Seismic Traveltime Inversion with Quantum Annealing

Hoang Anh Nguyen^{1,*} and Ali Tura¹

¹Department of Geophysics, Colorado School of Mines, Golden, 80401, Colorado, USA

ABSTRACT

This study demonstrates the application of quantum computing based quantum annealing to seismic traveltime inversion, a critical approach for inverting highly accurate velocity models. The seismic inversion problem is first converted into a Quadratic Unconstrained Binary Optimization problem, which the quantum annealer is specifically designed to solve. We then solve the problem via quantum annealing method. The inversion is applied on a synthetic velocity model, presenting a carbon storage scenario at depths of 1000-1300 meters. As an application example, we also show the capacity of quantum computing to handle complex, noisy data environments. This work highlights the emerging potential of quantum computing in geophysical applications, providing a foundation for future developments in high-precision seismic imaging.

Introduction

Quantum computing is an emerging field with significant promise for various scientific and engineering disciplines. As we stand at the frontier of this technological revolution, early-stage research in quantum computing is crucial for the advancement of geophysics. Numerous studies have begun to explore the integration of quantum computing within this field, highlighting its immense and revolutionary potential¹. For instance, quantum annealers can perform well in solving tomography optimization problems². The quantum computing is applied for binary-value full waveform inversion, addressing issues related to velocity variations³. In the frequency domain, the seismic wave equation can be reduced to a system of linear equations, allowing for the application of quantum annealing⁴. Furthermore, it has been shown that quantum annealing impedance inversion with L1 norm regularization can dramatically enhance accuracy and anti-noise capabilities⁵.

A quantum annealer is a specific type of quantum computer designed to solve optimization problems⁶. The quantum annealing process in quantum annealers can find the minimum energy state of a system, corresponding to the optimal solution of a given problem⁷. This process is achieved by utilizing quantum fluctuations, allowing the system to tunnel through energy barriers⁸. While there are various types of models in quantum computing⁹, this particular feature allows quantum annealing to efficiently explore complex energy landscapes, making them particularly well-suited for solving optimization problems.

Most previous attempts to address seismic problems using quantum annealers have primarily involved relatively simple models ^{10,11}. For conventional approach by classical computers, the cross-well seismic inversion between boreholes can be computationally expensive ¹², necessitating the development of new methods to tackle these challenges. Therefore, in this study, we aim to advance this line of research by applying quantum annealing to a complex problem: Seismic traveltime inversion of the velocity model between two boreholes. Our focus is on developing an inversion strategy that can accurately invert the velocity model with noisy data despite the limitation of the quantum hardware, specifically targeting carbon storage scenarios at depths of 1000-1300 meters. We use quantum annealer at D-Wave Advantage System, which has at least 5000 qubits ¹³. Clearly, this travel-time inversion method can be applied to other acquisition geometries and data such as surface seismic, vertical seismic profile (VSP), earthquake or micro seismic data.

Results

We start with exact traveltime data without noise and constant initial velocity model v_{ini} of 3475 m/s. The initial model and the results of the inverted model v_{inv} at each iteration obtained after the first 9 iterations indicate rapid convergence (Fig. 1). Notably, in the first iteration, the carbon storage area is immediately identified with high precision.

The component-wise relative errors e_{ij} between the true $v_{true,ij}$ and final inverted velocity model $v_{final,ij}$ after 10 iterations is shown in Fig. 2. The component-wise relative errors are calculated by $e_{ij} = |v_{inv,ij} - v_{true,ij}|/|v_{true,ij}|$. The most significant errors occurs in the shallow and deep regions with weakest constraints, yielding a maximum error value of about 0.00326. In contrast, the carbon storage area, spanning depths from 1100 to 1200 m, demonstrates exceptionally low errors due to the high

^{*}hoanganh_nguyen@mines.edu

Inversion Process

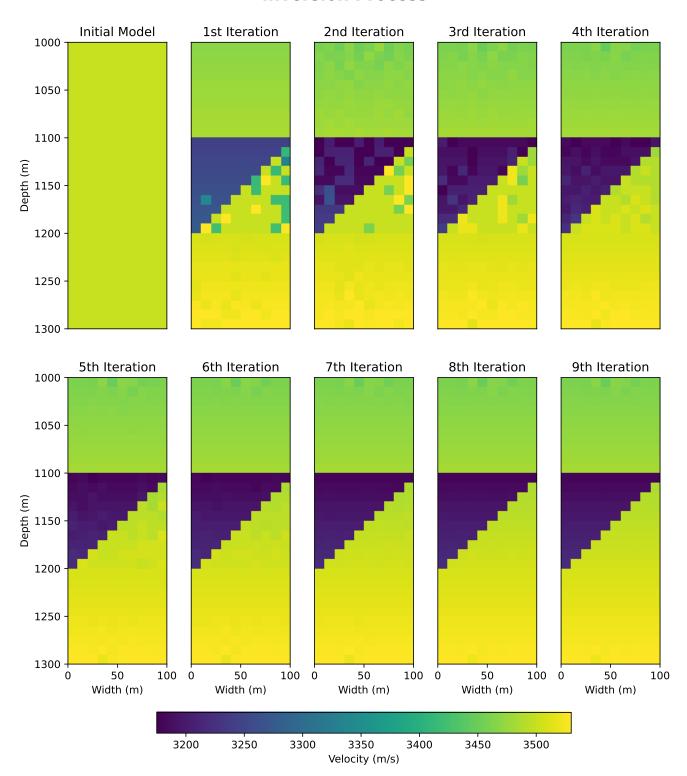


Figure 1. The starting model v_{ini} and the inverted velocity model v_{inv} over the first 9 iterations with exact, noise-free traveltime data.

ray coverage. The high-accuracy result underscores the effectiveness of the quantum annealing approach to the traveltime inversion.

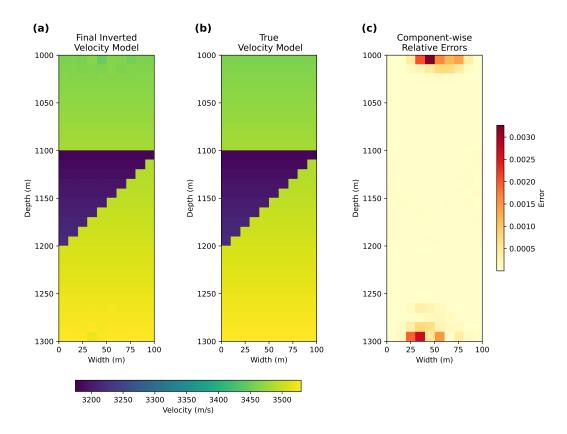


Figure 2. Velocity models and error: (a) final inverted model after 10 iterations $v_{inv,ij}$, (b) true model $v_{true,ij}$, and (c) the component-wise relative errors e_{ij} between the final inverted and true velocity model.

For seismic traveltime inversion problem, the quantum annealing method and the classical linear least squares approach produce similar levels of error under ideal conditions, where data was free of noise¹⁴. However, real-world data often contain random noise, making it essential to assess the robustness of these methods under more realistic conditions. Therefore, we introduce random noise into the synthetic data, varying the random noise levels from 0.01% to 1% of the traveltime, and evaluate the performance of both standard linear least squares and quantum annealing method.

The results reveal a stark contrast in the sensitivity of these methods to noise. The linear least squares method is highly sensitive to noise (Fig. 3a, b, c). At the noise level of 0.01% of the traveltime, while this method can identify the region of carbon storage, the deviation of the inverted model from true model is considerable. As the noise level increases to 0.1% and 1% of the traveltime, the linear least squares method fails to accurately determine the velocity model. This significant sensitivity limits its effectiveness in processing noisy seismic data, posing challenges for practical applications. In contrast, with the same noise, the quantum annealing method is more robust (Fig. 3d, e, f). At the 0.01% noise level, the differences between the noise-free model and the results obtained are negligible. Differences start to appear at the 0.1% noise level, primarily in shallow and deep areas where there is less constraint. Remarkably, at the 1% noise level, the quantum annealing method still effectively reproduces the velocity model. In areas with high ray coverage, these differences are small. This analysis underscores the potential of quantum annealing method for handling noisy seismic data more effectively than the classical linear least squares method.

Discussion

In this study, we demonstrate several advancements to enhance the robustness and accuracy of seismic traveltime inversion results using the quantum annealing method. A significant improvement involves incorporating non-uniform source and receiver spacing. By using non-uniform spacing, we achieve better ray coverage of both shallow and deep regions, thereby increasing the overall accuracy of the final inverted velocity model.

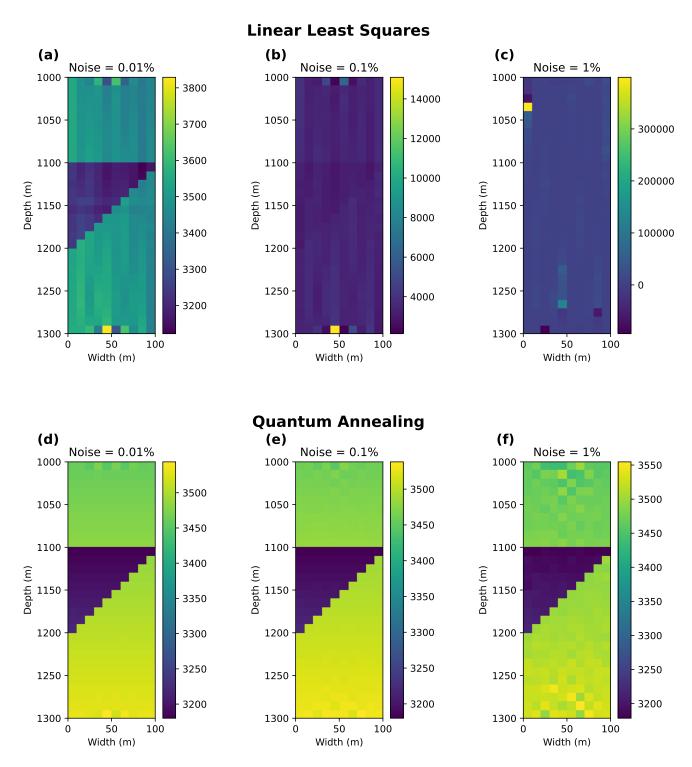


Figure 3. Traveltime inversion results with different noisy data using linear least squares (a, b, c) and quantum annealing method (d, e, f)

The initial problem is broken down into smaller sub-problems, allowing for better constraints on the boundary L during the inversion process and accommodating hardware limitations. Solving these sub-problems individually provides more reliable solutions and facilitates parallel processing, therefore it significantly reduces computational time.

For practical applications where noise is an unavoidable factor, quantum computing demonstrates a remarkable ability to

handle noisy data effectively. The quantum annealing method significantly outperforms classical linear least squares method under noisy conditions, maintaining high efficiency and accuracy.

The key advantage of quantum annealing method is its ability to locate the global minimum of the objective function more effectively than classical methods, which are often trapped in local minima. Quantum tunneling allows the quantum system to explore a broader solution space and tunnel barriers that would hinder classical optimization methods¹⁵. This method can result in more accurate and reliable inversion outcomes, particularly in complex and heterogeneous environments. The quantum annealer solves problems using principles of quantum mechanics, which inherently depend on the operating environment of the machine. Consequently, running the same problem multiple times may yield slightly different results because of the probabilistic nature of quantum computations^{16,17}. However, we expect the advancements in quantum technology will decrease this variability and also reduce computational time.

Conclusion

The integration of quantum annealing in seismic traveltime inversion represents a potentially major advancement in geophysics. We utilize the open access D-Wave Advantage quantum annealer to solve the inversion problem for a synthetic velocity model, representing a carbon storage scenario. The quantum annealing recursive method, can give a good inverted velocity result after 10 iterations. Notably, the quantum method outperforms classical linear least squares method in dealing with noisy data, where classical methods can sometimes struggle. Despite the promising results, the current state of quantum computing is not without its challenges. The slight variability in outcomes due to quantum noise ^{18–20} necessitates multiple runs to ensure reliable results. However, these challenges are expected to diminish as advancements in quantum technology continue. This study suggests that quantum annealing could revolutionize seismic inversion processes, offering more accurate solutions in scenarios where traditional methods are computationally intensive or even infeasible. As quantum computing matures, its applications in geophysics are likely to expand, encompassing more complex and larger-scale problems. This research not only underscores the potential of quantum annealing in seismic inversion but also sets the stage for future exploration into other areas of inverse problems. We expect that continued development of quantum computing technologies promises to unlock new capabilities, possibly making it an indispensable tool for new challenges.

Methods

Quantum Annealing

Quantum computing is rapidly emerging as a pivotal area of scientific and technological advancement, attracting considerable investments and interest due to its profound potential^{21–23}. Unlike classical computers that use bits, which exist only in states of 0 or 1, quantum computers employ quantum bits, or qubits. Qubits possess unique properties like superposition, entanglement, and interference^{24–26}, enabling them to perform certain complex computations beyond the capabilities of classical computers^{27,28}. Qubits can be constructed from various physical systems such as photons, trapped atoms, nuclear magnetic resonance, quantum dots, dopants in solids, and superconductors²⁹. Previous research^{30–34} has provided evidence that quantum computing possibly surpasses classical computers in terms of processing speed and efficiency of certain problems.

The quantum annealing process facilitates finding the global minimum of a cost function efficiently. This process can be described using the real-time Schrödinger equation³⁵:

$$i\hbar \frac{d}{dt}|\Psi(t)\rangle = H(t)|\Psi(t)\rangle$$
 (1)

where $|\rangle$ is the ket of the Diract notation³⁶, i is the imaginary unit, t is time, \hbar is reduced Planck's constant, $\Psi(t)$ is the wave function, $|\Psi(t)\rangle$ is the quantum state vector, H is the Hamiltonian represents the total energy of the quantum system^{37,38}. If \hbar is set as 1, the Eq. 1 becomes:

$$i\frac{d}{dt}|\Psi(t)\rangle = H(t)|\Psi(t)\rangle$$
 (2)

The Hamiltonian in quantum annealing can be composed of two components^{39,40}:

$$H(t) = A(t)H_0 + B(t)H_1$$
 (3)

where H_0 is the initial Hamiltonian, representing a system with an initial ground state. H_1 is the final Hamiltonian, whose ground state encodes the solution to the optimization problem. A(t) and B(t) are time-dependent coefficients. A(t) and B(t) are set in range from 0 to 1 so that $A(t_0) \gg B(t_0)$ at the initial time t_0 and $B(t_1) \gg A(t_1)$ at the final time t_0 . During the process, A(t) monotonically decreases while B(t) monotonically increases. At the start of the annealing process, $H(t) \approx H_0$. At the end

of the annealing process, $H(t) \approx H_1$. Thus, the system transitions from ground state of H_0 to the ground state of H_1 . If H(t) changes sufficiently slowly, the state evolves adiabatically⁴¹.

The problems are then often mapped onto a Quadratic Unconstrained Binary Optimization (QUBO) or Ising model⁴²:

QUBO:
$$\min_{x_i=0,1} \left(\sum_{i \le j} x_i Q_{ij} x_j + C_1 \right)$$
 (4)

Ising:
$$\min_{s_i = \pm 1} \left(\sum_i h_i s_i + \sum_{i < j} J_{ij} s_i s_j + C_2 \right)$$
 (5)

where i, j are indices, ranging over all qubits. In the QUBO model (Eq. 4), Q_{ij} is the QUBO matrix with values $Q_{ij} \in \mathbb{R}$. The binary variable vector is \mathbf{x} with $x_i \in \{0, 1\}$. In the Ising model (Eq. 5), the problem is defined by the biases $h_i \in \mathbb{R}$ and the couplers $J_{ij} \in \mathbb{R}$. and The binary variable vector is \mathbf{x} with $s_i \in \{-1, 1\}$. C_1 and C_2 are constants which do not affect the solution of the optimization problem. The Ising model and the QUBO model are mathematically equivalent, allowing them to be translated into each other. This equivalence provides a flexible approach to problem-solving, enabling the conversion of problems between these models based on the requirements and available tools. There are also tools, such as ToQUBO.jl, designed to convert standard problems into the QUBO format⁴³. In this paper, we utilize the quantum annealer from D-Wave Advantage Systems¹³ to employ direct quantum processing unit (QPU) methods for seismic travel inversion.

Seismic Data Acquisition

We construct a storage velocity model representing carbon storage applications, as shown in Fig. 4a. This model spans a depth range from 1000 m to 1300 m and extends 100 m horizontally. The size of the grid cell for this model is 10 x 10 m. Within this model, the carbon storage structure is depicted as a wedge, starting from 1100 m and reaching a maximum depth of 1200 m. The velocity model is constructed with varying velocities to reflect real-world geological conditions. The average velocity within the carbon storage area ranges from 3180 to 3220 m/s, which is about 11% lower than the surrounding background velocity, which ranges from 3530 to 3640 m/s. Furthermore, the velocities increase with depth.

The uniform placement (Fig. 4b) is commonly used in seismic data acquisition for simplicity in boreholes. However, this approach results in significantly lower ray coverage in the shallow and deep sections compared to the middle section. To address these limitations, in our survey, 20 pairs of sources and receivers are non-uniformly deployed within two boreholes (Fig. 4c). The non-uniform deployment is designed to introduce more constraints for the quantum annealing process, thereby improving the overall accuracy of the seismic inversion. The sources and receivers of non-uniform placement are distributed according to a quadratic polynomial distribution.

Transforming Ray Equations to QUBO

The set of ray equations can be represented as 12 :

$$Ds = T (6)$$

where **D** is the matrix of distance increments d_j , **s** is the slowness vector, and **T** is the travel time vector. The size of **D** can be very large, therefore solving for **s** through matrix operations on **D** is computationally intensive. This challenge is exacerbated by the relatively sparse and random distribution of elements within **D**. Consequently, alternative methods are required to solve these problems efficiently while maintaining accuracy. Quantum annealers can provide quantum metaheuristic algorithms to address this issue. Eq. 6 can be solved by minimizing the objective function:

$$f(\mathbf{s}) = \|\mathbf{D}\mathbf{s} - \mathbf{T}\|^2 \tag{7}$$

The objective function f(s) computes the difference between the observed travel times and those predicted by the model given a slowness vector s. Minimizing this function ensures that the model's predictions align as closely as possible with the observed data, thus achieving an optimal fit. Quantum annealers offer a direct approach to solving binary objective functions⁴⁴:

$$f(\mathbf{q}) = \left\| \mathbf{A}^d \mathbf{q} - \mathbf{b} \right\|^2 \tag{8}$$

In this formulation, \mathbf{q} is a binary vector, \mathbf{A}^d is a real-valued matrix, and \mathbf{b} is a real-valued vector. Because quantum computers are designed to solve QUBO problems, transforming real-valued variables to binary values is essential. However,

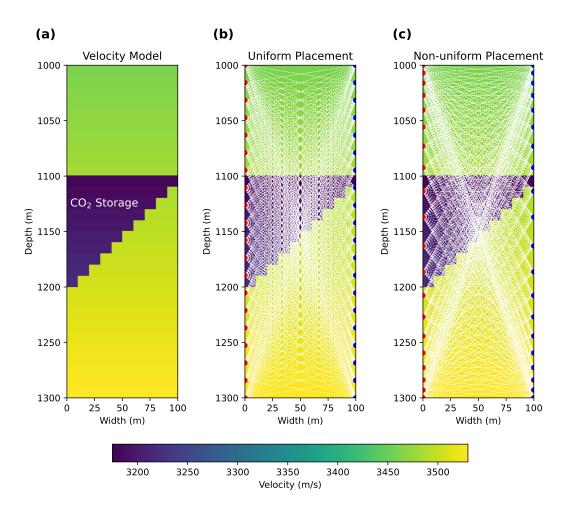


Figure 4. The carbon storage velocity model and ray coverage patterns. Red dots are sources, blue dots are receivers, and white lines represent the ray paths. (a) The synthetic velocity model with a wedge-shaped low-velocity carbon storage formation. (b) Ray coverage from sources and receivers placed in a uniform grid within two boreholes. (c) Ray coverage from sources and receivers placed in a non-uniform pattern, enhancing coverage and constraints for the quantum annealing process.

the number of binary variables n_{binary} increases with the number of bits R used for fixed-point approximation, and it is related to the number of real-valued variables n_{real} as: $n_{\text{binary}} = n_{\text{real}} \times R$. Higher values of R yield greater accuracy in representing the floating-point numbers, but the current limitations of quantum computer hardware restrict the number of qubits available. To address this issue without excessively increasing the number of binary variables, the initial velocity guess, variable boundaries ¹⁴ and recursive methods ⁴⁵ are employed. The initial guess and the boundaries are used to rescale the range of the slowness vector \mathbf{s} to a new vector \mathbf{x} such that $x_i \in [0,2]$, facilitating easier binary representation. Recursive methods are then applied to enhance the precision of floating-point divisions. These methods iteratively refine the estimate of \mathbf{s} , reducing the error at each step. The initial objective function Eq. 7 can be reformulated as:

$$f(\mathbf{x}) = \|\mathbf{D}\mathbf{x} - \mathbf{b}\|^2 \tag{9}$$

where $\mathbf{b} = (\mathbf{T} + L\mathbf{I} - \mathbf{D}\mathbf{s}_0)/L$, L is the variable boundaries, \mathbf{s}_0 is the initial guess of the slowness vector, and \mathbf{I} is the identity vector. The slowness vector \mathbf{s} is then in the range of $[\mathbf{s}_0 - L\mathbf{I}; \mathbf{s}_0 + L\mathbf{I}]$. This range ensures that the solution space is adequately covered. To express this as a binary objective function, x_i is represented in binary form using the R bit fixed-point approximation:

$$x_i = \sum_{r=0}^{R-1} 2^{-r} q_r \tag{10}$$

where $q_r \in \{0,1\}$ is the value of the r-th bit. This transformation is essential for harnessing the computational power of quantum annealers, which are inherently designed to solve binary optimization problems. The new matrix A^d in Eq. 8 is derived from D

and R such that $\mathbf{D}\mathbf{x} = \mathbf{A}^d\mathbf{q}$. The QUBO matrix Q_{ij} in Eq. 4 is then constructed from the given matrix \mathbf{D} and the calculated vector $\mathbf{b}^{44,46}$:

$$v_j = \sum_i A_{ij} (A_{ij} - 2b_j) \tag{11}$$

$$w_{jk} = 2\sum_{i} A_{ij} A_{ik} \tag{12}$$

The QUBO matrix is directly input into the Quantum Processing Unit (QPU). The system utilizes <code>DWaveSampler()</code> to employ a D-Wave system as the sampler. Subsequently, <code>EmbeddingComposite()</code> manages the mapping between the problem and the D-Wave system's numerically indexed qubits, a process known as minor-embedding⁴⁷.

In this study, we perform traveltime inversion using R = 3 for 10 iterations with quantum annealing. The total number of real-valued variables of the problem is 300. Due to quantum hardware limitations, we break down the model into 30 layers with 10 variables each. This division reduces the complexity of each sub-problem, making it manageable for the quantum processor and allowing for better control of the boundary L. Since the problem from Eq. 6 is depth-independent, we simplify the process by adjusting the system's coordinates at each iteration. The approach ensures that each layer is treated independently, reducing the overall complexity of the inversion. By implementing these techniques, we can efficiently solve large-scale traveltime inversion problems using quantum annealers.

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Author contributions statement

H. A. Nguyen contributed to the methodology, formal analysis, and writing of the manuscript. A. Tura was responsible for conceptualization and providing professional oversight.

Competing interests

The authors declare no competing interests

Data availability

Access to the D-Wave System in this study is available at https://cloud.dwavesys.com. The dataset, including the inputs and results, can be accessed and downloaded from the Zenodo repository at https://zenodo.org/records/13674667.