

1 Equations

1.1 Boussinesq Equations

The physical equations in the Boussinesq approximation are:

$$\vec{\nabla} \cdot \vec{u} = 0, \quad (1a)$$

$$\frac{\partial \rho_1}{\partial t} + (\vec{u} \cdot \vec{\nabla}) \rho_1 - \frac{\rho_0 u_z}{H} = 0, \quad (1b)$$

$$\frac{\partial \vec{u}}{\partial t} + (\vec{u} \cdot \vec{\nabla}) \vec{u} + \frac{\vec{\nabla} P_1}{\rho_0} + \frac{\rho_1 \vec{g}}{\rho_0} = 0. \quad (1c)$$

In these equations, we take $\rho_0(z) = \rho_0$ a constant reference density.

Including numerical terms and driving terms, my full equations are

$$\vec{\nabla} \cdot \vec{u}_1 = 0, \quad (2a)$$

$$\frac{\partial \rho_1}{\partial t} - \frac{\rho_0 u_z}{H} - \nu \nabla^6 \rho_1 = -\Gamma(z) \rho_1 - (\vec{u} \cdot \vec{\nabla}) \rho_1 + F e^{-\frac{(z-z_0)^2}{2\sigma^2}} \cos(k_x x - \omega t), \quad (2b)$$

$$\frac{\partial u_x}{\partial t} + \frac{\partial_x P}{\rho_0} - \nu \nabla^6 u_x = -\Gamma(z) u_x - (\vec{u} \cdot \vec{\nabla}) u_x, \quad (2c)$$

$$\frac{\partial u_z}{\partial t} + \frac{\partial_z P}{\rho_0} + \frac{\rho_1 g}{\rho_0} - \nu \nabla^6 u_z = -\Gamma(z) u_z - (\vec{u} \cdot \vec{\nabla}) u_z, \quad (2d)$$

$$\Gamma(z) = 5 \left[2 + \tanh \frac{z - z_T}{(z_{\max} - z_T)/2} + \tanh \frac{z_B - z}{z_B/2} \right], \quad (2e)$$

where $z_B = 0.07 z_{\max}$, $z_T = 0.93 z_{\max}$ are the boundaries of the damping zones, and my domain is $z \in [0, z_{\max}]$. The forcing term F is chosen to be weakly nonlinear, ω is chosen by inverting $\omega(k_x, k_z)$ dispersion relation for fixed $k_x = 2\pi/x_{\max}$ and some desired $k_z \ll H, z_{\max}$ (I've chosen $k_z \approx -\frac{2\pi}{z_{\max}/5}$ here), and $\sigma \lesssim \frac{1}{k_z}$ is used to excite a broad band of modes including the desired k_z mode.

1.2 Stratification

We simulate an incompressible, isothermal fluid, representative of degenerate matter in WD bulks. We assume a barotropic equation of state to simplify for the time being. The physical equations for an incompressible, barotropic fluid are:

$$\vec{\nabla} \cdot \vec{u} = 0, \quad (3a)$$

$$\frac{\partial \rho}{\partial t} + \vec{u} \cdot \vec{\nabla} \rho = 0, \quad (3b)$$

$$\frac{\partial \vec{u}}{\partial t} + (\vec{u} \cdot \vec{\nabla}) \vec{u} + \frac{\vec{\nabla} P}{\rho} + g \hat{z} = 0. \quad (3c)$$

We will introduce variables $T = P/\rho$. We then mandate ρ_0, T_0 backgrounds satisfy hydrostatic equilibrium $\vec{\nabla} T_0 + T_0 \vec{\nabla} \rho_0 + \vec{g} = 0$. Taking isothermal stratification, we find $T_0 = gH$. Making then substitution of variables $Y = \ln \rho - \ln \rho_0$ and $T_1 = T - T_0$ deviations from the background state, the exact

fluid equations in the new variables are:

$$\vec{\nabla} \cdot \vec{u} = 0, \quad (4a)$$

$$\frac{\partial Y}{\partial t} - \frac{u_z}{H} = 0, \quad (4b)$$

$$\frac{\partial u_x}{\partial t} + (\vec{u} \cdot \vec{\nabla}) u_x + \frac{\partial T}{\partial x} + gH \frac{\partial Y}{\partial x} + T_1 \frac{\partial Y}{\partial x} = 0, \quad (4c)$$

$$\frac{\partial u_z}{\partial t} + (\vec{u} \cdot \vec{\nabla}) u_z + \frac{\partial T}{\partial z} + gH \frac{\partial Y}{\partial z} + T_1 \frac{\partial Y}{\partial z} - \frac{T_1}{H} = 0. \quad (4d)$$

These are implemented as:

$$\vec{\nabla} \cdot \vec{u} = 0, \quad (5a)$$

$$\frac{\partial Y}{\partial t} - \frac{u_z}{H} - \nu \nabla^2 Y = -\Gamma(z)Y - (\vec{u} \cdot \vec{\nabla})Y + \frac{F}{\rho_0(z)} e^{-\frac{(z-z_0)^2}{2\sigma^2}} \cos(k_x x - \omega t), \quad (5b)$$

$$\frac{\partial u_x}{\partial t} + \frac{\partial T}{\partial x} + gH \frac{\partial Y}{\partial x} - \nu \nabla^2 u_x = -\Gamma(z)u_x - (\vec{u} \cdot \vec{\nabla})u_x - T_1 \frac{\partial Y}{\partial x}, \quad (5c)$$

$$\frac{\partial u_z}{\partial t} + \frac{\partial T}{\partial z} + gH \frac{\partial Y}{\partial z} - \frac{T_1}{H} - \nu \nabla^2 u_z = -\Gamma(z)u_z - (\vec{u} \cdot \vec{\nabla})u_z - T_1 \frac{\partial Y}{\partial z}. \quad (5d)$$

$\Gamma(z)$ is as before, but we use $k_z = -2\pi/H$ here, and ω in the forcing term accordingly.