

Global mode water detection and its representation in heat transport

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April 6th, 2022

Ocean heat uptake

- Ocean serves as a **heat reservoir** of the Earth system, accounting for over **90%** of the total warming that has occurred since 1955.
- Mode water plays a major role in **ventilating thermoclines** and **modulating SST signals**.
- In our study, a new algorithm is developed to determine the **mixed layer depth (MLD) and mode water (MW) thickness**, which is applied to the Argo array.
- By co-locating **mesoscale eddies** derived from satellite altimetry maps (Laxenaire et al., 2019) and Argo profiles, we also assess the role of eddies in mode water transport and subduction.

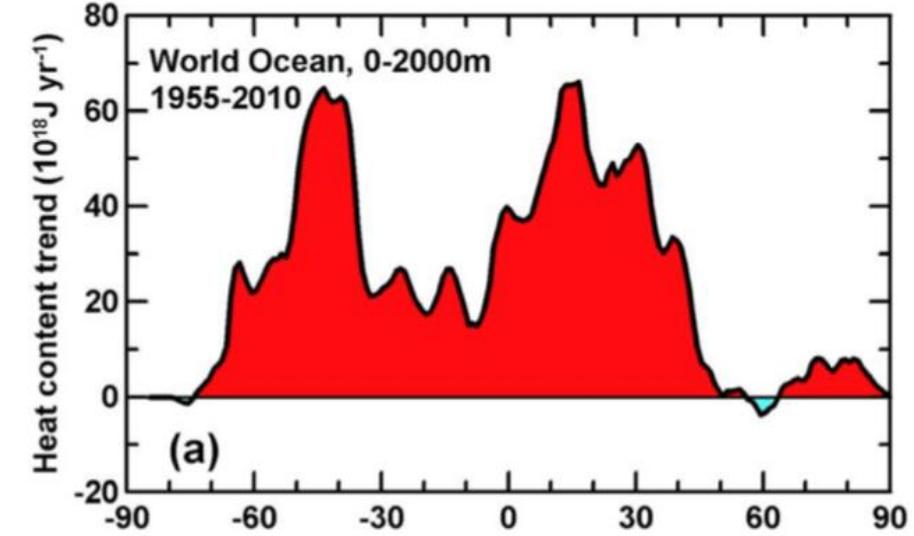
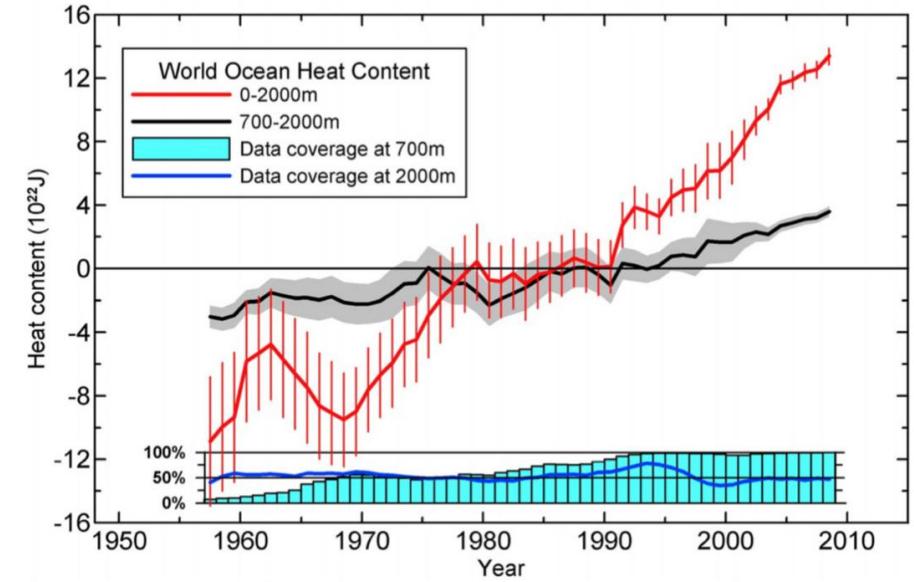


Figure: Ocean heat content. (Levitus et al., 2012)

The MLD and MW algorithm

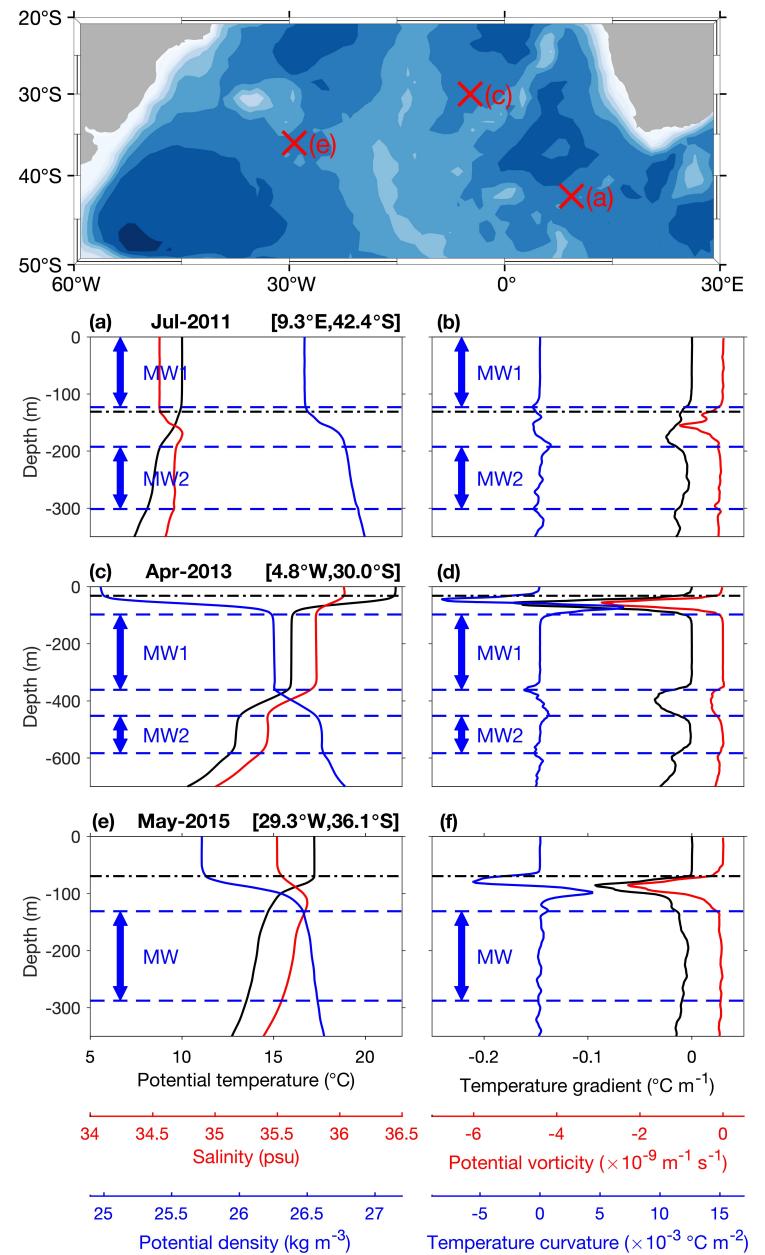
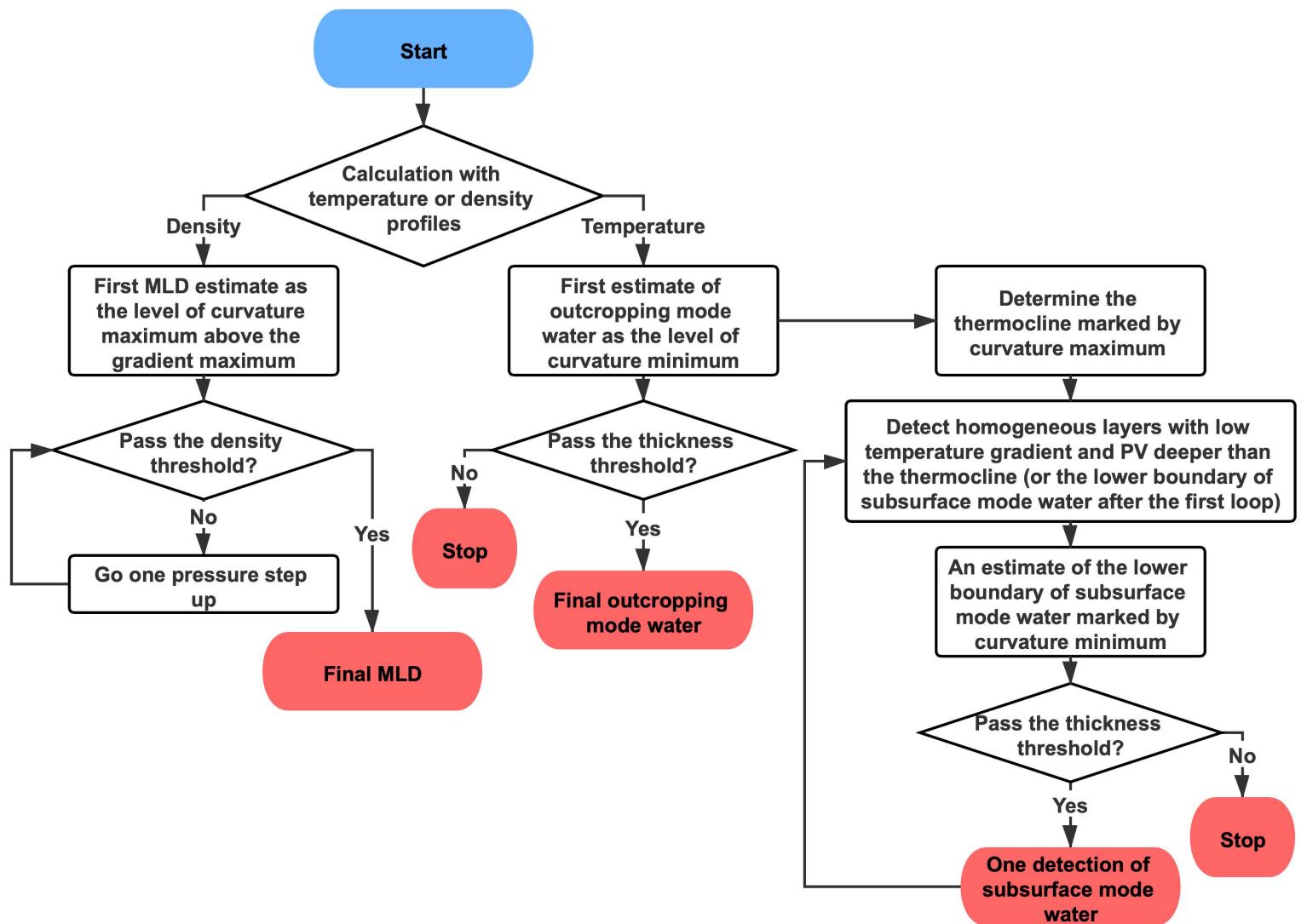


Figure: Three profile examples in the South Atlantic. (Chen et al., 2021)

Global MLD distribution

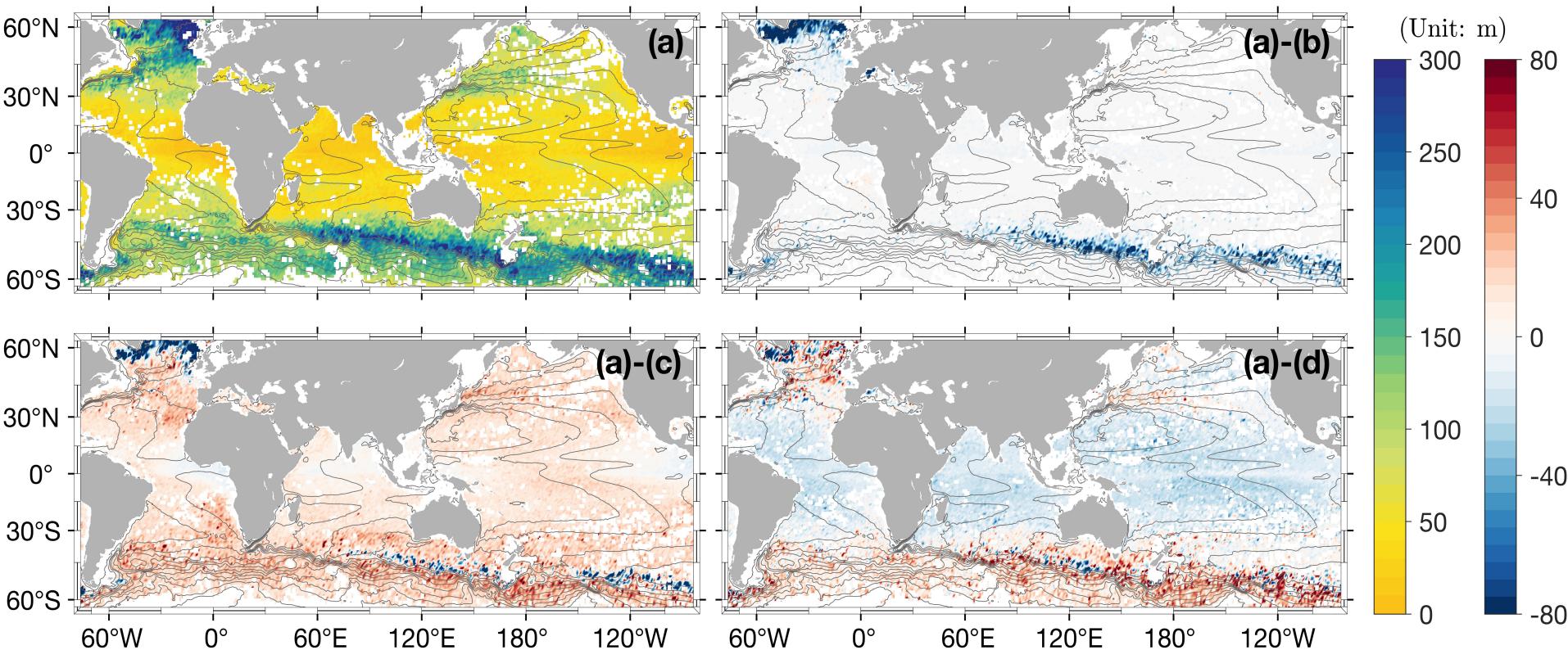


Figure: Winter-mean MLD.

Four methods of MLD detection are:

- The new algorithm
- Density threshold (0.03 kg m^{-3})
- Density gradient (0.0005 kg m^{-4})
- Hybrid method (Holte and Talley, 2009)

Mode Water T-S relation

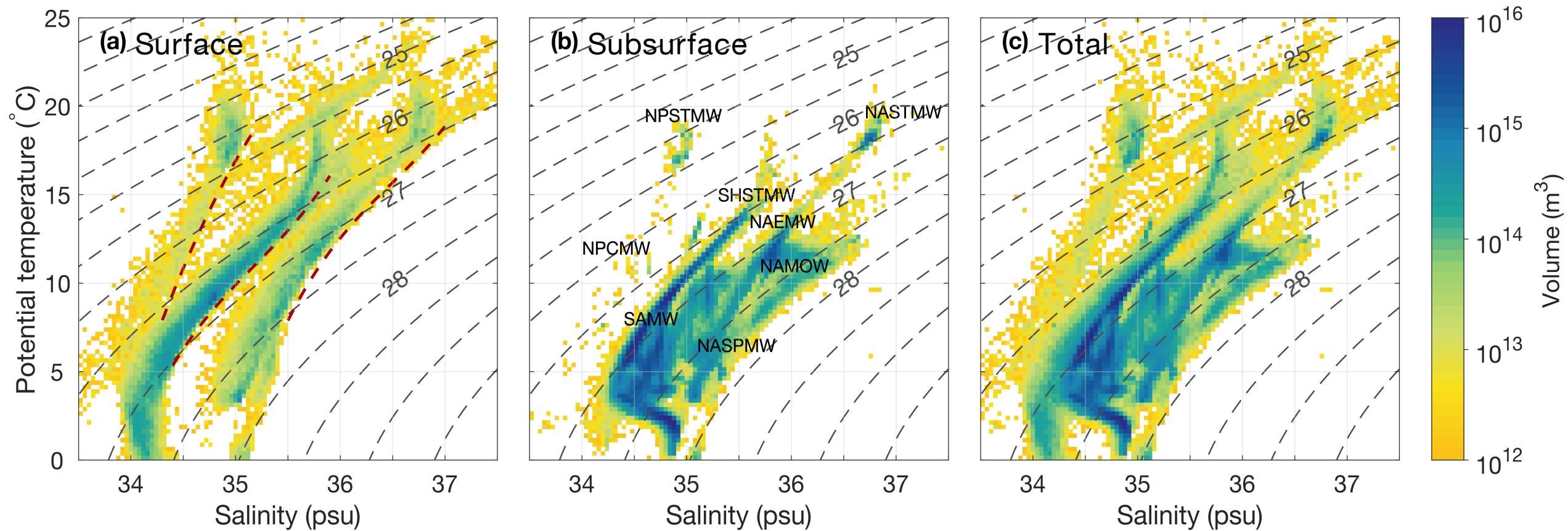
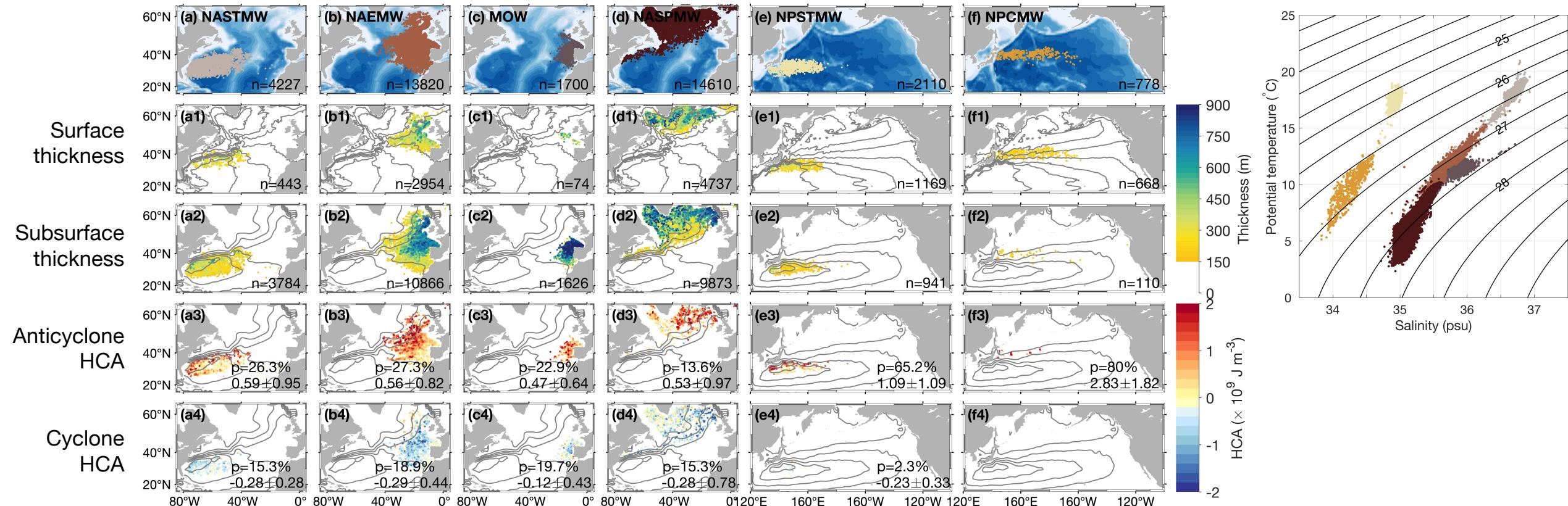


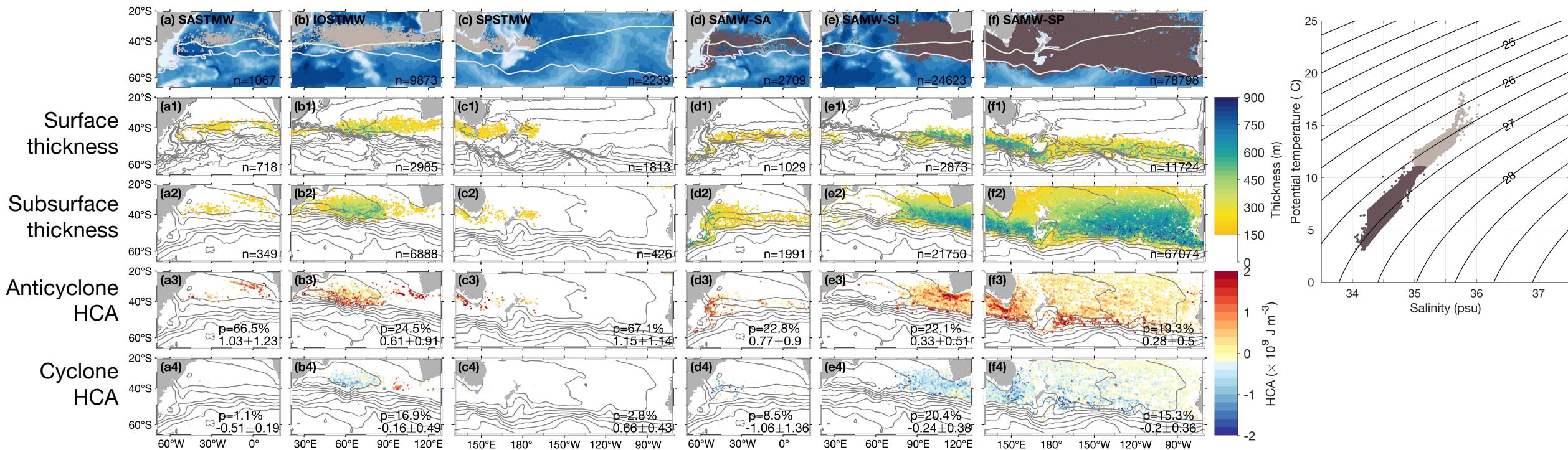
Figure: T-S diagram of surface and subsurface mode waters detected by the new algorithm.

Mode Water in the Northern Hemisphere



- 4 MW types in the North Atlantic, and 2 types in the North Pacific.
- Surface MWs for each group are retrieved from the entire pool with the same properties as subsurface MWs.
- The positive HCA inside MWs is related to co-location with anticyclones.

Mode Water in the Southern Hemisphere



- Cluster analysis is applied to divide all subsurface MWs into 2 types.
- STMWs originate at the northern periphery of STF, and SAMWs are formed insize the SAZ.
- Interbasin heat transport is associated with anticyclonic eddies.

Prospectives

- Evidence shows 60%-90% of excess heat is absorbed by the **Southern Ocean**, which draws attention to the ability of SAMWs in taking up heat.
- The detection of **subsurface eddies** (that are not detectable from satellite) needs to improve.
- By comparing the **trajectories** of anticyclonic eddies and the depths of these mode waters, we can assess the **ventilation process** associated with eddies.

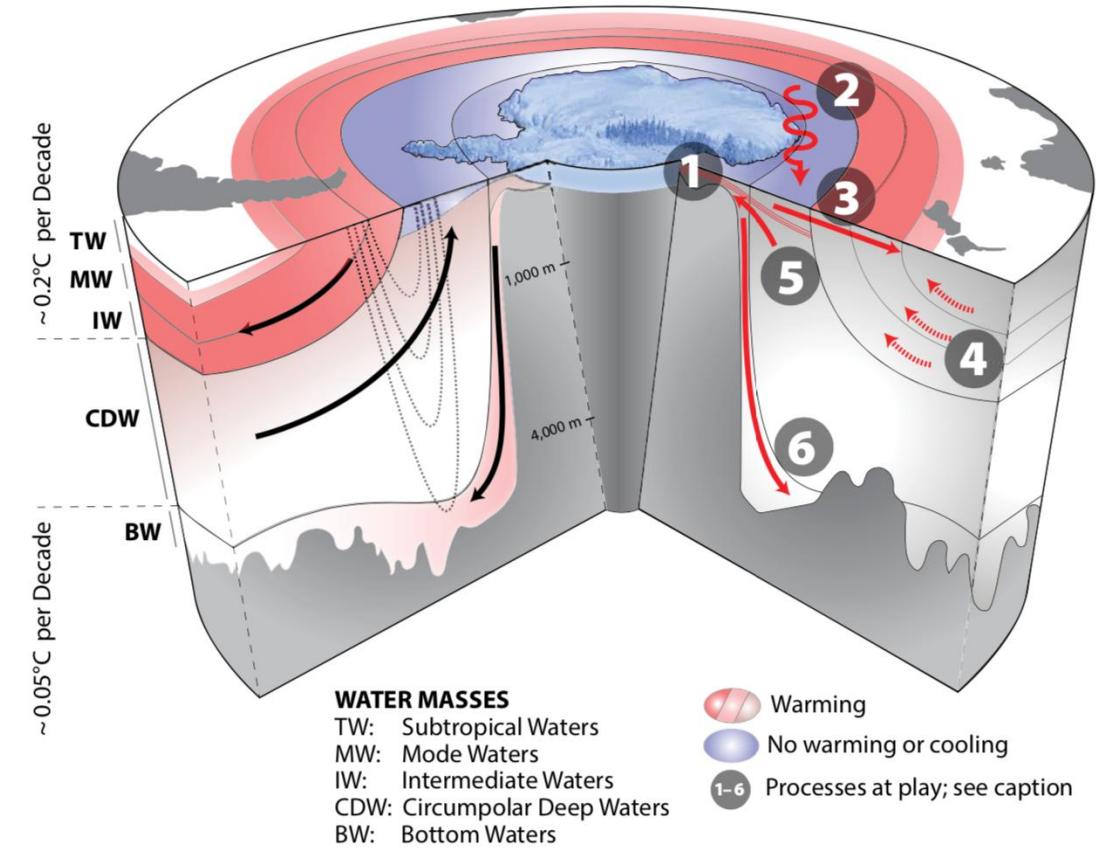
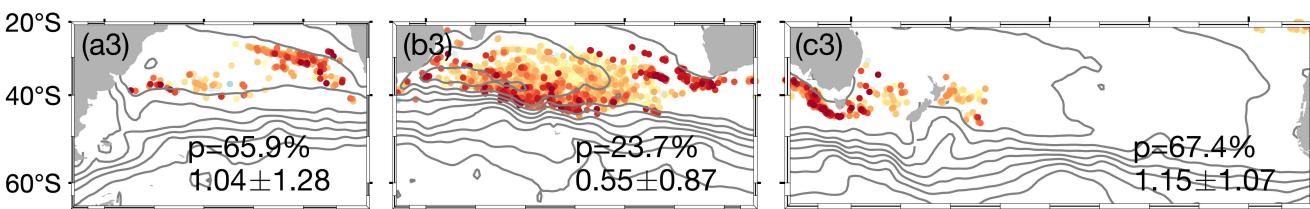


Figure: Temperature trends in different layers of the Southern Ocean. (Sallée, 2018)

Subduction rate

$$S_{ml}(t) = \frac{\partial H}{\partial t} + \nabla \cdot \left(\int_0^H \mathbf{u} dz \right), \quad (1)$$

$$S_{ml}(t) = \frac{\partial H}{\partial t} + \nabla \cdot \left(\int_0^{H(t)} \mathbf{u} dz \right) + \nabla \cdot \left(\int_{H(t)}^{H_{max}} \mathbf{u} dz \right). \quad (2)$$

$$S_{ml}(t) = \frac{\partial H}{\partial t} + \nabla \cdot (\mathbf{U}_{ek}) + \nabla \cdot [\mathbf{u}_{ml}(t)H(t)] + \nabla \cdot [\mathbf{u}_{sth}(t)(H_{max} - H(t))]. \quad (3)$$

$$\overline{S_{ml}} = \nabla \cdot (\overline{\mathbf{U}_{ek}}) + \nabla \cdot (\overline{\mathbf{u}_{ml}} \overline{H}) + \nabla \cdot [\overline{\mathbf{u}_{sth}} (\overline{H_{max}} - \overline{H})] + \nabla \cdot (\overline{\mathbf{u}_{ml}^*} \overline{H}) + \nabla \cdot [\overline{\mathbf{u}_{sth}^*} (\overline{H_{max}} - \overline{H})]. \quad (4)$$

$$\overline{S_{ml}} = \overline{S_{ek}} + \overline{S_{geo}} + \overline{S_{eddy}}, \quad (5)$$

$$\overline{S_{ek}} = \overline{w_{ek}},$$

$$\overline{S_{geo}} = \overline{\mathbf{u}_{ml}} \cdot \nabla \overline{H} + \overline{\mathbf{u}_{sth}} \cdot \nabla (\overline{H_{max}} - \overline{H}) + \beta \text{ correction},$$

$$\overline{S_{eddy}} = \nabla \cdot (\overline{\mathbf{u}_{ml}^*} \overline{H}) + \nabla \cdot [\overline{\mathbf{u}_{sth}^*} (\overline{H_{max}} - \overline{H})]. \quad (6)$$

$$\overline{S_{eddy}} = \nabla \cdot (\kappa \bar{s})_{z=H_{max}}. \quad (8)$$

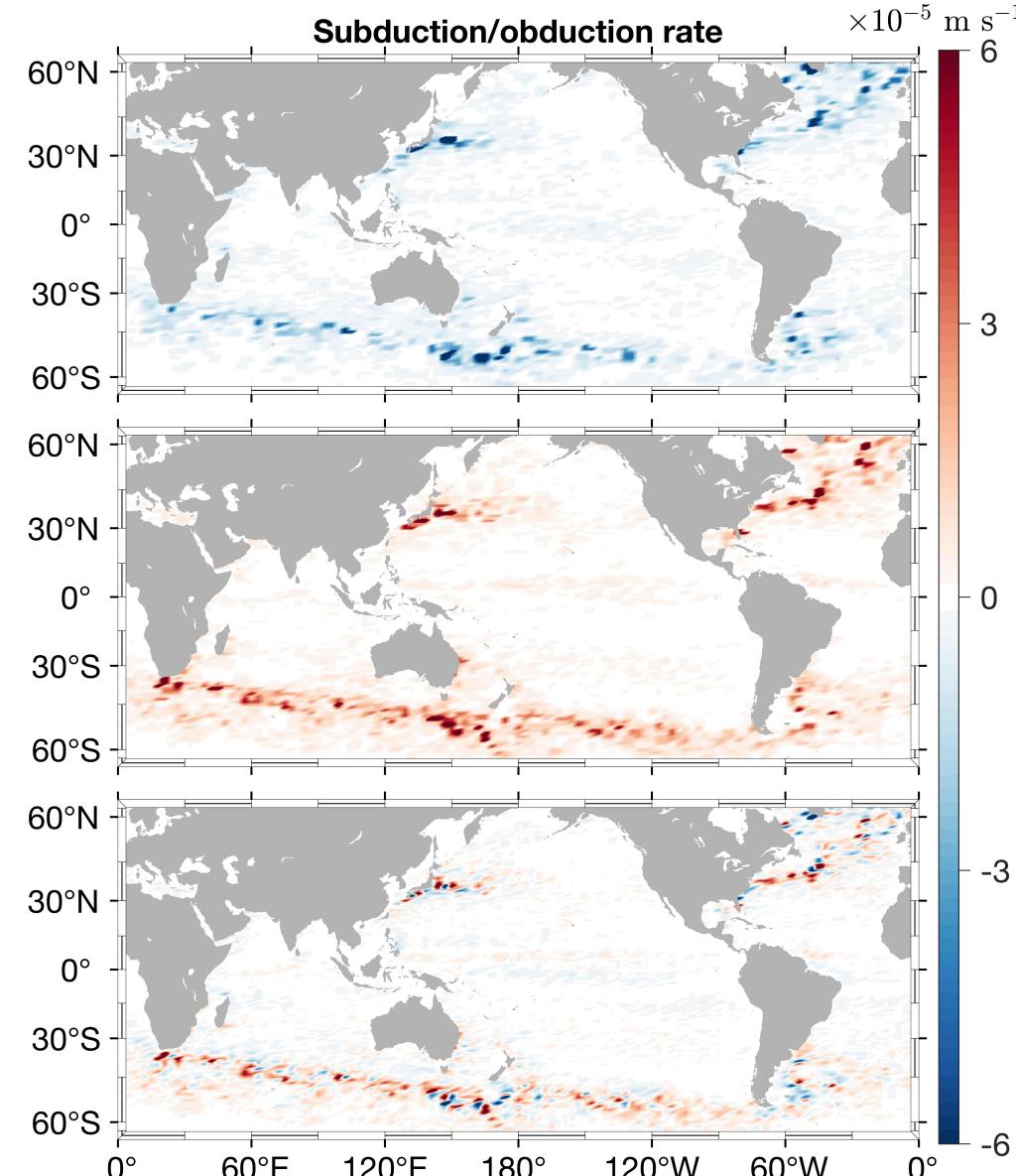


Figure: The geostrophic component of subduction and obduction rates based on the new MLDs.

Thanks for your attention.