



The SPAR Model: A New Paradigm for Multivariate Extremes: Application to Joint Distributions of Metocean Variables

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This paper presents the application of a new multivariate extreme value model for the estimation of metocean variables. The model requires fewer assumptions about the forms of the marginal distributions and dependence structure compared to existing approaches, and provides a flexible and rigorous framework for modeling multivariate extremes. The method involves a transformation of variables to polar coordinates. The tail of the radial variable is then modeled using the generalized Pareto distribution, with parameters conditional on angle, providing a natural extension of univariate theory to multivariate problems. The resulting model is referred to as the semi-parametric angular-radial (SPAR) model. We consider the estimation of the joint distributions of (1) wave height and wave period, and (2) wave height and wind speed. We show that the SPAR model provides a good fit to the observations in terms of both the marginal distributions and dependence structures. The use of the SPAR model for estimating long-term extreme responses of offshore structures is discussed, using some simple response functions for floating structures and an offshore wind turbine with monopile foundation. We show that the SPAR model is able to accurately reproduce response distributions, and provides a realistic quantification of uncertainty. [DOI: 10.1115/1.4065968]

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1 Introduction

Many design problems in offshore engineering require estimates of the joint extremes of metocean variables, such as winds, waves, currents, and water levels. Joint extremes of metocean variables are often quantified in terms of environmental contours [1]. Some types of environmental contours can be estimated without knowing the joint density of the variables [2,3]. However, in other applications, such as full long-term extreme response analysis [4], a model is needed for the joint density. Estimating the joint density in extreme regions of the variable space is subject to large uncertainties [5,6]. A wide range of methods have been proposed for estimating the joint density function of metocean variables; a relatively recent review is presented in Ref. [7]. At present, the most commonly used approaches are global hierarchical models and copula models. These models can be applied in any number of dimensions, but for simplicity we shall restrict the discussion to two-dimensional cases. In the global hierarchical model, the joint density function

$f_{X,Y}$ of variables X, Y is written in conditional form as

$$f_{X,Y}(x, y) = f_X(x)f_{Y|X}(y|x) \quad (1)$$

where $f_X(x)$ is the marginal density of X and $f_{Y|X}(y|x)$ is the density of Y conditional on $X = x$. Usually, a parametric form is assumed for both f_X and $f_{Y|X}$, with popular choices being Weibull and log-normal models [8,9]. The parameters of the conditional distribution are then modeled as a function of x . There are several drawbacks to this approach. Firstly, assuming a model for the bulk of the observations does not guarantee a good fit to the tail of the distribution, which is the region of interest when estimating extreme responses. Secondly, the model for the conditional dependence structure is not based on any physical or mathematical principles, and hence provides no rationale for extrapolating outside the range of observations. These limitations mean that fitted models obtained from this approach are often in poor agreement with observations [6].

In copula-based approaches, the joint density function is expressed as

$$f_{X,Y}(x, y) = f_X(x)f_Y(y)c(F_X(x), F_Y(y)) \quad (2)$$

where f_X, f_Y, F_X , and F_Y are the marginal density and distribution functions, and c is the copula density function [10,11]. As well as

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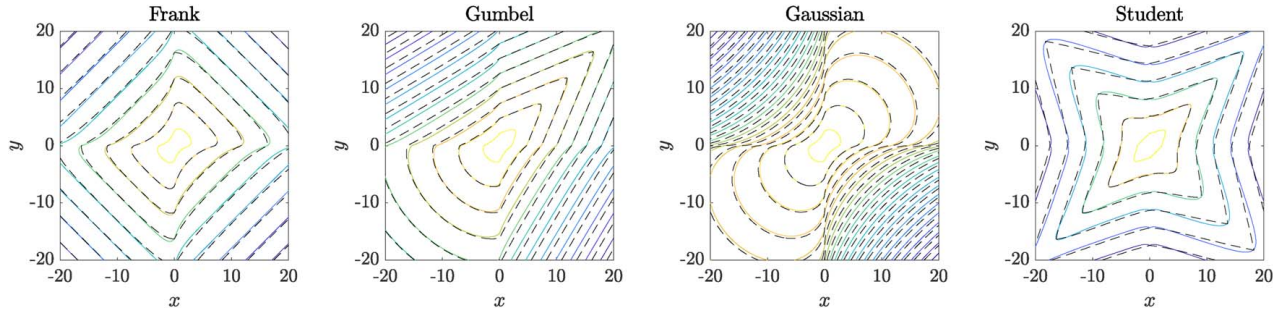


Fig. 1 Contour plots of joint densities of various copulas on Laplace marginal distributions. All copulas have Pearson correlation coefficient $\rho = 0.6$. Student- t copula has $\nu = 2$ degrees-of-freedom. Solid lines: true isodensity contours at logarithmic increments. Dashed lines: isodensity contours of SPAR approximations at the same density values.

assuming a form for the marginal distributions, a parametric form for the copula is also assumed. Common choices for the copula include Frank, Gumbel, Gaussian, and Student- t [12–14]. Examples of the joint density functions for these four copulas are shown in Fig. 1, with standard Laplace marginal distributions, i.e., $f_X(z) = f_Y(z) = (1/2)\exp(-|z|)$, $z \in \mathbb{R}$. In each case, the copula parameters have been selected so that the Pearson correlation coefficient is $\rho = 0.6$. It is evident that different choices of copula can lead to large differences in joint tail behavior. Moreover, usually there is no a priori reason to suppose that the dependence structure in observations follows a particular parametric form.

Various methods based on multivariate extreme value theory have been proposed; see Refs. [15,16] for an overview of the applications in an oceanographic context. Possibly the most popular choice for metocean variables is the conditional extremes model [17], which can be used to estimate the joint distribution of variables conditional on at least one variable being large. Example metocean applications include Refs. [18,19]. The key limitation of this approach is that it only characterizes the region of variable space where the conditioning variable is large, and inferences made using different conditioning variables are not necessarily consistent [20]. A further limitation of this method (and other methods in the multivariate extremes literature) is that it requires a transformation of the marginal distributions to a standard scale. This requires first estimating the marginal distributions for each variable—a process which is subject to uncertainty.

In this paper, we discuss the application of a new method, which overcomes the limitations of existing approaches and provides a general framework for modeling multivariate extremes. The model is referred to as the semi-parametric angular-radial (SPAR) model. The SPAR model provides a framework for estimating multivariate extremes that does not require any assumptions about the form of the marginal distributions or dependence structure, and provides a rational means for extrapolating outside the range of observations. Moreover, the model is only fitted to extreme observations, meaning that no assumptions are required about the bulk of the distribution. The SPAR model can either be applied on transformed marginal distributions or on the original marginal scale, removing the need to fit marginal models. Theoretical aspects of the SPAR model have been presented recently in Ref. [21], and a detailed discussion of SPAR model inference is presented in Ref. [22]. The purpose of this paper is to provide a more engineering-focused description of the SPAR model and demonstrate its application for estimating joint extremes of metocean variables and extreme responses of offshore structures.

The paper is organized as follows. A brief overview of the theory is presented in Sec. 2, and inference is discussed in Sec. 3. In Sec. 4, we discuss the application of the SPAR model to the estimation of the joint extremes of (1) wave height and wave period, and (2) wave height and wind speed. In Sec. 5, we consider the long-term extreme response distribution for some simple response functions, using the fitted SPAR models. We conclude with a discussion in Sec. 6.

2 Theory

2.1 Model Definition. The SPAR model is an extension of the univariate peaks-over-threshold (POT) method to the multivariate setting. The method involves a transformation of the variables to polar coordinates. In the present work, we will restrict our attention to standard polar coordinates in two dimensions, but more general polar coordinate systems can also be used (see Ref. [21] for a discussion). Define radial and angular variables as $R = (X^2 + Y^2)^{1/2}$ and $\Theta = \text{atan2}(X, Y)$, where atan2 is the four-quadrant inverse tangent function. Suppose that random vector (X, Y) has continuous joint density function $f_{X,Y}$, with simply connected support, containing the point $(0, 0)$. Then the joint density of (R, Θ) is given by

$$f_{R,\Theta}(r, \theta) = r \cdot f_{X,Y}(r \cos(\theta), r \sin(\theta)) \quad (3)$$

In a similar manner to global hierarchical models (1), the angular-radial joint density can be written in conditional form:

$$f_{R,\Theta}(r, \theta) = f_{\Theta}(\theta) f_{R|\Theta}(r|\theta) \quad (4)$$

In this form, the problem of modeling multivariate extremes is transformed to that of modeling an angular density $f_{\Theta}(\theta)$ and the tail of the conditional radial density $f_{R|\Theta}(r|\theta)$. For a given angle θ , the density $f_{R|\Theta}(r|\theta)$ is univariate. Univariate extreme value theory suggests that a suitable model for the tail of $f_{R|\Theta}(r|\theta)$ is the generalized Pareto (GP) distribution, with parameters conditional on angle [23]. This motivates the SPAR model, first introduced in Ref. [24], whereby parametric and nonparametric models are used to model the conditional radial and angular distributions, respectively. Define a threshold function $u(\theta) > 0$ to be the quantile of $R|\Theta = \theta$ at exceedance probability $\zeta \in (0, 1)$ (where ζ is close to 0), i.e., the solution of $\zeta = \Pr(R > u(\theta)|\Theta = \theta)$. Then the SPAR model can be written as

$$f_{R,\Theta}(r, \theta) = \zeta f_{\Theta}(\theta) f_{GP}(r; \xi(\theta), \sigma(\theta), u(\theta)), \quad r > u(\theta) \quad (5)$$

where f_{GP} is the GP density function, and $\xi(\theta) \in \mathbb{R}$ and $\sigma(\theta) > 0$ are the shape and scale parameters, given as functions of the angle θ . The GP density function is given by

$$f_{GP}(r; \xi, \sigma, u) = \begin{cases} \frac{1}{\sigma} \left(1 + \xi \frac{r-u}{\sigma}\right)^{-1/(1/\xi)}, & \xi \neq 0 \\ \frac{1}{\sigma} \exp\left(-\frac{r-u}{\sigma}\right), & \xi = 0 \end{cases} \quad (6)$$

The support is $0 \leq r \leq r_F$, where the upper end point is $r_F = \infty$ for $\xi \geq 0$ and $r_F = u - \sigma/\xi$ for $\xi < 0$.

For the purposes of inference, it is also assumed that both the angular density f_{Θ} and GP parameter functions $\xi(\theta)$, $\sigma(\theta)$, and $u(\theta)$ are finite and continuous with angle. It was shown in Ref. [21] that these assumptions are valid for a wide range of copulas with Laplace marginal distributions. For parametric copulas, the asymptotic values of the GP parameter functions can be derived and used to compare the SPAR model for the density of the true values. This is shown in Fig. 1 for the cases of the

Frank, Gumbel, Gaussian, and Student- t copulas with Laplace marginal distributions (see Ref. [21] for details). The SPAR representation is in good agreement with the theoretical values, showing that the model can represent a wide range of dependence structures, without having to make prior assumptions about a particular parametric form of the copula. Furthermore, the SPAR approach has advantage over the conditional extremes method, in that it can characterize all extreme regions of the variable space in a single inference. It can also be shown that various other methods for estimating multivariate extremes, such as Refs. [25–27], are special cases of the SPAR model, and that SPAR provides a more flexible framework than existing techniques [21].

2.2 Use of the Model. Once the angular density and GP parameter functions have been estimated, Eqs. (3) and (5) can be combined to obtain the SPAR estimate of the joint density in the original variable space for observations satisfying $r > u(\theta)$:

$$f_{X,Y}(r \cos(\theta), r \sin(\theta)) = \frac{\zeta}{r} f_{\Theta}(\theta) f_{GP}(r; \xi(\theta), \sigma(\theta), u(\theta)) \quad (7)$$

The SPAR model also provides an explicit means for calculating a contour which has exceedance probability $\beta \leq \zeta$. The radius of this contour at angle θ is simply the quantile of the GP distribution at exceedance probability β/ζ , given by

$$r_{\beta}(\theta) = u(\theta) + \frac{\sigma(\theta)}{\xi(\theta)} ((\beta/\zeta)^{-\xi(\theta)} - 1) \quad (8)$$

This contour is defined in terms of the probability of an observation falling anywhere outside the region, or the “total exceedance probability.” As such, these contours are more conservative than those defined in terms of marginal exceedance probabilities, such as Inverse First Order Reliability Method (IFORM) contours, or variants thereof [28]. The SPAR model could also be used to estimate other types of contours, such as “Direct Sampling” [29] or “Highest Density Region” [30] contours. Estimation of contours defined in terms of the Rosenblatt transformation, such as IFORM [1] or Inverse Second Order Reliability Method (ISORM) [31], is less straightforward. The Rosenblatt transformation requires an estimate of the conditional density $f_{Y|X}$, which is not estimated directly by the SPAR model. The SPAR model only estimates the density in the region where $r > u(\theta)$. SPAR could be combined with a nonparametric estimate for the density in the region where $r \leq u(\theta)$, inferred using, for example, kernel density (KD) estimation, to produce an estimate of the density of the entire variable range. However, if the primary interest of the analysis is to estimate environmental contours, then the use of the SPAR model is not necessary. Instead, the authors recommend the use of the Direct-IFORM method [2,3], which does not require a model for the joint density or any assumptions about the dependence structure between the variables.

Calculating marginal and joint probabilities from the SPAR model involves integrating the joint density over various angular and radial ranges. However, probabilities can be estimated empirically by simulating from the estimated model. Simulation under the SPAR model is straightforward. We start by generating a random number p , uniformly distributed in $[0, 1]$. A random angle $\theta \in [0, 2\pi)$ can then be calculated by applying the probability integral transform so that $\theta = F_{\Theta}^{-1}(p)$. A corresponding radial value r is then simulated as a random value from the GP distribution with parameter vector $(\xi(\theta), \sigma(\theta), \mu(\theta))$. The pair (θ, r) is then a random sample from the SPAR model. This can be converted back to the original variable space using the inverse transformation $(x, y) = (r \cos(\theta), r \sin(\theta))$.

3 Inference

Inference for the SPAR model can be viewed as a non-stationary POT analysis, for which there are many examples in the literature [32–36]. In the present work, the parameter functions of the GP model for the tail of the conditional radial density are estimated

using the Extreme Value Generalised Additive Model (EVGAM) package [34]. This approach uses generalized additive models (GAMs) to represent the variation of the GP threshold, scale and shape parameters as functions of angle. GAMs are a flexible class of regression models that allow for complex, non-linear relationships between response and predictor variables [37]. Full details of this inference method are given in Ref. [22] and a GitHub repository of the associated code is available.² In this section, we provide high-level details only. We start by discussing inference for the angular density in Sec. 3.1, and then discuss inference for the conditional radial density in Sec. 3.2. The selection of tuning parameters for the model is discussed in Sec. 3.3.

3.1 Angular Density. As the angular density is assumed to be finite and continuous, it is readily amenable to nonparametric estimation methods. In the present work, we use KD estimation. Given a sample $\{\theta_1, \dots, \theta_n\}$, the KD estimate of the angular density at angle $\theta \in [0, 2\pi)$ is

$$\hat{f}_{\Theta}(\theta) = \frac{1}{n} \sum_{i=1}^n K_h(\theta, \theta_i) \quad (9)$$

where K_h denotes some kernel defined on a circular domain, with bandwidth parameter h . The bandwidth controls the smoothness of the estimate with smoothness increasing with bandwidth. For the examples considered in Sec. 4, we have used a von Mises kernel, given by

$$K_h(\theta, \theta_i) = \frac{1}{2\pi I_0(1/h)} \exp\left(\frac{\cos(\theta - \theta_i)}{h}\right)$$

where I_0 is the zero-order modified Bessel function of the first kind. As $h \rightarrow 0$, the von Mises kernel converges to a Gaussian kernel with variance h , and as $h \rightarrow \infty$ the kernel tends to the uniform distribution on $[0, 2\pi)$. This choice of kernel was shown to work well for the theoretical examples considered in Ref. [22].

We note that the angular density could also be estimated using GAMs, using a similar approach to that in Refs. [32,33]. While this is more elegant in that the same approach is used to model both the angular and radial components, we have opted to use a KD model in the present work due to its simplicity.

3.2 Conditional Radial Density. In the GAM framework, an arbitrary function $g(\theta)$ (which could represent the threshold, scale, or shape parameter) is represented as a sum of smooth basis functions specified at a finite number of locations, known as knots:

$$g(\theta) = \sum_{j=1}^k \beta_j B_j(\theta), \quad \theta \in [0, 2\pi) \quad (10)$$

where $B_j(\theta)$ are the basis functions, $\beta_j = g(\phi_j)$ is the value of the function at knot location $\phi_j \in [0, 2\pi)$, and $k \in \mathbb{N}$ is the number of knots (also referred to as the basis dimension). A wide variety of smooth basis functions exist [38]. In the present application, we have used cyclic cubic regression splines, which ensure that the GP parameter function estimates are periodic with angle. In practice, the values of β_j are not known and are estimated as part of the inference. EVGAM uses a penalized maximum likelihood method to estimate the coefficients β_j for the GP parameter functions. The penalty terms are defined in terms of the roughness of the solution, thus avoiding over-fitting. The optimal choice of roughness penalties are estimated using a cross-validation procedure. As with univariate POT analysis, the threshold is selected prior to fitting the GP distribution. In the multivariate case, threshold function is estimated using quantile regression [39,40], with a GAM representation for the model parameters.

²<https://github.com/callumbarltrop/SPAR>

The GP parameter functions estimated from EVGAM can be compared to estimates from a local stationary inference to check their plausibility. The local stationary inference is conducted for a grid of angular values. At each angle, a stationary GP model is fitted to the m nearest observations (in terms of angle). The assumption that the local distribution is stationary is an approximation. The choice of m is a bias-variance trade-off: smaller values lead to higher estimation variance, while larger values make the assumption of stationarity less valid. In practice, the local stationary inference is only used for verification of the GAM estimates.

3.3 Selection of Tuning Parameters. Application of the SPAR model requires the selection of the following tuning parameters:

- (1) kernel bandwidth, h , for the angular density;
- (2) threshold exceedance probability, ζ ;
- (3) spline basis dimension, k , and knot locations ϕ_1, \dots, ϕ_k .

In principle, optimal values of all these tuning parameters could be estimated from the data. However, to simplify the inference, we have opted to select these manually in the present implementation. For the kernel bandwidth, as with any application of kernel density estimation, the goal is to select h as small as possible, while avoiding over-fitting. In the examples presented in Sec. 4, it was found that a bandwidth of $h = 0.02$ was sufficient to represent the observed angular densities. The choice of threshold exceedance probability is directly analogous to the case in univariate analyses. That is, the threshold must be high enough such that the GP model is a reasonable approximation. Using too low a threshold will risk the model being misspecified, while too high a threshold will reduce the number of exceedances, increasing variance in the parameter estimates. A wide range of methods have been proposed for threshold selection; see Ref. [41] for a review and Ref. [42] for more recent developments. In the present work, the SPAR model was fitted for various threshold choices and diagnostic plots (discussed below) were checked to assess goodness of fit. For the examples considered here, we found that a threshold exceedance probability of $\zeta = 0.3$ was reasonable.

Selecting an appropriate basis dimension is essential for ensuring accuracy and flexibility in spline modeling procedures. Selecting too few knots may result in functional estimates that do not capture the underlying covariate relationships. Provided the basis dimension is sufficiently large, the resulting functional estimates should be relatively insensitive to the exact value. This is due to the roughness penalties, which prevent over-fitting, thus dampening the effect of adding additional basis knots to the spline formulations [37]. For the examples considered here, we found that using $k = 35$ knots was sufficient for the threshold and scale functions, and $k = 12$ knots was used for the shape parameter functions. The knot locations have been defined at equally spaced empirical quantiles of Θ . This ensures that the knot spacing is closer in regions where more data have been observed. Moreover, assuming the basis dimension is large enough, and model fit should be relatively insensitive to the precise location of knots [37]. We note that other spline representations and inference techniques allow the estimation of the optimal number of spline knots and their locations [35], although this is not considered further here.

4 Joint Extremes of Metocean Variables

In this section, we demonstrate the application of the SPAR model to the datasets provided as part of the recent benchmarking exercise for environmental contours [5]. This comprises three datasets of wave buoy measurements of significant wave height (H_s) and zero-crossing period (T_z) from wave buoys around the US coastline, and three datasets of wind speed (W_{spd}) and significant wave height for locations in the North Sea from the coastDat-2 hindcast [43]. All datasets have 1 hour time-steps and observations

will therefore be serially correlated. The effect of this is discussed further below. Dataset C from [5] was from a buoy in the Gulf of Mexico, an area affected by hurricanes. These types of datasets typically require careful treatment for hurricane-generated waves (see, e.g., Ref. [44]). Therefore, for the present study, we have replaced Dataset C with another wave buoy record from the US West coast, National Data Buoy Center (NDBC) buoy number 46014. Details of the datasets are given in Table 1. As different variables can have very different scales, we begin by centering and scaling the variables. For random variables X and Y , define normalized variables $(\tilde{X}, \tilde{Y}) = ((X - m_X)/s_X, (Y - m_Y)/s_Y)$, where m_X , m_Y , s_X , and s_Y are the sample mean and standard deviation of X and Y , respectively. We define our polar coordinates in terms of the normalized variables, thus ensuring that the origin is within the body of the data. As this is a rescaling of the data, rather than a marginal transformation, it is subject to lower uncertainties, and makes no assumptions about the distribution of the data.

For each dataset, the SPAR model was fitted using the inference procedure described in Sec. 3. To quantify the uncertainty, the datasets were resampled 200 times, and the SPAR model was fitted to each resampled dataset. As the hourly observations are serially correlated, a block bootstrap [45] was used, with a block length of 4 days. Here we consider various diagnostic plots to assess the quality of the fitted models. Comparisons of the observations with isodensity contours from the fitted SPAR models are shown in Fig. 2. In general, the isodensity contours provide a good description of the location of the observations. Although no information about physical constraints has been used to inform the model fitting, the models are able to capture various features of the data. First, the lower bounds on all variables are zero—this is captured in datasets A–E, but the model for dataset F does predict slightly negative wind speeds at higher values of H_s . This could potentially be prescribed as a hard constraint in the model. However, given that this part of the distribution is of less interest, and any negative data can easily be removed from estimates, we have not attempted to apply this constraint in the present work. Second, the SPAR models for the wave height-period data capture the limiting effect of the wave steepness, defined as $s = 2\pi H_s / (g T_z^2)$, where g is acceleration due to gravity. When the steepness exceeds a certain value, waves break, limiting the wave height at a given period. The limiting value of s is dependent on wind speed and water depth (among other factors), but values exceeding $s = 0.1$ are rare. The ability of the SPAR models to capture the limiting steepness is evident in the tight grouping of isodensity contours on the upper left side of the distributions.

A more localized assessment of the quality of the fitted model can be made by comparing the tails of the observed and modeled radial distributions over small angular ranges. Figure 3 shows an example for dataset B. Five angles have been selected to illustrate various features of the data. For each angle, observations are selected within a ± 2 deg range. At 13.5 deg and 40.5 deg, the model agrees well with the observations, within the estimated uncertainty. At 90 deg and 135 deg, there is some small discrepancy between the model and observations, although the uncertainty range is much smaller, due to the limiting effects of wave steepness at these angles. At 67.5 deg, which corresponds to the largest wave heights, there is a larger discrepancy between the model and largest observations. Note that all observations exceeding the upper 95% bound at this angle come only from two storms. This

Table 1 Description of metocean datasets

Dataset	Source	Variables	Start	End
A	NDBC buoy 44007	H_s, T_z	16/02/1982	31/12/2022
B	NDBC buoy 44024	H_s, T_z	29/04/1991	31/12/2022
C	NDBC buoy 46014	H_s, T_z	01/04/1981	31/03/2023
D, E, F	coastDat-2 hindcast	H_s, W_{spd}	01/01/1965	31/12/1989

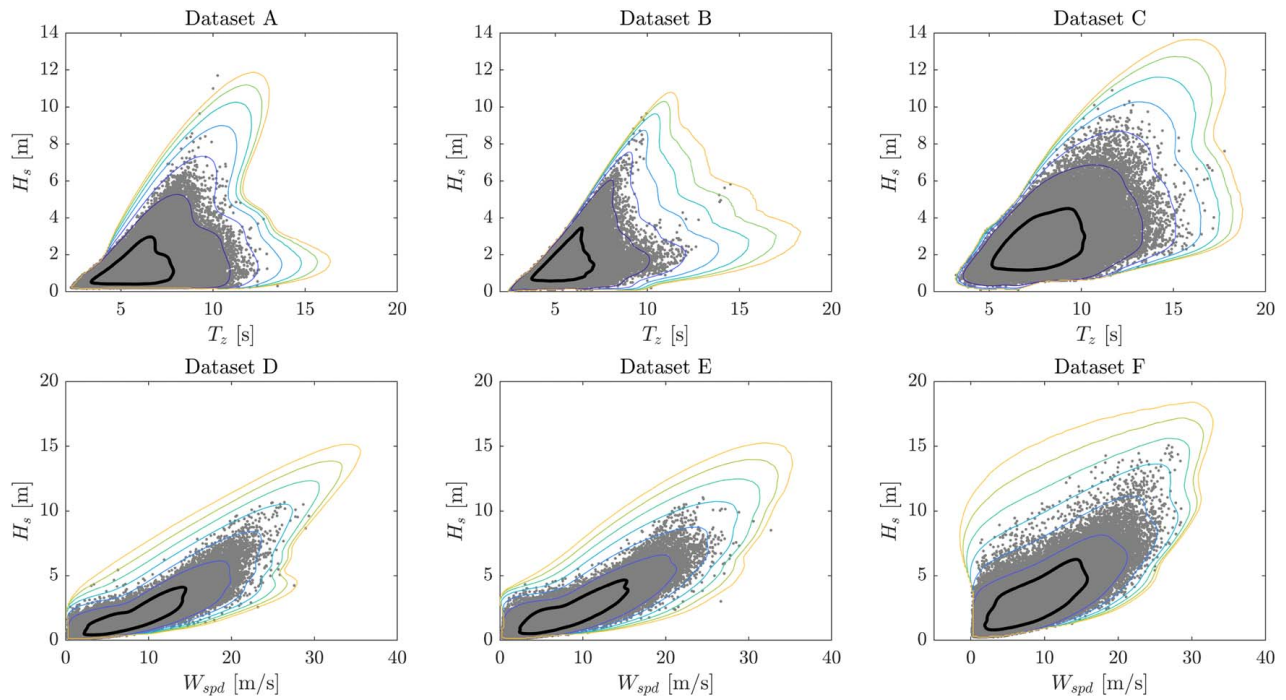


Fig. 2 Comparison of observations (dots) and isodensity contours from fitted SPAR models (lines). Thresholds are shown in thick black line. Isodensity contours are at equal logarithmic increments: 10^{-3} , 10^{-4} , ..., 10^{-8} .

is illustrated in Fig. 4, which highlights two 12-h periods (blue and red points), centered on these two storm peaks. Similarly, the largest observations of wave period all come from a single event (highlighted green). For comparison, the observations with these three 12-h periods removed are shown in the upper right plot of Fig. 4. As the model uncertainty is estimated using a block bootstrap, it accounts for these hourly observations coming from discrete events. The upper left plot of Fig. 4 shows exceedance contours

calculated using (8) at probability levels 10^{-2} and 10^{-4} together with 95% bounds. There is very low uncertainty on the high-steepness side of the contour, while on the upper side and right side the uncertainty range is wider, reflecting the longer tails of the conditional radial distribution in these ranges. Similar behavior was observed for datasets A and C (not shown).

Finally, we compare the marginal exceedance probabilities from the SPAR models and observations. To ensure consistency, both

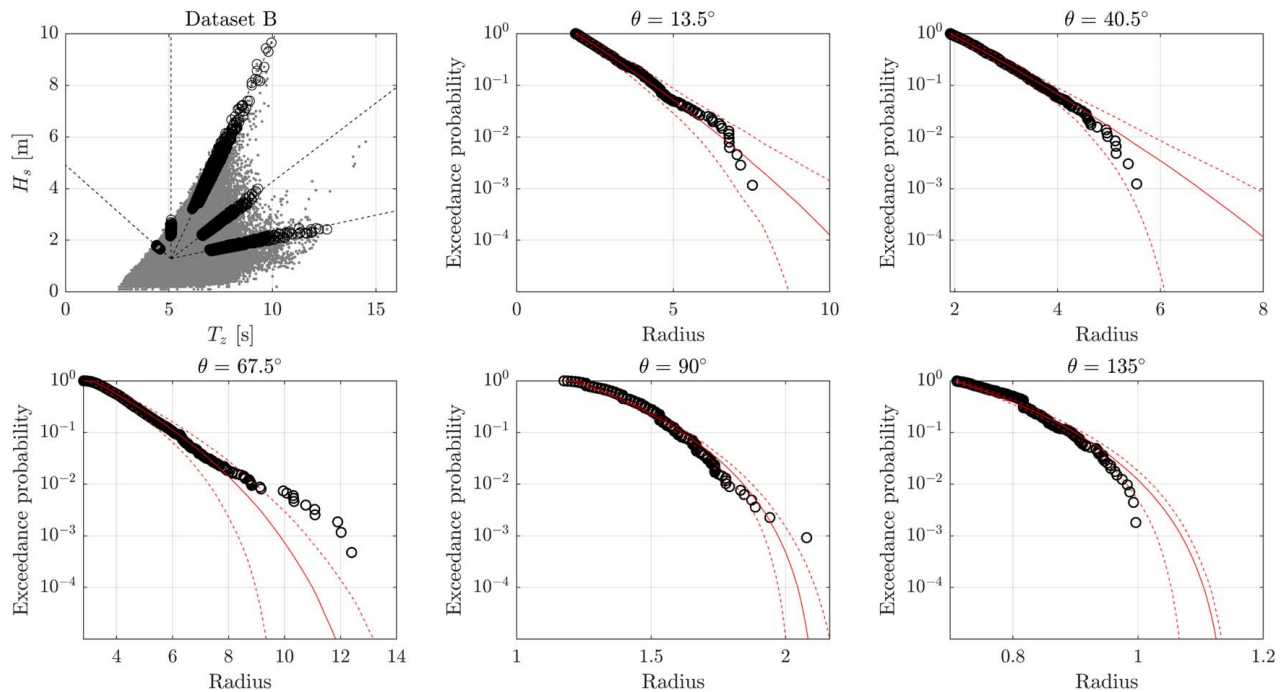


Fig. 3 Comparison of conditional radial exceedance probabilities at various angles, from observations (circles) and SPAR models (solid lines) for dataset B. Dashed lines indicate 95% bounds for model quantiles. Top left plot shows all observations (dots) and threshold exceedances used at each angle (circles).

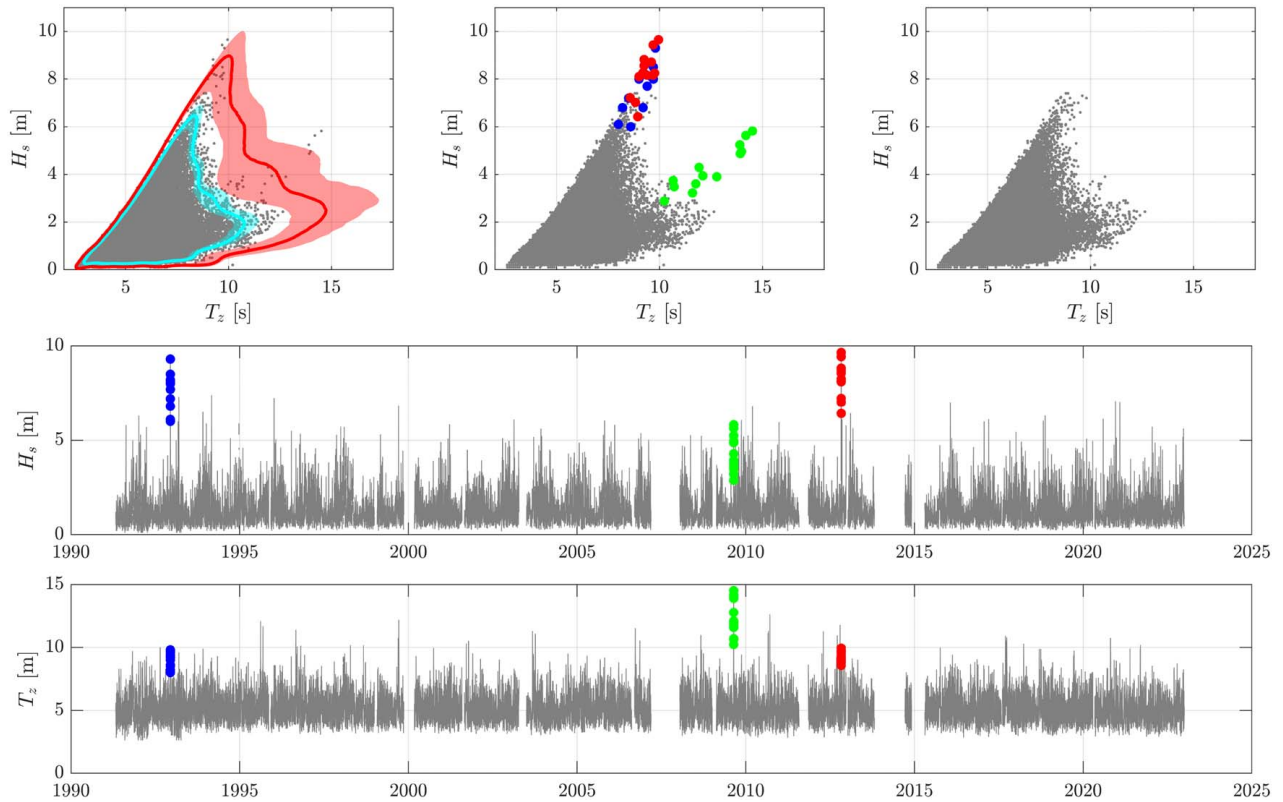


Fig. 4 Analysis of uncertainties in SPAR contours for dataset B. Top left: contours (solid lines) and 95% bounds for exceedance levels of 10^{-2} and 10^{-4} . Top middle: observations with three 12-h events highlighted (larger circles). Top right: observations with three 12-h events excluded. Lower plots: time series of H_s and T_z with three 12-h events highlighted, emphasizing the serial correlation and the need to account for this when estimating uncertainties.

observed and simulated data are restricted to exceedances of the highest value along the threshold function in the relevant dimension. That is, for the x -direction, observations are restricted to those for which $x > \max_{\theta \in (-\pi, \pi]} (u(\theta) \cdot \cos(\theta))$. Figure 5 shows exceedance probabilities from observations and fitted models for H_s , T_z , and s for each dataset. Generally, the models and observations are in good agreement. We emphasize that the marginal distributions have not been estimated directly, but are outputs from the fitted SPAR models (for the tail regions). Despite not fitting the marginal tails directly, the SPAR models still perform well. As discussed above, for dataset B, all observations with $H_s > 7.5$ m come from two storms. Similarly, for dataset A, the observations with $H_s > 10$ m come from a single storm. The SPAR models are in good agreement with the observations for the largest values of T_z for all datasets. The model slightly over-predicts the steepness for dataset A, although the model values are still realistic. For dataset C, the model slightly under-predicts the largest values of wave steepness. However, for this dataset, the largest values of steepness occur in smaller wave heights, with $H_s < 4$ m, so these conditions are less likely to affect structural design. Figure 6 shows marginal exceedance probabilities for datasets D–F. In these cases, the agreement is generally very good, with the fitted models agreeing with observations to within the estimated uncertainty.

5 Long-Term Extreme Responses

The analysis in the previous section has considered the fit of the SPAR model to observations. From an engineering perspective, it is usually the response of a structure that is of primary interest, rather than the environmental conditions themselves. To assess the SPAR method further, we consider how well the fitted model can capture response distributions for some simple response functions. In the case of the joint distributions of H_s and T_z , we consider three

response functions used in Ref. [6] for the vertical bending moment (VBM) on various vessels. The normalized responses are shown in Fig. 7(a) as functions of T_z , and responses are assumed to vary linearly with H_s . The peak responses occur at approximately 5.9 s, 8.5 s, and 11.8 s and are relatively broad-banded. Although the responses are for VBM on particular vessels, they can be considered as representative of a wider range of responses of offshore structures. For the joint distributions of H_s and W_{spd} , we consider the mudline overturning moment on a 5 MW offshore wind turbine using the response function derived in Ref. [46]; this is illustrated in Fig. 7(b). The response function given in Ref. [46] is also dependent on wave period, but we have removed this dependence by assuming a constant steepness of 0.03. The response function is also stochastic, so to avoid adding additional random effects to the comparison, we have used the median value for each environmental condition. There is a discontinuity in the response surface at $W_{spd} = 25$ m/s, which corresponds to the cut-out wind speed of the turbine. For winds above this speed, the turbine is shut down and the blades are feathered to reduce loading. At lower wave heights, this causes a reduction in loading. However, at very high wave heights, an increased response is observed above the cut-out speed, as the rotor is no longer damping the wave-induced motions of the tower.

Figure 8 compares exceedance probabilities for responses calculated from observations and simulations from the fitted SPAR models for datasets A–C. As the SPAR model is only fitted to threshold exceedances, the observed sample is restricted to the same range. We have also restricted both samples to values exceeding the mean H_s , to avoid including smaller responses in the comparison. For dataset C, the responses from models and observations are in good agreement. For datasets A and B, the agreement is generally good, but with some discrepancies, which are broadly inline with those shown for the marginal quantities in Fig. 5. Figure 9 compares exceedance probabilities for the wind turbine response

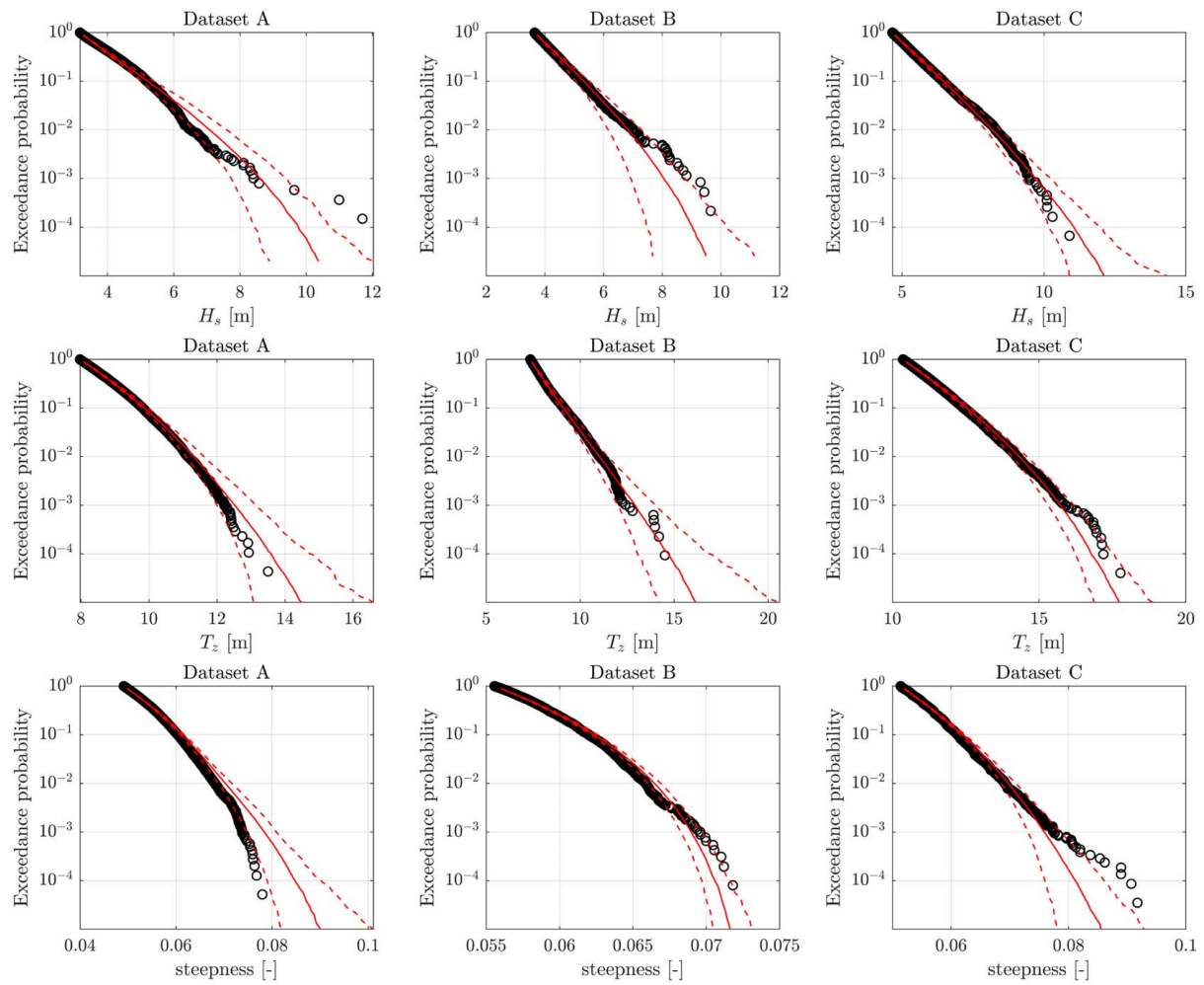


Fig. 5 Comparison of marginal exceedance probabilities from observations (circles) and SPAR models (solid lines) for wave height-period datasets. Dashed lines indicate 95% bounds on model quantiles.

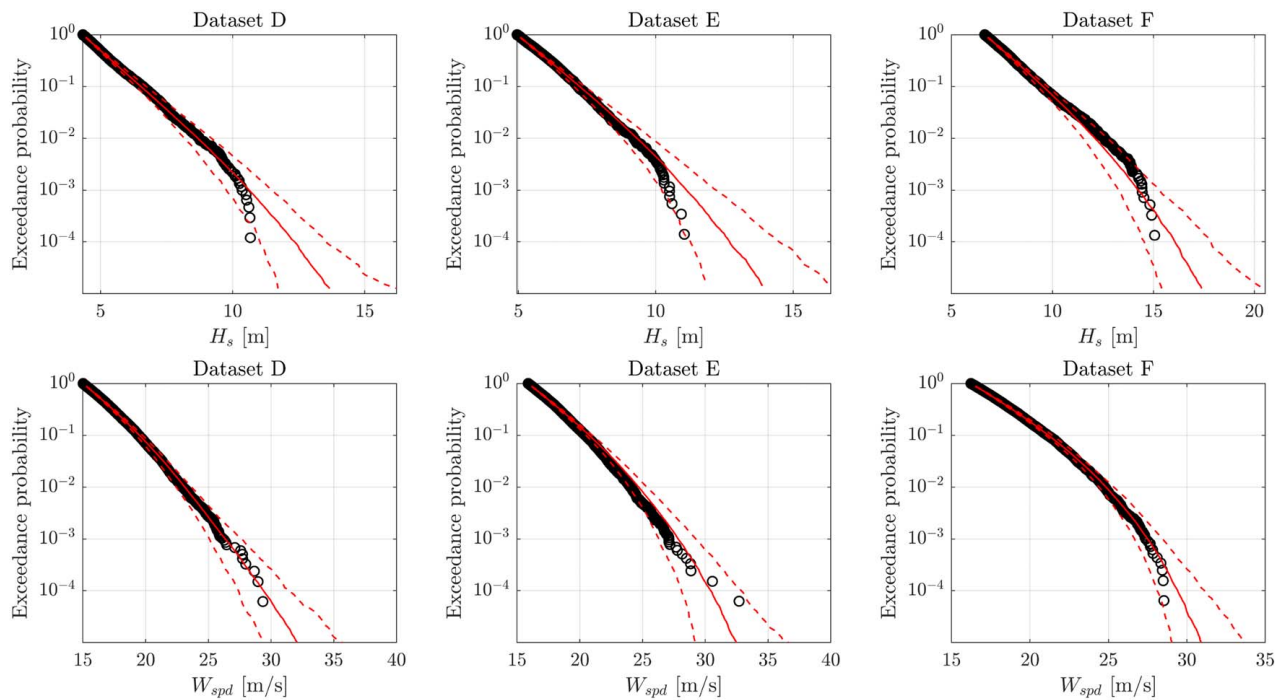


Fig. 6 Comparison of marginal exceedance probabilities from observations (circles) and SPAR models (solid lines) for wind-wave datasets. Dashed lines indicate 95% bounds on model quantiles.

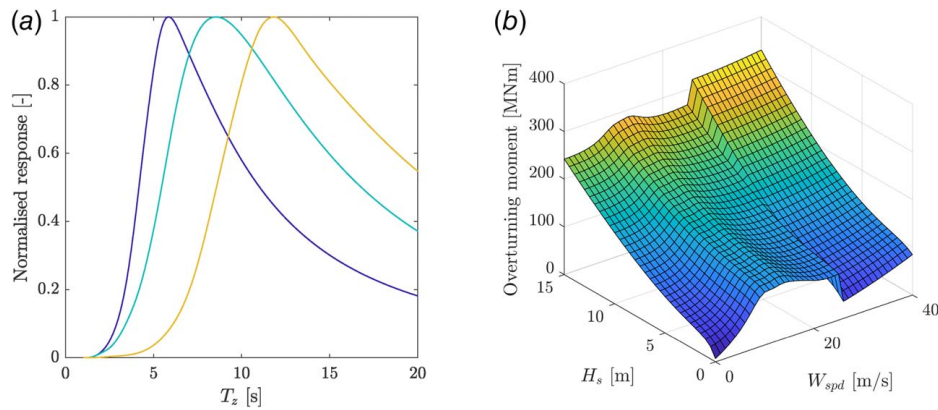


Fig. 7 Example response functions. Left: normalized responses for floating structures (responses increase linearly with H_s). Right: response surface for offshore wind turbine.

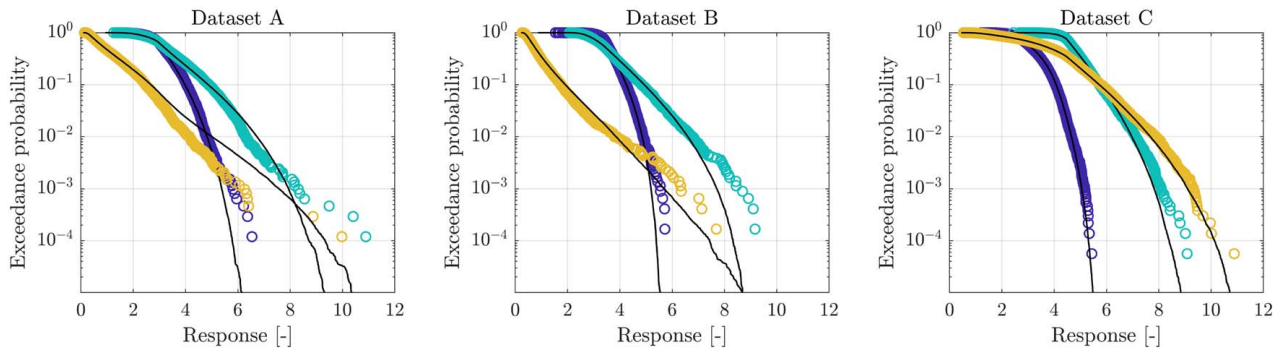


Fig. 8 Comparison of response exceedance probabilities from observations (circles) and SPAR models (lines) for wave height-period data

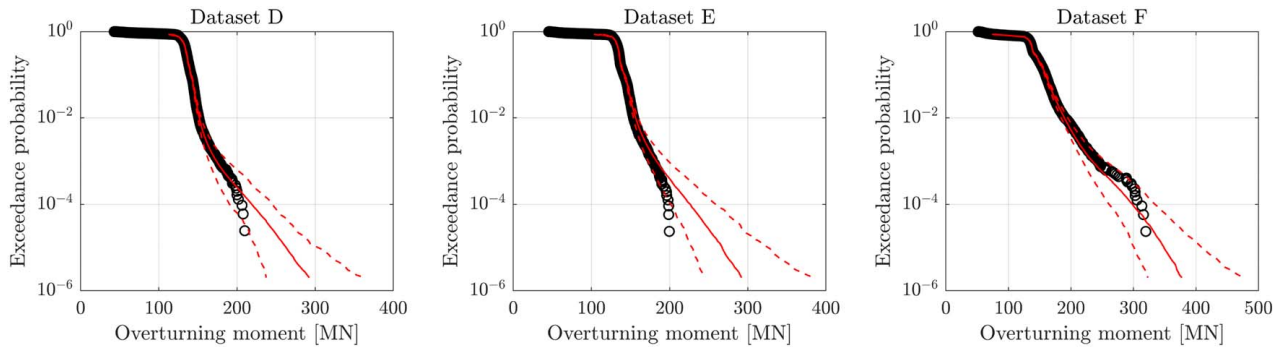


Fig. 9 Comparison of wind turbine response exceedance probabilities from observations (circles) and SPAR models (solid lines). Dashed lines indicate 95% bounds on model quantiles.

function calculated from observations and models for datasets D–F. In all cases, the models agree well with observations, indicating that they provide an accurate representation of the joint distribution in the relevant areas of the variable space.

6 Discussion and Conclusions

This paper has demonstrated the application of the SPAR model for estimating the joint distribution of metocean variables. The SPAR model has several advantages over existing methods. First, no assumptions are required about either the form of the marginal distributions or dependence structure, and the model can be applied on the original scale of the data, without a marginal

transformation. Second, the model provides a rigorous basis for extrapolating outside the range of observations, based on extreme value theory. And third, in contrast to existing methods, the same model formulation is applicable in different applications, removing the need for ad hoc assumptions. By transforming the variable space to polar coordinates, the SPAR approach reframes multivariate extreme value modeling as a natural extension of univariate extremes, with angular dependence.

The SPAR model was shown to provide a good representation of the joint extremes of wave heights and periods, and wave heights and wind speeds, using data for several locations. Accuracy of the fitted SPAR models was also assessed using simple response functions, relevant to offshore engineering applications. In the examples considered, the response distributions calculated from the models were in

good agreement with those from observations, giving confidence in the use of the models in engineering applications.

As discussed in Sec. 3.3, the SPAR model requires the selection of three tuning parameters: (1) the bandwidth for the kernel density estimate, (2) the threshold exceedance probability, and (3) the spline basis dimension and corresponding knot locations. While a level of user experience is required in selecting these tuning parameters, the modeling approach itself is not fundamentally new, since the inference is a non-stationary POT analysis. There is a sizeable body of literature on the application of these types of models to oceanographic variables, e.g., Refs. [33,35,36,47,48]. We note that the selection of tuning parameters for multivariate distribution modeling is not unique to the SPAR model. For example, in global hierarchical models, the parameters of the conditional distribution $Y|X$ are modeled as a function of the conditioning variable, X . In practice, it is assumed that the conditional distribution is piecewise-stationary over small intervals of X . The observations are divided into bins and the conditional distribution model is fitted to data in each bin. The bin width is a tuning parameter. The choice of bin width is a bias-variance trade-off: wider bins give more data to fit to but assume stationarity over a wider interval, and vice-versa. Fully parametric approaches, e.g., assuming parametric models for both the copula and marginal distributions, do not require selection of tuning parameters. However, as discussed in the introduction, the disadvantage for these approaches is that they require the selection of particular parametric models, for which there is no mathematical justification, other than convenience.

In future work, it will be useful to conduct a sensitivity study to assess the influence of the choice of tuning parameters. Moreover, it will be useful to conduct comparisons of the SPAR model with competitor approaches, such as global hierarchical models or copula models. However, given the results which have already been presented in a recent benchmarking study [5,6], which showed large uncertainties in these models, the results presented here using the same datasets demonstrate that SPAR offers a significant improvement in terms of the quality of fit.

The examples considered here are for two-dimensional problems. The SPAR approach is also applicable in higher dimensions. However, extending the inference to higher dimensions becomes more computationally challenging due to the number of parameters involved. The present work has not considered including covariate effects such as seasonality or directionality. However, as inference for the SPAR model can be viewed as non-stationary POT, where radius is the response variable and angle is the covariate, periodic covariates such as season and direction can be included as additional “angles,” treated in the same way as in existing models [47,48].

A limitation of the present model is that it does not provide information about serial correlation in the observations (although serial correlation is accounted for in the uncertainty calculations). This is usually accounted for by “declustering” the data, and only modeling peak values. However, in multivariate applications, what constitutes a “peak value” depends on which variable is of interest, since extremes of each variable do not necessarily occur simultaneously; see Refs. [3,49] for further discussion. Modeling the distribution of other values, conditional on a peak value in a similar manner to Ref. [50], may provide a suitable solution, but this will require further investigation.

Finally, the sensitivity of the inference to the choice of origin and polar coordinate system has not been considered in the present study. Comparisons of return level sets from inferences in two different coordinate systems were presented in Ref. [22] and shown to agree very well, indicating that the results are insensitive to the choice of coordinates.

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Conflict of Interest

There are no conflicts of interest.

Data Availability Statement

The datasets generated and supporting the findings of this article are obtainable from the corresponding author upon reasonable request.

References

- [1] Winterstein, S. R., Ude, T. C., Cornell, C. A., Bjerager, P., and Haver, S., 1993, “Environmental Parameters for Extreme Response: Inverse FORM With Omission Factors,” 6th International Conference on Structural Safety & Reliability (ICOSSAR), Innsbruck, Austria, Aug. 9–13.
- [2] Derbanne, Q., and de Hauteclercq, G., 2019, “A New Approach for Environmental Contour and Multivariate De-clustering,” 38th International Conference on Ocean, Offshore and Arctic Engineering, Glasgow, Scotland, June 9–14.
- [3] Mackay, E., and de Hauteclercq, G., 2023, “Model-Free Environmental Contours in Higher Dimensions,” *Ocean Eng.*, **273**, p. 113959.
- [4] Naess, A., and Moan, T., 2013, *Stochastic Dynamics of Marine Structures*, Cambridge University Press.
- [5] Haselsteiner, A. F., Coe, R. G., Manuel, L., Chai, W., Leira, B., Clarindo, G., Soares, C. G., et al., 2021, “A Benchmarking Exercise for Environmental Contours,” *Ocean Eng.*, **236**, p. 109504.
- [6] de Hauteclercq, G., Mackay, E., and Vanem, E., 2022, “Quantitative Comparison of Environmental Contour Approaches,” *Ocean Eng.*, **245**, p. 110374.
- [7] Ross, E., Astrup, O. C., Bitner-Gregersen, E., Bunn, N., Feld, G., Gouldby, B., Huseby, A., Liu, Y., Randell, D., Vanem, E., and Jonathan, P., 2020, “On Environmental Contours for Marine and Coastal Design,” *Ocean Eng.*, **195**, p. 106194.
- [8] Haver, S., 1985, “Wave Climate Off Northern Norway,” *Appl. Ocean Res.*, **7**(2), pp. 85–92.
- [9] Mathisen, J., and Bitner-Gregersen, E., 1990, “Joint Distributions for Significant Wave Height and Wave Zero-Up-Crossing Period,” *Appl. Ocean Res.*, **12**(2), pp. 93–103.
- [10] Nelsen, R. B., 2006, *An Introduction to Copulas*, Springer.
- [11] Joe, H., 2015, *Dependence Modeling With Copulas*, CRC Press.
- [12] Silva-González, F., Heredia-Zavoni, E., and Montes-Iturrizaga, R., 2013, “Development of Environmental Contours Using Nataf Distribution Model,” *Ocean Eng.*, **58**, pp. 27–34.
- [13] Montes-Iturrizaga, R., and Heredia-Zavoni, E., 2015, “Environmental Contours Using Copulas,” *Appl. Ocean Res.*, **52**, pp. 125–139.
- [14] Montes-Iturrizaga, R., and Heredia-Zavoni, E., 2017, “Assessment of Uncertainty in Environmental Contours Due to Parametric Uncertainty in Models of the Dependence Structure Between Metocean Variables,” *Appl. Ocean Res.*, **64**, pp. 86–104.
- [15] Jonathan, P., and Ewans, K., 2013, “Statistical Modelling of Extreme Ocean Environments for Marine Design: A Review,” *Ocean Eng.*, **62**, pp. 91–109.
- [16] Vanem, E., Zhu, T., and Babanin, A., 2022, “Statistical Modelling of the Ocean Environment—A Review of Recent Developments in Theory and Applications,” *Marine Struct.*, **86**, p. 103297.
- [17] Heffernan, J. E., and Tawn, J. A., 2004, “A Conditional Approach for Multivariate Extreme Values,” *J. R. Stat. Soc. B*, **66**, pp. 497–546.
- [18] Gouldby, B., Wyncoll, D., Panzeri, M., Franklin, M., Hunt, T., Hames, D., Tozer, N., Hawkes, P., Dornbusch, U., and Pullen, T., 2017, “Multivariate Extreme Value Modelling of Sea Conditions Around the Coast of England,” *Proc. Inst. Civil Eng. Maritime Eng.*, **170**, pp. 3–20.
- [19] Ross, E., Sam, S., Randell, D., Feld, G., and Jonathan, P., 2018, “Estimating Surge in Extreme North Sea Storms,” *Ocean Eng.*, **154**, pp. 430–444.
- [20] Liu, Y., and Tawn, J. A., 2014, “Self-consistent Estimation of Conditional Multivariate Extreme Value Distributions,” *J. Multivariate Anal.*, **127**, pp. 19–35.
- [21] Mackay, E., and Jonathan, P., 2023, “Modelling Multivariate Extremes Through Angular-Radial Decomposition of the Density Function,” *arXiv*, 10.
- [22] Murphy-Barltrop, C., Mackay, E., and Jonathan, P., 2023, “Inference for Multivariate Extremes Via a Semi-parametric Angular-Radial Model,” *arXiv*.
- [23] Coles, S., 2001, *An Introduction to Statistical Modeling of Extreme Values*, Springer.
- [24] Mackay, E., 2022, “Improved Models for Multivariate Metocean Extremes,” Tech. Rep., Supergen ORE Hub, 1.
- [25] Coles, S. G., and Tawn, J. A., 1991, “Modelling Extreme Multivariate Events,” *J. R. Stat. Soc. Ser. B (Methodol.)*, **53**(2), pp. 377–392.
- [26] Ledford, A. W., and Tawn, J. A., 1997, “Modelling Dependence Within Joint Tail Regions,” *J. R. Stat. Soc. Ser. B (Methodol.)*, **59**(2), pp. 475–499.

- [27] Wadsworth, J. L., Tawn, J. A., Davison, A. C., and Elton, D. M., 2017, "Modelling Across Extremal Dependence Classes," *J. R. Stat. Soc. Ser. B Methodol.*, **79**(1), pp. 149–175.
- [28] Mackay, E., and Haselsteiner, A. F., 2021, "Marginal and Total Exceedance Probabilities of Environmental Contours," *Marine Struct.*, **75**, p. 102863.
- [29] Huseby, A. B., Vanem, E., and Natvig, B., 2013, "A New Approach to Environmental Contours for Ocean Engineering Applications Based on Direct Monte Carlo Simulations," *Ocean Eng.*, **60**, pp. 125–135.
- [30] Haselsteiner, A. F., Ohlendorf, J.-H., Wosniok, W., and Thoben, K.-D., 2017, "Deriving Environmental Contours From Highest Density Regions," *Coastal Eng.*, **123**, pp. 42–51.
- [31] Chai, W., and Leira, B. J., 2018, "Environmental Contours Based on Inverse SORM," *Marine Struct.*, **60**, pp. 34–51.
- [32] Chavez-Demoulin, V., and Davison, A. C., 2005, "Generalized Additive Modelling of Sample Extremes," *J. R. Stat. Soc. Ser. C: Appl. Stat.*, **54**(1), pp. 207–222.
- [33] Randell, D., Turnbull, K., Ewans, K., and Jonathan, P., 2016, "Bayesian Inference for Nonstationary Marginal Extremes," *Environmetrics*, **27**(7), pp. 439–450.
- [34] Youngman, B. D., 2019, "Generalized Additive Models for Exceedances of High Thresholds With an Application to Return Level Estimation for U.S. Wind Gusts," *J. Am. Stat. Assoc.*, **114**(528), pp. 1865–1879.
- [35] Zanini, E., Eastoe, E., Jones, M., Randell, D., and Jonathan, P., 2020, "Flexible Covariate Representations for Extremes," *Environmetrics*, **31**, p. e2624.
- [36] Barlow, A. M., Mackay, E., Eastoe, E., and Jonathan, P., 2023, "A Penalised Piecewise-Linear Model for Non-stationary Extreme Value Analysis of Peaks Over Threshold," *Ocean Eng.*, **267**, p. 113265.
- [37] Wood, S. N., 2017, *Generalized Additive Models: An Introduction With R*, CRC Press.
- [38] Perperoglou, A., Sauerbrei, W., Abrahamowicz, M., and Schmid, M., 2019, "A Review of Spline Function Procedures in R," *BMC Med. Res. Methodol.*, **19**(1), pp. 1–16.
- [39] Koenker, R., 2005, *Quantile Regression*, Cambridge University Press.
- [40] Northrop, P. J., and Jonathan, P., 2011, "Threshold Modelling of Spatially Dependent Non-stationary Extremes With Application to Hurricane-Induced Wave Heights," *Environmetrics*, **22**(7), pp. 799–809.
- [41] Scarrott, C., and MacDonald, A., 2012, "A Review of Extreme Value Threshold Estimation and Uncertainty Quantification," *REVSTAT-Stat. J.*, **10**(1), pp. 33–60.
- [42] Murphy, C., Tawn, J. A., and Varty, Z., 2023, "Automated Threshold Selection and Associated Inference Uncertainty for Univariate Extremes," *arXiv*.
- [43] Groll, N., and Weisse, R., 2017, "A Multi-decadal Wind-Wave Hindcast for the North Sea 1949–2014: CoastDat2," *Earth Syst. Sci. Data*, **9**(2), pp. 955–968.
- [44] Wada, R., Rohmer, J., Krien, Y., and Jonathan, P., 2022, "Statistical Estimation of Spatial Wave Extremes for Tropical Cyclones From Small Data Samples: Validation of the STM-E Approach Using Long-Term Synthetic Cyclone Data for the Caribbean Sea," *Natl. Hazards Earth Syst. Sci.*, **22**(2), pp. 431–444.
- [45] Kunsch, H. R., 1989, "The Jackknife and the Bootstrap for General Stationary Observations," *Ann. Stat.*, **17**(3), pp. 1217–1241.
- [46] Haselsteiner, A. F., Frieling, M., Mackay, E., Sander, A., and Thoben, K.-D., 2022, "Long-Term Extreme Response of an Offshore Turbine: How Accurate Are Contour-Based Estimates?," *Renew. Energy*, **181**, pp. 945–965.
- [47] Randell, D., Feld, G., Ewans, K., and Jonathan, P., 2015, "Distributions of Return Values for Ocean Wave Characteristics in the South China Sea Using Directional-Seasonal Extreme Value Analysis," *Environmetrics*, **26**, pp. 442–450.
- [48] Hansen, H. F., Randell, D., Zeeberg, A. R., and Jonathan, P., 2020, "Directional-Seasonal Extreme Value Analysis of North Sea Storm Conditions," *Ocean Eng.*, **195**, p. 106665.
- [49] Mackay, E., de Haeteclocque, G., Vanem, E., and Jonathan, P., 2021, "The Effect of Serial Correlation in Environmental Conditions on Estimates of Extreme Events," *Ocean Eng.*, **242**, p. 110092.
- [50] Tendijck, S., Jonathan, P., Randell, D., and Tawn, J., 2023, "Temporal Evolution of the Extreme Excursions of Multivariate k th Order Markov Processes With Application to Oceanographic Data," *Environmetrics*.