# Research Notes

Yubo Su

March 6, 2018

# **Contents**

1	Preliminary Problems			2
	1.1	Incom	pressible, No Gravity	2
		1.1.1	Dirichlet BCs	3
		1.1.2	Things to Note	3
	1.2	Incom	pressible, Stratified w/ Gravity	4
		1.2.1	Eigenfunctions	4
		1.2.2	Analytically Solving an IVP, Dirichlet + Driving BCs	6
		1.2.3	Computationally Solving an IVP, Dirichlet + Driving BCs	7
		1.2.4	Waveform Properties, Analytical	7
		1.2.5	Extracting Information from the Waveform	8
		1.2.6	Supressing Reflection	8
	1.3	(Algeb	ora/Optional) The Anelastic/Boussinesq Approximations	9
		1.3.1	Developing the Anelastic/Boussinesq Approximations	9
		1.3.2	Anelastic Solution to Stratified Atmosphere	10

# Chapter 1

# **Preliminary Problems**

To get an intuition for how Dedalus and fluid mechanics works, we will solve some toy problems. Recall fluid equations in the presence of a uniform gravitational field  $\vec{g} = -g\hat{z}$ :

$$\frac{\mathrm{d}\rho}{\mathrm{d}t} + \vec{\nabla} \cdot \vec{u} = 0,$$

$$\frac{\mathrm{d}\vec{u}}{\mathrm{d}t} + \frac{\vec{\nabla}P}{\rho} - \vec{g} = 0.$$
(1.1)

In the incompressible limit,  $\frac{\mathrm{d}\rho}{\mathrm{d}t}=0$ , which implies  $\vec{\nabla}\cdot\vec{u}=0$ . We use subscripts to indicate perturbed quantities,  $Q_0$  is background and  $Q_1$  is perturbed. We will generally use  $\vec{u}_0=0$  unless otherwise noted. We will also generally assume symmetry along all axes except z the vertical axis.

In the incompressible limit, the fluid equations become

$$\vec{\nabla} \cdot \vec{u}_1 = 0,$$

$$\frac{\partial \rho_1}{\partial t} + u_{1z} \frac{\partial \rho_0}{\partial z} = 0,$$

$$\frac{\partial \vec{u}_1}{\partial t} + \frac{1}{\rho_0} \vec{\nabla} P_1 + \frac{\rho_1 g \hat{z}}{\rho_0} = 0.$$
(1.2)

We have used  $\vec{\nabla}P_0 = -\rho_0 g\hat{z}$  in the absence of perturbations.

# 1.1 Incompressible, No Gravity

We note that in the no gravity limit that  $\rho_1$  does not have an effect on other dynamical variables, so the equations of motion we must solve are

$$\vec{\nabla} \cdot \vec{u}_1 = 0,$$

$$\frac{\partial \vec{u}_1}{\partial t} + \frac{\vec{\nabla} P_1}{\rho_0} = 0.$$
(1.3)

We can take the divergence of the momentum equation and substitute the continuity equation to get  $\nabla^2 P = 0$ .

#### 1.1.1 Dirichlet BCs

This is a Laplace equation, which we've solved countless times. Imposing periodic boundary conditions in the x direction and  $P_1(z=L)=0, P_1(z=0)=\mathcal{P}(x,t)$ , we obtain eigenfunctions

$$P_{1,n}(x,z,t) = \frac{\mathscr{P}_n(t)}{\sinh(k_n L)} e^{ik_n x} \sinh(k_n (L-z)),$$

$$u_{1x,n}(x,z,t) = \int_0^t -\frac{1}{\rho_0} \frac{\partial P_{1,n}}{\partial x} dt,$$

$$u_{1z,n}(x,z,t) = \int_0^t -\frac{1}{\rho_0} \frac{\partial P_{1,n}}{\partial z} dt.$$
(1.4)

We define  $k_n = \frac{2\pi n}{L}$ ,  $n \ge 0$  and  $\mathscr{P}(x,t) = \sum_{n} \mathscr{P}_n(t) e^{ik_n x}$ .

Thus, if we impose BCs  $\mathcal{P}(x,t) = \sin \frac{2\pi x}{L}$  and start with initial conditions such that all quantities are zero, we would expect after transients die out that

$$P(x,z,t) = \frac{\sin\frac{2\pi x}{L}}{\sinh 2\pi} \sinh\left(2\pi \frac{L-z}{L}\right),$$

$$u_{1x}(x,z,t) = -\frac{2\pi t}{L\rho_0} \frac{\cos\frac{2\pi x}{L}}{\sinh 2\pi} \sinh\left(2\pi \frac{L-z}{L}\right),$$

$$u_{1z}(x,z,t) = +\frac{2\pi t}{L\rho_0} \frac{\sin\frac{2\pi z}{L}}{\sinh 2\pi} \cosh\left(2\pi \frac{L-z}{L}\right).$$
(1.5)

This is in good agreement with the results, presented in Fig. 1.1. Note that P is constant while  $\vec{u}$  increases linearly in time, and we observe the expected  $\sim \sin x \sinh \frac{L-z}{z}$  dependence. In fact,  $u_{1x}, u_{1z}$  are exactly  $\frac{2\pi}{10}$  at t=1.

### 1.1.2 Things to Note

It is worth noting that, since our Eq. 1.3 reduced to a Laplace equation, we needed two z BCs and two x BCs (periodic BCs amount to equating the value and derivative of the function). This is in agreement with the observation that the original Eq. 1.3 had two derivatives in x,z apiece, so we needed two BCs each.

It is also worth seeing from our solution that P immediately goes to the equilibrium solution. This is not surprising since in the incompressible limit, sound speed goes to infinity which is the timescale on which the pressure field adjusts to net forces. Thus, the dynamics solely arise from the static pressure field pushing the velocity field to equilibrium.

We chose time-independent  $\mathcal{P}(x,t)$ , but it is clear that whatever  $\mathcal{P}(x,t)$  we choose, the time dependence propagates to the velocities by way of an integral. If we had instead chosen to take Fourier



**Figure 1.1:** P,  $u_x$ ,  $u_z$  at t = 0.5 and t = 1 for  $\rho_0 = 1$ . We choose a L = 10 square domain.

transform  $t \to \omega$ , we would have had to integrate the boundary condition against the eigenfunctions for each of the  $\omega$ , which is still easily computable, to get the full  $u_{1x}(x,z,t)$ . We consider this in preparation for when we can only solve for a set of  $\vec{k}, \omega$  in the next problem.

The next thing we would have wanted to do is solve a problem with radiative BCs, but we need to have wave solutions, which are missing in the absence of gravity. We thus move on to the next configuration.

# 1.2 Incompressible, Stratified w/ Gravity

## 1.2.1 Eigenfunctions

Let's restore the  $\rho_1 g$  term now. For funsies, we begin by solving for arbitrary stratification  $\rho_0(z)$  first. The fluid equations to first order reduce to

$$\begin{split} \frac{\partial \rho_1}{\partial t} + \left( \vec{u}_1 \cdot \vec{\nabla} \right) \rho_0 &= 0, \\ \vec{\nabla} \cdot \vec{u}_1 &= 0, \\ \frac{\partial \vec{u}_1}{\partial t} &= -\frac{\vec{\nabla} P_1}{\rho_0} - \frac{\rho_1 g}{\rho_0} \end{split} \tag{1.6}$$

We expect there to be some z dependence in the amplitude, so we substitute variables of form  $e^{i(kx-\omega t)}$  and do not specify the z dependence. This gives us

$$-i\omega\rho_{1} - u_{1z}\frac{\partial\rho_{0}}{\partial z} = 0,$$

$$iku_{1x} + \frac{\partial u_{1z}}{\partial z} = 0,$$

$$-iwu_{1x} + \frac{ik_{x}P_{1}}{\rho_{0}} = 0,$$

$$-iwu_{1z} + \frac{1}{\rho_{0}}\frac{\partial P_{1}}{\partial z} + \frac{\rho_{1}g}{\rho_{g}} = 0.$$

$$(1.7)$$

We substitute  $N^2 = -\frac{g}{\rho_0} \frac{\partial \rho_0}{\partial z}$  to obtain

$$-i\omega\rho_1 - u_{1z}\frac{\rho_0 N^2}{g} = 0, (1.8a)$$

$$iku_{1x} + \frac{\partial u_{1z}}{\partial z} = 0, \tag{1.8b}$$

$$-iwu_{1x} + \frac{ik_x P_1}{\rho_0} = 0, (1.8c)$$

$$-iwu_{1z} + \frac{1}{\rho_0} \frac{\partial P_1}{\partial z} + \frac{\rho_1 g}{\rho_0} = 0.$$
 (1.8d)

Eliminating  $u_{1x}$  by substituting (1.8b) into (1.8c) and  $\rho_1$  by substituting (1.8a) into (1.8d) give

$$i\omega \frac{\partial u_{1z}}{\partial z} + \frac{k_x^2 P_1}{\rho_0} = 0, \tag{1.9a}$$

$$\left(\omega^2 - N^2\right)u_{1z} + \frac{i\omega}{\rho_0}\frac{\partial P_1}{\partial z} = 0. \tag{1.9b}$$

Finally, we multiply (1.9a) with  $\rho_0$  and differentiate dz and combine with (1.9b) to give

$$\frac{\mathrm{d}^2 u_{1z}}{\mathrm{d}z^2} + \frac{1}{\rho_0} \frac{\partial \rho_0}{\partial z} \frac{\partial u_{1z}}{\partial z} + k_x^2 \left(\frac{N^2}{\omega^2} - 1\right) u_{1z} = 0. \tag{1.10}$$

Let's now pick stratification  $\rho \propto e^{-z/H}$  Eq. 1.10 clearly has exponential solutions  $e^{\kappa z}$  for

$$\kappa^2 - \frac{\kappa}{H} + k_x^2 \left( \frac{N^2}{\omega^2} - 1 \right) = 0. \tag{1.11}$$

We permit complex  $\kappa = \frac{1}{2H} + ik_z$ , and from the above clearly

$$k_z^2 = -\frac{1}{4H^2} + k_x^2 \left(\frac{N^2}{\omega^2} - 1\right),$$

$$\omega^2 = \frac{N^2 k_x^2}{k_x^2 + k_z^2 + \frac{1}{4H^2}}.$$
(1.12)

Thus the eigenfunctions are

$$\begin{split} u_{1z} &= e^{z/2H} e^{i(k_z z + k_x x - \omega t)}, \\ u_{1x} &= -\frac{k_z + i/2H}{k_x} u_{1z}, \\ \rho_1 &= \frac{i\rho_0}{H\omega} u_{1z}, \\ P_1 &= -\frac{\rho_0 \omega}{k_x^2} (k_z + i/2H) u_{1z}. \end{split} \tag{1.13}$$

## 1.2.2 Analytically Solving an IVP, Dirichlet + Driving BCs

We will analyze everything in terms of  $u_{1z}$  since it has the simplest form; note that when actually choosing the BCs we will have to consider the gauge freedom of P and some considerations we defer to the computational section.

Currently, we have a set of eigenfunctions

$$u_{1z}(x,z,t|\vec{k},\omega) = e^{z/2H}e^{i(k_zz+k_xx-\omega t)}.$$
 (1.14)

Note that  $u_{1z}$  is really only a function of two parameters rather than the three  $(k_x, k_z, \omega)$ , since the three are related by dispersion relation Eq. 1.12.

Now, we implement BCs. Consider domain  $x, z \in [0, L]$ . We will use periodic BCs again in x, so then  $k_{x,n} = \frac{2\pi n}{L}$ ,  $n \ge 0$ . Then, we will require  $u_{1z,n}(x,L,t) = 0$ , a Dirichlet condition at the top boundary, which restricts us to eigenfunctions of form

$$u_{1z,n}(x,z,t|\vec{k},\omega) = e^{z/2H}e^{i(k_x x - \omega t)}\sin(k_z(L-z)).$$
 (1.15)

Finally, we must choose a BC at z = 0. We will choose a general function  $u_{1z}(x,0,t) = F(x,t)$  where we can decompose

$$F(x,t) = \int \sum_{n} \mathcal{F}(k_{x,n},\omega) e^{i(k_{x,n}x - \omega t)} d\omega.$$
 (1.16)

Matching BCs then gives us general solution for  $u_{1z}$  given an arbitrary driving function

$$u_{1z}(x,z,t|\vec{k},\omega) = \int \sum_{n} \mathcal{F}(k_{x,n},\omega) \frac{e^{z/2H} e^{i(k_{x,n}x-\omega t)} \sin(k_{z}(L-z))}{\sin k_{z}L} d\omega.$$
 (1.17)

For ease of computation, let's pick  $F(x,t) = \cos(\frac{2\pi x}{L} - \omega_0 t)$ , so our full expected solution is (note

$$A + \epsilon = Ae^{\epsilon/A} + \mathcal{O}(\epsilon^2)$$

$$u_{1z}\left(x,z,t\left|\vec{k},\omega\right) = e^{z/2H} \frac{\cos\left(\frac{2\pi x}{L} - \omega_{0}t\right)\sin(k_{z}(L-z))}{\sin k_{z}L},$$

$$u_{1x}\left(x,z,t\left|\vec{k},\omega\right) \approx -\frac{k_{z}}{k_{x}}e^{z/2H} \frac{\cos\left(\frac{2\pi x}{L} - \omega_{0}t + \frac{1}{2Hk_{z}}\right)\sin(k_{z}(L-z))}{\sin k_{z}L},$$

$$\rho_{1}\left(x,z,t\left|\vec{k},\omega\right) \approx \frac{\rho_{0}}{H\omega}e^{z/2H} \frac{-\sin\left(\frac{2\pi x}{L} - \omega_{0}t\right)\sin(k_{z}(L-z))}{\sin k_{z}L},$$

$$P_{1}\left(x,z,t\left|\vec{k},\omega\right) \approx -\frac{\rho_{0}\omega k_{z}}{k_{x}^{2}}e^{z/2H} \frac{-\cos\left(\frac{2\pi x}{L} - \omega_{0}t + \frac{1}{2Hk_{z}}\right)\sin(k_{z}(L-z))}{\sin k_{z}L},$$

$$(1.18)$$

where  $k_z : \omega(k_x, k_z) = \omega_0$  by the dispersion relation Eq. 1.12. Note that H contributes both to the overall exponential profile and to the phase lag of  $u_{1x}$ .

### 1.2.3 Computationally Solving an IVP, Dirichlet + Driving BCs

To solve this computationally with the aforementioned BCs, periodic in x, Dirichlet 0 at z=L and  $\cos(\frac{2\pi x}{L}-\omega_0 t)$  at z=0, we must address the gauge freedom in P. This arises because for  $k_x=0$ , the divergence-free condition  $\nabla \cdot \vec{u}_1 = \frac{\partial u_z}{\partial z} = 0$  specificies  $u_{1z}$  up to a constant already, so the bottom BC will fix the value of  $u_z$  when  $k_x=0$ .

A different way of phrasing the same argument is as follows. Consider the discrete  $N \times N$  (square for notational simplicity) grid. At the boundary, there is a list of values  $f(\{x_i\}, z_0)$  that lives in an N dimensional space. We can thus pick a spanning set of N basis vectors, and for each of these N vectors, by enforcing  $\vec{\nabla} \cdot \vec{u} = 0$  at the boundary we fix the allowed  $f(\{x_i\}, z_{-1})$  boundary conditions we can implement. But there exists a choice of basis vectors for which one of the basis vectors is constant  $f_i(\{x_i\}, z_0) = C$ . For this basis vector, the boundary condition is fully determined, so we only have N-1 dimensions from which to choose the BCs  $f(\{x_i\}, z_{-1})$ . Since the dimensionality of a space cannot depend on the choice of basis vector, the divergence free condition actually yields an extra degree of freedom.

We can use this extra degreee of freedom to specify P(z=L)=0 so the oscillations should have zero mean. Then we can just simulate away!

From here we can just tweak the parameters until we get something useful, can look at the videos. Note that we should simulate until at least  $T > \frac{2L_z}{v_{p,z}}$  the z phase space velocity, so we can capture any reflections off the boundary. Turns out both Dirichlet and Neumann BCs give strong reflections that produce standing waves in the z and traveling waves in the x.

### 1.2.4 Waveform Properties, Analytical

Let's figure out the analytical forms for the phase, group velocity and the energy density/power flux. We will try to recover these from waveforms.

Let's first consider the phase velocity. We traditionally think about the phase velocity in 1D

 $v_{ph} = \frac{\omega(k)}{k}$ , but in general it is the function such that  $\vec{k} \cdot \vec{r} - \omega t = \vec{k} \cdot (\vec{r} - \vec{v}_{ph} t)$  is constant (the phase of the wave). Thus, it must satisfy  $\vec{k} \cdot \vec{v}_{ph} = \omega$ , and so one sensible choice is

$$\vec{v}_{ph} = \frac{\omega \hat{k}}{\left|\vec{k}\right|} = \frac{\omega \vec{k}}{\left|\vec{k}\right|^2}.$$
(1.19)

For our stratified atmosphere problem,  $\omega^2 = \frac{N^2 k_x^2}{k^2 + \frac{1}{4H^2}}$ . Plugging in our parameters  $k_z = -k_x = 2pi, N^2 = 10$  we can find  $\omega \approx 2.23$  Hz, so  $\vec{v}_{ph} \approx 0.1777(-\hat{x} + \hat{z})$  s<sup>-1</sup>.

Since we use a height of 5, we should expect the initial disturbance to reach the top boundary at  $t \sim 30$  s.

On the other hand the group velocity is given  $\vec{v}_g = \frac{\partial \omega}{\partial k_i} \hat{\iota}$ . In our problem,  $v_{g,z} = -\frac{N^2 k_x^2 k_z}{(k_x^2 + k_z^2 + 1/4H^2)^{3/2}}$  while  $v_{g,x} = \frac{2N^2 k_x}{\sqrt{k_x^2 + k_z^2 + 1/4H^2}} - \frac{N^2 k_x^3}{(k_x^2 + k_z^2 + 1/4H^2)^{3/2}}$ .

We should note that the kinetic energy density is just  $E = \frac{1}{2}\rho v^2$ , the kinetic energy. Using then  $\frac{dE}{dt} = 0$  shows that the corresponding power flux is  $\vec{P} = \frac{\vec{v}}{2}\rho v^2$ .

## 1.2.5 Extracting Information from the Waveform

Let's make sure we can extract information from the linearized problem. Notable things we should be interested in: phase/group velocity, energy density and energy flux.

Based on our phase velocity expression Eq. 1.20, if we have some waveform f(x,z,t), we can estimate its phase velocity  $\vec{v}_{ph}(x,z,t)$  by

$$\vec{v}_{ph}(x,z,t) = \frac{\partial_t f \vec{\nabla} f}{\left| \vec{\nabla} f \right|^2}.$$
 (1.20)

This can be validated in the  $f(x,z,t) \propto e^{i(\vec{k}\cdot\vec{x}-\omega t)}$  case to give  $\vec{v}_{ph} = \frac{\omega\vec{k}}{k^2}$ . Note the sign in Eq. 1.20

### 1.2.6 Supressing Reflection

Sourced from https://people.maths.ox.ac.uk/trefethen/6all.pdf. There are largely two ways to do this. The first is to add a damping zone, a region in which  $\dot{q}=-q/\tau_d(z)$  where  $\tau_d(0)=0$  and increases to some non-small number for some  $z_0$  beyond which we want damping. One form, chosen in Ryan and Dong's paper, is to use multiplicative factor  $f(z)=\max\left[0,1-\frac{(z-z_b)^2}{(z_d-z_b)^2}\right]$  where  $z_b$  is where one begins supression and  $z_b$  is the boundary. Then a dynamical variable q can be supressed via  $\dot{q}=(\ldots)-f(z)\frac{q-q_0}{\tau}$  where  $\tau$  is the dynamical timescale on which to damp and  $q_0$  is the value to damp to.

This requires we compute a ton of extra points though. Ideally, we want to use BCs. TODO

# 1.3 (Algebra/Optional) The Anelastic/Boussinesq Approximations

### 1.3.1 Developing the Anelastic/Boussinesq Approximations

Let's relax the incompressibility constraint (we will expand the continuity equation to first order, but the momentum equation will merit a separate treatment):

$$\begin{split} \frac{\partial \rho_1}{\partial t} + \vec{\nabla} \cdot \left( \rho_0 \vec{u}_1 \right) &= 0, \\ \frac{\partial \vec{u}_1}{\partial t} &= -\frac{\vec{\nabla} P}{\rho} - \vec{g}. \end{split} \tag{1.21}$$

Suppose we are interested in phenomena with characteristic length scale L and time scale  $\tau$ . Let's first examine the relative magnitudes of the terms in the continuity equation

$$\frac{\rho_1}{\tau} + \frac{\rho_0 |u_1|}{L} = 0.$$

Thus, if we are interested in time scales  $\tau \gg \frac{\rho_1}{\rho_0} \frac{L}{|u_1|}$  then we neglect the first term, the time derivative. This corresponds to making the perturbation incompressible; note that  $\frac{\partial \rho_1}{\partial t} \approx \frac{\mathrm{d}\rho_1}{\mathrm{d}t}$  to first order, so we drop the high frequency restoring forces in the perturbation.

For the momentum equation, we instead first manipulate to first order

$$-\frac{\vec{\nabla}P}{\rho} - \vec{g} = -\frac{\vec{\nabla}P_0}{\rho} - \frac{\vec{\nabla}P_1}{\rho_0} - \vec{g},$$

$$= -\frac{\vec{\nabla}P_1}{\rho_0} + \left(\frac{\rho_0}{\rho} - 1\right)\vec{g},$$

$$= -\vec{\nabla}\left(\frac{P_1}{\rho_0}\right) - \frac{P_1}{\rho_0^2}\vec{\nabla}\rho_0 - \frac{\rho_1}{\rho_0}\vec{g}.$$
(1.22)

We now have three equations for four variables,  $\vec{u}_1, \rho_1, P_1$ . We must introduce a fourth equation, a thermodynamic equation. For an adiabatic process  $P\rho^{-\gamma} \propto P^{1-\gamma}T^{\gamma}$  is constant. We thus introduce the concept of the *potential temperature* 

$$\theta = T \left(\frac{P_0}{P}\right)^{\kappa}.\tag{1.23}$$

For an adiabatic process,  $\frac{d\theta}{dt} = 0$ . Motivated by this, we use

$$\begin{split} \frac{\partial 1}{\partial \rho_0} \frac{\partial \rho_0}{\partial z} &= \frac{1}{\gamma P_0} \frac{\partial P_0}{\partial z} - \frac{1}{\theta_0} \frac{\partial \theta_0}{\partial z}, \\ \frac{\rho_1}{\rho_0} &= \frac{1}{\gamma} \frac{P_1}{P_0} - \frac{\theta_1}{\theta_0}, \end{split} \tag{1.24}$$

to give the momentum equation form

$$\frac{\mathrm{d}\vec{u}_1}{\mathrm{d}t} = -\vec{\nabla} \left(\frac{P_1}{\rho_0}\right) + \frac{P_1}{\rho_0} \left(\frac{1}{\theta_0}\vec{\nabla}\theta_0\right) + \vec{g}\frac{\theta_1}{\theta_0}. \tag{1.25}$$

We also recognize  $N^2 = \frac{g}{\theta_0} \frac{\partial \theta_0}{\partial z}$ . We now do the same trick where we consider dynamics on length scale D and compare the first and second terms in Eq. 1.25. Their ratio is  $\frac{N^2D}{g}$ , and so as  $N^2 \ll \frac{g}{D}$  the freefall time we neglect the second term.

The anelastic fluid equations thus read

$$\vec{\nabla} \cdot (\rho_0 \vec{u}) = 0,$$

$$\frac{\partial \vec{u}_1}{\partial t} + \vec{\nabla} \left(\frac{P_1}{\rho_0}\right) - \vec{g} \frac{\theta_1}{\theta_0} = 0,$$

$$\frac{\partial \theta_1}{\partial t} + \left(\vec{u} \cdot \vec{\nabla}\right) \theta_0 = 0.$$
(1.26)

The Boussinesq equations are obtained from these in the limit where  $H \gg D$  the relevant length scale, thus we allow  $\rho_0$  to be approximately constant.

### 1.3.2 Anelastic Solution to Stratified Atmosphere

We simply substitute  $e^{i(\vec{k}\cdot\vec{r}-\omega t)}$  into Eq. 1.26 with  $ho_0\propto e^{-z/H}$  and obtain

$$\begin{bmatrix} 0 & 0 & ik_{x}\rho_{0} & ik_{z}\rho_{0} - \frac{\rho_{0}}{H} \\ 0 & -i\omega & 0 & \frac{N^{2}\theta_{0}}{g} \\ \frac{ik_{x}}{\rho_{0}} & 0 & -i\omega & 0 \\ \frac{ik_{z}}{\rho_{0}} + \frac{1}{\rho_{0}H} - \frac{g}{\theta_{0}} & 0 & -i\omega \end{bmatrix} \begin{bmatrix} P_{1} \\ \theta_{1} \\ u_{1x} \\ u_{1z} \end{bmatrix} = 0.$$
 (1.27)

Taking the determinant of this matrix produces

$$\begin{split} -k_x^2 \left( -N^2 + \omega^2 \right) + \left( ik_z - \frac{1}{H} \right) & \left( ik_z + \frac{1}{H} \right) \omega^2 = 0, \\ & \frac{N^2 k_x^2}{k_x^2 + k_z^2 + \frac{1}{4H^2}} = \omega^2. \end{split} \tag{1.28}$$