# Tidal Theory with Newtonian Cooling

By Richard S. Lindzen¹) and Donald J. McKenzie²)

Summary – Processes representable by Newtonian cooling models are found to be of importance for tides in various planetary atmospheres. The equations of tidal theory are, therefore, rederived to include a rather general Newtonian cooling, and some of the effects of this inclusion are investigated.

#### 1. Introduction

The classical theory of linearized, adiabatic tides has been developed with fair thoroughness (see, for example Siebert [1],3) Lindzen [2]). Moreover, some attention has been devoted to the breakdown of this theory due to the growing importance of neglected nonlinear terms (Pekeris and Alterman [3]) and molecular diffusion (YANOWITCH [4]). Little attention has been given, however, to the influence of radiative processes—apart from their role in establishing thermal drives for tides (BUTLER and SMALL [5], LINDZEN [2], SIEBERT [1]). It turns out, however, that radiative cooling is in several cases sufficiently rapid to play a discernible role in motions with tidal periods. One such case occurs in the vicinity of the mesopeak of the Earth's atmosphere. Here the combined effects of CO<sub>2</sub> infrared cooling and O<sub>3</sub> photochemistry lead to cooling rates whose time scales are on the order of one day (LINDZEN and GOODY [6]). Another case concerns the Martian atmosphere which is composed largely of CO<sub>2</sub> (Prabhakara and Hogan [7]). Here too, radiative responses can be very rapid. Moreover, numerical experiments suggest a very large diurnal tide on Mars (Leovy and MINTZ [8]). In both these cases it appears possible to describe the effect of radiation (combined in the first case with photochemistry) upon deviations of temperature away from some equilibrium approximately with Newtonian cooling models; i.e., the radiative cooling is proportional to some function of space times the temperature deviation,  $\delta T$ . When the function of space depends only on altitude, it is a simple matter to extend classical tidal theory to allow for the inclusion of Newtonian cooling. This paper describes that extension. Apart from the potential practical importance of this extension in the above mentioned cases, it also displays how a dissipative process may affect the gross structure of tidal waves.

<sup>1)</sup> National Center for Atmospheric Research, Boulder, Colorado, USA.

<sup>2)</sup> University of Washington, Seattle, Washington, USA.

<sup>3)</sup> Numbers in brackets refer to References, page 96.

## 2. Mathematical development

The basic dynamic and thermodynamic equations used are those described by Siebert [1]. The inclusion of Newtonian cooling alters only the expression of the first law of thermodynamics (equation (3.17) of Siebert [1]) which now becomes

$$\frac{DT}{Dt} = \frac{gH}{\rho_0} \frac{M(\gamma - 1)}{R} \frac{D\varrho}{Dt} + \frac{M(\gamma - 1)}{R} J - a(z) \delta T, \qquad (1)$$

where

T = temperature,

o = density,

 $\delta T$  = tidal temperature perturbation from basic temperature,  $T_0$ ,

J = heating per unit mass per unit time,

 $\gamma = C_p/C_v = 1.40$ ,

M =molecular weight,

R = gas constant,

 $o_0$  = basic density,

a(z) = rate coefficient for Newtonian cooling,

 $H = R T_0/M g,$ 

g = acceleration of gravity.

Also,

$$K = (\gamma - 1)/\gamma.$$

The remaining equations are

$$i \omega u - 2 \Omega v \cos \theta = -\frac{1}{a} \frac{\partial}{\partial \theta} \left( \frac{\partial p}{\rho_0} \right),$$
 (2)

$$i \omega v + 2 \Omega u \cos \theta = -\frac{1}{a \sin \theta} \frac{\partial}{\partial \varphi} \left( \frac{\delta p}{\varrho_0} \right),$$
 (3)

$$\frac{\partial \delta p}{\partial z} = -g \, \delta \varrho \,, \tag{4}$$

$$i\,\omega\,\delta\varrho + w\,\frac{d\varrho_0}{dz} + \varrho_0\,\chi = 0\,,\tag{5}$$

and

$$\frac{\delta p}{p_0} = \frac{\delta \varrho}{\varrho_0} + \frac{\delta T}{T_0} \,, \tag{6}$$

where a time dependence of the form  $e^{iwt}$  has been assumed and

u =westerly velocity,

v = northerly velocity,

w = vertical velocity,

 $\delta p$  = tidal pressure perturbation,

 $P_0$  = basic pressure,

 $\delta \varrho = \text{tidal density perturbation,}$ 

 $\Theta$  = colatitude,

 $\varphi$  = longitude,

 $\Omega$  = rotation rate =  $2 \pi/1$  day,

a = radius of planet,

 $\chi$  = velocity divergence.

Gravitational excitation has been neglected.

From (2) and (3) one obtains

$$\chi - \frac{\partial w}{\partial z} = \frac{i \, \omega}{4 \, a^2 \, \Omega^2} \, F\left(\frac{\delta p}{\varrho_0}\right), \tag{7}$$

where F is the following operator

$$F = \frac{1}{\sin\theta} \frac{\partial}{\partial\theta} \left( \frac{\sin\theta}{f^2 - \cos^2\theta} \frac{\partial}{\partial\theta} \right) + \frac{1}{f^2 - \cos^2\theta} \left[ \frac{i}{f} \frac{f^2 + \cos^2\theta}{f^2 - \cos^2\theta} \frac{\partial}{\partial\varphi} + \frac{1}{\sin^2\theta} \frac{\partial^2}{\partial\varphi^2} \right],$$
(8)

and

$$f = \frac{\omega}{2 \Omega}$$
.

In general, J,  $\chi$ , w,  $\delta p$ ,  $\delta \varrho$  and  $\delta T$  may be expanded in terms of the eigenfunctions of the equation

$$F(\psi_n) + \frac{4 a^2 \omega^2}{g h_n} \psi_n = 0 , \qquad (9)$$

where  $h_n$  = equivalent depth.

(9) is Laplace's Tidal Equation. If we restrict ourselves to a given n, we must obtain an equation describing the vertical (z) structure of that mode. In adiabatic tidal theory the equation has its simplest form for the dependent variable

$$Y_n = \chi_n - \frac{K J_n}{g H}, \tag{10}$$

where  $\chi_n$ ,  $J_n$ , etc. are the coefficients in the expansions

$$\chi = \sum_n \chi_n(z) \; \psi_n( heta, \, arphi) \; e^{iwt} \; ,$$
  $J = \sum_n J_n(z) \; \psi_n( heta, \, arphi) \; e^{iwt} \; , \;\; ext{etc.}$ 

Moreover, it is especially easy to relate the various fields to  $Y_n$  (see Siebert [1]). We should like to determine such a simplifying dependent variable for the present case as well. The procedure is implicitly contained in Lindzen [9]. There, it was found that for the solar diurnal tide, there is a mode with a horizontal dependence of the form  $\sin\theta \cos\theta$ , and for this mode the quantity  $\chi - (\kappa J/g H)$  is identically equal to

zero. Redoing the analysis of that paper, using equation (1) of this paper, one finds that for the same horizontal dependence, the quantity

$$\left(1+\frac{a(z)}{i\,\omega\,\gamma}\right)\chi-\frac{\kappa\,J}{g\,H}-\frac{a(z)}{i\,\omega\,\gamma}\,\frac{w}{H}\,\frac{dH}{dz}$$

is identically equal to zero. This same quantity proves to be a simple dependent variable. However, it proves even simpler to use the above quantity multiplied by  $(1 + \lceil a(z)/i \omega \rceil)^{-1}$ . Thus, as our dependent variable we will take

$$L_n = \left(1 + \frac{a}{i\,\omega}\right)^{-1} \left\{ \left(1 + \frac{a}{i\,\omega\,\gamma}\right) \chi_n - \frac{\kappa J_n}{g\,H} - \frac{a}{i\,\omega\,\gamma} \,\frac{w}{H} \,\frac{dH}{dz} \right\}. \tag{11}$$

Using (11), together with (1), (4), (5), (6) and (9) one obtains after some manipulation

$$\frac{d^{2}L_{n}}{dx^{2}} - \left\{1 + \frac{a}{i\omega\gamma}\left(1 + \frac{a}{i\omega\gamma}\right)\frac{1}{H}\frac{dH}{dx}\right\}\frac{dL_{n}}{dx} + \left\{\frac{1}{h_{n}}\left(\frac{dH}{dx} + KH\right)\frac{a}{i\omega\gamma}\frac{1}{H}\frac{dH}{dx}\right\}\left(1 + \frac{a}{i\omega\gamma}\right)^{-1}L_{n}\right\} = \left(1 + \frac{a}{i\omega\gamma}\right)^{-1}\frac{\kappa}{v_{n}h_{n}}J_{n}, \qquad (12)$$

where

$$x = \int_{0}^{z} \frac{dz'}{H(z')} \ . \tag{13}$$

From equations (1) to (7), and (9) we obtain the following equations relating  $u_n$ ,  $v_n$ ,  $w_n$ ,  $\delta p_n$  and  $\delta T_n$  to  $L_n(z)$  and  $\Psi_n(\theta, \varphi)$ :

$$u_n = \frac{\gamma g h_n}{4 a \Omega^2 (f^2 - \cos^2 \theta)} \left( \frac{dL_n}{dx} - L_n \right) \left( \frac{\partial}{\partial \theta} - \frac{i}{f} \omega t \theta \frac{\partial}{\partial \psi} \right) \psi_n , \qquad (14)$$

$$v_n = \frac{i \gamma g h_n}{4 a \Omega^2 (f^2 - \cos^2 \theta)} \left( \frac{dL_n}{dx} - L_n \right) \left( \frac{\cos \theta}{f} \frac{\partial}{\partial \theta} - \frac{i}{\sin \theta} \frac{\partial}{\partial \varphi} \right) \psi_n, \qquad (15)$$

$$w_n = \gamma h_n \left( \frac{dL_n}{dx} + \left( \frac{H}{h_n} - 1 \right) L_n \right) \psi_n , \qquad (16)$$

$$\delta p_n = \frac{\varrho_0 \, \gamma \, g \, h_n}{i \, \omega} \left( \frac{dL_n}{dx} - L_n \right) \psi_n \,, \tag{17}$$

$$\delta T_{n} = \frac{M}{i \omega R} \left( 1 + \frac{a}{i \omega \gamma} \right)^{-1} \left[ \gamma g \left( \frac{dH}{dx} \left( \frac{h_{n}}{H} - 1 \right) - K H \right) L_{n} - \frac{\gamma g h_{n}}{H} \frac{dH}{dx} \frac{dL_{n}}{dx} + K J_{n} \right] \psi_{n} .$$
(18)

Reference to SIEBERT [1] will show that (14) to (18) in  $L_n(x)$  are identical to the equations for the adiabatic case in terms of  $Y_n$  [see (10)]—with the exception of the factor  $[1 + (a/i \omega \gamma)]^{-1}$  in the equation for  $\delta T_n$ .

We see from the equations, that the effects of Newtonian cooling fall into two categories:

- 1. To the extent that  $L_n$ , as determined from equation (12), differs from  $Y_n$ , its adiabatic counterpart, the vertical structure of the various fields will differ.
- 2. Insofar as equations (14) to (17) in  $L_n$  are identical to the adiabatic equations in  $Y_n$ , while equation (18) differs by the factor  $[1 + (a/i\omega\gamma)]^{-1}$  from the adiabatic

equation for  $\delta T_n$ , the relative phases between the temperature and the velocity fields will be altered by Newtonian cooling.

The last matter may prove particularly important since the relative phase between  $\delta T$  and the velocity fields determines the mean tidal heat advection, and anything affecting this phase relation can result in mean tidal heating. The practical implications of this will be explored in subsequent papers. It should be noted, however, that the effect is simple to deal with since it enters primarily through a multiplicative factor in equation (18).

## 3. Asymptotic behavior

In most potential applications of the above theory, H is not constant, and the ratio  $|a/\omega|$  is of order unity. However, taking H a constant vastsy simplifies the problem. Moreover, the consideration of extreme values of  $|a/\omega|$  makes clear certain qualitative effects of this particular dissipative process.

In proceeding it proves simpler to transform (12) to the following

$$\frac{d^{2}\Phi_{n}}{dx^{2}} + \left(1 + \frac{a}{i\,\omega\,\gamma}\right)^{-1} \left\{ \frac{1}{h_{n}} \frac{dH}{dx} + \frac{\kappa\,H}{h_{n}} - \frac{1}{4} \right\} 
+ \frac{1}{2} \frac{a}{i\,\omega\,\gamma} \left[ \frac{1}{H} \frac{dH}{dx} - \frac{1}{2} + \frac{1}{H} \frac{dH}{dx} \frac{1}{a} \frac{da}{dx} + \frac{1}{H} \frac{d^{2}H}{dx^{2}} \right] 
- \frac{1}{H^{2}} \left( \frac{dH}{dx} \right)^{2} - \frac{a}{i\,\omega\,\gamma} \left( 1 + \frac{a}{i\,\omega\,\gamma} \right)^{-1} \frac{1}{H} \frac{dH}{dx} \left( \frac{1}{a} \frac{da}{dx} + \frac{1}{2} \frac{1}{H} \frac{dH}{dx} \right) \right\} \Phi_{n}$$

$$= \frac{\kappa\,J_{n}}{\gamma\,g\,h_{n} \left( 1 + \frac{a}{i\,\omega\,\gamma} \right)} e^{-x/2} \exp\left\{ -\frac{1}{2} \int_{0}^{x} \left[ \frac{\frac{a}{i\,\omega\,\gamma} \frac{1}{H} \frac{dH}{dx}}{1 + \frac{a}{i\,\omega\,\gamma}} \right] dx \right\},$$
(19)

where

$$L_n = \exp\left[\frac{x}{2} + \frac{1}{2} \int_0^x \left\{ \frac{\frac{a}{i\,\omega\,\gamma} \cdot \frac{1}{H} \cdot \frac{dH}{dx}}{1 + \frac{a}{i\,\omega\,\gamma}} \right\} dx \right] \Phi_n. \tag{20}$$

Assuming H = constant (i.e., an isothermal atmosphere) (19) becomes

$$\frac{d^2 \Phi_n}{dx^2} + \left(1 + \frac{a}{i \omega \gamma}\right)^{-1} \left\{ \frac{KH}{h_n} - \frac{1}{4} - \frac{1}{4} \frac{a}{i \omega \gamma} \right\} \Phi_n = \frac{\kappa J_n}{\gamma g h_n \left(1 + \frac{a}{i \omega \gamma}\right)} e^{-x/2}, \quad (21)$$

where

$$L_n = e^{x/2} \Phi_n . (22)$$

Consider now the case where  $|a(z)/\omega \gamma| \gg 1$  for all z, and where, moreover,  $|a/\omega \gamma|$  $\gg |\kappa H/h_n|$ . Equation (21) becomes

$$\frac{d^2\Phi_n}{dx^2} - \frac{1}{4}\Phi_n = \frac{i \overset{\circ}{\omega}}{\gamma g \overset{\circ}{h_n}} \kappa J_n e^{-x/2}. \tag{23}$$

Let  $J_n$  go to zero above some height. The solution of (23) above that height will be of the form

$$\Phi_n = A e^{-x/2} + B e^{x/2}. \tag{24}$$

The condition of boundedness requires, moreover, that B=0. Thus, using (22), we have that  $L_n$  and consequently  $u_n$ ,  $v_n$ ,  $w_n$ , and  $\delta T_n$ , are independent of height. This contrasts with the adiabatic case where  $Y_n$  would increase with height if  $h_n > 0$  and decrease with height if  $h_n < 0$ . Both these cases lead to constancy with height for sufficiently rapid Newtonian cooling.

Another pedagogically interesting case arises when not only H, but also a is a constant, and where, furthermore  $|a/\omega \gamma| \le 1$ . (21) becomes

$$\frac{d^2\Phi_n}{dx^2} + \left(\frac{KH}{h_n} - \frac{1}{4} - \frac{a}{i\ \omega\ \gamma} \frac{KH}{h_n}\right)\Phi_n = \frac{KJ_n}{\gamma\ g\ h_n} \left(1 - \frac{a}{i\ \omega\ \gamma}\right)e^{-x/2}. \tag{25}$$

Again, let  $J_n$  go to zero above some height. The solution of (25) above that height will be of the form

$$\Phi_n = A e^{\lambda_n X} + B e^{-\lambda_n X}, \qquad (26)$$

where

$$\lambda_n = \sqrt{-\frac{KH}{h_n} + \frac{1}{4} + \frac{a}{i \omega \gamma} \frac{KH}{h_n}}$$

or

$$\lambda_n \simeq \sqrt{-\frac{KH}{h_n} + \frac{1}{4}} \left( 1 - \frac{1}{2} - \frac{\frac{a}{i \omega \gamma} - \frac{KH}{h_n}}{\frac{KH}{h_n} - \frac{1}{4}} \right).$$
 (27)

Let us now restrict ourselves to the case where  $KH/h_n > 1/4$ . Then (27) becomes

$$\lambda_n \simeq i \sqrt{\frac{KH}{h_n} - \frac{1}{4}} \left( 1 - \frac{1}{2} \frac{\frac{a}{i \omega \gamma} \frac{KH}{h_n}}{\frac{KH}{h_n} - \frac{1}{4}} \right). \tag{28}$$

Boundeness, as  $x \to \infty$ , now leads to B = 0 in (26) which is, in fact, the traditional radiation condition (WILKES [10]).

#### 4. Concluding remarks

We have, in this paper, rederived the equations of tidal theory including the effects of Newtonian cooling. It was shown that the inclusion of Newtonian cooling affects 1. the altitude distribution of the amplitude and phase of all tidal fields, and 2. the relative phase between the tidal temperature and velocity fields. The latter may result in net tidal heating.

Certain asymptotic effects of Newtonian cooling are also investigated. Very rapid Newtonian cooling leads to a constancy of wave amplitude with height in cases both where the adiabatic wave amplitude would increase with altitude and where it would decrease with altitude. Very small constant Newtonian cooling leads to the traditional radiation condition. Concerning the last matter, a word of caution is necessary. Consider the following situation: we have a thermal drive concentrated entirely below

some height  $z_1$ ; Newtonian cooling is negligible below some height  $z_2 > z_1$ ; above  $z_2$  Newtonian cooling increases to an appreciable value, but eventually decreases to a small constant value above some height  $z_3$ . According to the analysis of section 3, the radiation condition would apply above  $z_3$ . However, there is no reason to assume it holds for the region above the drive region between  $z_1$  and  $z_3$ . That is to say, the existence of a dissipation mechanism at high levels does not imply that the radiation condition is appropriate directly above some lower level drive.

As a final note, one must point out that in practical situations processes other than those considered here are certainly operative, and the application of the above equations is limited by these omissions.

## 5. Acknowledgement

One of the authors, D. J. McKenzie, wishes to thank NCAR for its hospitality during Summer 1966.

#### References

- [1] M. SIEBERT, Atmospheric tides, Advances in Geophysics (1961), 105.
- [2] R. S. LINDZEN, Thermally driven diurnal tide in the atmosphere, Submitted to Quarterly Jour. Roy. Met. Soc. (1966).
- [3] C. L. Pekeris and Z. Alterman, A method of solving nonlinear equations of atmospheric tides with applications to an atmosphere of constant temperature (The atmosphere and the sea in motion, Rockefeller Institute Press, New York 1959).
- [4] M. YANOWITCH, The effect of viscosity on gravity waves and the upper boundary condition, NCAR manuscript No. 147 (1966).
- [5] S. T. BUTLER and K. A. SMALL, The excitation of atmospheric oscillations, Proc. Roy. Soc. [A], 274 (1963), 91.
- [6] R. S. LINDZEN and R. M. GOODY, Radiative and photochemical processes in mesospheric dynamics. Part I, Models for radiative and photochemical processes, Jour. Atmos. Sci. 22 (1965), 341.
- [7] C. Prabhakara and J. S. Hogan, Ozone and carbon dioxide heating in the martian atmosphere, Jour. Atmos. Sci. 22 (1965), 97.
- [8] C. LEOVY and Y. H. MINTZ, A numerical general circulation experiment for the atmosphere of mars, RAND report (1966), available from RAND Corp., Santa Monica, California, USA.
- [9] R. S. Lindzen, On the asymmetric diurnal tide, Pageoph 62/III (1965), 142.
- [10] M. V. Wilkes, Oscillations of the earth's atmosphere (Cambridge University Press, 1949).

(Received 30th August 1966)