Seismic Inversion

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Contents

About the Author xiii
Preface
Acknowledgments
Notation Convention
Abbreviations
Part I: Iterative Optimization Methods
Chapter 1: Introduction to Seismic Inversion
1.1 Notation31.2 Inverse problem41.3 Types of seismic inversion71.4 Inverse crimes81.5 Summary9Appendix 1A: Basics of exploration seismology9Seismic sources10Nonzero-offset seismic experiment11Reflection amplitudes11Seismic processing11Reflection imaging12Key difficulty with migration images12
Chapter 2: Introduction to Gradient Optimization
2.1 Mathematical definitions162.2 Gradient optimization, Taylor series, and Newton's method172.3 Geometric interpretation of the gradient and Hessian182.4 Eigenvalues of the Hessian determine shape of contours192.5 MATLAB examples of Newton's method192.6 Summary212.7 Exercises22
2.8 Computational labs

Cha	pter 3: Steepest-Descent Method	• •	23
3.1	Steepest-descent method		23
	3.1.1 Convergence rate		23
3.2	Step-length calculation		25
	3.2.1 Exact line search		
	3.2.2 Inexact Newton method and inexact line search		
	3.2.3 Numerical line search		
	3.2.4 2D plane minimization		
	Steepest-descent method and linear systems of equations		
	3.3.1 Regularized steepest descent		
	3.3.2 MATLAB steepest-descent code		
	3.3.3 Preconditioned steepest descent		
	Summary		
	Exercises		
	Computational labs		
	endix 3A: Levenberg-Marquardt regularization		
App	endix 3B: Choosing a value for the regularization parameter		31
Cha	pter 4: Conjugate-Gradient and Quasi-Newton Methods		33
	Conjugate-gradient method		
	4.1.1 Conjugate-gradient algorithm		
	4.1.2 Conjugate-gradient method and linear systems of equations		
	4.1.3 MATLAB conjugate-gradient code		
	4.1.4 Convergence rate		
	4.1.5 Preconditioned and regularized conjugate gradients		
	Quasi-Newton methods		
	Nonlinear functionals		
	What really works?		
4.5	Summary		40
	Exercises		
	Computational labs		
	endix 4A: Successive line-minimization methods		
	Successive-eigenvector and conjugate-direction methods		41
David	t II. Tuovoltima Toma avanku		
rarı	t II: Traveltime Tomography		
Cha	pter 5: Raypath Traveltime Tomography		45
	Perturbed traveltime integral		
	Raypath traveltime tomography		
	5.2.1 Normal equations		
	5.2.2 Poorly conditioned equations and regularization		
	5.2.3 Reweighted least squares		
	Iterative steepest-descent solution		
	Reflection tomography		
	Field-data example		
	*		
	Summary		
	Exercises		
	Computational labs		
	endix 5A: Eikonal equation derivation		
App	endix 5B: l_p misfit gradient		29

Chapter 6: Traveltime Tomography: Assessing model accuracy	61
6.1 Introduction	
6.2 Model covariance matrix	
6.2.1 Numerical estimation of the model covariance matrix	
6.3 Model resolution matrix	
6.4 Analytic model covariance matrices	
6.4.1 VSP transmission data	
6.4.2 VSP reflection data	
6.4.3 CMP reflection data	
6.4.4 Uncertainty principle	
6.6 Projection-slice theorem: Traveltime tomography	
6.7 Summary	
6.8 Exercises	
Appendix 6A: Null-space properties of L^TL	
rippendix of i. I tall space properties of E. E	70
Part III: Numerical Modeling	
Chapter 7: Traveltime Calculation by Solution of the Eikonal Equation	81
7.1 Finite-difference solution of the eikonal equation	
7.1 Finite-difference solution of the cikonal equation	
7.3 Exercises	
Appendix 7A: Efficient sorting of traveltimes	
Chapter 8: Numerical Solutions to the Wave Equation	85
8.1 Finite-difference method	85
8.1.1 Finite-difference approximation to the wave equation	85
8.1.2 Stability and accuracy analysis	
8.1.3 MATLAB code for FD solution of the acoustic wave equation	
8.2 Pseudospectral solution of the wave equation	
8.2.1 MATLAB code for pseudospectral solution of the acoustic wave equation	
8.2.2 Stability and accuracy analysis	
8.3 Spectral element solution of the wave equation	
8.4 Staggered-grid FD solution of the wave equation	
8.4.1 Staggered-grid FD of the first-order acoustic equations	
8.4.2 Staggered-grid FD of the first-order elastodynamic equations	
8.4.3 MATLAB code for staggered-grid FD of the first-order elastic equations	
8.5 Modeling in the oil and gas industry	
8.6 Summary	
8.7 Exercises	
8.8 Computational labs	
Appendix 8A: Absorbing boundary conditions	
8A.1 Sponge zone	
8A.2 PDE absorbing boundary conditions	
67.5 Hyond I DE ausoronig conditions	UU
Chapter 9: The Viscoacoustic Wave Equation	01
9.1 Introduction to linear viscoelasticity	01
9.2 Viscoacoustic wave equation	
9.3 Summary	05

9.4 Exercises	105 105
Part IV: Reflection Migration	
Chapter 10: Forward and Adjoint Modeling Using Green's Functions	109
10.1 Integral-equation forward modeling $10.1.1 \text{ Green's functions}$ $10.1.2 (\nabla_{\mathbf{x}'}^2 + k^2)^{-1} \text{ by Green's theorem}$ $10.1.3 \text{ Lippmann-Schwinger solution}$ $10.1.4 \text{ Neumann series solution}$ $10.1.5 \text{ Born approximation}$ $10.1.6 \text{ Matrix operator notation}$	109 110 111 112 112 114
10.2 Integral-equation adjoint modeling	
10.3 Summary 10.4 Exercises 10.5 Computational labs Appendix 10A: Causal and acausal Green's functions Appendix 10B: Generalized Green's theorem	119 120 120
Chapter 11: Reverse Time Migration	123
11.1 Introduction 11.2 General imaging algorithm 11.2.1 RTM = generalized diffraction-stack migration 11.2.2 Reverse time migration 11.3 Numerical examples of RTM 11.4 Practical implementation of RTM 11.5 Summary 11.6 Exercises 11.7 Computational labs Appendix 11A: MATLAB RTM code	123 125 125 127 129 132 132 134 134
Chapter 12: Wavepaths	
12.1 Traveltime wavepaths 12.1.1 Computing traveltime wavepaths 12.2 Pressure wavepaths 12.2.1 Computing pressure wavepaths 12.3 Summary 12.4 Exercises	138 139 140 140
Chapter 13: Generalized Diffraction-stack Migration and Filtering of Coherent Noise	143
Abstract	143 144 145 145 148 148
15.0 110HP110H DAULDIDO	101

		Contents	vii
	13.5.1 Directional filtering results		. 151
	13.5.2 Anti-aliasing filtering results		
13.6	Conclusions		. 155
	Acknowledgements		
	ences		
	ndix A Migration as a Pattern Matching Operation		
	ndix B Computation and Compression of the Migration Kernel		
Exerc	ises		. 158
Chap	ter 14: Resolution Limits for Wave Equation Imaging	· • • • • • • •	159
1	Abstract		. 159
	Introduction		
	14.1.1 Resolution limits for traveltime tomography		160
	14.1.2 Resolution limits for reflection imaging		161
14.2]	Born forward and adjoint modeling		163
	14.2.1 Born forward modeling		
	14.2.2 Born adjoint modeling		
	Model resolution function and FWI resolution limits		
	14.3.1 Model resolution equation: $\mathbf{m}^{mig} = \mathbf{L}^{\dagger} \mathbf{L} \mathbf{m} \dots \dots$		
	14.3.2 Wavelength imaging at the diffractor		
	Filling in the model spectrum with multiples		
	14.4.1 Lower wavenumber resolution with prism waves and free-surface multiples		
	Discussion and summary		
	owledgment		
	ndix A: Resolution properties of Fresnel volume in constant and layered media		
	ndix B: Resolution limits for imaging diving wave residuals		
	ndix C: Determinant of a Jacobian matrix		
	ences		
Part V	V: Least-Squares Migration		
	ter 15: Iterative Least-Squares Migration		177
_			
15.1	Least-squares migration theory		
15.2	Overdetermined and underdetermined iterative LSM		
15.3	Implementation of LSM		
	15.3.2 Diffraction-stack LSM		
15.4	Problems with LSM		
13.1	15.4.1 LSM sensitivity to velocity errors		
	15.4.2 Computational cost of LSM		
15.5	Numerical results		
	15.5.1 3D point-scatterer model		
	15.5.2 Partial compensation for poor source and receiver sampling		
	15.5.3 Poststack migration of Gulf of Mexico data		
	15.5.4 Prestack migration of Gulf of Mexico data		
15.6	LSM with a crosscorrelation objective function		
15.7	LSRTM with internal multiples		
15.8	Trim statics and LSM		
15.9	Artifact reduction with LSM		
	15.9.1 Numerical results		. 189

15.11 15.12 Apper	Summary	91 92 92 93
Chap	ter 16: Viscoacoustic Least-Squares Migration	97
16.2 I 16.3 S	Theory of viscoacoustic least-squares migration 19 16.1.1 Viscoacoustic Born modeling equations 19 16.1.2 Viscoacoustic adjoint equations 19 16.1.3 Viscoacoustic gradient 19 16.1.4 Algorithm for Q LSRTM 19 Numerical results 20 Summary 20 Computational labs 20	97 98 98 99 00
Chap	ter 17: Least-Squares Migration Filtering	03
17.2	Least-squares migration filtering 20 Numerical results 20 17.2.1 LSMF of PS and PP reflections for a Graben model 20 17.2.2 LSMF of Valhall data 20 Problems with LSMF 20 17.3.1 Encoded multisource LSMF 20 Summary 20	04 04 05 06 06
Chap	ter 18: Migration Deconvolution	11
18.2 18.3 18.4 18.4 18.5 18.6 18.6 1Appen	Migration Green's function2Approximations to $[\mathbf{L}^{\dagger}\mathbf{L}]^{-1}$ 218.2.1 Hessian inverse by $\delta_{ij}/[\mathbf{L}^{\dagger}\mathbf{L}]_{ij}$ 218.2.2 Hessian inverse by a nonstationary matching filter218.2.3 Migration deconvolution2Iterative migration deconvolution2Numerical tests218.4.1 Point-scatterer model218.4.2 Meandering-stream model218.4.3 Converted-wave marine field data218.4.4 3D Alaska field data2Summary2Exercises2ndix 18A: Numerical implementation of MD2	12 12 15 16 16 16 18 18 18 20
	VI: Waveform Inversion	
-	ter 19: Acoustic Waveform Inversion and its Numerical Implementation	
Exam Exam Exam 19.1	ple 19.0.1 Pseudolinear traveltime misfit function with quasimonotonic character 22 ple 19.0.2 Highly nonlinear waveform misfit function 23 ple 19.0.3 Mildly nonlinear waveform misfit function 25 ple 19.0.4 Finite-difference solution 26 Numerical implementation of waveform inversion 27 Expediting convergence 26 19.2.1 Conjugate-gradient and quasi-Newton gradient methods 27	25 26 27 27 29

	19.2.2 Starting model	. 229
	19.2.3 Preconditioning	
	19.2.4 Subspace decomposition	
	19.2.5 Adaptive multiscale FWI	
	19.2.6 Estimation of the source wavelet	
	19.2.7 Ignoring amplitudes	
19.3	Numerical tests	
	Summary	
	Exercises	
	Computational labs	
~		
	oter 20: Wave-Equation Inversion of Skeletonized Data	
	Implicit function theorem	
20.2	Examples of skeletonized inversion	. 236
	20.2.1 Wave-equation traveltime tomography	
	20.2.2 Early-arrival waveform tomography	. 238
	20.2.3 Wave-equation inversion of surface waves	
20.3	Alternative objective functions	. 244
	Summary	
	Exercises	
Appe	endix 20A: Gradient of the traveltime misfit function	. 248
Appe	endix 20B: Implementation of WT	. 250
	Forward modeling	. 250
	Backward propagation	. 250
	Direction of updating the model	. 250
	Calculation of the step length	. 251
Cha	oter 21: Acoustic Waveform Inversion: Case histories	253
21.1	Early-arrival waveform inversion applied to land data	
	21.1.1 Data acquisition	
	21.1.2 Data processing	
	21.1.3 Estimating and correcting for Q	
	21.1.4 EWT of the Wadi Qudaid data	
	21.1.5 Synthetic data sanity test	
21.2	21.1.6 Key points	
21.2	Acoustic FWI applied to Gulf of Mexico marine data	
	21.2.1 Hybrid linear and nonlinear FWI	
	21.2.2 Synthetic two-box model data	
	21.2.3 Gulf of Mexico data	
21.2	21.2.4 Key points	
21.3	Rolling-offset FWI	
	21.3.1 Workflow for rolling-offset FWI	
	21.3.2 Rolling-offset FWI: Synthetic data	
	21.3.3 Rolling-offset FWI: 2D Gulf of Mexico data	
	21.3.4 FWI with macro windows: 3D marine data	
	21.3.5 Key points	
21.4	Acoustic FWI applied to crosswell data	
	21.4.1 Synthetic crosshole data	
	21.4.2 Friendswood crosshole data	
	21.4.3 Key points	
21.5	Summary	. 275

ix

Contents

21.6 Exercises				
Chapter 22: Elastic and Viscoelastic Full-Waveform Inversion				
22.1 Elastic FWI	277			
22.2 FWI of crosswell hydrophone records				
22.2.1 Acoustic FWI applied to acoustic synthetic data				
22.2.2 Elastic and viscoelastic FWI applied to synthetic viscoelastic data				
22.3 FWI applied to McElroy crosswell data	. 281			
22.3.1 Data processing				
22.3.2 Elastic and viscoelastic waveform inversion	. 283			
22.4 Summary				
22.5 Exercises				
Appendix 22A: Misfit gradient for λ				
Appendix 22B: Source-wavelet inversion				
Appendix 22D: Borehole pressure-field simulation				
Appendix 22D: Estimation of Q_P and Q_S				
Appendix 22F: Elastic FWI gradient				
Appendix 221. Elastic I w I gradient	. 290			
Chapter 23: Vertical Transverse Isotropy FWI				
23.1 Theory	. 293			
23.2 Numerical results				
23.2.1 Synthetic VSP data				
23.2.2 3D Gulf of Mexico data				
23.3 Extension to TTI media				
23.4 Summary				
23.5 Exercises	. 299			
Part VII: Image-Domain Inversion				
Chapter 24: Classical Migration Velocity Analysis	. 303			
24.1 Image-domain inversion	. 303			
24.2 Classical ray-based MVA	. 303			
24.3 Angle-domain CIGs	. 306			
24.4 Trim statics MVA				
24.5 Ray-based tomography				
24.6 Summary	. 309			
Chapter 25: Generalized Differential Semblance Optimization	. 311			
25.1 Introduction	. 311			
25.2 Theory of generalized differential semblance optimization				
25.2.1 Wave-equation traveltime and waveform inversion	. 312			
25.2.2 Differential semblance optimization	. 312			
25.2.3 Generalized differential semblance optimization				
25.3 Numerical examples				
25.4 Summary				
25.5 Exercises				
Appendix 25A: Migration images in the subsurface offset domain				
Appendix 25B: H-DSO Fréchet derivative and gradient	320			

	Contents	хi
		222
Chapter 26: Generalized Image-Domain Inversion	• • • • • • •	323
26.1 Introduction		323
26.2 Theory of generalized image-domain inversion		324
26.2.1 Interpretation of the gradient functions		325
26.3 Numerical results		326
26.4 Summary		330
References		331
Index		345

About the Author

Gerard Schuster is currently a professor of geophysics at King Abdullah University Science and Technology (KAUST) and an adjunct professor at University of Utah and University of Wyoming. He was the founder and director of the Utah Tomography and Modeling/Migration consortium from 1987 to 2009 and is now the co-director and founder of the Center for Fluid Modeling and Seismic Imaging at KAUST. Schuster helped pioneer seismic interferometry and its practical applications in applied geophysics through his active research program and through his extensive publications. He also has extensive experience in developing innovative migration and inversion methods for both exploration and earthquake seismology.

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gil Kauffman Gold Medal in 2010 for his work in seismic interferometry. He was the SEG Distinguished Lecturer in 2013.

Preface

This book describes the theory and practice of inverting seismic data for the subsurface rock properties of the earth. The primary application is for inverting reflection and/or transmission data from engineering or exploration surveys, but the methods described also can be used for earthquake studies. I have written this book with the hope that it will be largely comprehensible to scientists and advanced students in engineering, earth sciences, and physics. It is desirable that the reader has some familiarity with certain aspects of numerical computation, such as finite-difference solutions to partial differential equations, numerical linear algebra, and the basic physics of wave propagation (e.g., Snell's law and ray tracing). For those not familiar with the terminology and methods of seismic exploration, a brief introduction is provided in the Appendix of Chapter 1. Computational labs

are provided for most of the chapters, and some field data labs are given as well.

MATLAB and Fortran labs at the end of some chapters are used to deepen the reader's understanding of the concepts and their implementation. Such exercises are introduced early and geophysical applications are presented in every chapter. For the non-geophysicist, geophysical concepts are introduced with intuitive arguments, and their description by rigorous theory is deferred to later chapters.

The lab exercises in the Computational Toolkit can be found at http://csim.kaust.edu.sa/web/SeismicInversion and http://utam.gg.utah.edu/SeismicInversion/; the exercises can be accessed using the login Paulina and the password Brozina.

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Pelletier, Mrinal Sinha, Ahmad Tarhini, Xin Wang, Han Yu, Ge Zhan, Zhendong Zhang, and Sanzong Zhang. The last two chapters on image-domain inversion are modified versions of Sanzong Zhang's dissertation, and the description of the viscoacoustic gradient in Chapter 15 is a modified version of an appendix in Gaurav Dutta's dissertation. The derivation of the VTI adjoint equations are from Mrinal Sinha and Bowen Guo. I also thank Yue Wang and Ken Bube for their derivation of the adjoint of the viscoelastic wave equation in Chapter 22. Bowen Guo also helped generalize the gradient equation for inverting trim statics with the wave equation. Zongcai Feng scrupulously checked the accuracy of the equations in Chapters 20 through 23.

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Notation Convention

- \mathbb{R}^N denotes the *N*-dimensional real vector space.
- \mathbb{C}^N denotes the N-dimensional complex vector space.
- A column vector will be denoted by boldface lower-case letters. For example, $\mathbf{x} = [x_1, x_2, ..., x_N]^T$ represents the $N \times 1$ vector where x_i is the i^{th} element of \mathbf{x} .
- A matrix will be denoted by boldface upper-case letters. For example, $\mathbf{A} \in \mathbb{R}^{M \times N}$ represents an $M \times N$ real matrix whose ij^{th} element is denoted by A_{ij} .
- An order-of-magnitude estimate of a variable whose precise value is unknown is an estimate rounded to the nearest power of ten.
- A scalar will be denoted by lower-case letters.
- Subscripts are usually used to denote the element index of a vector or matrix.
- Superscripts with parentheses are used to denote an iterate of a vector or matrix. For example, x^(k) denotes the kth iterate of an iterative scheme.
- $\mathbf{x} \cdot \mathbf{y} = \mathbf{x}^T \mathbf{y} = \langle \mathbf{x}, \mathbf{y} \rangle = \sum_{i=1}^N x_i^* y_i$ represents a dot product or an inner product between finite-dimensional vectors \mathbf{x} and \mathbf{y} .
- MATLAB syntax is sometimes used to represent vectors or matrices. For example, [a b; c d] denotes the matrix

$$\left[\begin{array}{cc} a & b \\ c & d \end{array}\right].$$

• $||\mathbf{x}||_1$ denotes the 1-norm of the $N \times 1$ vector \mathbf{x} which is equal to

$$||\mathbf{x}||_1 = \sum_{i=1}^N |x_i|.$$

 |x| = ||x||₂ denotes the 2-norm or Euclidean norm of the N × 1 vector x which is equal to

$$||\mathbf{x}||_2 = \sqrt{\sum_{i=1}^N x_i^2}.$$

If the subscript is missing then the 2-norm is indicated, l_2 for a discrete vector and L_2 for a well-behaved function of a continuous variable.

- The length of a vector \mathbf{x} will often be denoted as $|\mathbf{x}|$ rather than $||\mathbf{x}||_2$.
- $||\mathbf{x}||_p$ denotes the *p*-norm or Euclidean norm of the $N \times 1$ vector \mathbf{x} which is equal to

$$||\mathbf{x}||_p = \left(\sum_{i=1}^N x_i^p\right)^{\frac{1}{p}}.$$

- $\hat{\mathbf{x}} = \frac{\mathbf{x}}{|\mathbf{x}|}$ denotes the unit vector.
- A* denotes the complex conjugate of the matrix A.
- A^T denotes the transpose of matrix A. We will often insist it also means the transpose and complex conjugated matrix A.
- A^{\dagger} denotes the conjugated and transposed matrix A.
- * denotes temporal convolution. For example, assuming f(t) and g(t) are real continuous functions of the scalar variable t and are square integrable then

$$f(t) \star g(t) = \int_{-\infty}^{\infty} f(t - \tau)g(\tau)d\tau = \int_{-\infty}^{\infty} f(\tau)g(t - \tau)d\tau.$$
(1)

 \oint denotes temporal correlation. For example, assuming
 f(t) and g(t) are real continuous functions of the scalar
 variable t and are square integrable then

$$f(t) \otimes g(t) = f(-t) \star g(t) = \int_{-\infty}^{\infty} f(\tau)g(t+\tau)d\tau. \quad (2)$$

• $\mathcal{F}[f(t)] = F(\omega)$ denotes the Fourier transform of f(t) and $\mathcal{F}^{-1}[F(\omega)]$ is the inverse Fourier transform of $F(\omega)$. For example,

$$F(\omega) = \mathcal{F}[f(t)] = \int_{-\infty}^{\infty} f(t)e^{-i\omega t}dt,$$

$$f(t) = \mathcal{F}^{-1}[F(\omega)] = \frac{1}{2\pi} \int_{-\infty}^{\infty} F(\omega)e^{i\omega t}d\omega,$$

$$\frac{d^n f(t)}{dt^n} = \frac{1}{2\pi} \int_{-\infty}^{\infty} (i\omega)^n F(\omega)e^{i\omega t}d\omega,$$

$$f(-t) = \mathcal{F}^{-1}[F(\omega)^*],$$

$$\mathcal{F}[f(t) \star g(t)] = F(\omega)G(\omega),$$

$$\mathcal{F}[f(t) \otimes g(t)] = \mathcal{F}[f(-t) \star g(t)] = F(\omega)^* G(\omega),$$

$$f(t) \star g(t) = \frac{1}{2\pi} \int_{-\infty}^{\infty} F(\omega) G(\omega) e^{i\omega t} d\omega,$$

$$f(t) \otimes g(t) = \frac{1}{2\pi} \int_{-\infty}^{\infty} F(\omega)^* G(\omega) e^{i\omega t} d\omega,$$

$$f(t) \otimes g(t)|_{t=0} = \int_{-\infty}^{\infty} f(\tau) g(\tau) d\tau$$

$$= \frac{1}{2\pi} \int_{-\infty}^{\infty} F(\omega)^* G(\omega) d\omega,$$

$$f(t) \otimes f(t)|_{t=0} = \int_{-\infty}^{\infty} f(\tau)^2 d\tau = \frac{1}{2\pi} \int_{-\infty}^{\infty} |F(\omega)|^2 d\omega.$$
(3)

 The Dirac delta function δ(t) is a generalized function (Zemanian, 1965) that is zero everywhere on the real line, except at t = 0. The Dirac delta function has a broadband spectrum with the constant amplitude 1:

$$\delta(t - t') = \frac{1}{2\pi} \int_{-\infty}^{\infty} e^{i\omega(t - t')} d\omega. \tag{4}$$

For a smooth function $f(\tau)$, the delta function has the sifting property:

$$f(t) = \int_{-\infty}^{\infty} f(\tau)\delta(\tau - t)d\tau.$$
 (5)

Abbreviations

ABC	absorbing boundary condition	KM	Kirchhoff migration
ADCIG	angle-domain common image gather	LSM	least squares migration
CAG	common angle gather	LSRTM	least squares reverse time migration
CFL	Courant-Friedrichs-Lewy	MD	migration deconvolution
CG	conjugate gradient	MVA	migration velocity analysis
CIG	common image gather	NLCG	nonlinear conjugate gradient
CMG	common midpoint gather	NMO	normal moveout
COG	common offset gather	PDE	partial differential equation
CSG	common shot gather	PML	perfectly matched layer
DFP	Davidon-Fletcher-Powell	PSTM	prestack time migration
DM	diffraction-stack migration	QN	quasi-Newton
DOD	domain of dependence	RTM	reverse time migration
DSO	differential semblance optimization	SD	steepest descent
EWT	early arrival wave equation tomography	SE	spectral element
FD	finite difference	SLS	standard linear solid
FE	finite element	SPD	symmetric positive definite
FWI	full waveform inversion	SSP	surface seismic profile
GCV	generalized cross validation	SV	singular vector
GDM	generalized diffraction-stack migration	VSP	vertical seismic profile
GDSO	generalized differential semblance	WT	wave equation traveltime tomography
	optimization	WTW	wave equation traveltime and waveform
GIDI	generalized image domain inversion		tomography
GOM	Gulf of Mexico	ZO	zero offset
IDI	image domain inversion		

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