**The effect of Bed Thickness on Hillslope Morphology and Sediment Size in Last Chance Canyon, New Mexico**

Sam Anderson1, Nicole Gasparini1, Joel Johnson2

1Earth and Environmental Science, Tulane University, New Orleans, 70118, USA

2Jackson School of Geosciences, University of Texas at Austin, Austin, 78712, USA

*Correspondence to*: Sam Anderson (sanderson@tulane.edu)

**Abstract.** In this paper we explore the effect that variance in bed thickness has on grain size distributions, channel, and hillslope form in Last Chance canyon, New Mexico, USA. Here, the landscape is composed of horizontally to near horizontally bedded rock, have relatively little regolith, yet some hillslopes appear diffusive. We posit that differences in bed thickness influence hillslope shape here in Last Chance canyon, where thinner bedrock generates a more diffusive geometry resembling the idealised convex hillslope shape endemic of soil mantled landscapes. We used drone photos to construct high resolution orthomosaics and digital elevation models (DEMs) of seven hillslope transects and stream channels at the base of each transect from two proximal watersheds. Using these DEMs and orthomosaics, we measured bed thicknesses, slope, and curvature on the hillslopes and grain size distributions at the base of each hillslope. We find that hillslopes are steeper and less diffuse where there is more thickly bedded rock and become shallow and diffuse in thinly bedded rock. Furthermore, sediment input to channels is controlled by bed thickness on hillslopes and affects channel morphology. Thickly bedded rock units on proximal hillslopes contribute larger sized colluvial sediment to the channels, which steepen as a result. ADD CURVATURE AND SOIL PIT STUFF

# 1 Introduction (NOT MY DOG)

Variance in relevant bedrock properties influences the production and size of sediment on hillslopes (Johnstone and Hilley, 2015), erosion rates (Dixon et al, 2012), and landscape form (Glade, 2017; Hurst, 2013). Rock properties, specifically fracture spacing, has been shown to influence rock surface slope (e.g., Brook and Tippett, 2002; Matasci et al., 2015; Moore et al., 2009; Selby, 1980), erosion, and imprints its signature into the topography (Molnar et al., 2007; Clarke and Burbank, 2011; St. Clair et al., 2015; Voigtlander et al., 2017; Eppes and Keanini, 2017; Eppes et al., 2018). Soil depth, a function of both soil production and erosion rate, has been shown to affect hillslope convexity (Roering, 2008). Soil mantled hillslopes are thought to generate convex hillslopes, however, in Last Chance canyon, sandstone hillslopes with predominantly exposed bedrock are convex in shape and resemble soil mantled hillslopes. But what about these bedrock hillslopes causes them to look like soil mantled hillslopes? Why do predominantly sandstone hillslopes appear diffusive while carbonate hillslopes are steep and shear?

In this study, we seek to understand the effect that changes in rock properties have on sediment production and erosion. More specifically, we ask what controls weathering, soil production, and sediment movement on hillslopes with high spatial variability in rock properties? how is sediment removed so that diffusive looking landscapes are generated in thinly bedded sandstone and not in the more thickly bedded carbonates? To answer these questions, we will couple a hillslope sheet wash component with the overland flow model to explore sediment delivery from hillslopes to channels. We will inform our model runs with realistic hydrographs reconstructions and field measurements of sediment size. and distance between bedrock beds. We will also measure sediment depth, bulk lithology of less than gravel sized colluvium, and the lithology of larger sized sediment.

# 2 Field Area

Because of their differing morphology and accessibility, we use data gathered from different sections of two, first to second order watersheds, called LC1 and LC3 (map figure 1). Last Chance Canyon has horizontally to near horizontally bedded bedrock and is currently tectonically inactive (Hill, 1987; Hill, 2006). During Permian time, a shallow lagoon existed behind a reef complex to the south and deposited what would become interbedded carbonate and siliciclastic bedrock of various thicknesses (Hill, 2000; Phelps et al., 2008; Kerans et al., 2017). The Guadalupe mountains were uplifted during basin and range extension beginning 27 million years ago, exposing the previously buried bedrock (Chapin and Cather, 1994; Ricketts et al.., 2014, Hoffman, 2014; Decker et al., 2018).

# 3 Methods

## 3.1 Field Survey

We used a DJI Mavik 2 pro to take photos of seven hillslope transects in each of the two watersheds from elevations of approximately 20 meters above LC1 and LC3. We then used Agisoft Metashape software to process the drone images and to produce orthomosaics and DEM’s with a sub cm scale spatial resolution.

We measured the depth of regolith above bedrock by digging soil pits at 20m intervals, when possible, down each of the seven transects.

## 3.2 Sediment Size Measurements

We measured the diameter of alluvium in channels at the base of the seven hillslope transects with the PebbleCounts image analysis package (Purinton and Bookhagen, 2019). PebbleCounts is a Python based product for the detection and sizing of sediment grains from drone images. We used the k means with manual (KMS) method, which allows an operator to validate measured grains to mitigate error. Due to the relatively large size of the orthomosaics we utilized the PebbleCounts-Application, which subset the images to manageable sizes (photo sieve fig). To further streamline the process, we developed a application which loops through the subset images and then compiles the all of the data for a channel section into a single file. All of which is in the GitHub repository for this paper.

## 3.3 DEM and Orthomosaic analysis

We measured the bed thickness, the vertical distance between the upper and lower horizontal bedding planes, of every exposed bed on the seven hillslope transects using the high-resolution orthomosaics in the ArcScene program (figure showing Sophia method).

At 20m intervals beginning at the channel and moving up each hillslope transect, we measured slope and curvature over a 80m window (40m upslope and 40m downslope of each point) using the 3DEP 1m DEM (hillslope schematic figure). We also measured all beds within the same window and found average bed thickness, the sum of all thicknesses, and the largest bed thickness. The window distance was reduced in proximity to the channel or ridgeline so that measurements were only taken from the hillslope in question.

We used a 10 m digital elevation model (DEM) of Last Chance canyon along with TopoToolBox to generate ksn maps of all surveyed channels (Schwanghart and Scherler, 2014). The channel steepness index, or ksn, is a measure of channel gradient normalized for drainage area and allows for the comparison of slope along a single channel or among multiple channels to isolate erosional and/or bedrock erodibility patterns (Kirby & Whipple, 2012).

# 4 Results

## 4.1 Last Chance Canyon Morphology

Last Chance canyon tributaries have upstream sections with relatively shallow channels and lower gradient hillslopes, and a knickzone downstream which has steep channels and hillslopes (slope map). **Here will be how curvature changes up channel, compare LC1 and 3. 3 should be more strait forward. 1 will be weird. Maybe more reflective of whats happening in the channel.**

## 4.2 Bed Thickness affect on hillslopes

As the average, max, and sum of all bed thickness within the 80 m window increases slope also increases (local slope vs bed thickness figure). Furthermore, average bed thickness across the entire hillslope transects increases with the slope of the entire transect (total slope figure). Shallower hillslopes correspond to the upstream sections of the landscape, and the data show that beds are thinner there.

In Last Chance canyon, hillslopes appear more diffuse in the more thinly bedded sandstone rock, but not in the more thickly bedded rock (curve vs distance fig). Diffusive hillslopes are convex, with a negative curvature closer to the ridgeline and then transition to a positive curvature downslope closer to the channel. We plotted curvature values by distance from the channel in the seven hillslopes. We then plotted the slope values and the r squared values for the seven transects against max bed thickness and found that where the max bed thickness is lower, the slope of the distance vs curvature function is more negative and has a higher r-squared value, both indicating a more diffusive shape.

## 4.3 Bed thickness affects channel morphology

Max bed plane thickness of each of the transects affects both the size of the sediment in the channel below as well as the channel steepness of the channel below.

## 5 Discussion

FOR NOW: its ideas…

I wonder if we can separate the effect bed thickness has on where a hs looks diffusive or not by tracking changes up channel and comparing between lc1 and 3. Worst comes to worst if its not just distance up channel we can say that it’s a combo of bed thickness and drainage area.

How do I explain away slope change with bed thickness? Is it too duh?

# 6 Conclusions

We present several observations about the effect of bed thickness on landscape morphology in Last Chance canyon. We suggest that bed thickness influences hillslope geometries. hills steepen across units with thicker beds. Conversely, hills are shallower in thinly bedded units. Furthermore, hillslopes with smaller max bed thickness appear more diffusive.

The size of boulders in the channel also impacts channel morphology. More thickly bedded rock on the hillslopes contributes larger sized, and more geomorphically relevant, alluvium to the channel. This coarse sediment armors the channel and causes channel steepness to increase. In Last Chance canyon, channel sections that contain larger alluvium are generally steeper.

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