{u}$$

Compare the ratio of inertial and the viscous term

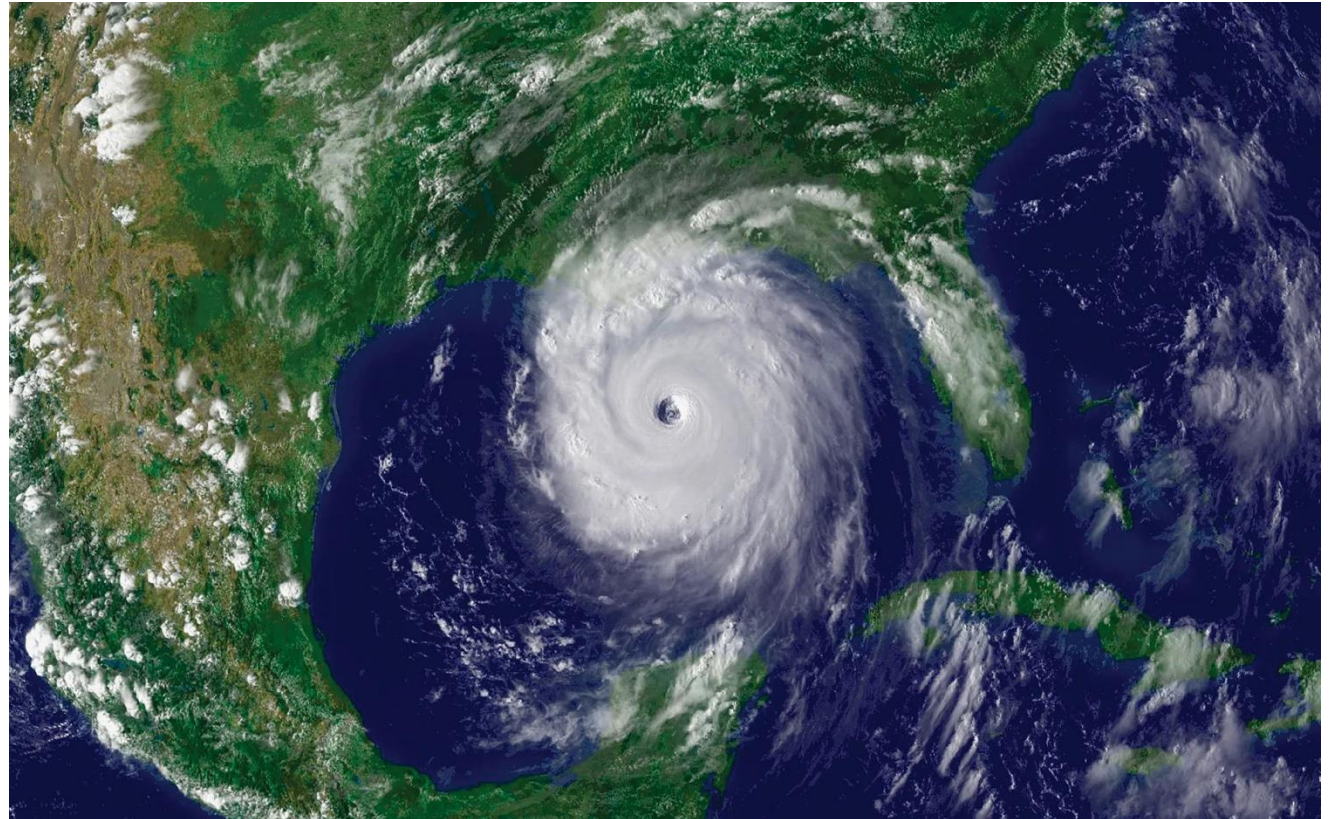
$$\frac{(\mathbf{u} \cdot \nabla) \mathbf{u}}{\nu \nabla^2 \mathbf{u}}$$

The ratio of the two terms is a dimensionless quantity, the *Reynolds number*

$$\text{Re} = \frac{UL}{\nu}$$

## Reynolds number

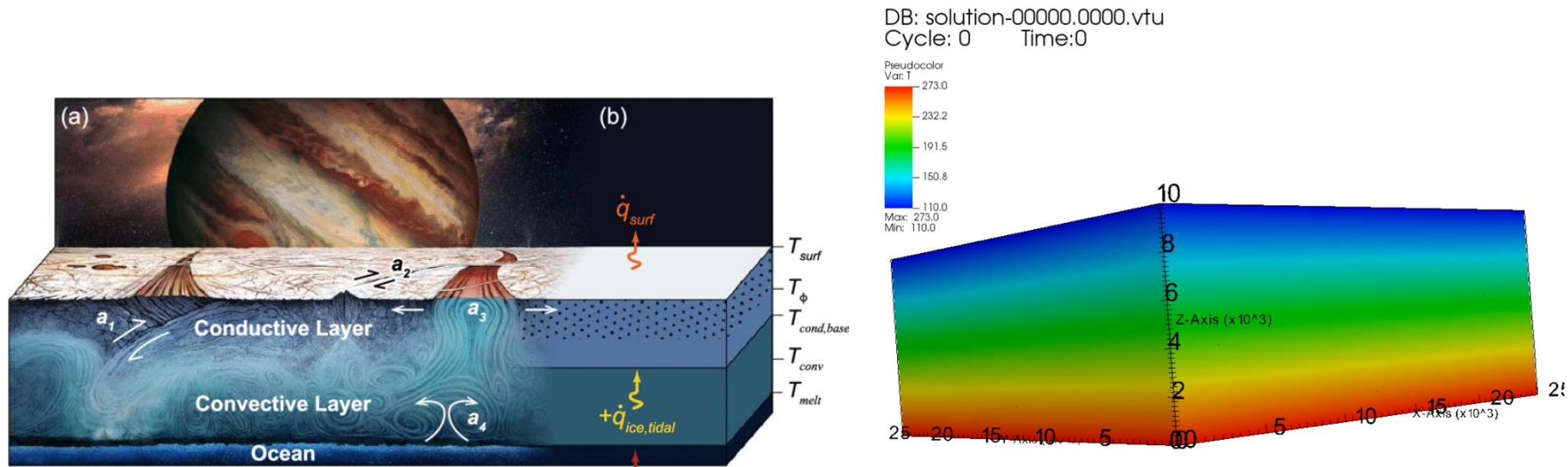
The viscosity of air is  $10^{-1} \text{ cm}^2/\text{s}$ .  
Usual reasonable values for the speed  
and length and estimate the Reynolds  
number of a hurricane.





# Most fluids in the Universe are high Reynolds number

## Notable exception: geophysics!



$$0 = \nabla \cdot \mathbf{u},$$

$$0 = \nabla \cdot \boldsymbol{\sigma} - \nabla p + \text{Ra } T \hat{\mathbf{z}},$$



Equations of motion:

$$\frac{\partial T}{\partial t} + (\mathbf{u} \cdot \nabla) T = \nabla^2 T + q,$$

$$\sigma_{ij} = 2\eta \dot{\epsilon}_{ij} = \eta \left( \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right),$$

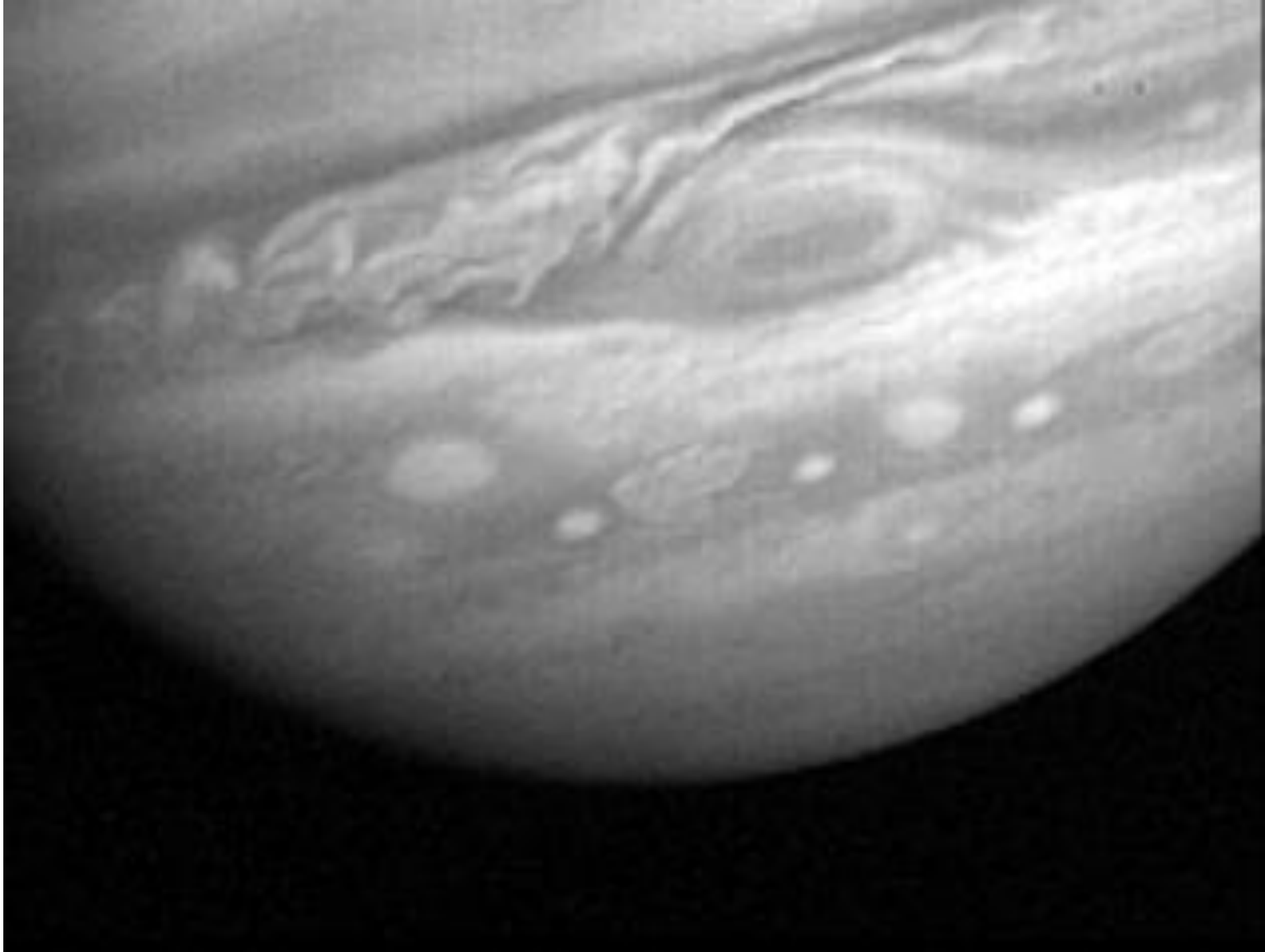
## Vorticity

Vortices are like entities plastered on the fluid, and moving with the flow.

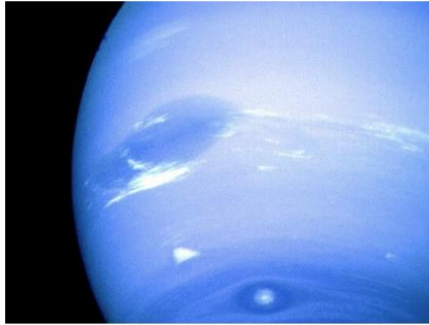
We can define a quantity, the *vorticity*, the curl of the velocity

$$\boldsymbol{\omega} = \boldsymbol{\nabla} \times \boldsymbol{u}$$

and show that under some circumstances that are often realized in nature, the vorticity is conserved.



## Vortices – an ubiquitous fluid mechanics phenomenon



## Vorticity Equation

$$\frac{\partial \omega}{\partial t} = \underbrace{-(\mathbf{u} \cdot \nabla) \omega}_{\text{advection}} - \underbrace{\omega (\nabla \cdot \mathbf{u})}_{\text{compression}} + \underbrace{(\omega \cdot \nabla) \mathbf{u}}_{\text{stretching}} + \frac{1}{\rho^2} \underbrace{\nabla \rho \times \nabla p}_{\text{baroclinicity}} + \underbrace{\nu \nabla^2 \omega}_{\text{dissipation}}$$

The diagram illustrates the Vorticity Equation with terms labeled by color-coded arrows. The equation is:  $\frac{\partial \omega}{\partial t} = -(u \cdot \nabla) \omega - \omega (\nabla \cdot u) + (\omega \cdot \nabla) u + \frac{1}{\rho^2} \nabla \rho \times \nabla p + \nu \nabla^2 \omega$ . The terms are labeled as follows: 

- advection**: points to  $-(u \cdot \nabla) \omega$  (grey arrow)
- compression**: points to  $-\omega (\nabla \cdot u)$  (blue arrow)
- stretching**: points to  $(\omega \cdot \nabla) u$  (red arrow)
- baroclinicity**: points to  $\frac{1}{\rho^2} \nabla \rho \times \nabla p$  (green arrow)
- dissipation**: points to  $\nu \nabla^2 \omega$  (orange arrow)



## The baroclinic term

$$\frac{1}{\rho^2} \nabla \rho \times \nabla p$$

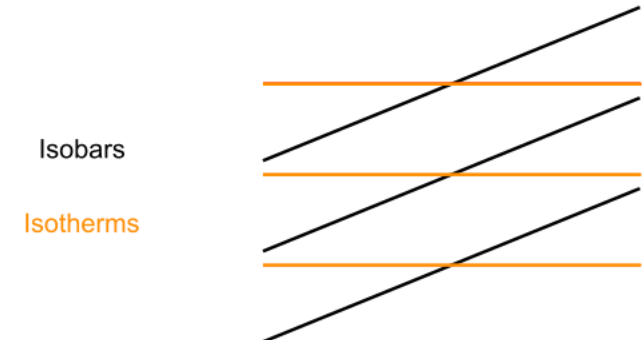
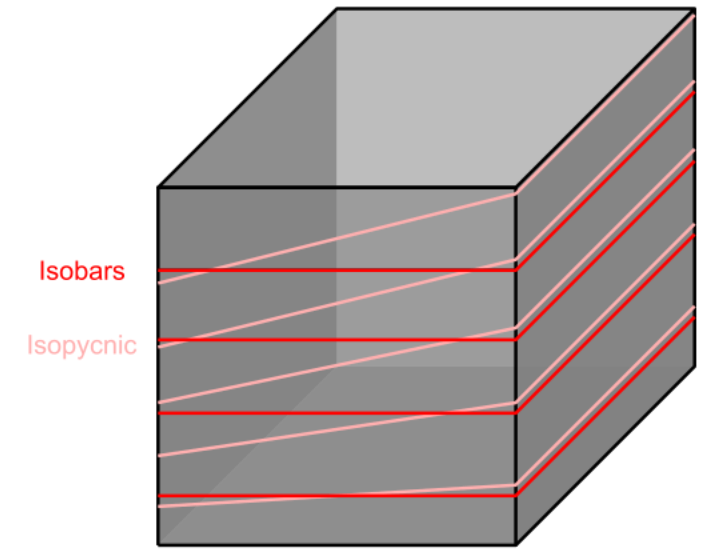
A **barotropic** fluid has  $p = p(\rho)$ . The surfaces of constant pressure and constant density are aligned, so the gradients point in the same direction.

The word barotropic comes from *baro* (pressure) + *tropos* (direction). Meaning pressure gradient is in the same direction as the density gradient.

A **baroclinic** fluid has  $p = p(\rho, T)$ . The surfaces of constant pressure and constant density can be inclined, with the gradients pointing in different directions.

The word baroclinic comes from *baro* (pressure) + *inclination* (direction). Meaning pressure and density gradients are misaligned.

## Baroclinic fluid



## The baroclinic term

$$\frac{1}{\rho^2} \nabla \rho \times \nabla p$$

Notice that the polytropic equation of state  $p = K \rho^\gamma$  is barotropic. The interior of stars and giant planets are well-approximated by polytropes.

$K$  is the entropy. For barotropic fluids, the entropy is constant.

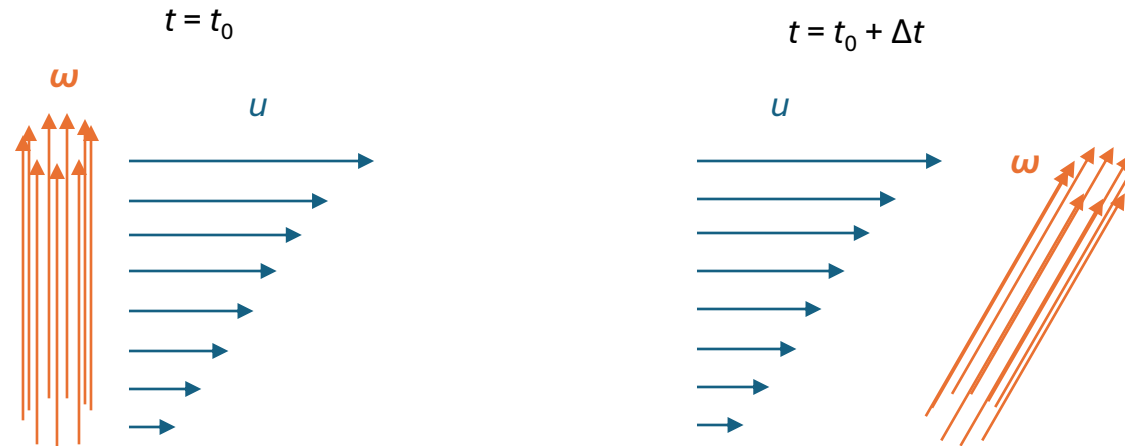
The condition of baroclinicity then translates simply into: *varying entropy*.

An entropy gradient leads to vortex generation by *convection*.



## The stretching term

$$\frac{\partial \omega}{\partial t} = (\omega \cdot \nabla) u$$

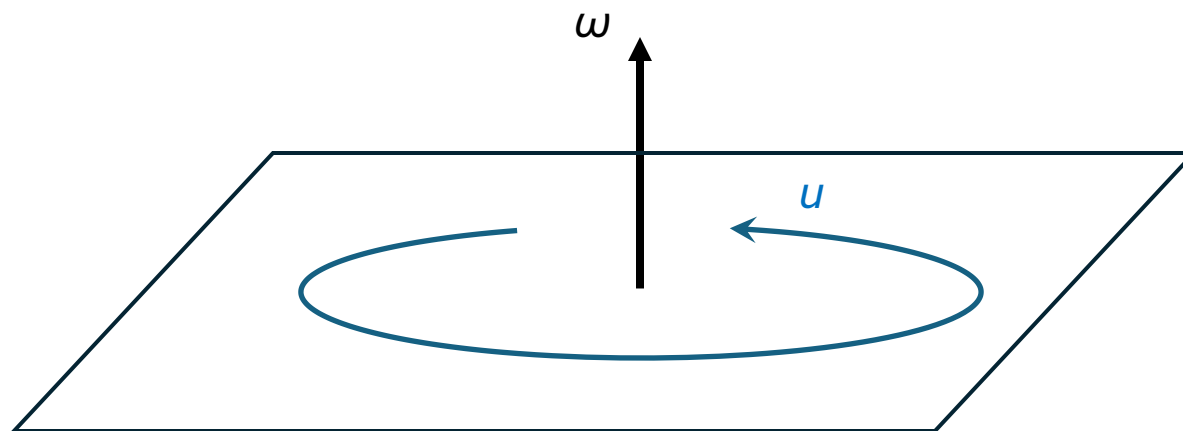


Stretching modifies vorticity along the direction of shear.  
In essence, stretching is differential advection.

In components, one starts from pure y vorticity, and builds an x vorticity component.

## The stretching term

$$\frac{\partial \omega}{\partial t} = (\omega \cdot \nabla) u$$



For 2D, the vorticity always points out of the flow plane, so the stretching is zero.

# Vorticity Equation

Cancels when:

$$\frac{\partial \omega}{\partial t} = \underbrace{-(\mathbf{u} \cdot \nabla) \omega}_{\text{advection}} - \underbrace{\omega (\nabla \cdot \mathbf{u})}_{\text{compression}} + \underbrace{(\omega \cdot \nabla) \mathbf{u}}_{\text{stretching}} + \underbrace{\frac{1}{\rho^2} \nabla \rho \times \nabla p}_{\text{baroclinicity}} + \underbrace{\nu \nabla^2 \omega}_{\text{dissipation}}$$

*incompressible*      *2D*      *barotropic*      *Re >> 1*

Under these conditions:

$$D\omega/Dt = 0$$

Vorticity is conserved!