

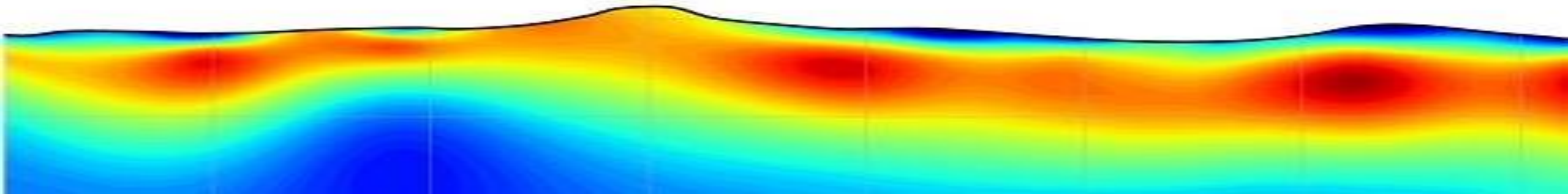
# ESS302 Applied Geophysics II

Gravity, Magnetic, Electrical, Electromagnetic and Well Logging

## Ground Penetrating Radar

Instructor: Dikun Yang

Feb – May, 2021



well logging  
(everything in borehole)

# Maxwell Equations

$$\nabla \cdot \mathbf{D} = \rho$$

$$\nabla \cdot \mathbf{B} = 0$$

$$\nabla \times \mathbf{E} = -\frac{\partial \mathbf{B}}{\partial t}$$

$$\nabla \times \mathbf{H} = \mathbf{J} + \frac{\partial \mathbf{D}}{\partial t}$$

zero frequency

low frequency

high frequency

steady state

quasi-static state

EM wave

mechanical wave

magnetic

gravity

potential field

electrical

electromagnetic (induction)

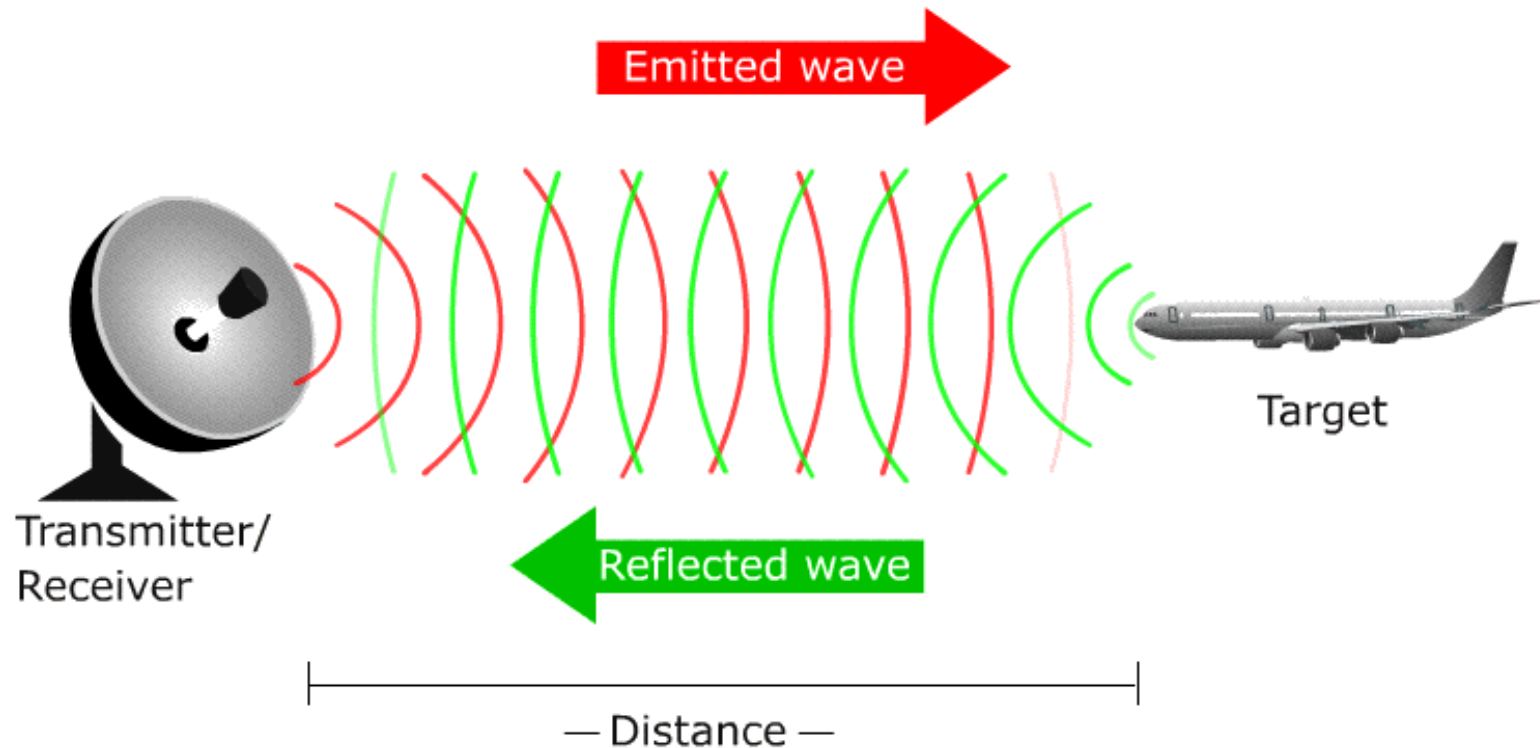
electrical conductivity/resistivity

electromagnetic (geo-radar)

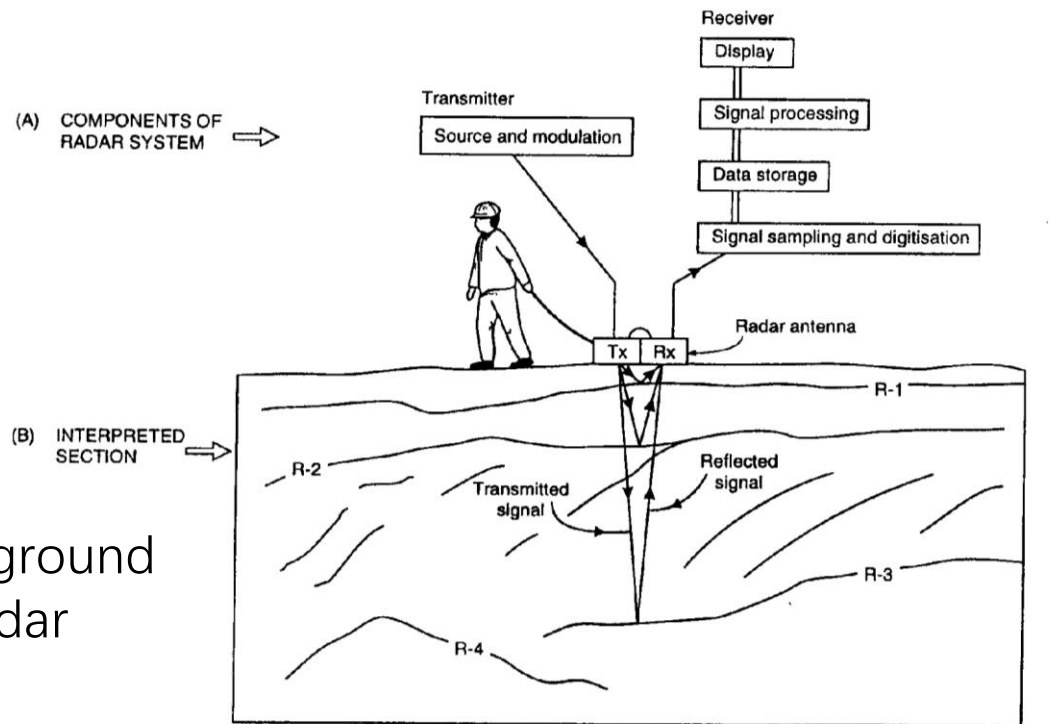
seismic

wave phenomena

# Radar

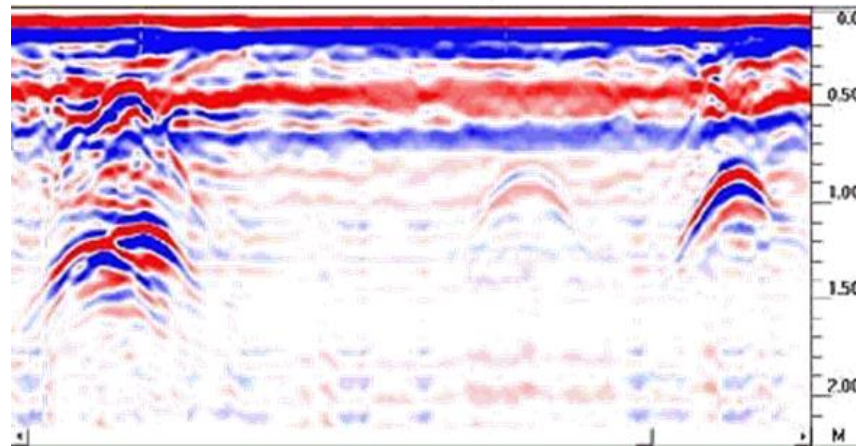


- Can we do the same thing to the subsurface?
- What are the differences between finding an object in the air and underground using EM waves?

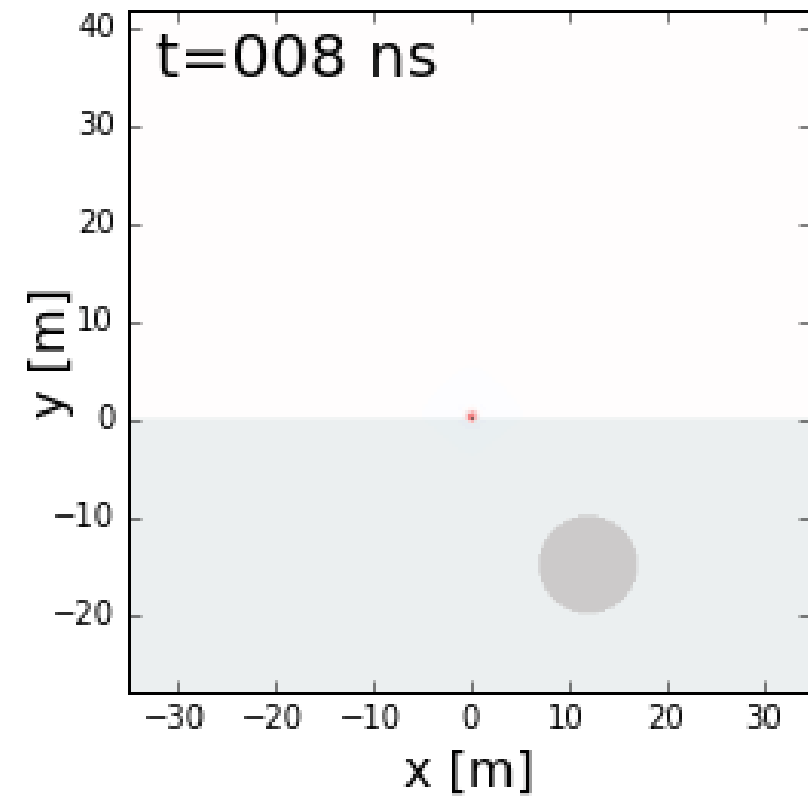
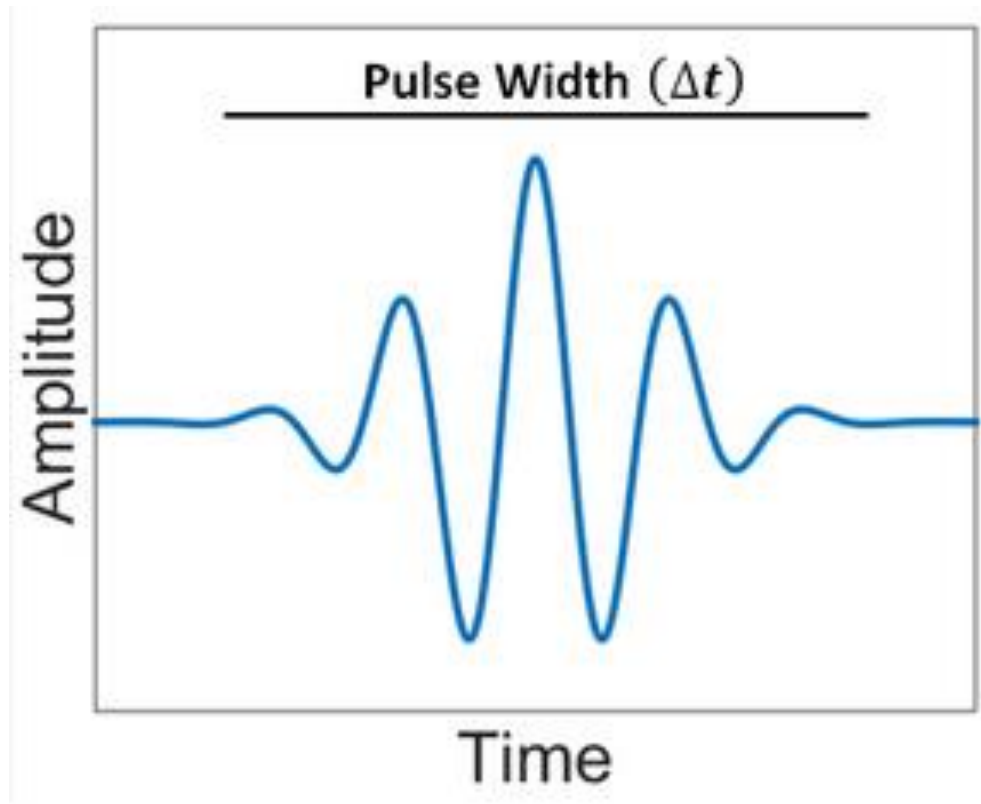


Geo-radar or ground  
penetrating radar  
(GPR)

Radargram

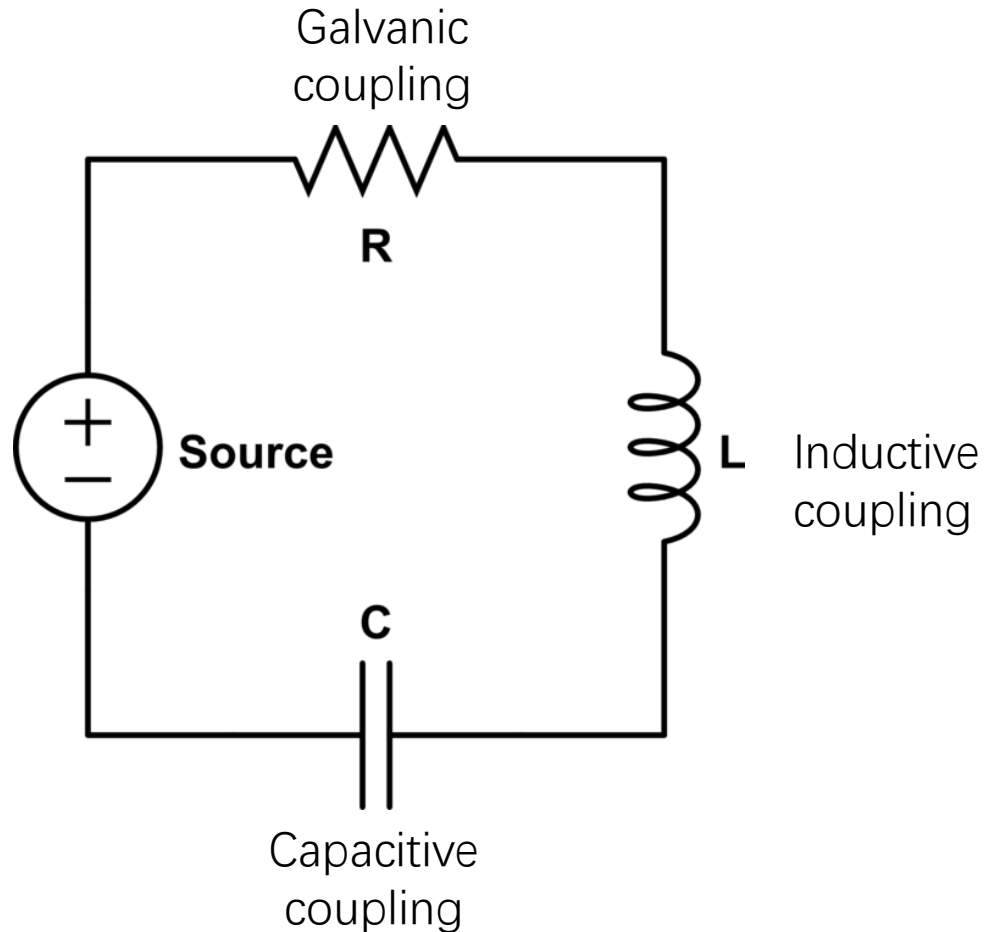


# EM Field at High Frequencies – Wave

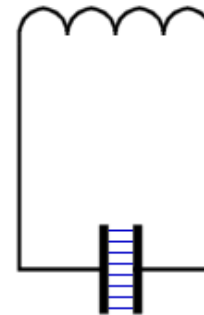
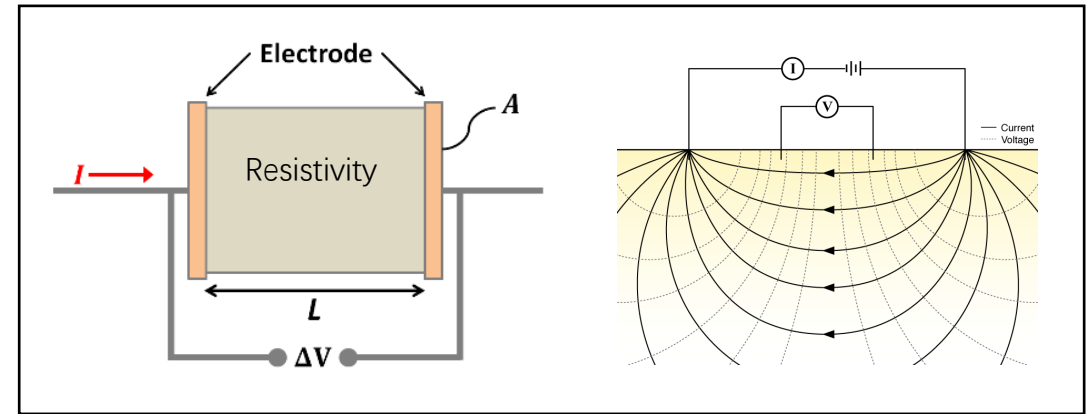




# Ground Penetrating Radar (GPR)



Recall...



## Capacitive coupling

- High frequency EM field
- Dielectric constant ( $\epsilon_r$ )
- Wave phenomenon

# Wave Propagation

Medium characterized by three physical properties:  
 $\sigma$  (electrical conductivity),  $\epsilon$  (electrical permittivity),  
 $\mu$  (magnetic permeability)

In general:

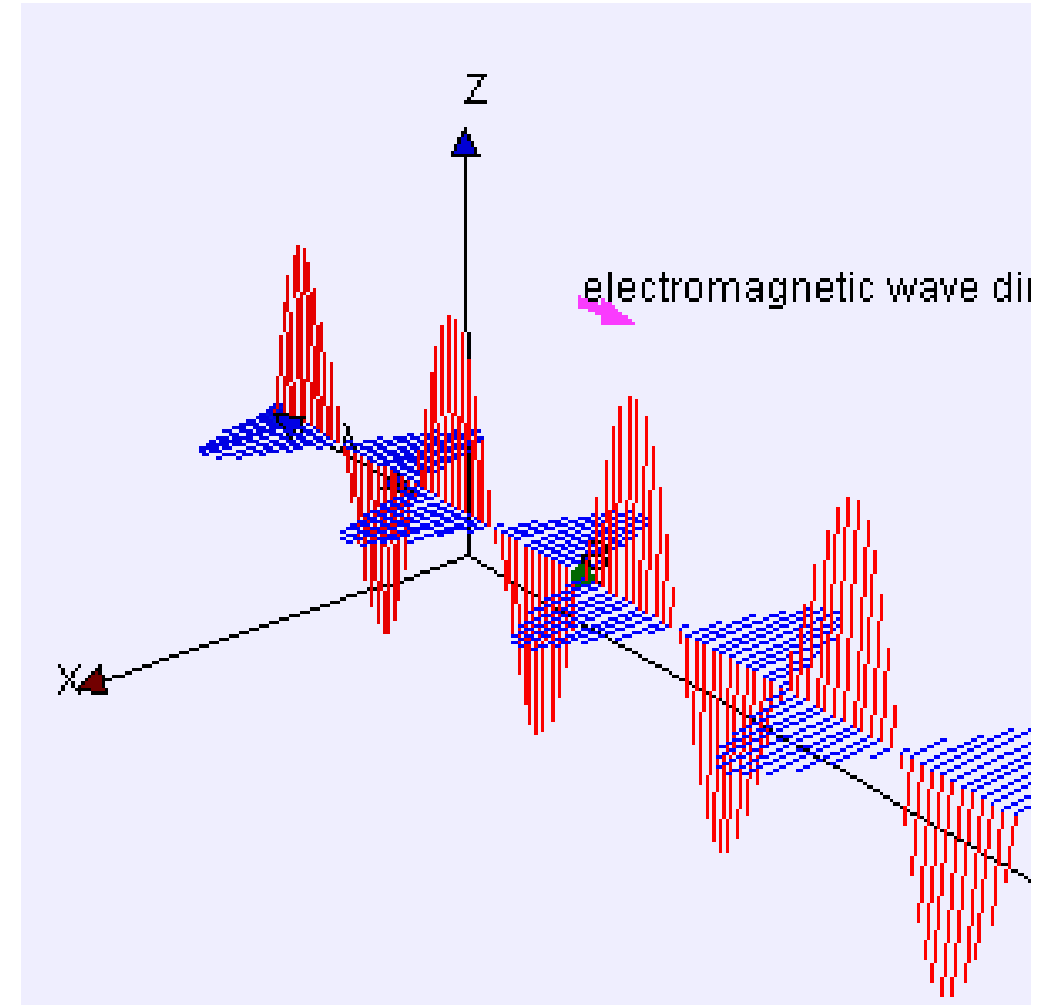
$$V = \sqrt{\frac{2}{\mu\epsilon} \left[ \left( 1 + \left( \frac{\sigma}{\omega\epsilon} \right)^2 \right)^{1/2} + 1 \right]^{-1/2}}$$

Wave regime  
( $\sigma \ll \omega\epsilon$ ):

$$V = \frac{1}{\sqrt{\mu\epsilon}} = \frac{c}{\sqrt{\mu_r\epsilon_r}}$$

Non-magnetic  
approximation  
( $\mu_r = 1$ ):

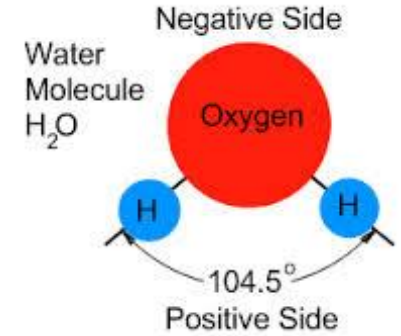
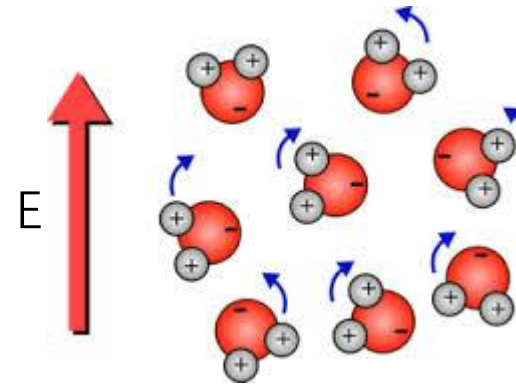
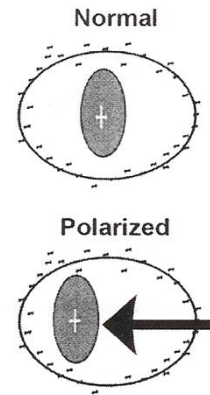
$$V = \frac{c}{\sqrt{\epsilon_r}}$$



Question: How does EM wave propagate in perfect conductors?

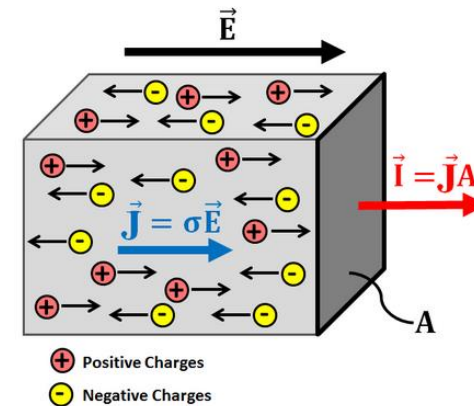
## Dielectric Permittivity ( $\epsilon$ ):

How easily a material is electrically polarized



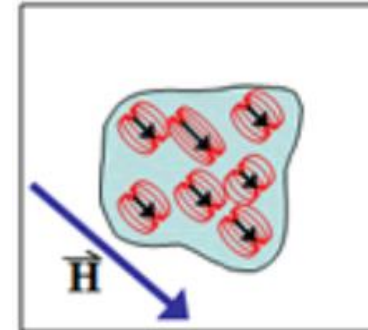
## Electrical Conductivity ( $\sigma$ ):

How easily electrical charges flow through a material



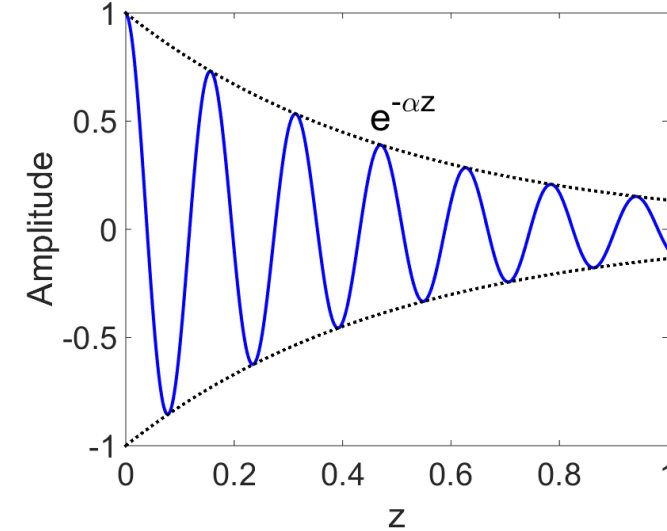
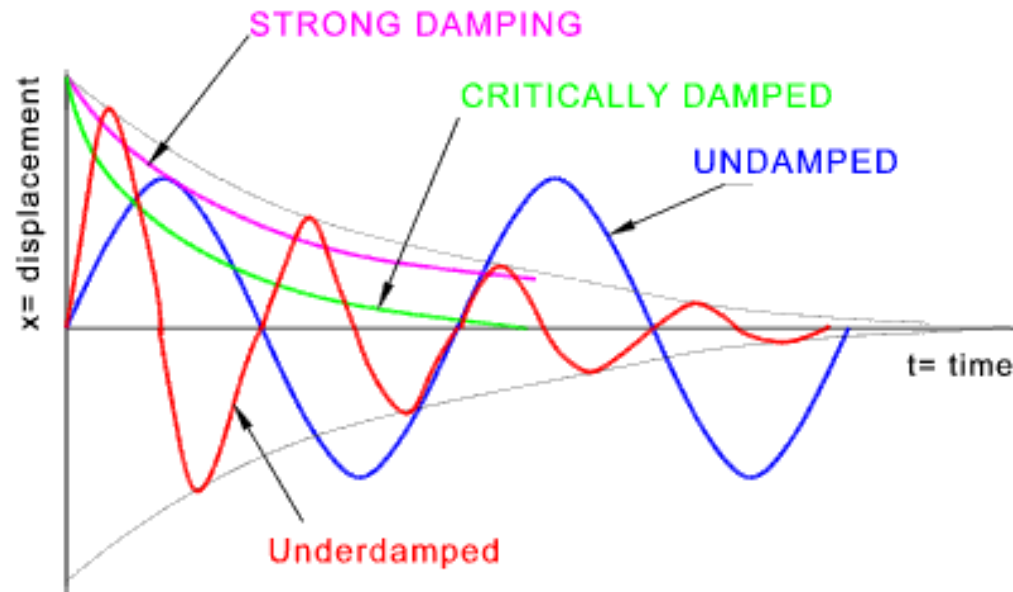
## Magnetic Permeability ( $\mu$ ):

How strongly a material supports magnetism





# Wave Attenuation

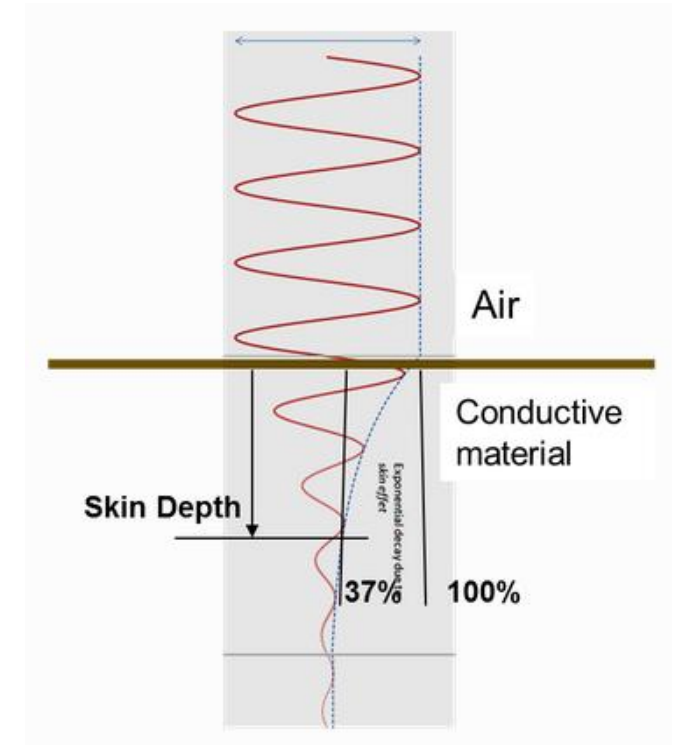
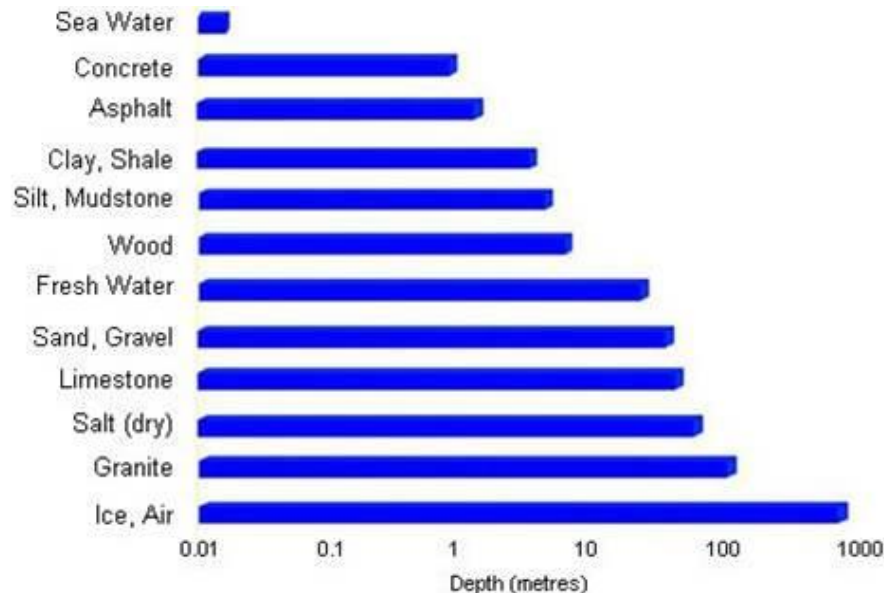


$$\alpha = \omega \sqrt{\frac{\mu \epsilon}{2}} \left[ \left( 1 + \left( \frac{\sigma}{\omega \epsilon} \right)^2 \right)^{1/2} - 1 \right]^{1/2} \approx \begin{cases} \sqrt{\frac{\omega \mu \sigma}{2}} & \text{for } \omega \epsilon \ll \sigma \\ \frac{\sigma}{2} \sqrt{\frac{\mu}{\epsilon}} & \text{for } \sigma \ll \omega \epsilon \end{cases}$$

- Quasi-Static ( $\omega \epsilon \ll \sigma$ ): Conductive/Low-frequency
- Wave Regime ( $\sigma \ll \omega \epsilon$ ): Resistive/High-frequency

# Skin Depth and Probing Distance

- Skin Depth: Distance at which a wave is reduced to 37% (1/e) of its original amplitude
- The probing distance is approximated 3 skin depths.



$$\delta \approx \begin{cases} 503 \sqrt{\frac{1}{\sigma f}} & \text{for } \omega\epsilon \ll \sigma \\ 0.0053 \frac{\sqrt{\epsilon_r}}{\sigma} & \text{for } \sigma \ll \omega\epsilon \end{cases}$$

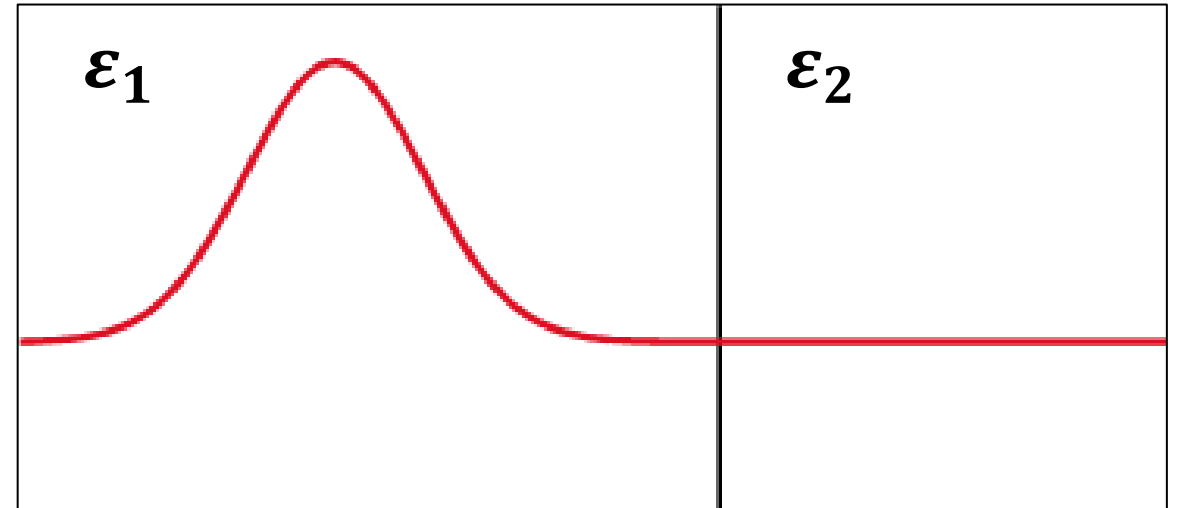
**Table of relative dielectric permittivity ( $\epsilon_R$ ), electrical conductivity ( $\sigma$ ), and velocity.**

Material	$\epsilon_R$	$\sigma$ (mSeimens/m)	$V$ avg (m/ns)
Air	1	0	.3
Distilled water	<b>80</b>	<b>0.01</b>	0.033
Fresh water	80	0.5	0.033
Sea water	80	<b>3000</b>	<b>0.01</b>
Dry sand	<b>3</b> - 5	0.01	0.15
Saturated sand	20-30	0.1-1.0	0.06
Limestone	4-8	0.5-2.0	0.12
Shales	5-15	1-100	0.09
Silts	5-30	1-100	0.07
Clays	5-40	2- 1000	0.06
Granite	4-6	0.01-1.0	0.13
Dry salt	5-6	0.01-1.0	0.13
Ice	3-4	0.01	<b>0.16</b>

# Reflection and Transmission

$$R = \frac{\text{Reflected Amplitude}}{\text{Incident Amplitude}} = \frac{\sqrt{\epsilon_1} - \sqrt{\epsilon_2}}{\sqrt{\epsilon_1} + \sqrt{\epsilon_2}}$$

$$T = \frac{\text{Transmitted Amplitude}}{\text{Incident Amplitude}} = \frac{2\sqrt{\epsilon_2}}{\sqrt{\epsilon_1} + \sqrt{\epsilon_2}}$$



- If  $\epsilon_1 \approx \epsilon_2$ , most of the wave is transmitted
- If  $\epsilon_1 \ll \epsilon_2$  or  $\epsilon_1 \gg \epsilon_2$ , most of the wave is reflected

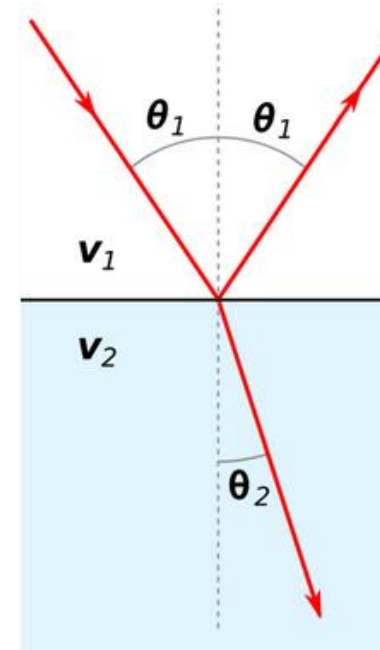
# Refraction

- Snell's Law:

$$\frac{\sin\theta_1}{V_1} = \frac{\sin\theta_2}{V_2}$$

$$V = c/\sqrt{\epsilon_r}$$

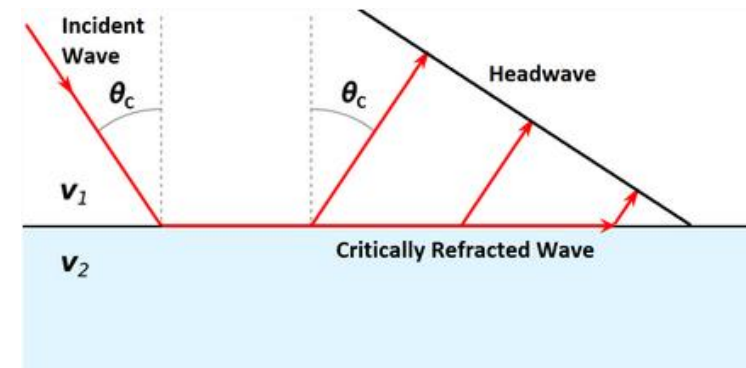
$$\sqrt{\epsilon_1} \sin\theta_1 = \sqrt{\epsilon_2} \sin\theta_2$$



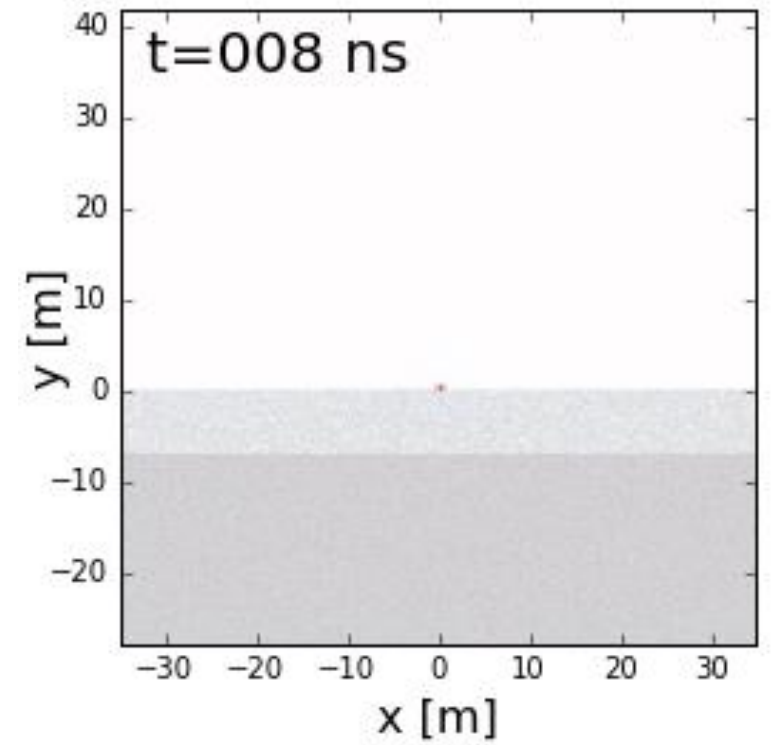
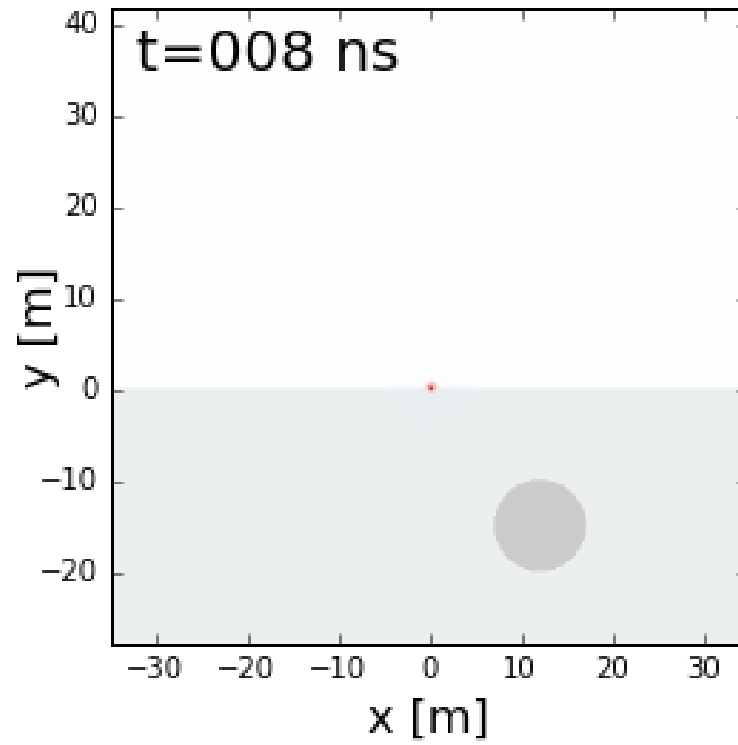
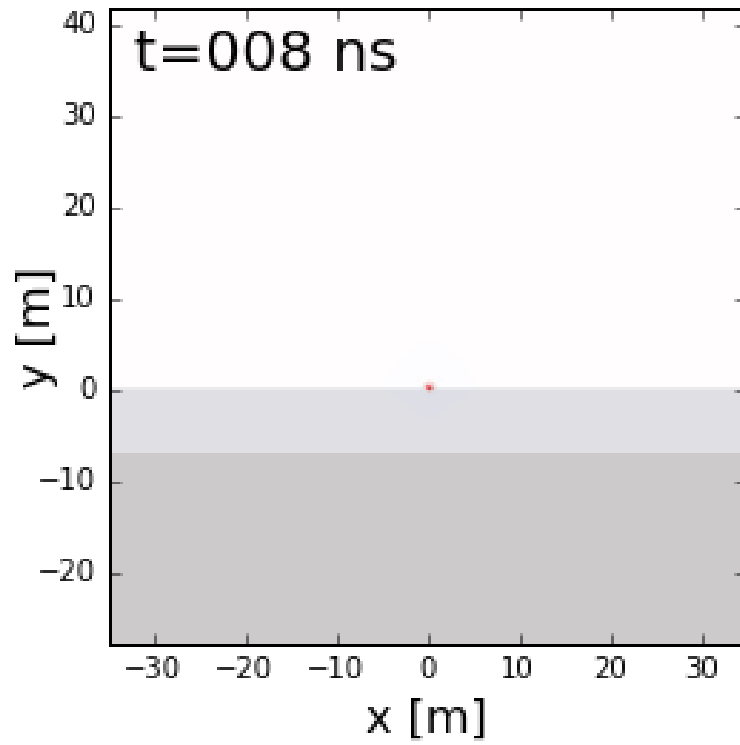
- Critical refraction

$$\sin\theta_c = \frac{V_1}{V_2}$$

Requires  $V_1 < V_2$



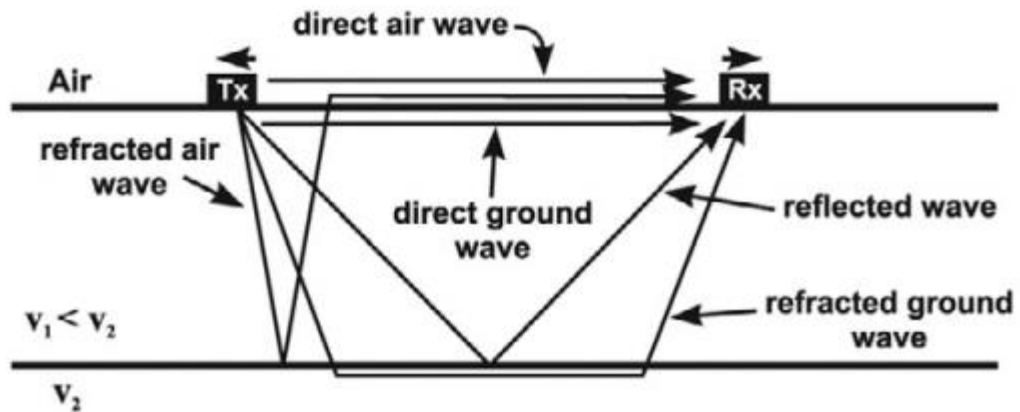
# Reflection, Transmission, Refraction, Scattering



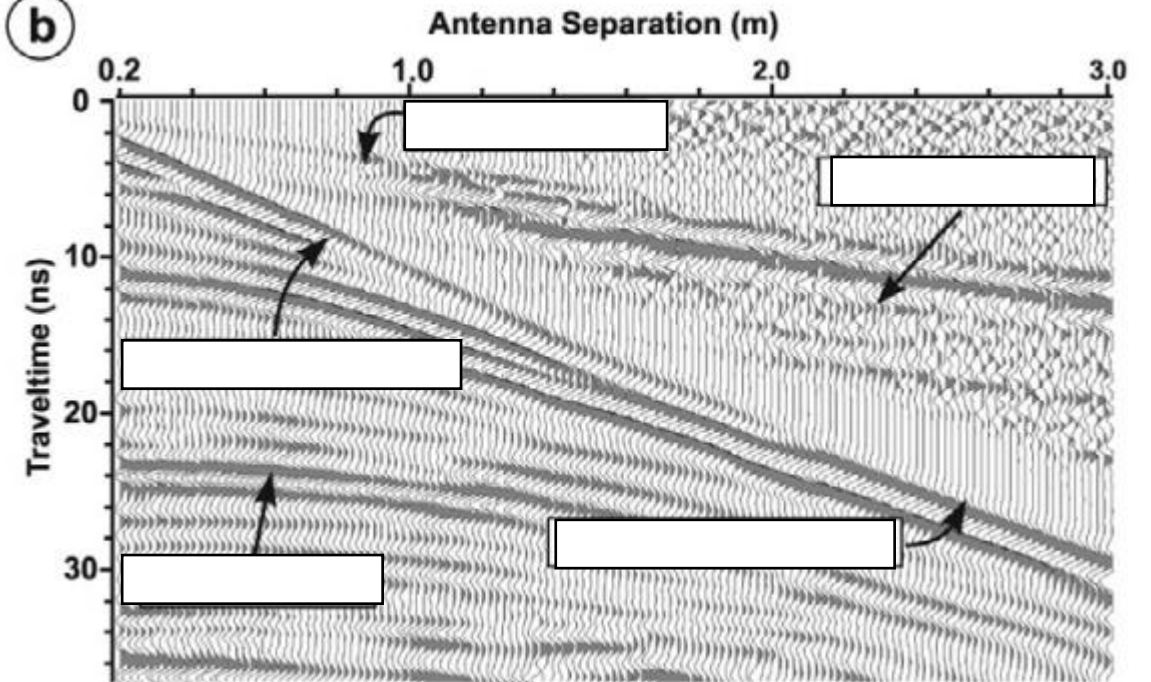


# EM Wave Propagation in a Two-layer Earth

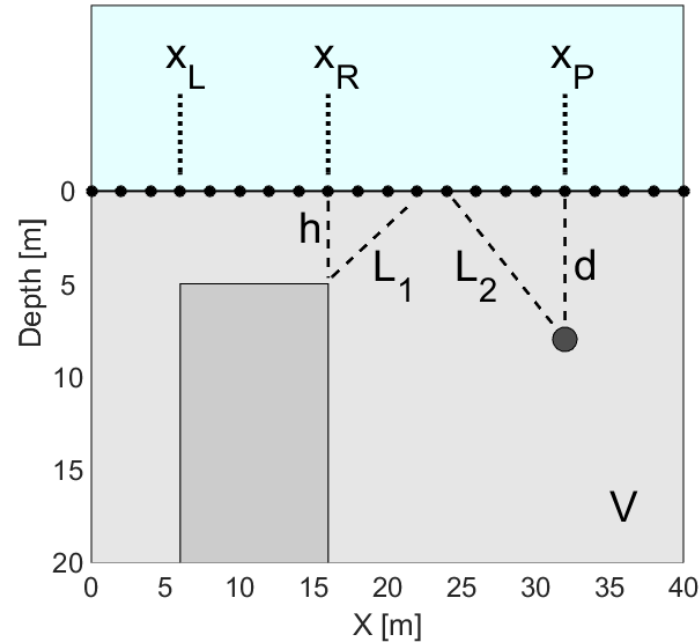
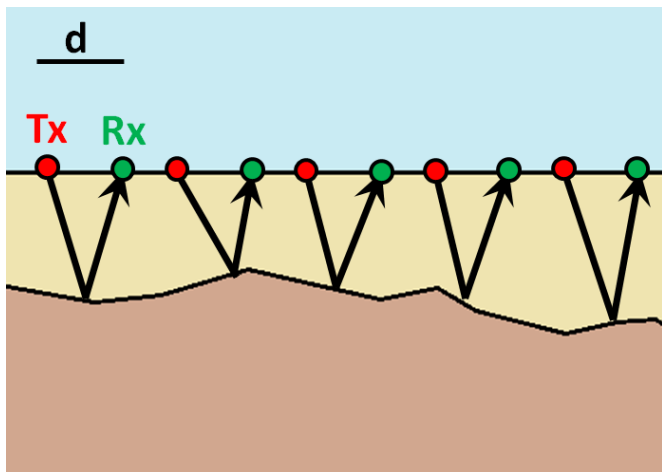
(a)



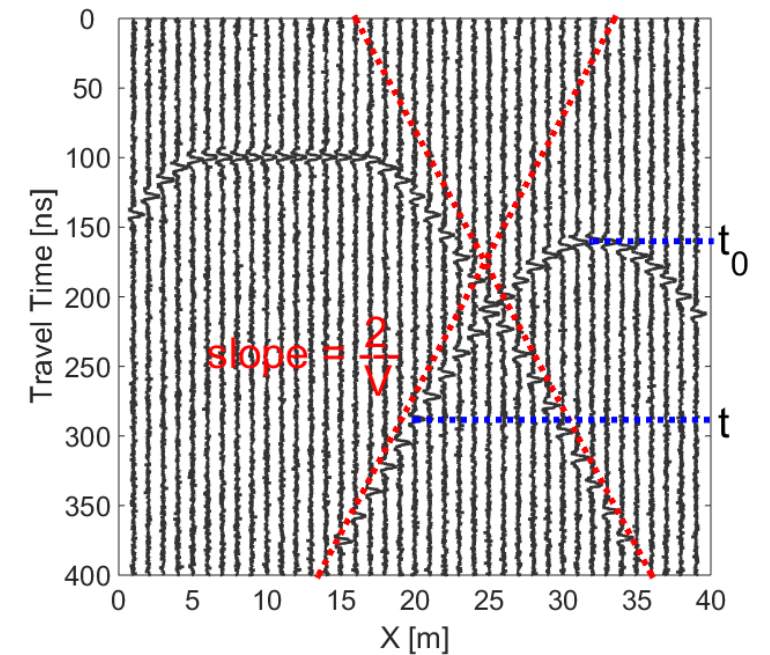
(b)



# GPR Anomaly on Radargram



$$t_p = \frac{2L_2}{V} = \frac{2\sqrt{(x - x_p)^2 + d^2}}{V}$$



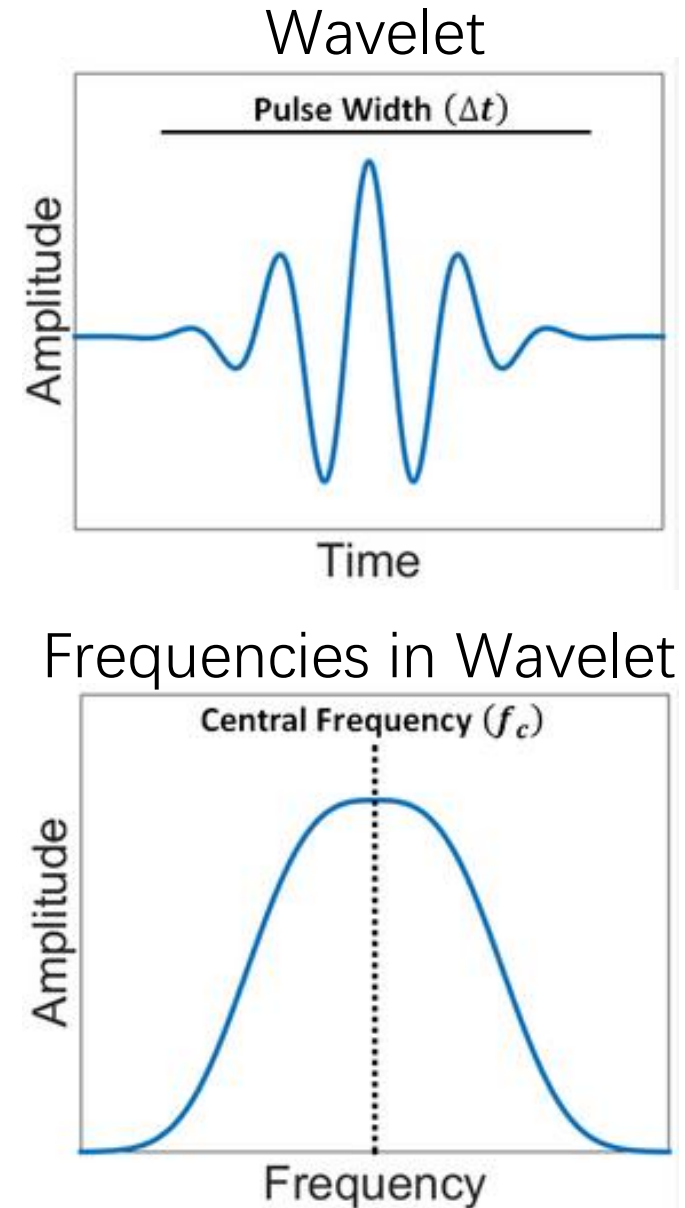
- Determine background medium velocity
- Determine the depth of burial
- Determine the size of extended objects

# GPR Source Signal

- Wavelet: A wave-like oscillation of short duration
- Bandwidth: Range of frequencies in the wavelet
- Pulse Width: Time-duration of wavelet
- Spatial Length: Wavelength of the wavelet
- Central Frequency: Operating frequency of GPR survey

$$f_c = \frac{1}{\Delta t}$$



Typically 50 MHz to 1 GHz



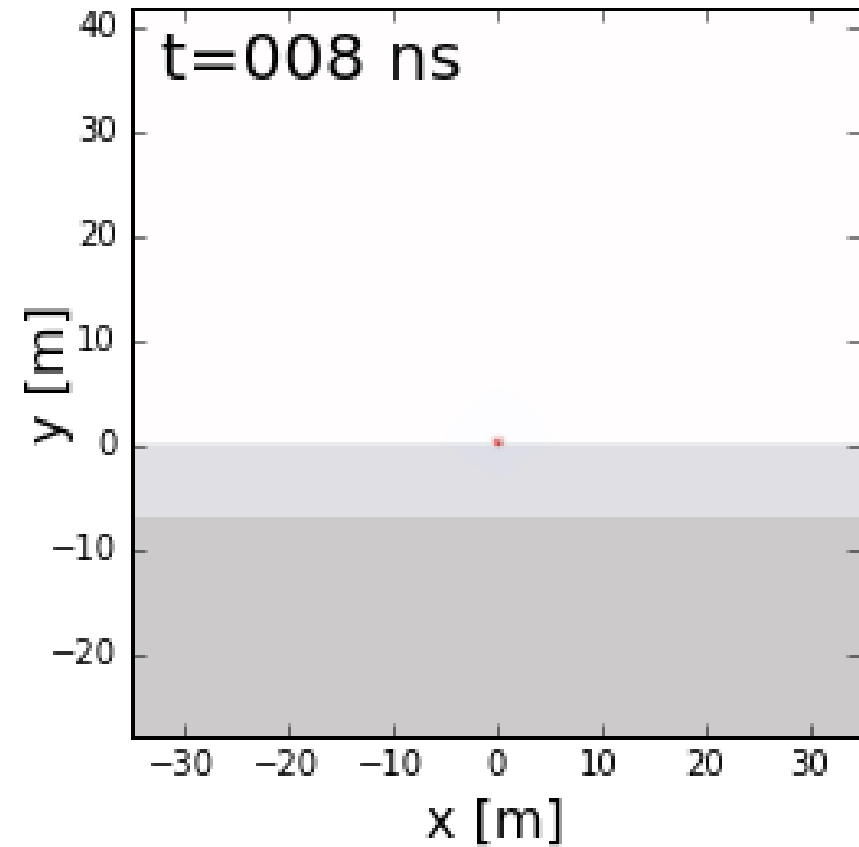
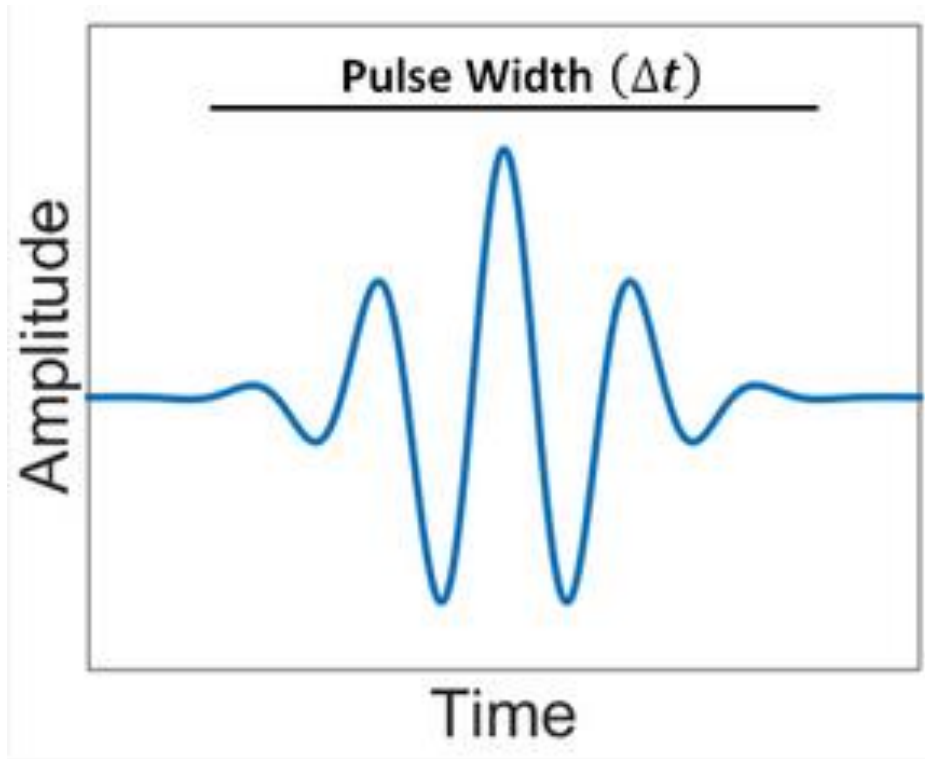
# GPR Source Signal: Spatial Length

- The spatial length (wavelength) of the GPR pulse is dependent on the central frequency and velocity

$$\lambda = \frac{V}{f_c} = \frac{c}{f_c \sqrt{\epsilon_r}} = \frac{c \Delta t}{\sqrt{\epsilon_r}}$$

- When the GPR signal at some frequency is transmitted across an interface, it can be stretched or contracted
- Lower velocity  Shorter spatial length
- Lower frequency  Larger spatial length

# Signal Stretched or Contracted?



# Resolution of GPR

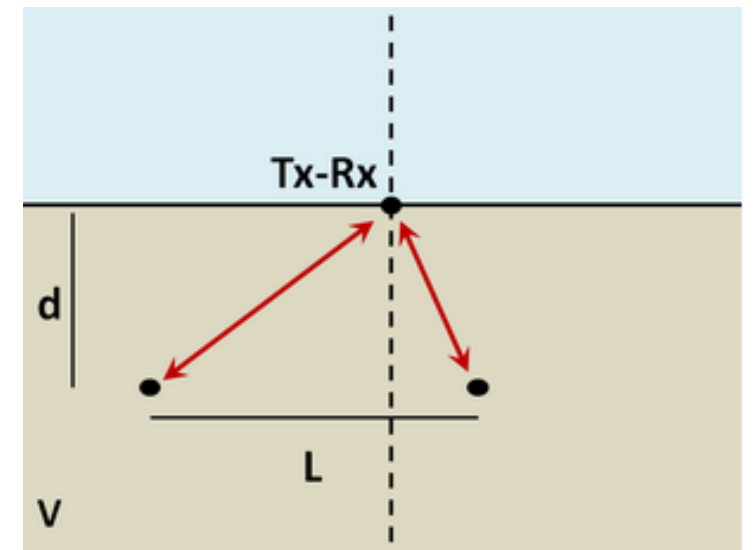
- $\frac{1}{4}$  wavelength rule:

The thickness of a layer must be at least  $\frac{1}{4}$  the wavelength of the GPR signal.

$$L > \frac{c}{4f_c\sqrt{\epsilon_r}} = \frac{c\Delta t}{4\sqrt{\epsilon_r}}$$

- For zero offset survey

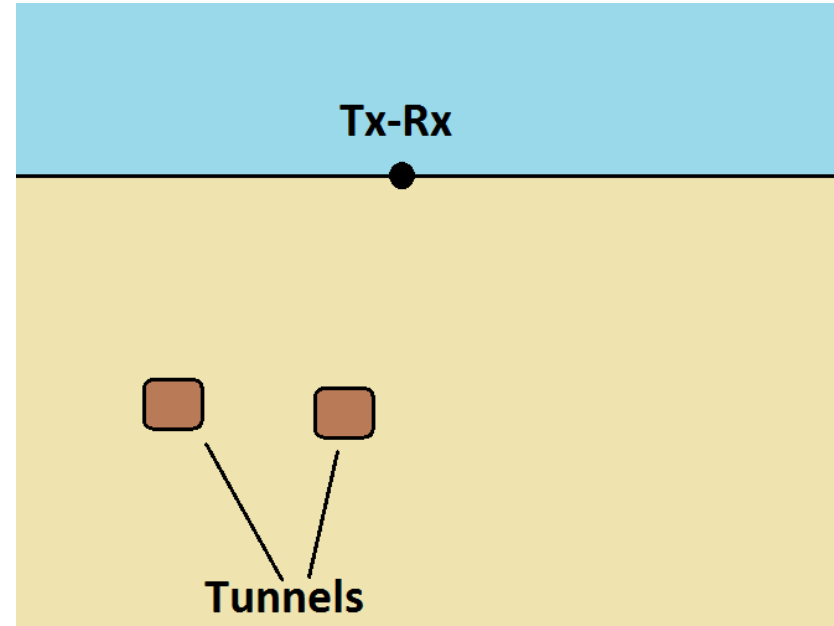
$$L > \sqrt{\frac{Vd}{2f_c}}$$






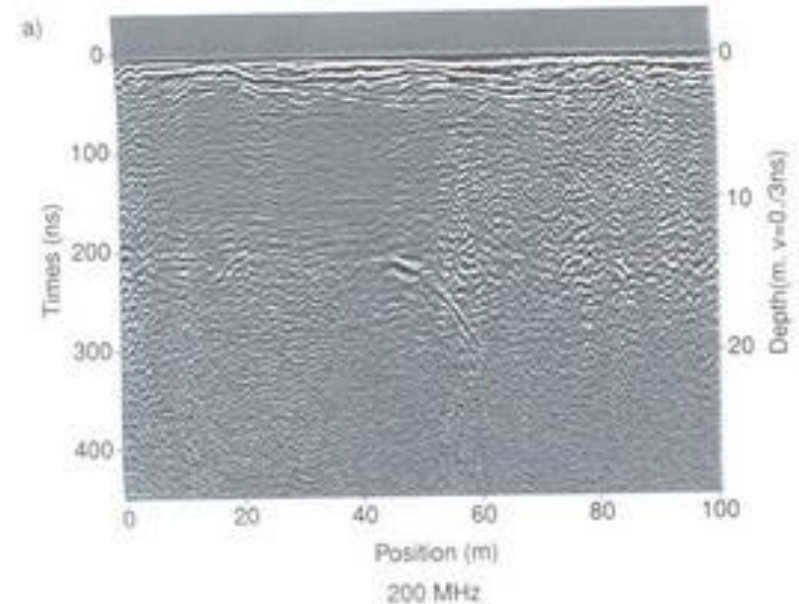
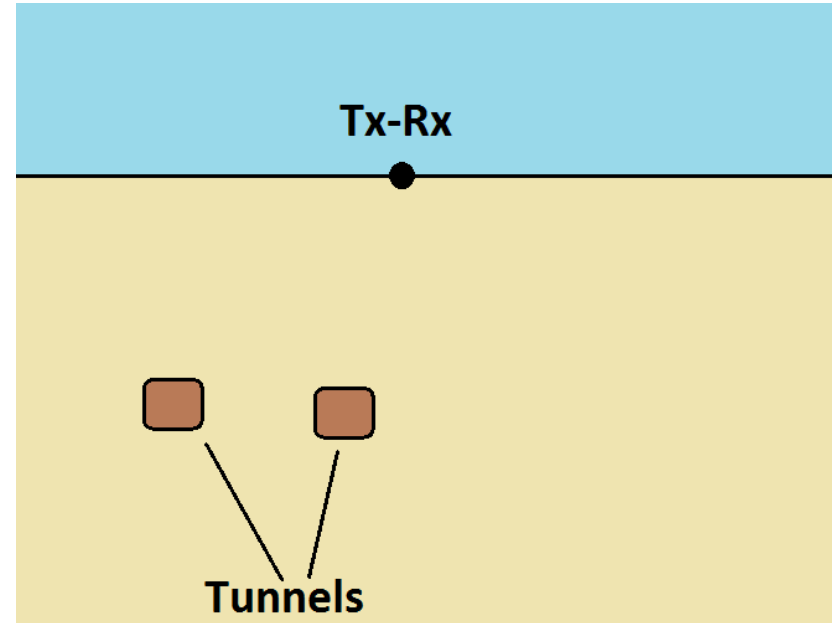
# Probing Distance vs. Resolution

- Want to find two buried tunnels.
- Using a zero offset survey configuration.
- Higher frequencies give better resolution
- Lower frequencies give larger probing distance



# Radargram 200 MHz

- Little to no useful signal after 200 ns
  - Can't see features from the tunnels
- 
- Too much attenuation of signal
  - Probing distance insufficient

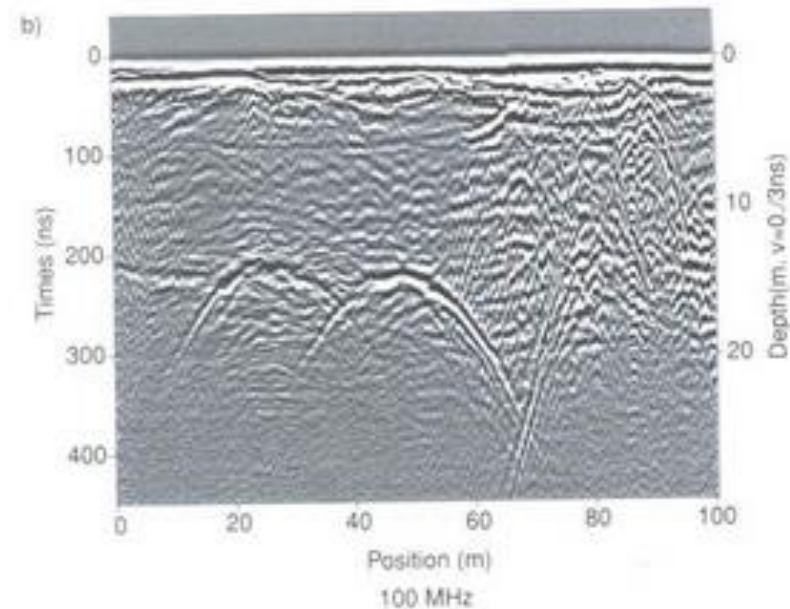
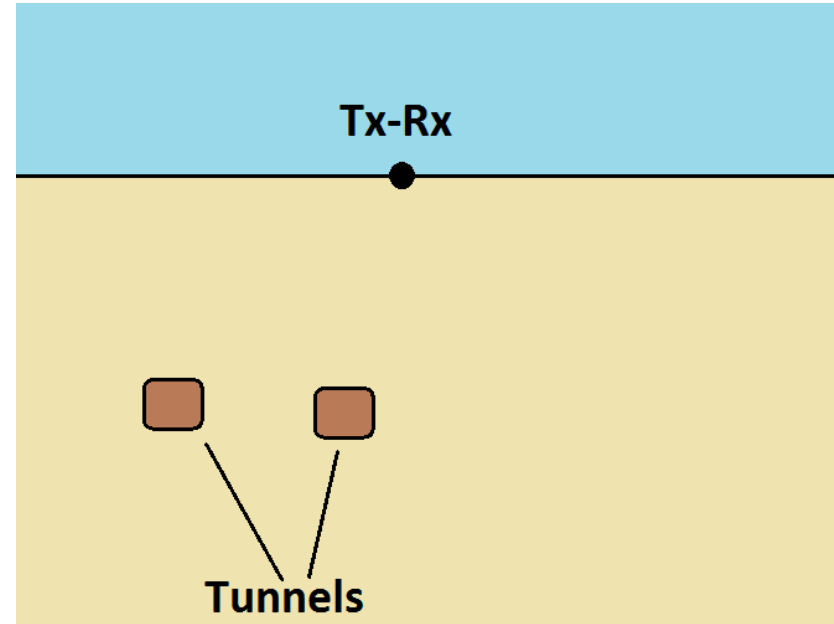


# Radargram 100 MHz

- Useful signals up to 300 ns
- See top of hyperbolas from tunnels



- Lower resolution
- Can see tunnels

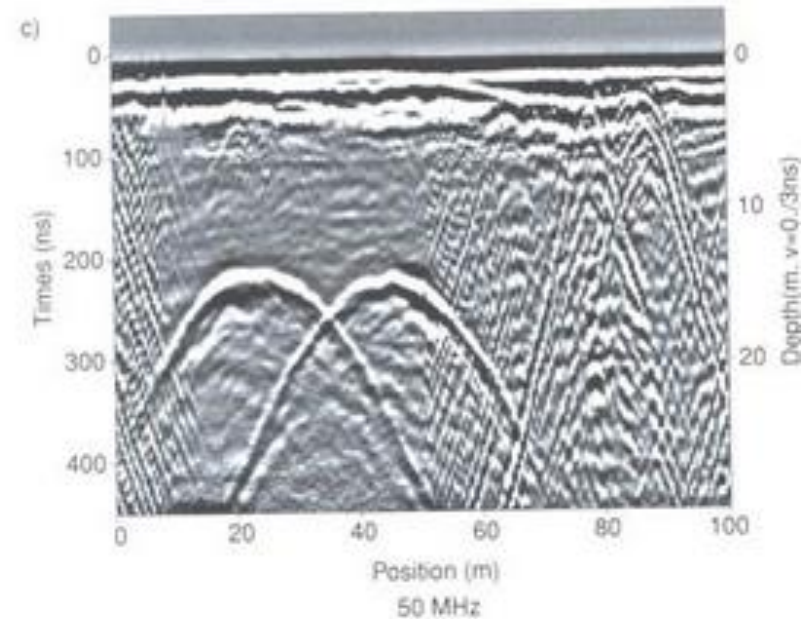
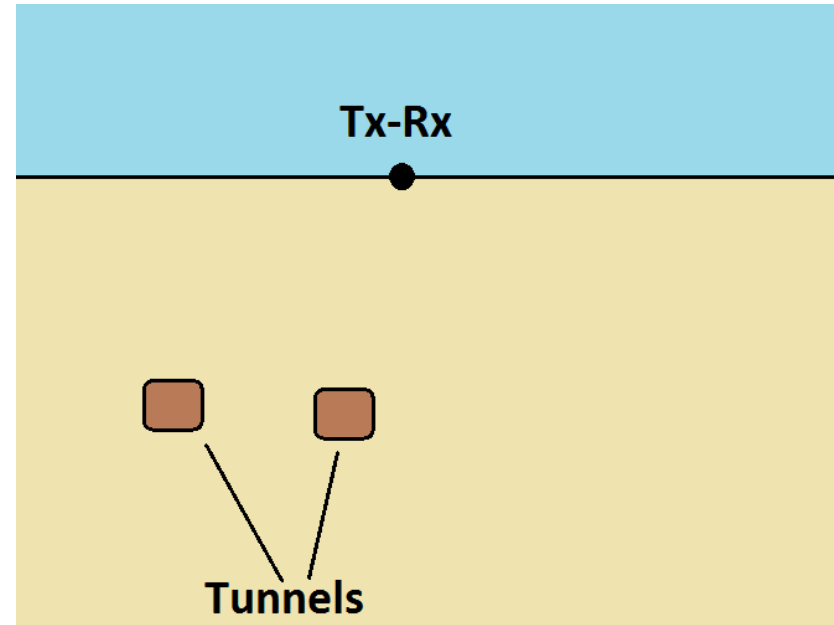


# Radargram 50 MHz

- Useful signals through 400 ns
- Well-defined hyperbolas from tunnels

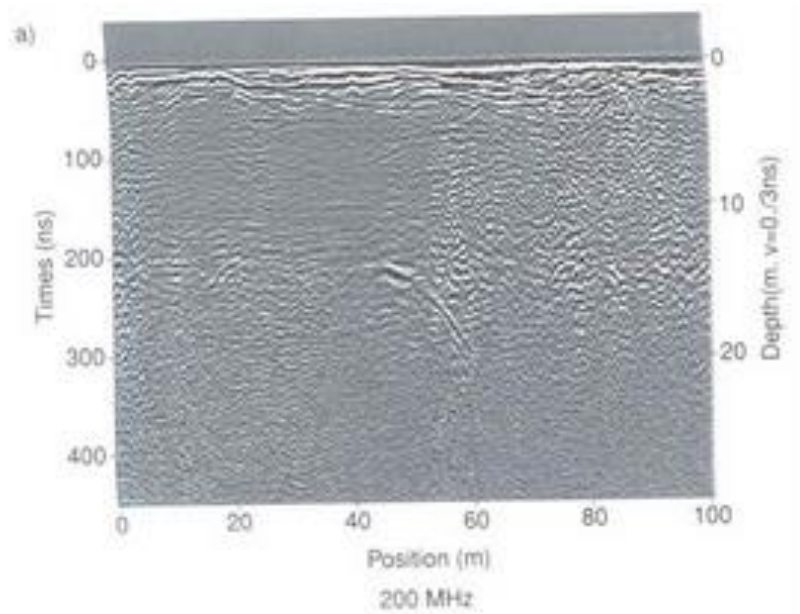


- Lower resolution image
- Best frequency for what we want to observe

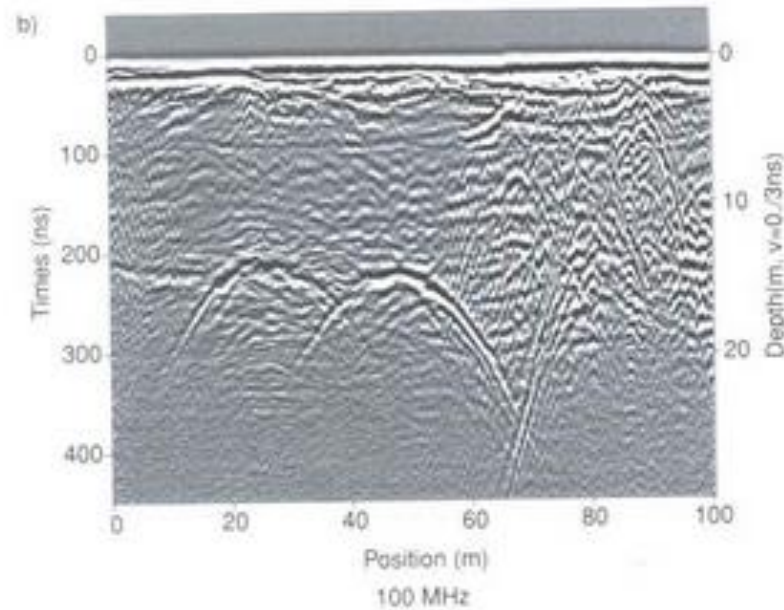


# Depth vs. Resolution

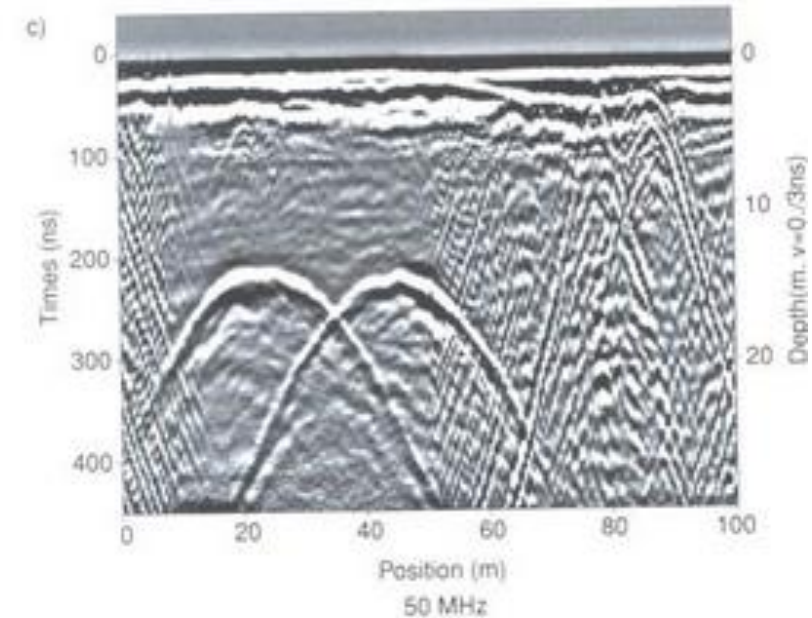
200 MHz



100 MHz



50 MHz

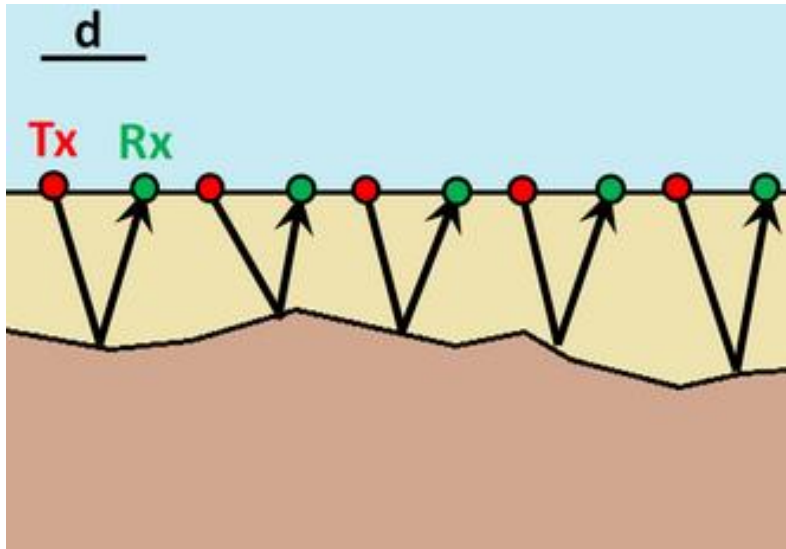


# Quiz

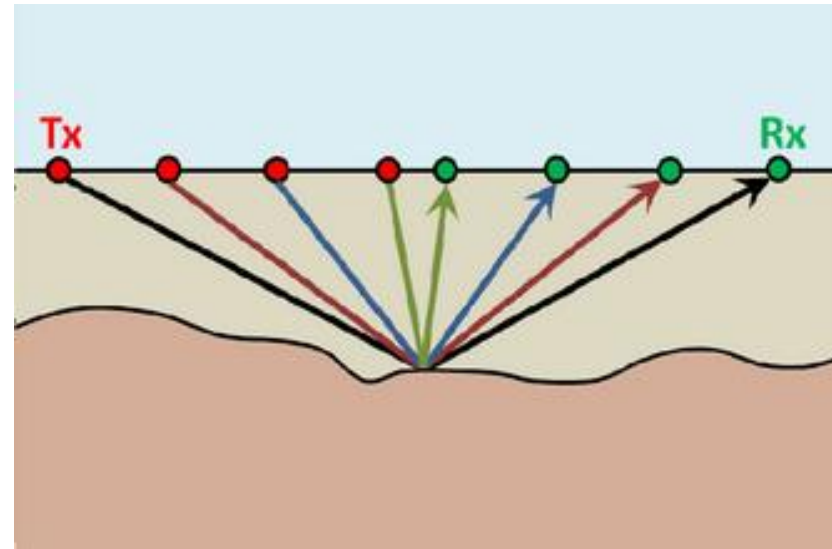
- True or false and why: While dc resistivity is only sensitive to the electrical resistivity, GPR data only response to the variation of electrical permittivity.
- Both dc resistivity and GPR can use electrical dipole sources. In a dc survey the dipole electrodes need to be in contact with the earth, but the GPR dipole source can be suspended in the air. Why?
- Which survey parameters determine the depth of investigation (DOI) in dc resistivity and GPR?



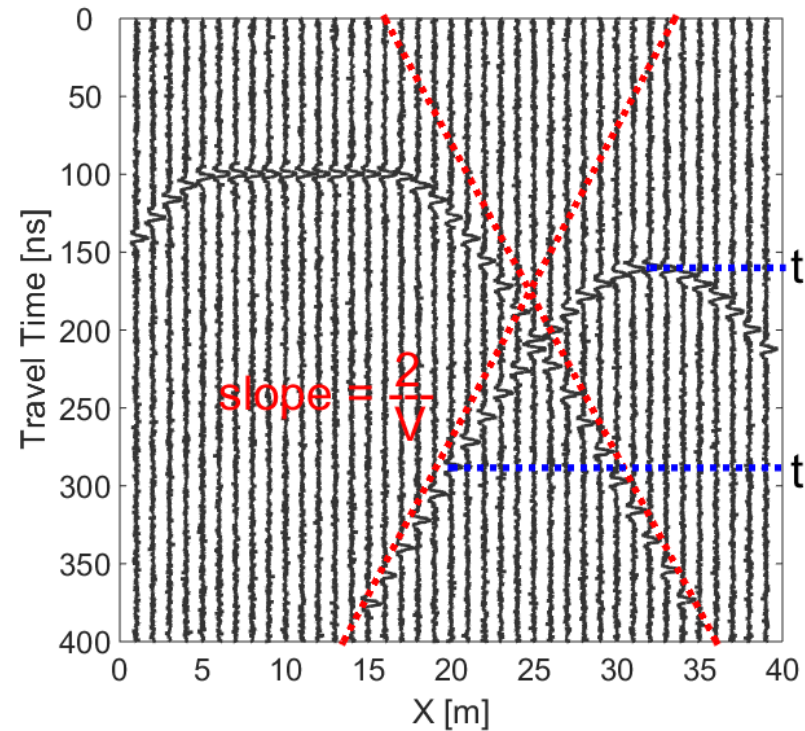
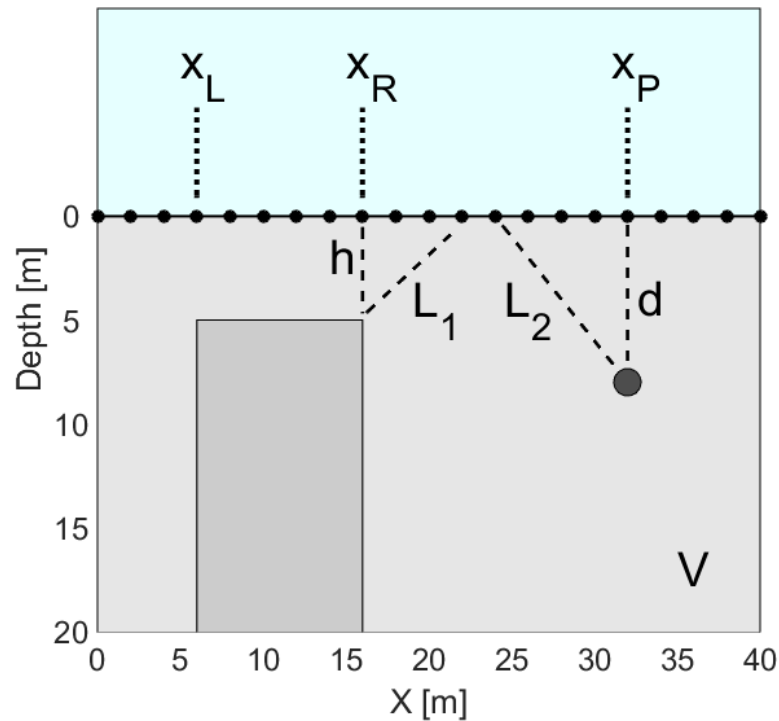
Common Offset



Common Midpoint



# Zero Offset: Finding Buried Objects



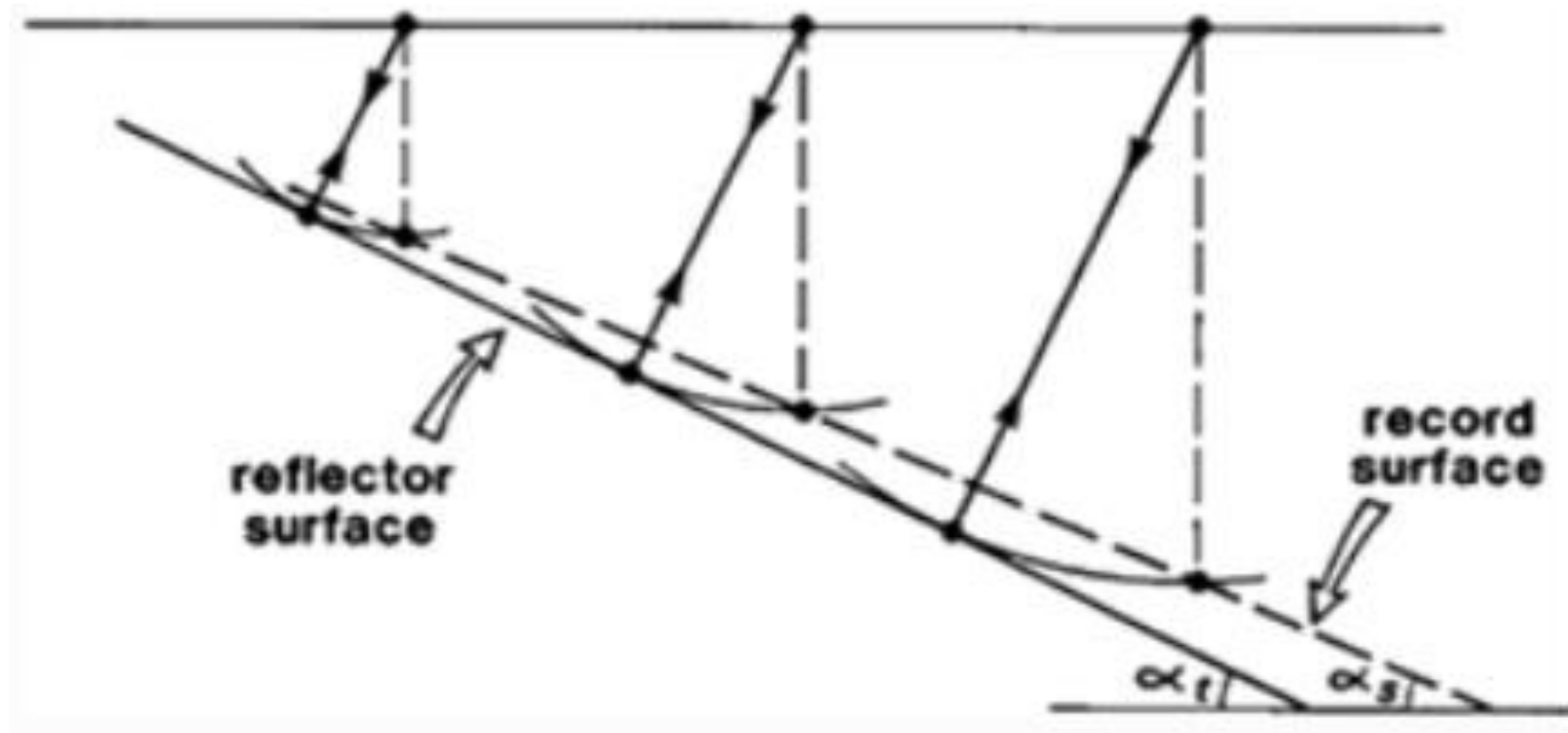
Two-way travel time  
for a point scatter

$$t_p = \frac{2L_2}{V} = \frac{2\sqrt{(x - x_p)^2 + d^2}}{V}$$

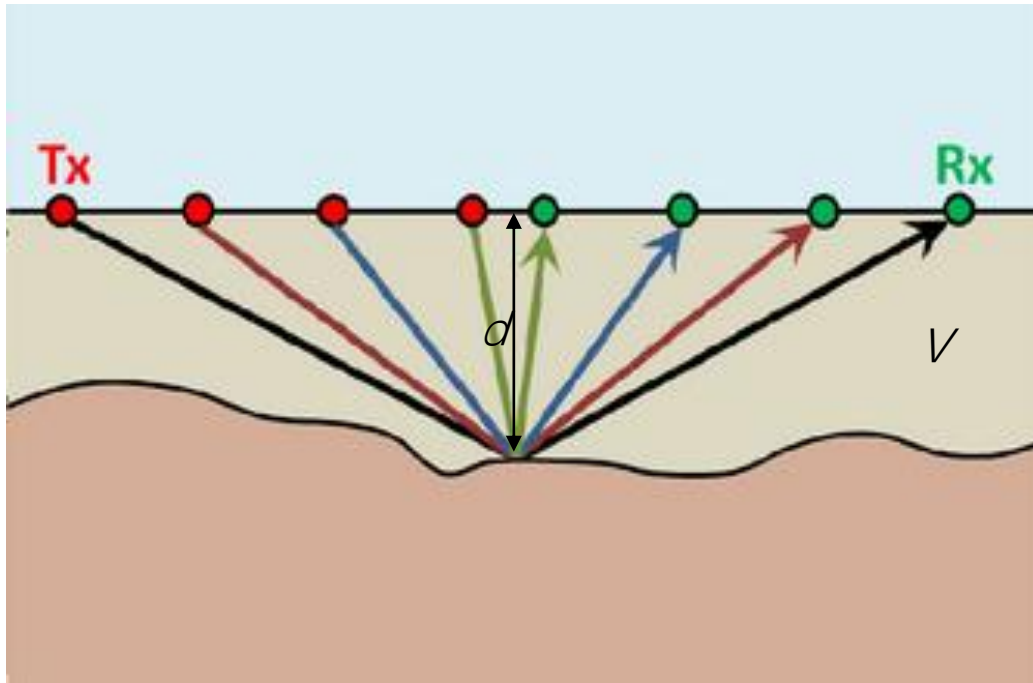
- (1) Estimate the velocity  $V$ . Can you think of two methods?
- (2) Calculate the depth of burial  $d$  or  $h$

# Migration

Zero offset survey along lines

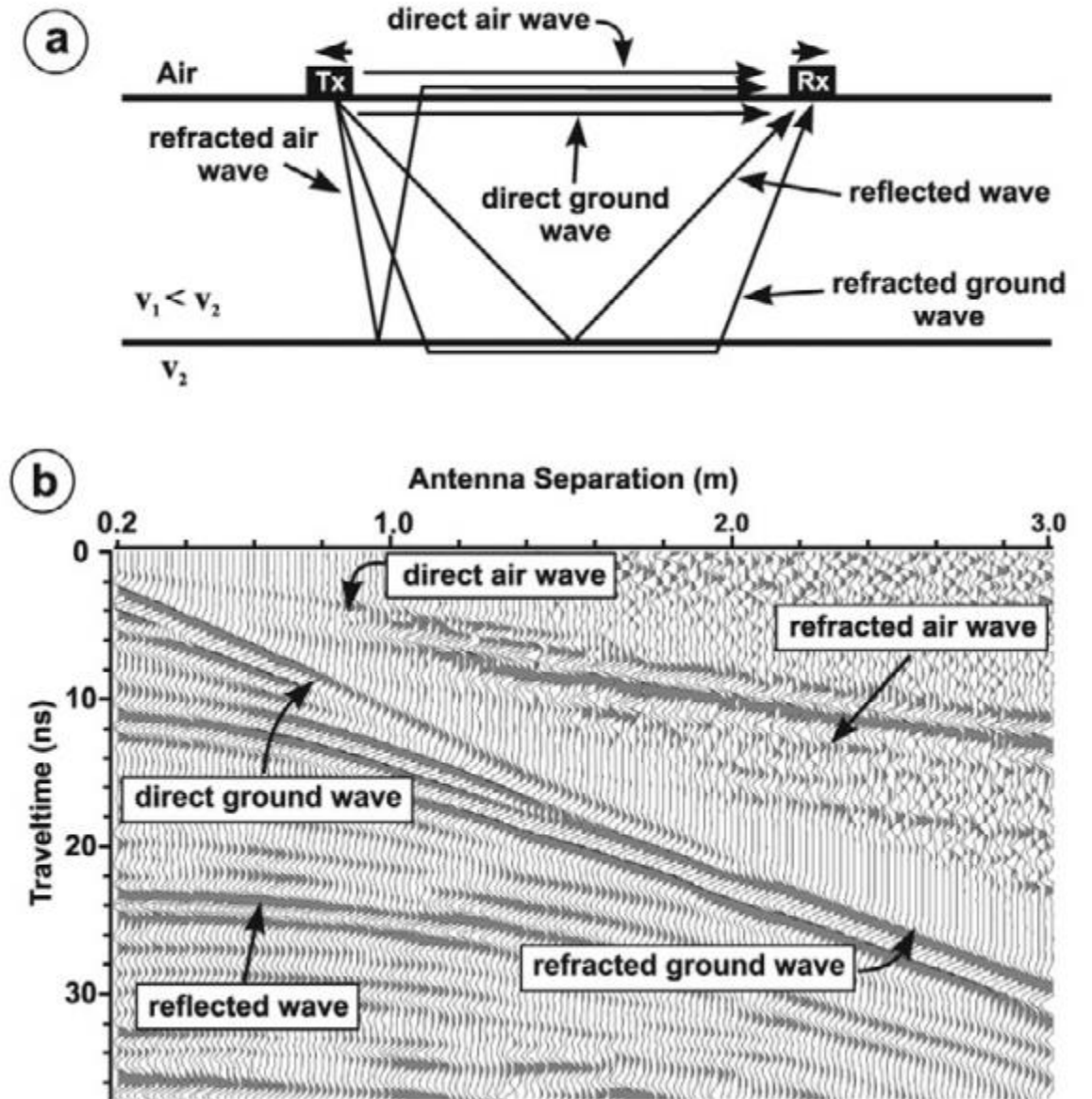


# Common Midpoint



$$t = \frac{2\sqrt{x^2 + d^2}}{V}$$

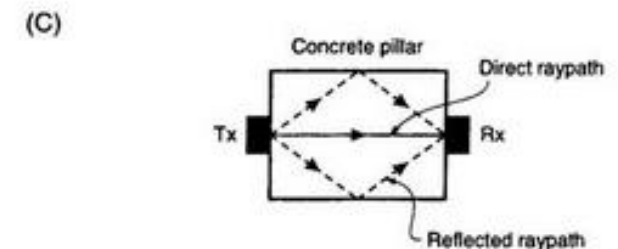
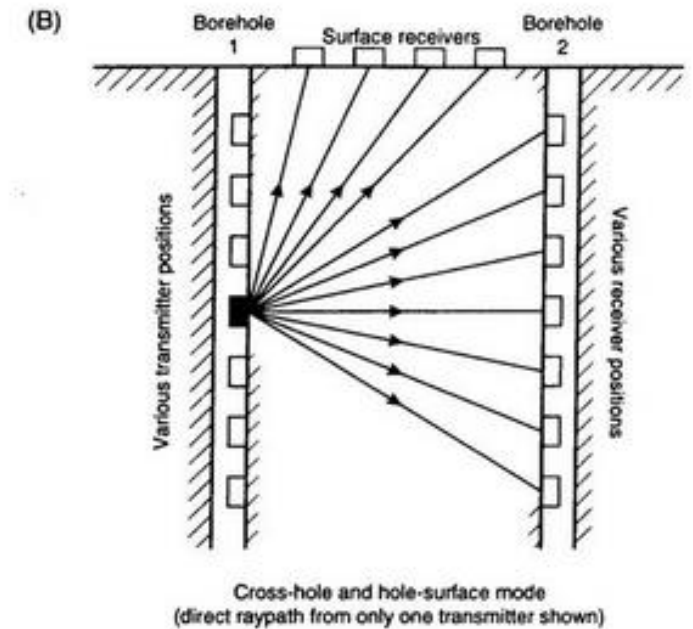
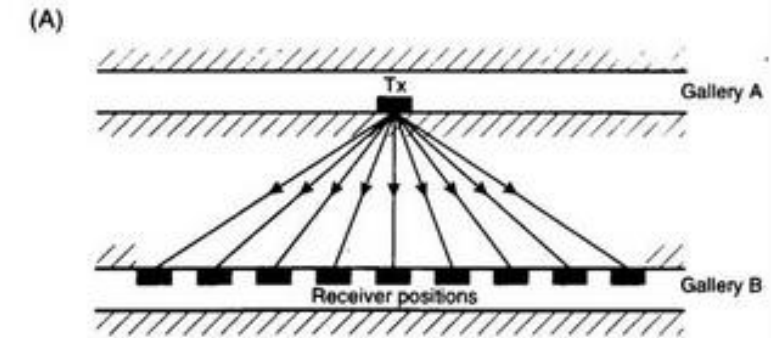
Solve for  $V$  and  $d$



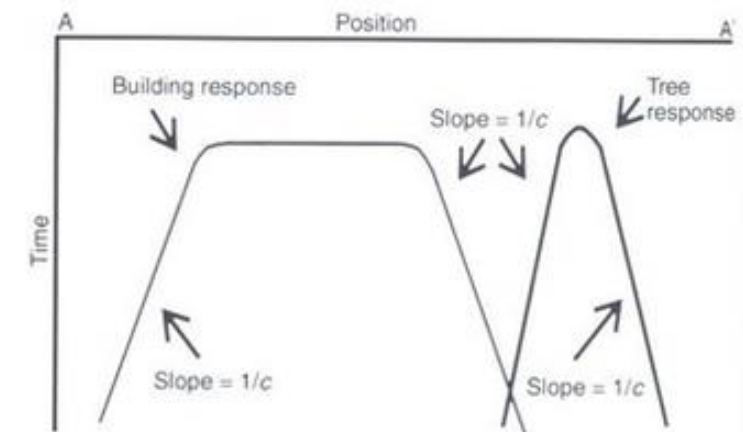
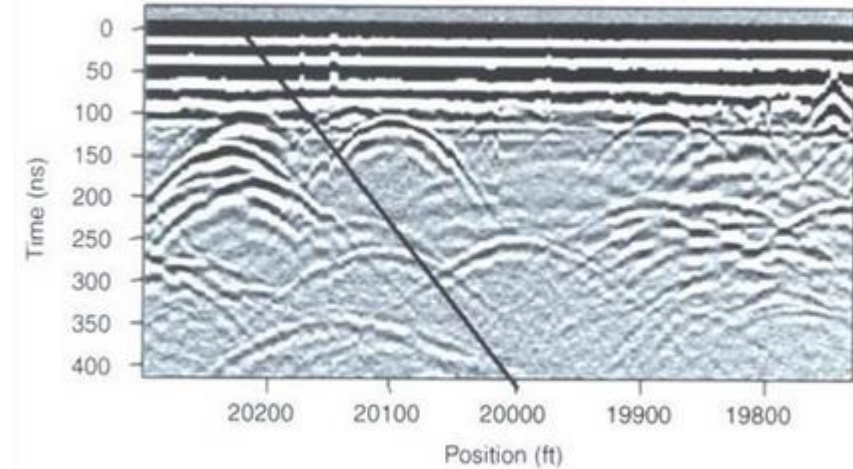
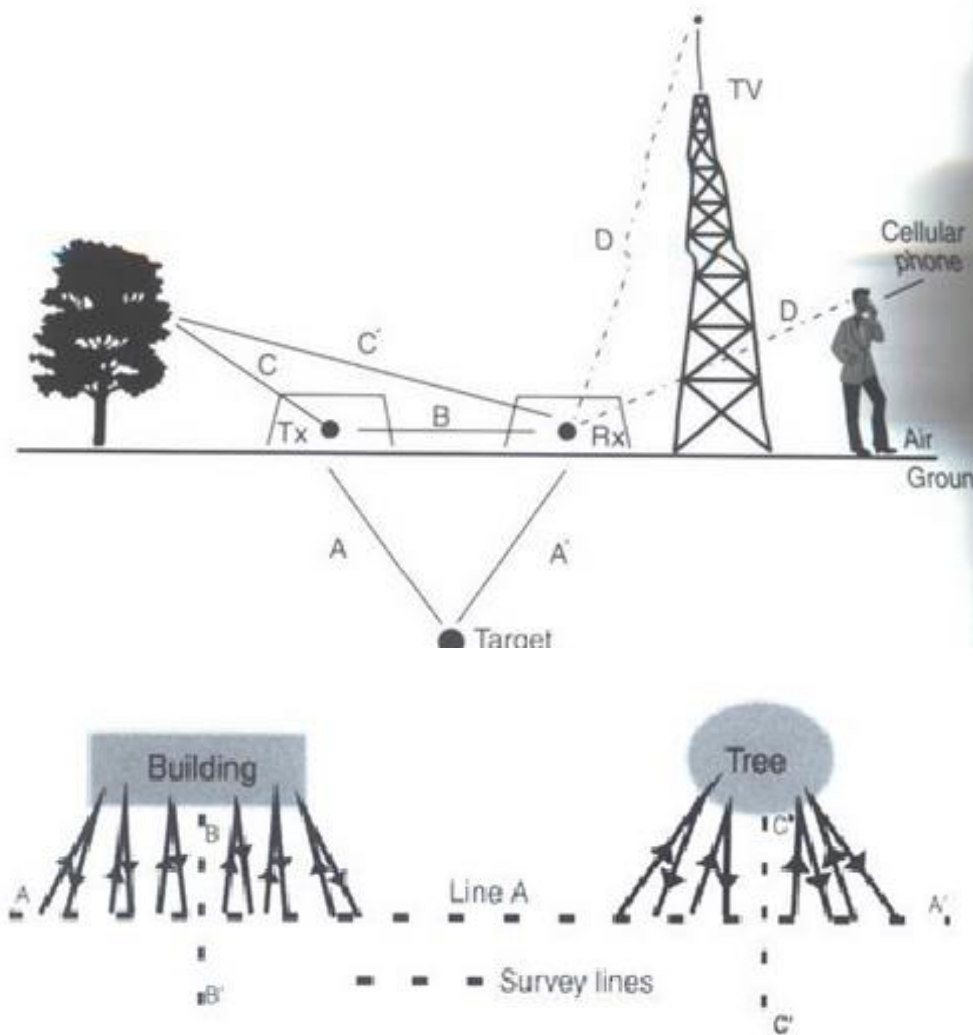


# Transillumination Surveys

- Tx and Rx are placed on opposing sides of a target.
- Sometimes many Tx and Rx
- Used for:
  - Structural integrity of mine shafts
  - Borehole surveys
  - Finding internal structures within objects

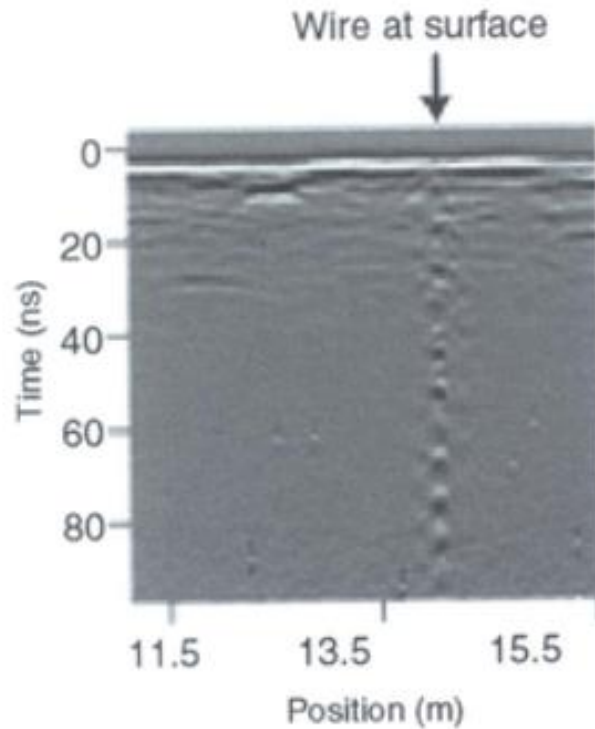


# Noise – External Radiowave or Above Ground Reflection





# Noise - “Ringing”



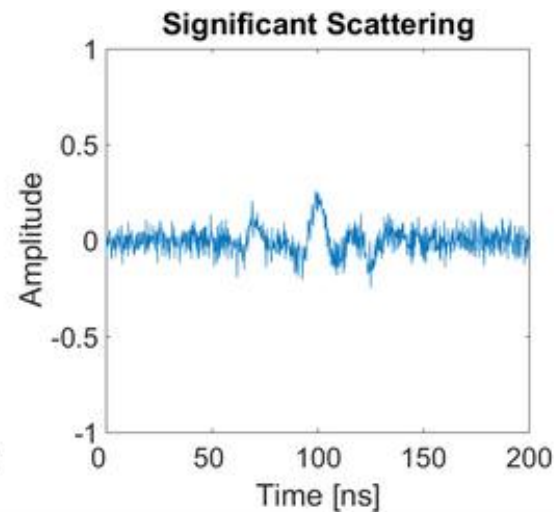
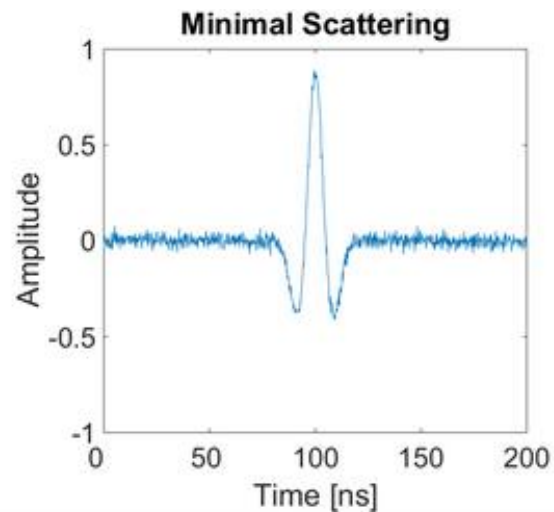
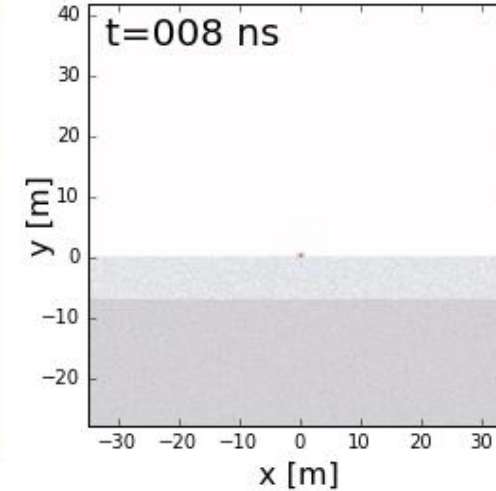
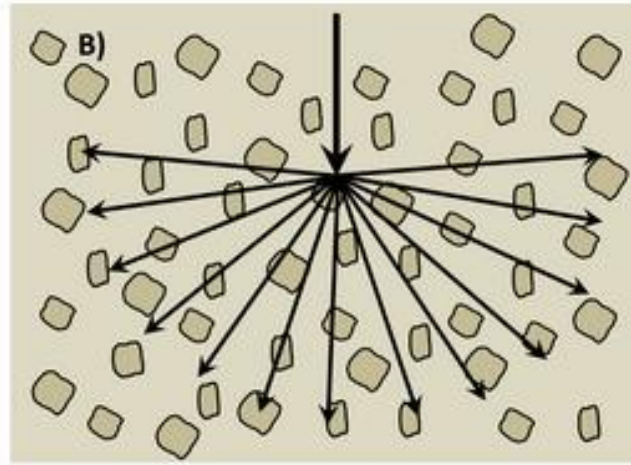
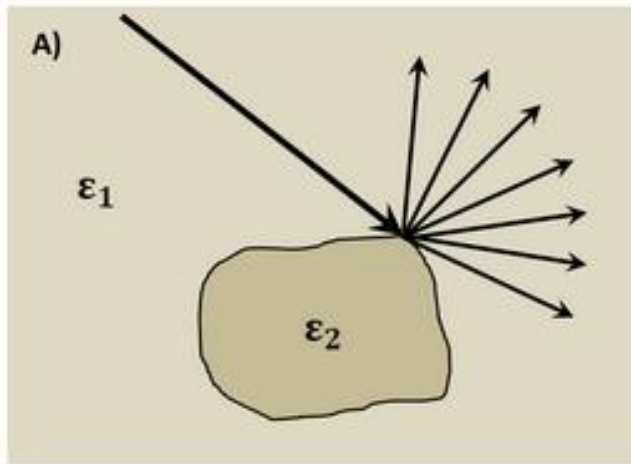
Wire below surface



2 nearby objects

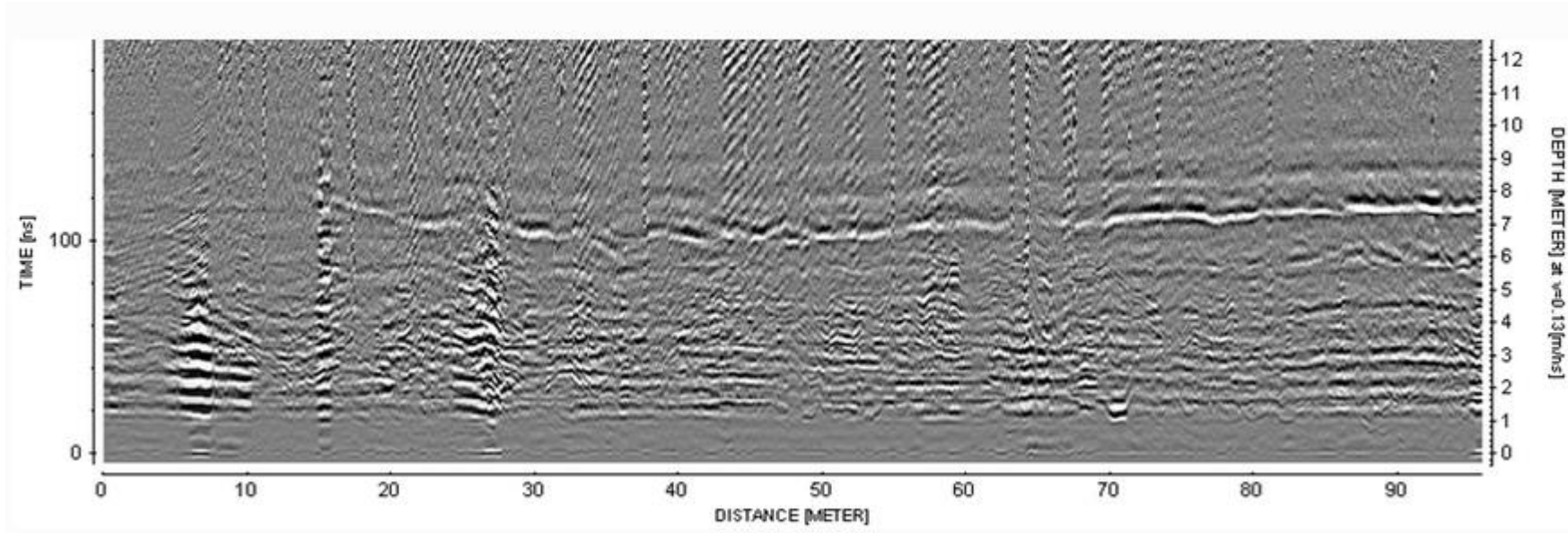
- Caused when signals reverberate in regular fashion
- Signal repeatedly bounces within a layer or between objects.

# Noise – Scattering



- Deviations in signal path due to localized non-uniformities.
- Reduces amplitude of usable signal and increases noise.

# Processing – Time-depth Conversion

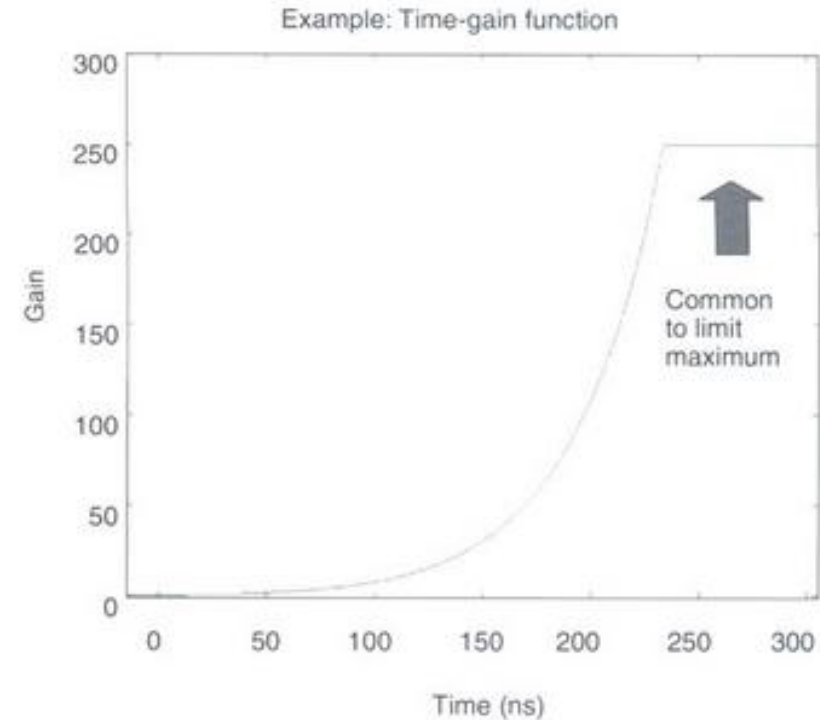
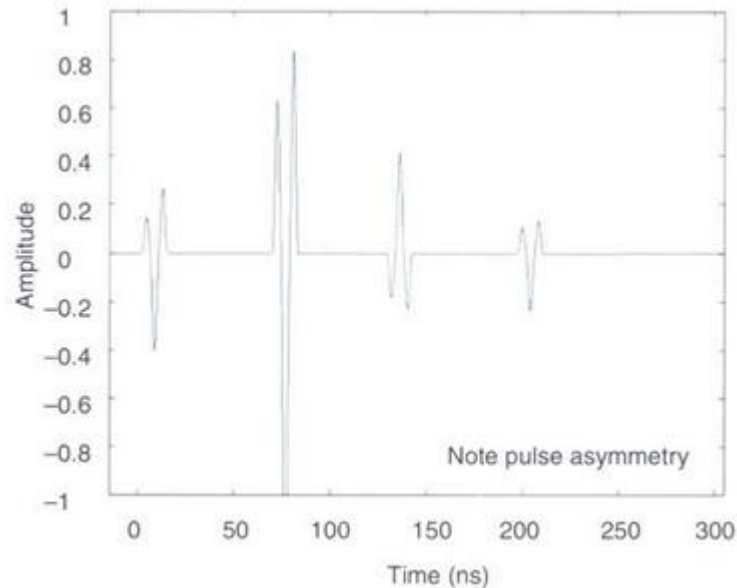
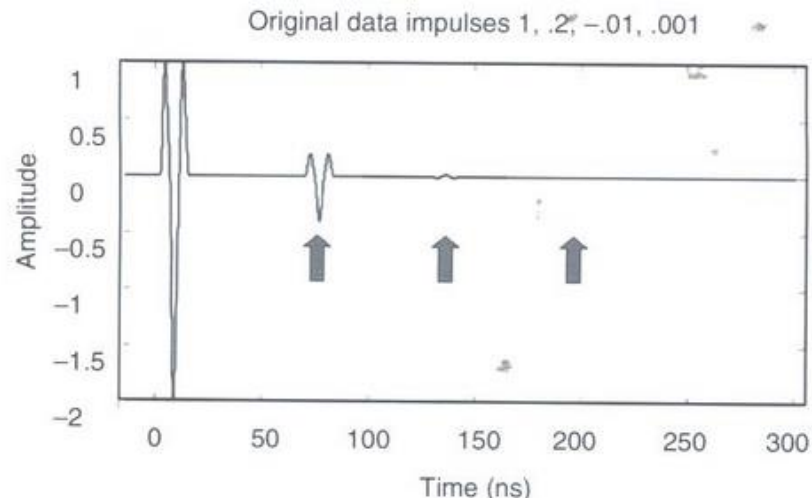


Apparent depth:

$$d_a = \frac{Vt}{2}$$

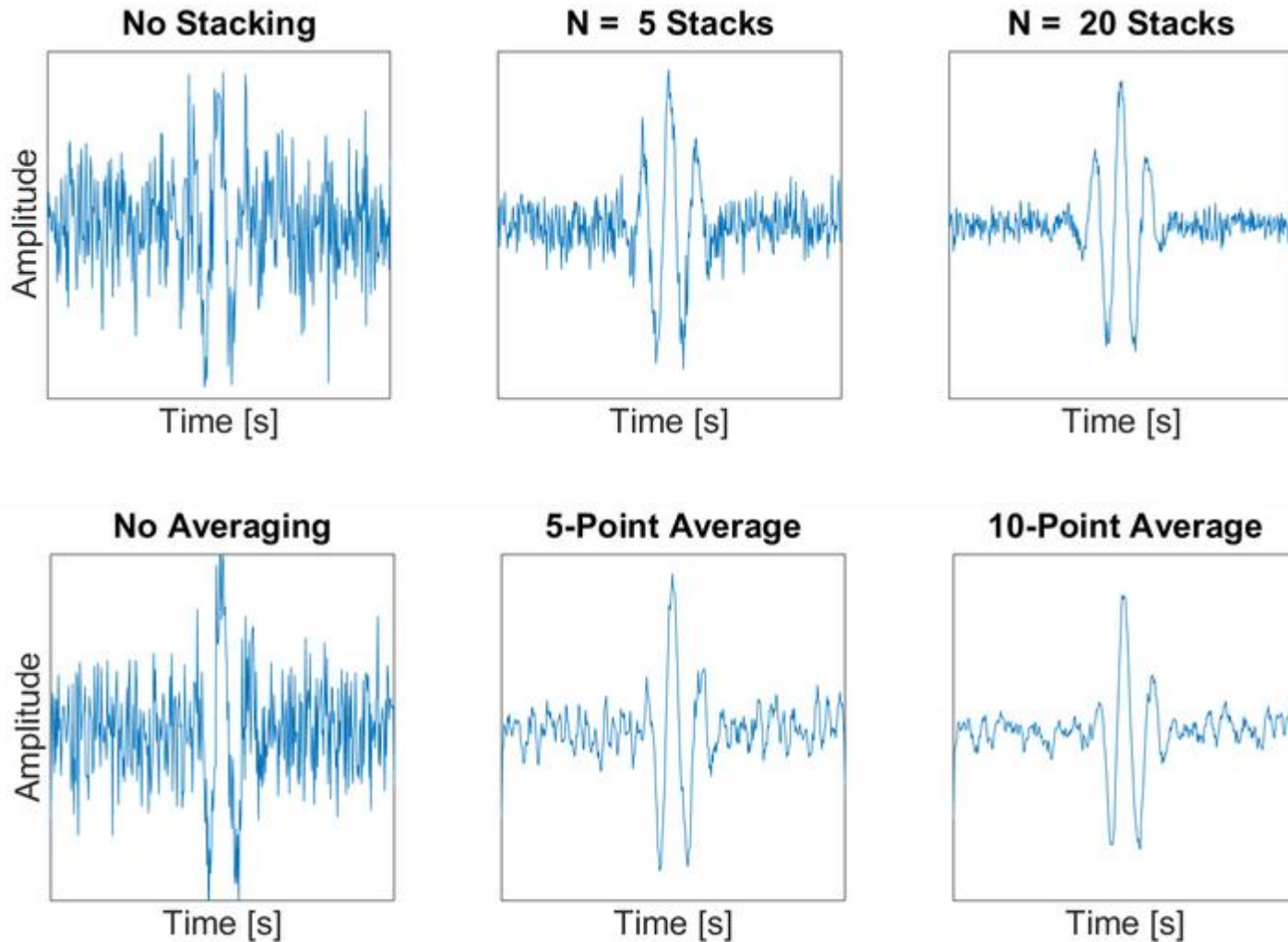
- Vertical axis usually 2-way travel time [ns]
- Get velocity first, then get an apparent depth

# Processing – Gain Correction



- Multiply raw data by a gain factor so that late signals can be recognized.
- Gain factor generally counteracts exponential decay in amplitude

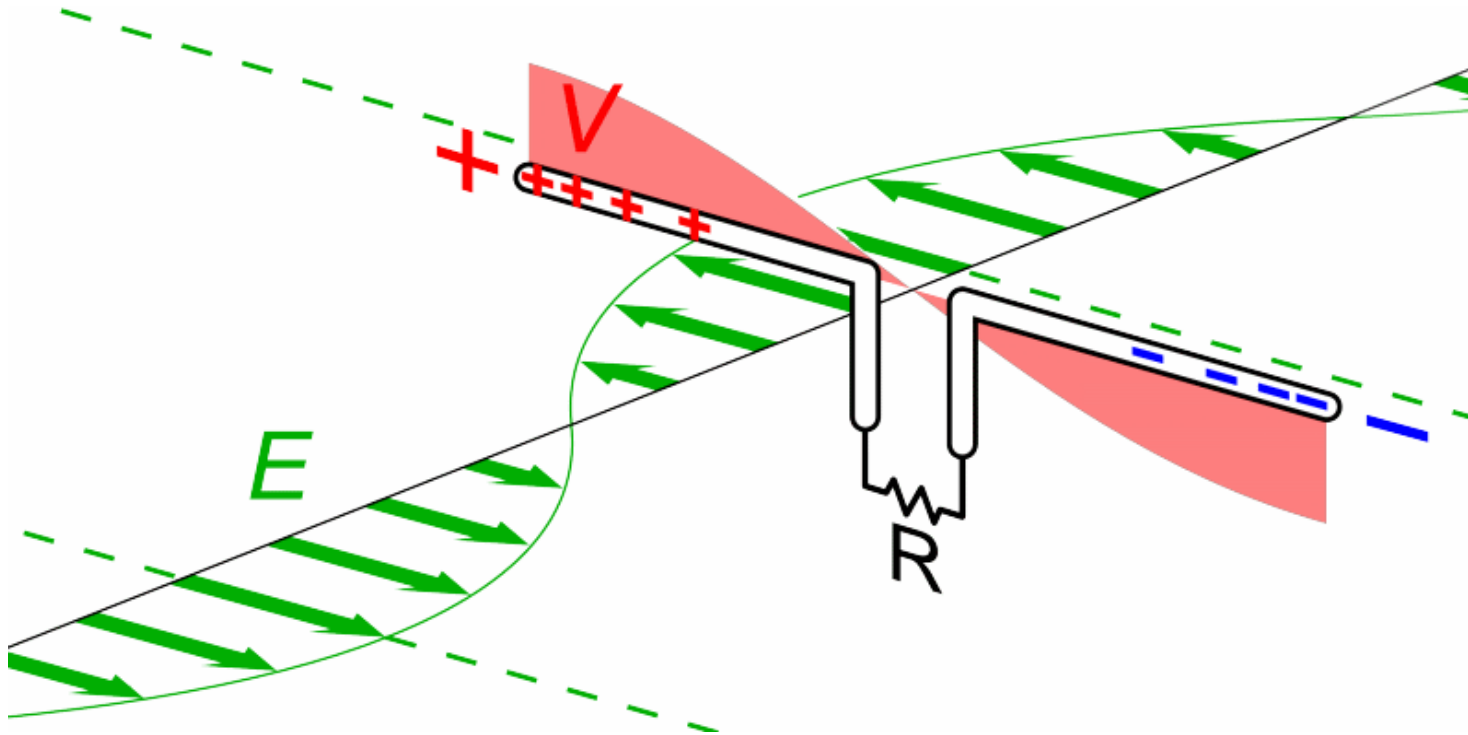
# Processing – Stacking and Averaging



- Data from repeated shots are averaged (stacked)
- Stacking reduces the amplitude of incoherent noise
- Wavelet signal is smooth whereas incoherent noise is random
- Smoothing decreases amplitude of random noise relative to returning signals.



# GPR Antenna



**Half-wave dipole antenna:**  
Length is determined by the intended wavelength (or frequency) of operation



# GPR Antenna

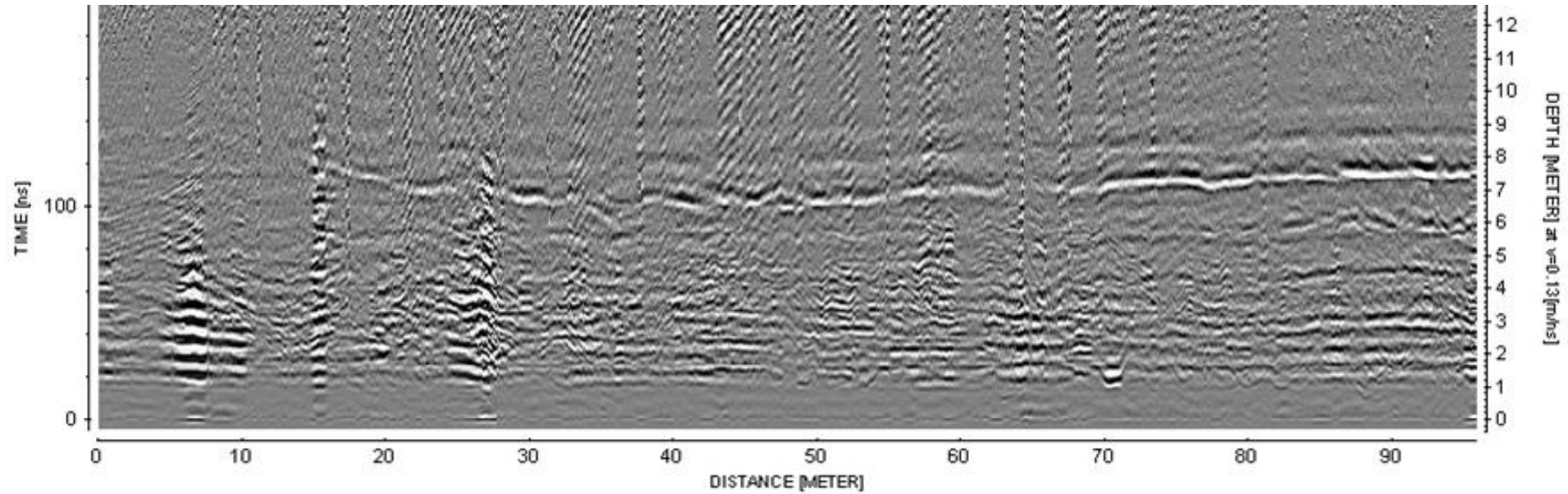




# Water Hazard in Potash Mine

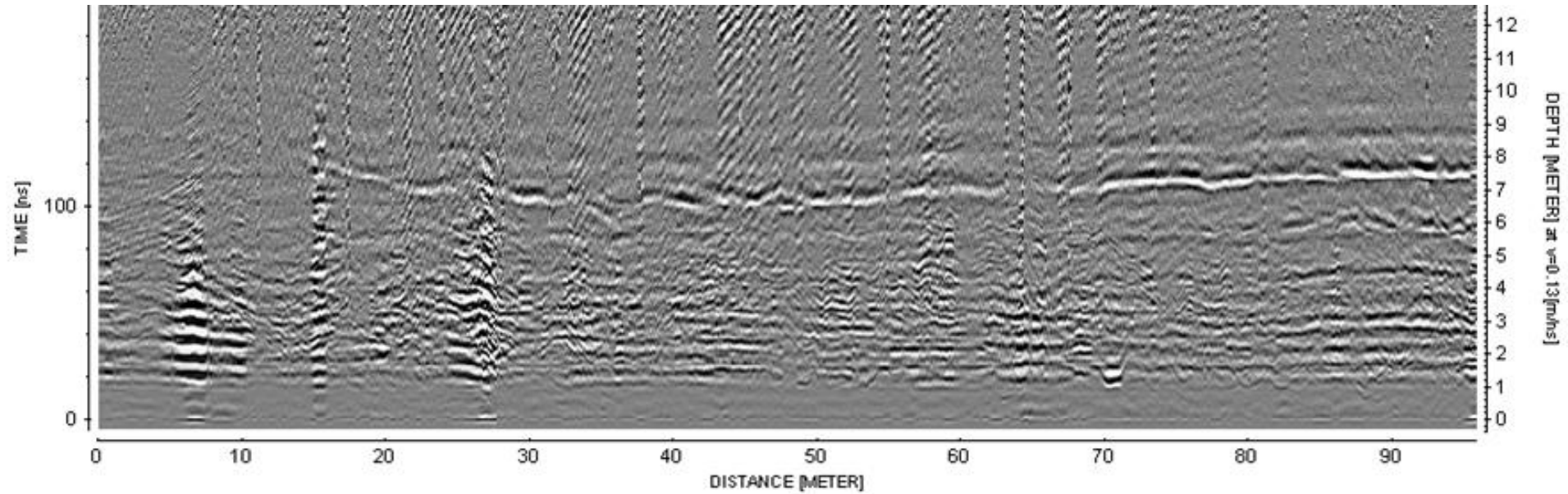






- Zero offset GPR survey performed.
- Arrival time to depth conversion performed

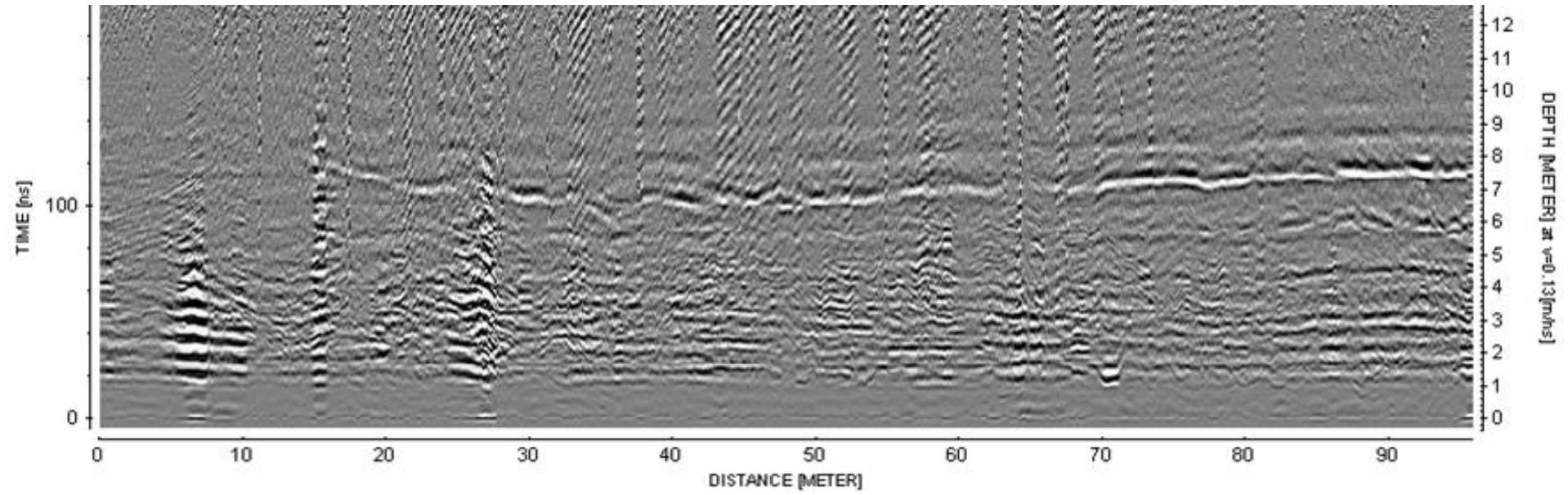
Q: Without a direct ground wave measurement or hyperbola to obtain propagation speed, how could they do conversion?



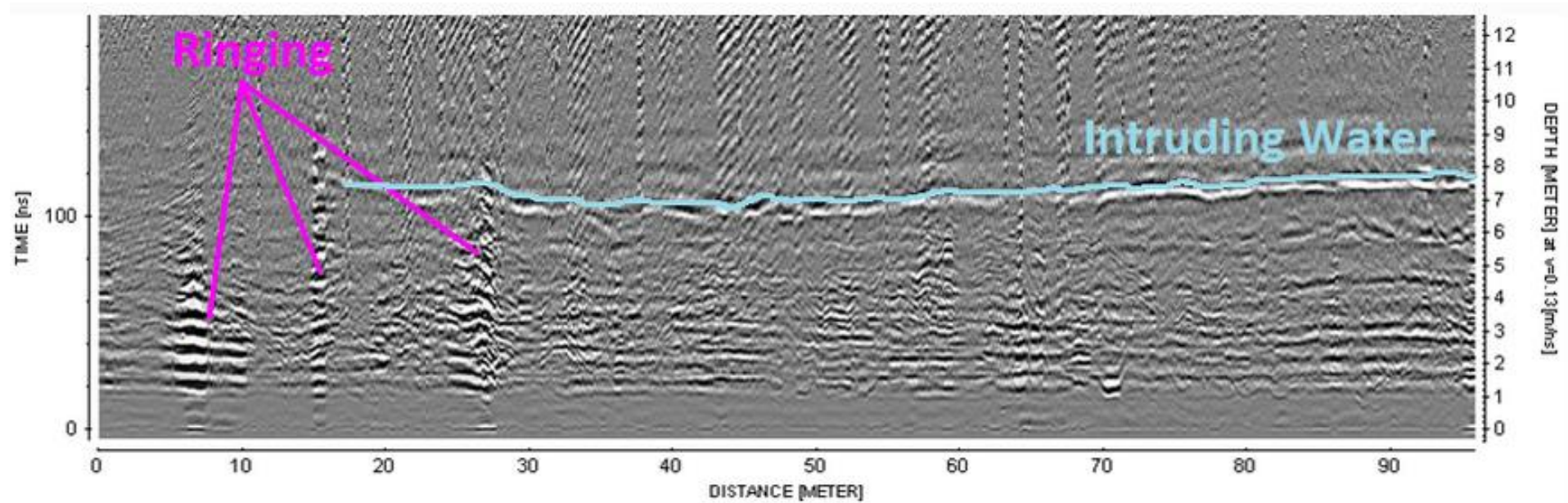
A: Potash in an anhydrite mineral.

From known physical properties,  $V \sim 0.13 \text{ m/ns}$

Apparent depth  $d_a = V t / 2$



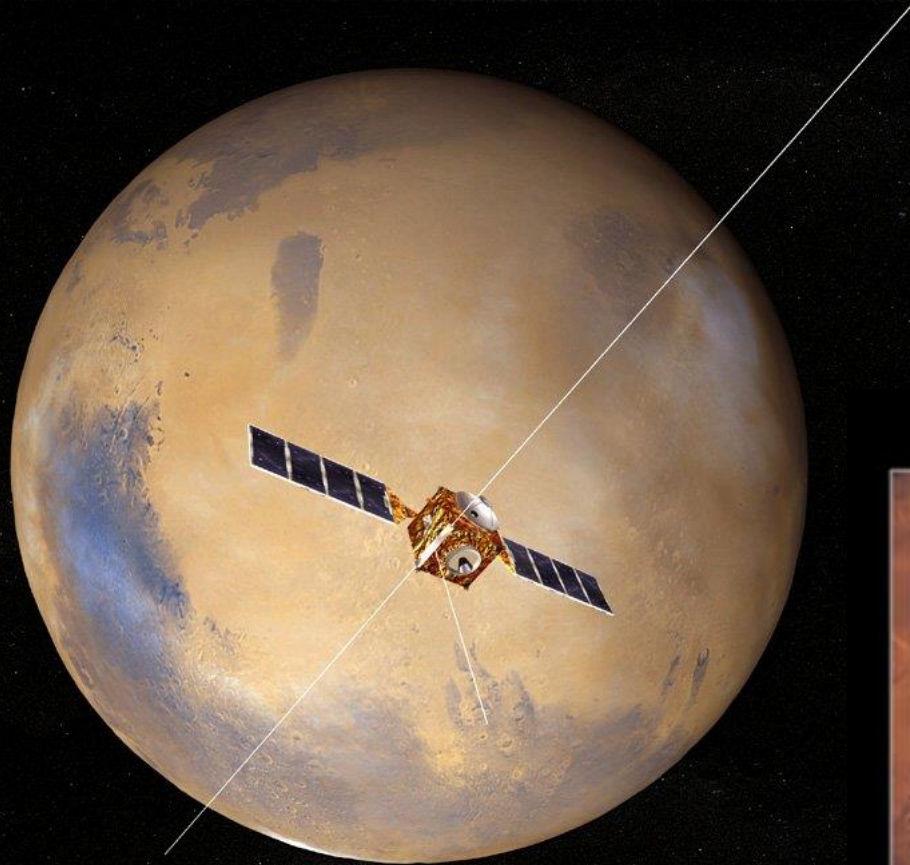
Q: What kinds of features do you see in the data?



- Strong reflector from intruding water (7 - 8 m into the wall)
- Water is delineated and seems to be coming from the right
- Ringing from mine infrastructure

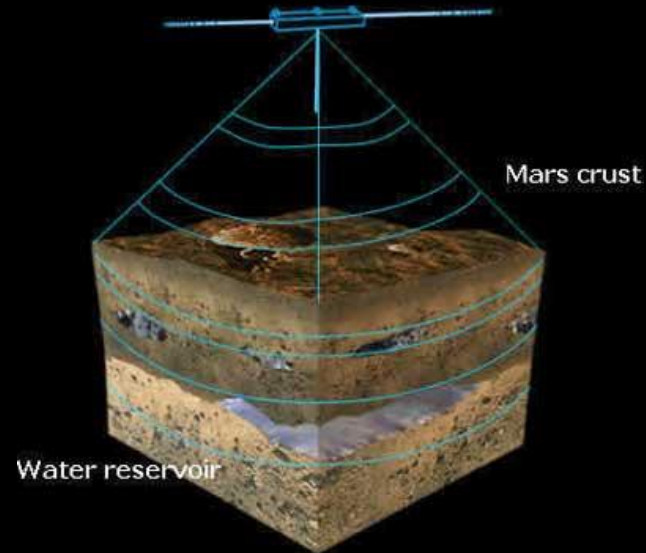


# Mars Radar



40 m dipole antenna  
1.8 ~ 5.0 MHz

MARSIS antenna beam



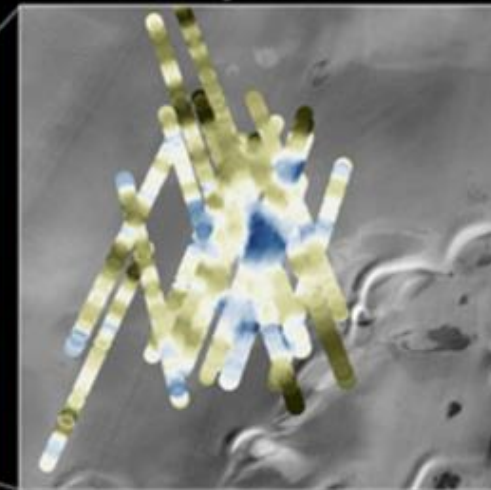
Mars Advanced Radar for  
Subsurface  
and Ionosphere Sounding  
MARSIS mission

Liquid water beneath ice cap

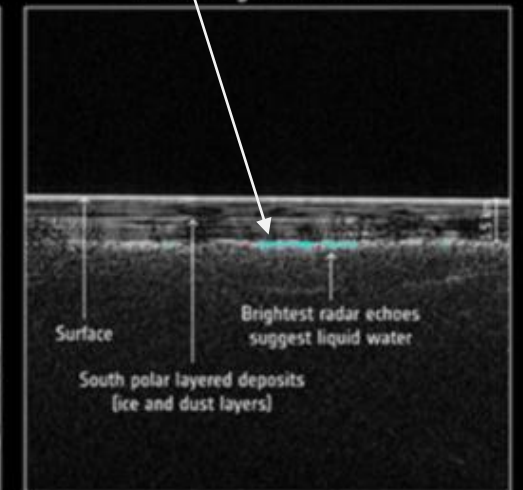
Mars south polar region



Mars Express radar footprints  
(blue = brightest radar echo)



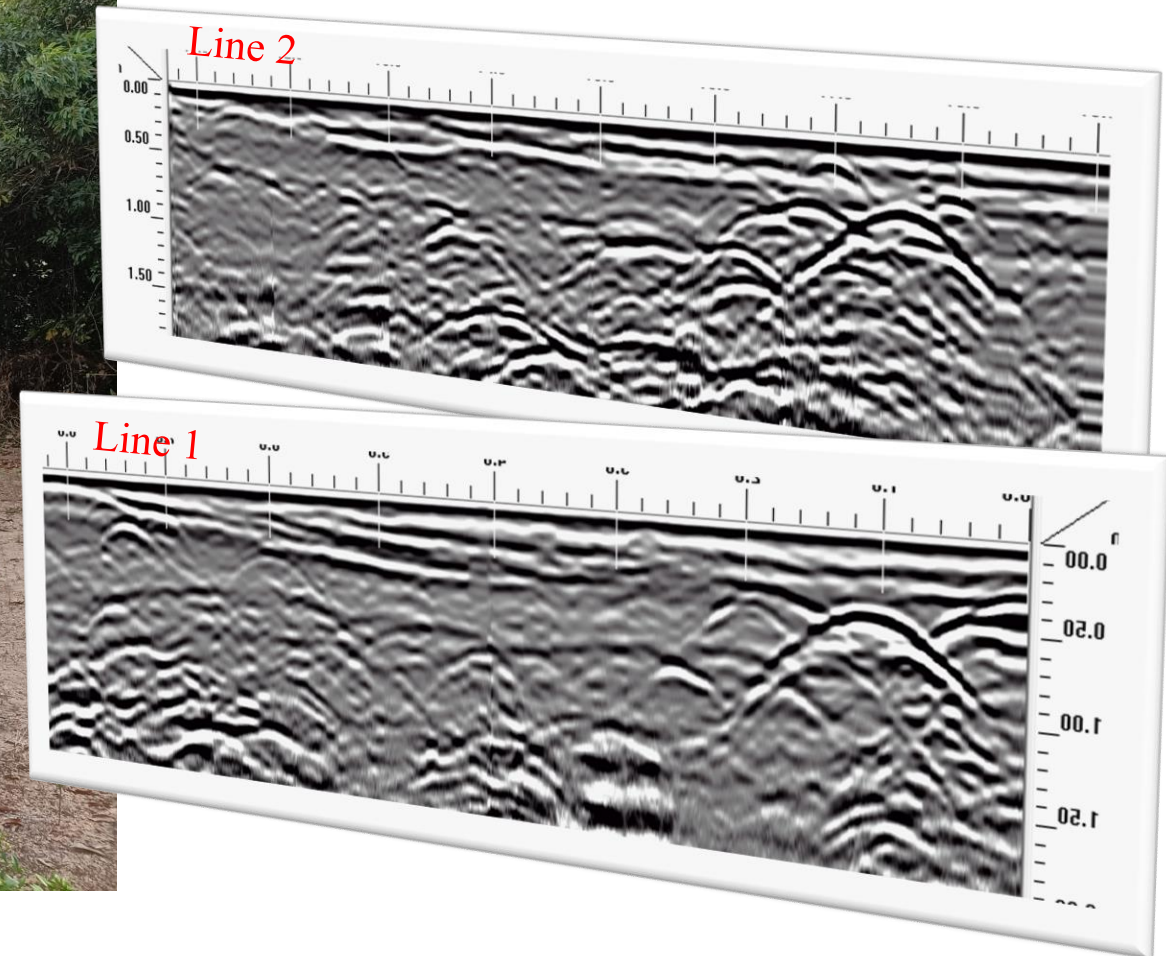
Radar image of subsurface





# GPR on SUSTech Campus

What are in the tomb?



# Summary of GPR

- EM at high frequency: Wave regime
- Physical properties utilized by EM/GPR
- Reflection, transmission, refraction and scattering
- Signal length scale and resolution
- Depth vs. spatial resolution
- GPR survey types
- GPR data analysis: velocity and depth
- GPR data processing
- GPR noise in practice
- GPR instruments: Antenna
- Applications: Water gushing in potash mines, MARS radar, Searching tombs.