

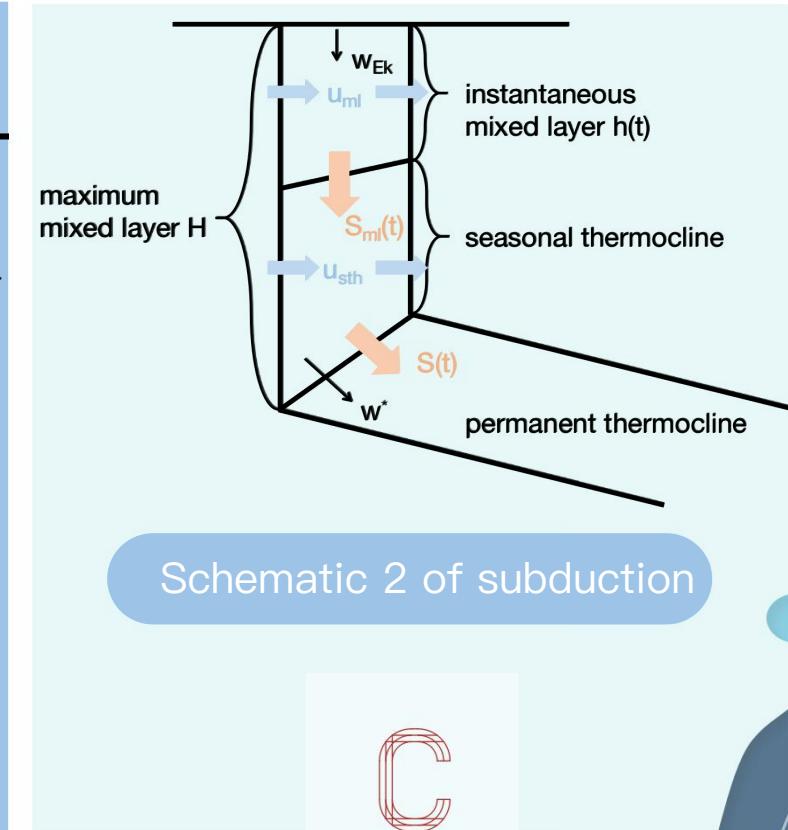
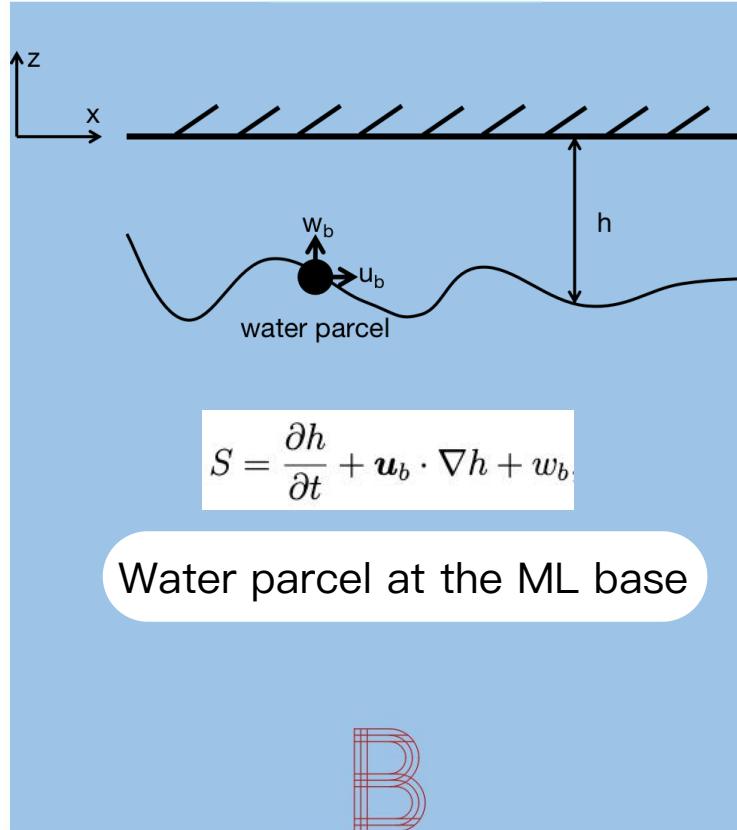
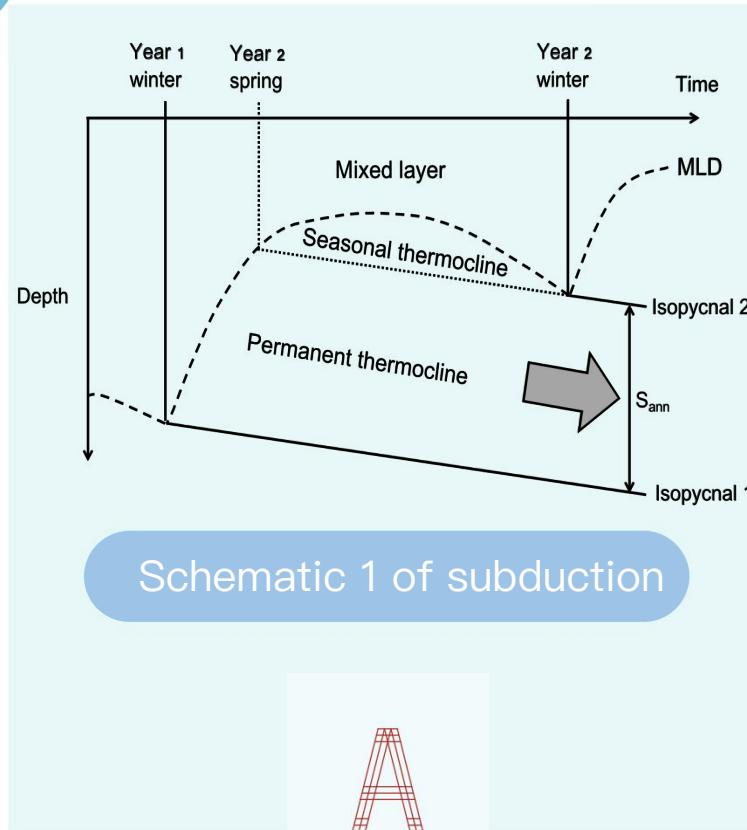
Water mass subduction in the isopycnic coordinate

Yanxu Chen and Sabrina Speich
LMD–ENS

Contents

- 1 Classic theories of subduction
- 2 Subduction estimated at the migrating isopycnal
- 3 “Eddy” component of subduction
- 4 Conclusions and future work

Classic theories of subduction



Subduction estimated at the migrating isopycnal

$$S = \frac{\partial h}{\partial t} + \mathbf{u}_b \cdot \nabla h + w_b \quad (1)$$

$$\frac{\partial h}{\partial t} \Big|_{\sigma} \Delta t = h(r_0 + \Delta r, t_0 + \Delta t) - h(r_0, t_0) = \frac{\partial h}{\partial t} \Big|_r \Delta t + (\mathbf{c} \Delta t) \cdot \nabla h \quad (2)$$

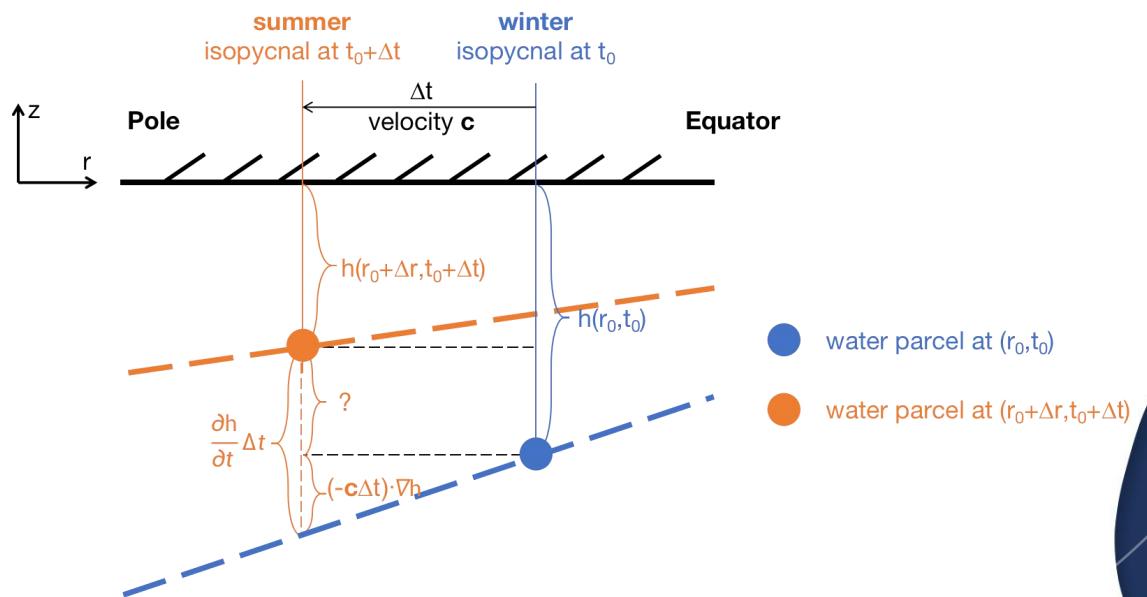
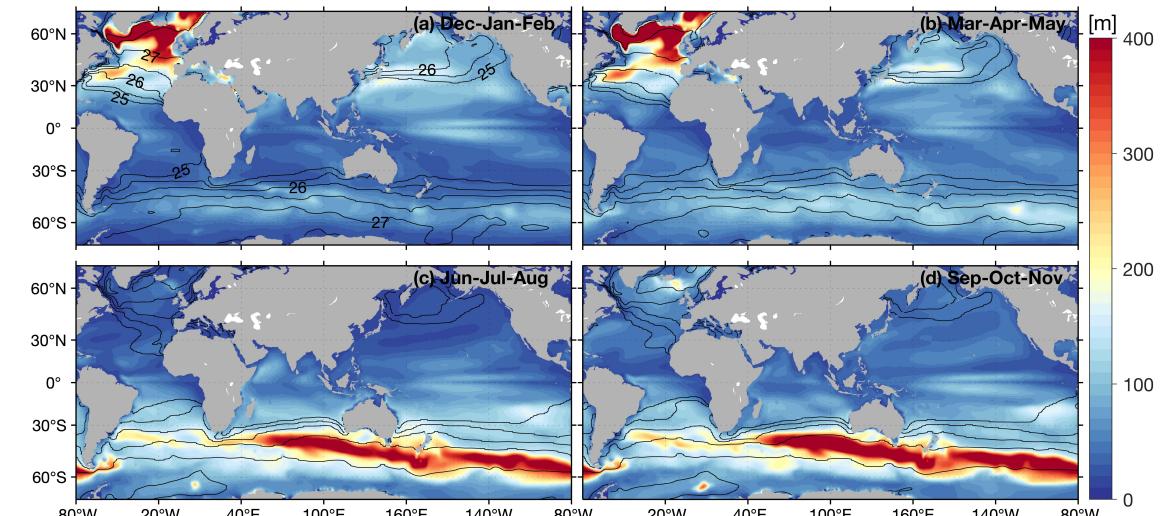
$$S = \frac{\partial h}{\partial t} \Big|_{\sigma} + \mathbf{u}_b \Big|_{\sigma} \cdot \nabla h + w_b = \frac{\partial h}{\partial t} \Big|_r + \mathbf{c} \cdot \nabla h + (\mathbf{u}_b - \mathbf{c}) \cdot \nabla h + w_b \quad (3)$$

$$S_t = \frac{\partial h}{\partial t} \Big|_{\sigma}, \quad \text{temporal term}$$

$$S_h = S_{h1} + S_{h2} = -\mathbf{c} \cdot \nabla h + \mathbf{u}_b \cdot \nabla h, \quad \text{lateral induction}$$

$$S_v = w_b. \quad \text{migration of isopycnal}$$

$$\frac{\partial h}{\partial t} \Delta t \quad \text{vertical velocity}$$



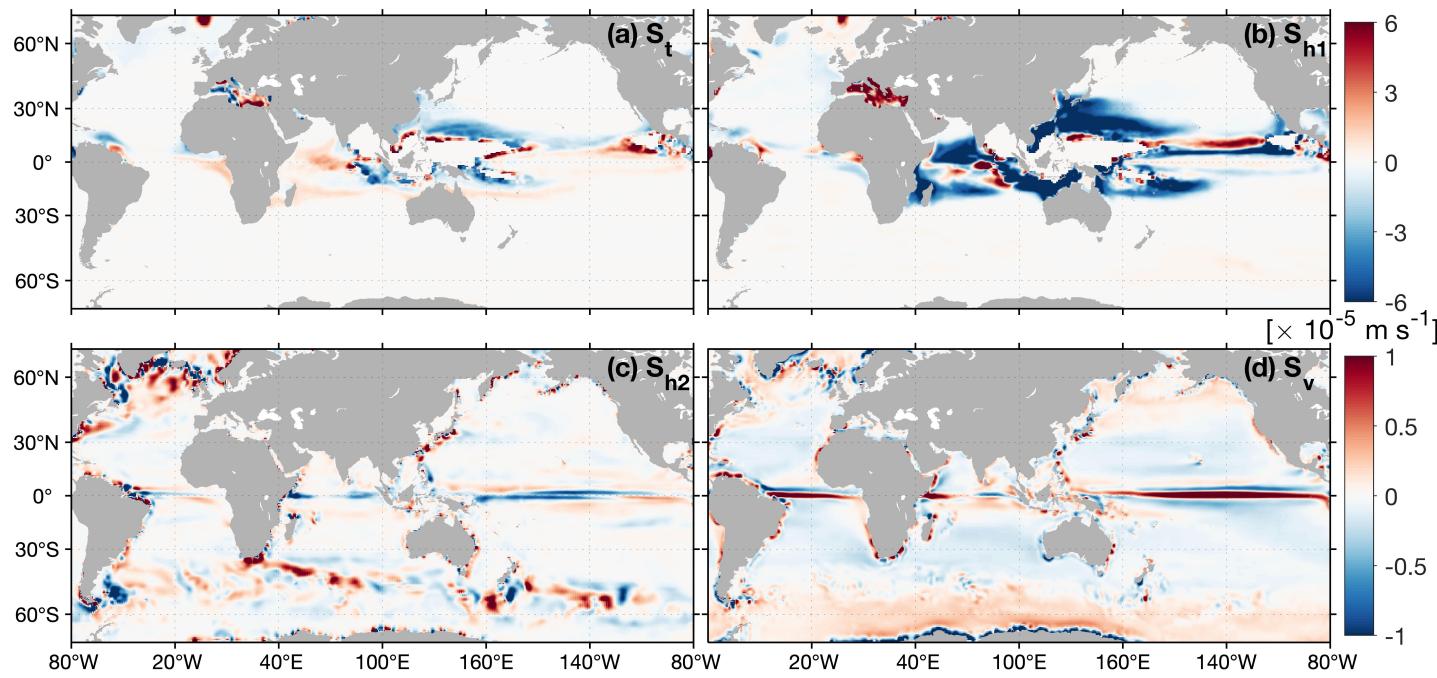
$$S_t = \frac{\partial h}{\partial t} \Big|_{\sigma}, \quad \text{temporal term}$$

$$S_h = S_{h1} + S_{h2} = -\mathbf{c} \cdot \nabla h + \mathbf{u}_b \cdot \nabla h, \quad \text{lateral induction}$$

$$S_v = w_b. \quad \text{migration of isopycnal}$$

↓

vertical velocity

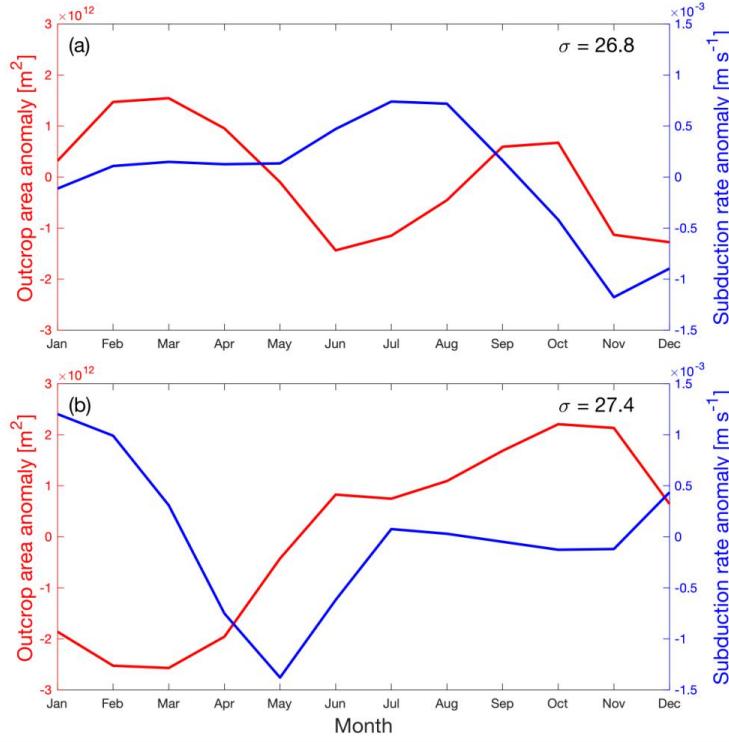


Subduction estimated at the migrating isopycnal

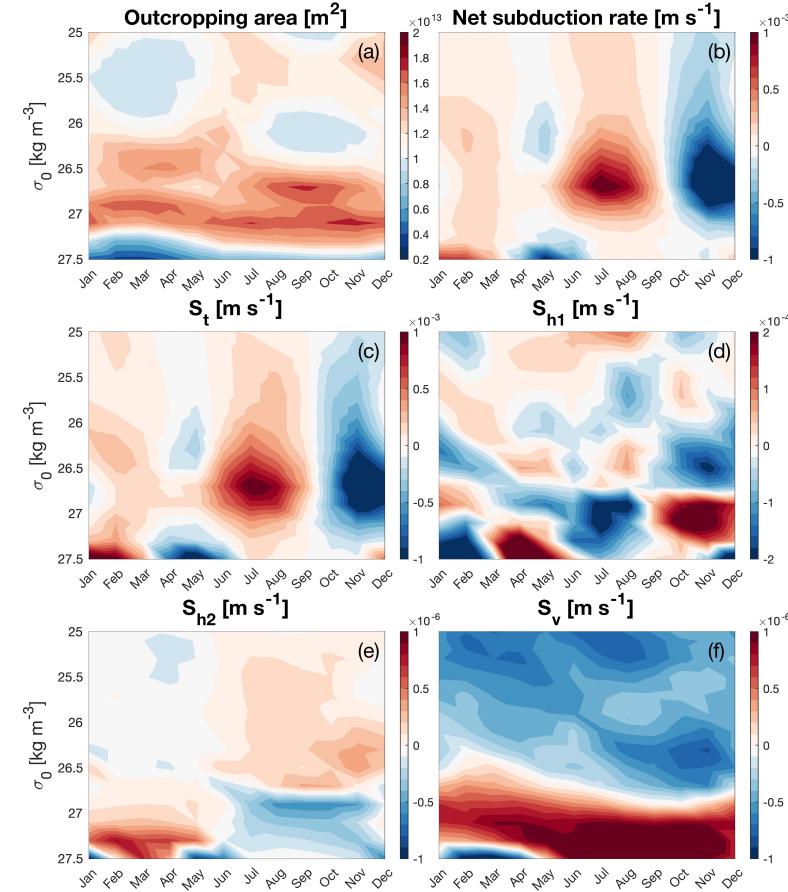
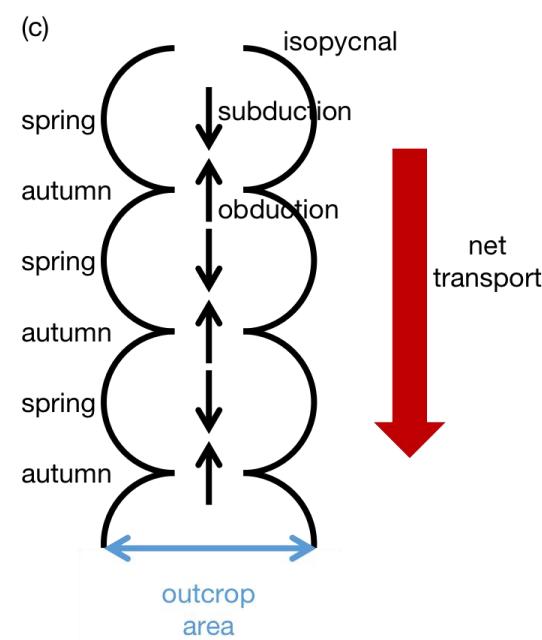
- 1) Large scale is dominated by the vertical velocity at the ML base, i.e., Ekman pumping.
- 2) Spatial patterns along the ACC and in the polar North Atlantic are controlled by lateral induction.
- 3) Migration of isopycnals matters in the tropical and subtropical regions.
- 4) The temporal term does not vanish to zero as assumed in the theory of Stommel's demon.

“Eddy” component of subduction

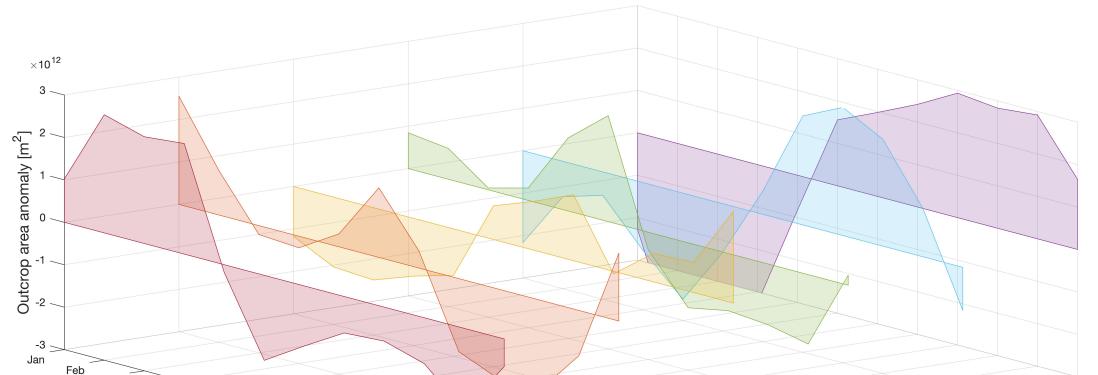
Mean + Eddy: $\bar{u} + u' \longrightarrow \overline{u' \cdot u'}$



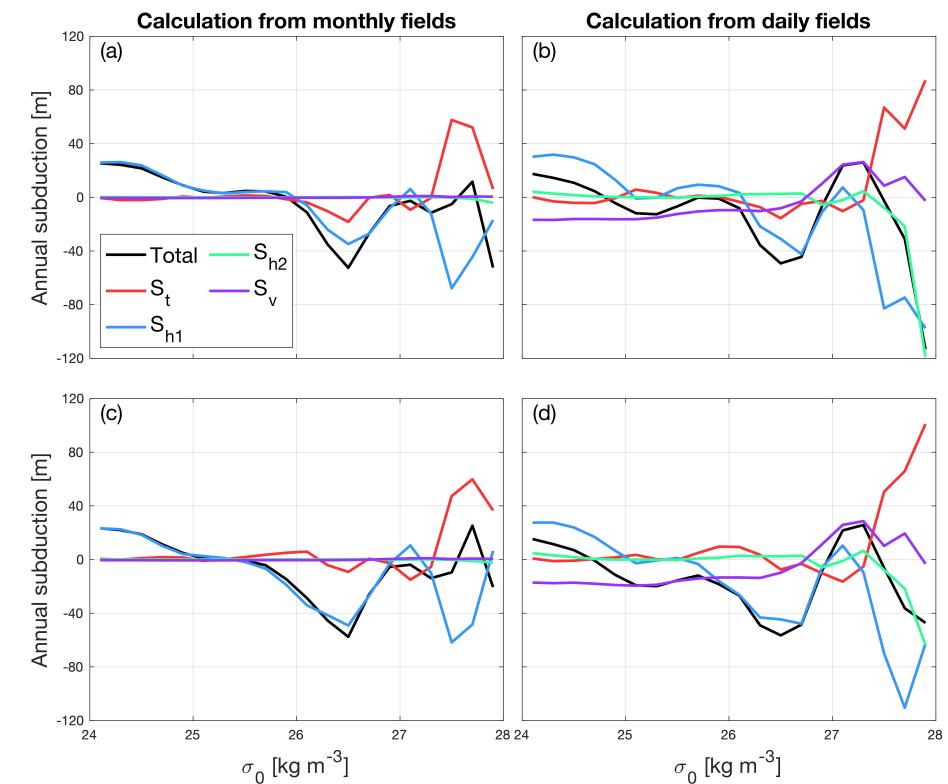
$$\overline{S} = \frac{1}{T} \int_0^T \left(\frac{1}{A} \int_0^A S dA \right) dt$$



“Eddy” component of subduction



Acronyms	Full name	Density range (ECCO)	Density range (Argo)
NPSTMW	North Pacific Subtropical Mode Water	$25.2 \leq \sigma < 26.4$	$25.1 \leq \sigma < 25.5$
NASTMW	North Atlantic Subtropical Mode Water	$25.2 \leq \sigma < 26.4$	$26.4 \leq \sigma < 26.6$
SHSTMW	Southern Hemisphere Subtropical Mode Water	$25.2 \leq \sigma < 26.4$	$26.3 \leq \sigma < 26.8$
SAMW	Subantarctic Mode Water	$26.4 \leq \sigma < 27.1$	$26.8 \leq \sigma < 27.2$
AAIW	Antarctic Intermediate Water	$27.1 \leq \sigma < 27.6$	$26.8 \leq \sigma < 27.4$



1. The outcropping area is dependent on time and density.
2. Nonlinearity leads to modifications of subduction rates at different isopycnals.

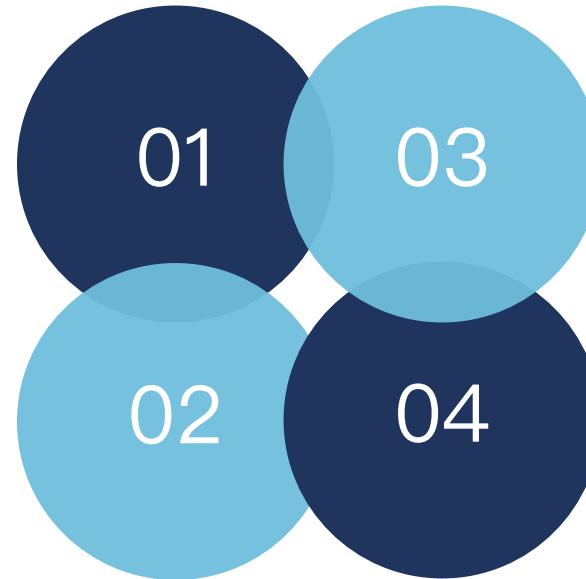
Conclusions and future work

Conclusion 1

It is necessary to switch to the isopycnic coordinate.

Conclusion 2

Subduction rates are greatly modified by the eddy component.

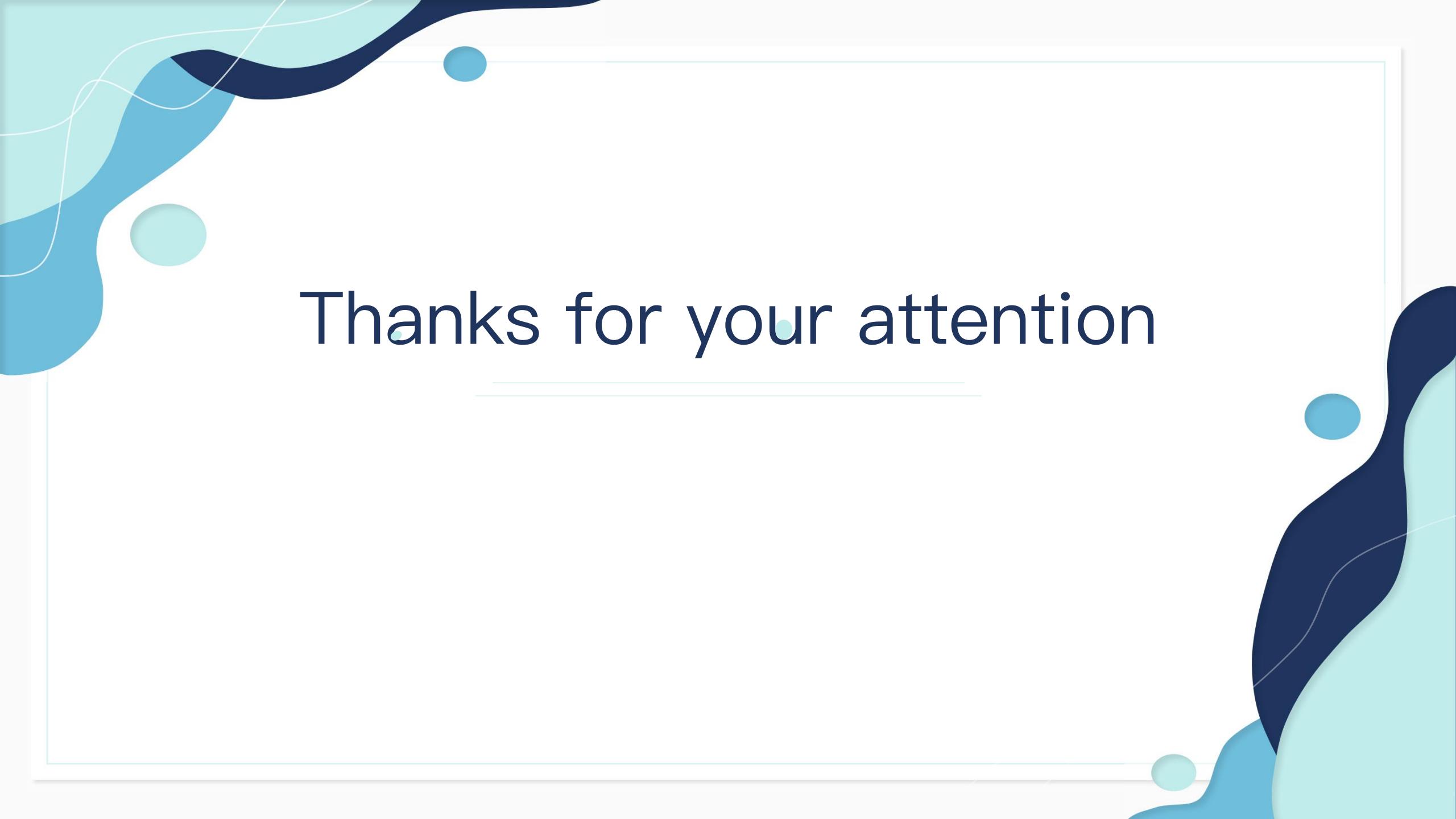


Prospective 1

It is of interest to consider PV fluxes across the ML base.

Prospective 2

Another interesting calculation will be to connect water mass subduction to transformation at the surface.



Thanks for your attention
