



## Seafloor Geodesy Lab 4



Ausarbeitung im Studiengang  
**Geodäsie und Geoinformatik**  
an der Universität Stuttgart

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# Chapter 1

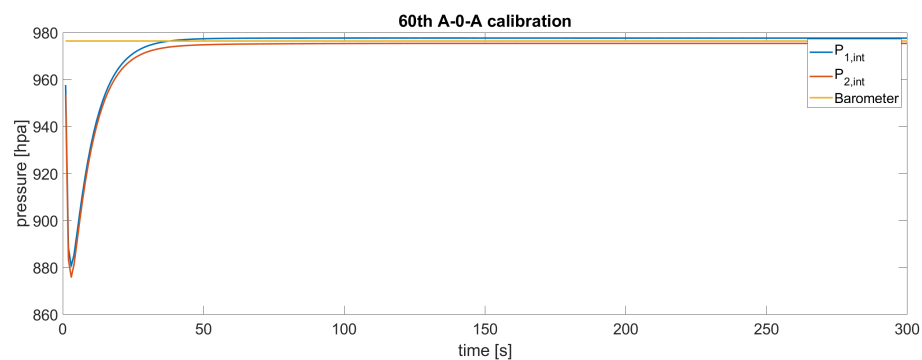
## Report

### 1.1 Introduction

The A-0-A method is a new approach for correcting drift in seafloor pressure measurement. In this work 2 A-0-A calibrated pressure gauges and a three-component accelerometer were set on the MARS cabled observatory in Monterey Bay and 30-months data were collected.

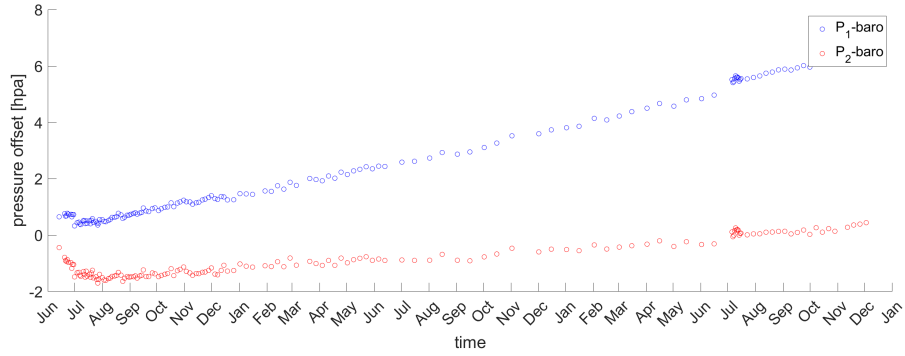
### 1.2 A-0-A Correction

A total of 154 A-0-A calibrations were performed with each lasting 5 minutes, an Example is shown in Figure 1.1, The measured pressure drop initially overshoots to pressure lower than atmospheric because of short term viscoelastic transients and because the adiabatic expansion of oil leads to cooling that takes some time to equilibrate. After removing the first and last 30 seconds, the mean value of difference. In this example,  $mean(P_{1,int} - baro) = 1.14$  hpa and  $mean(P_{2,int} - baro) = -1.19$  hpa



*Figure 1.1: Example A-0-A Calibration*

The results of all calibrations are in Figure 1.2

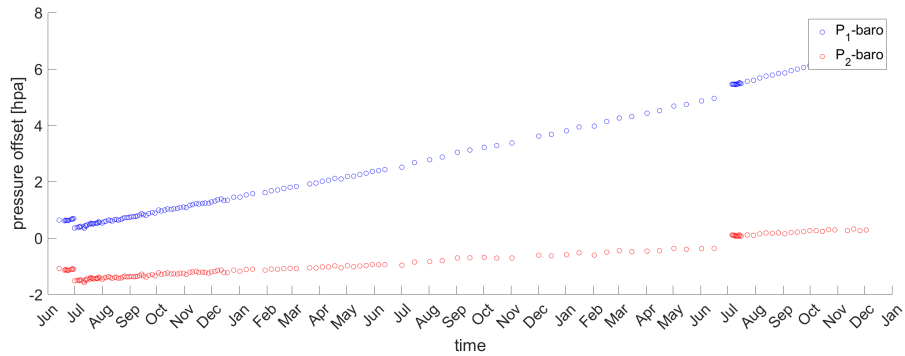


**Figure 1.2:** Result A-0-A Calibration

This series follow a form of function

$$p(t) = a \cdot \exp(-t/t_0) + bt + c + d(T - T_{ref}) \quad (1.1)$$

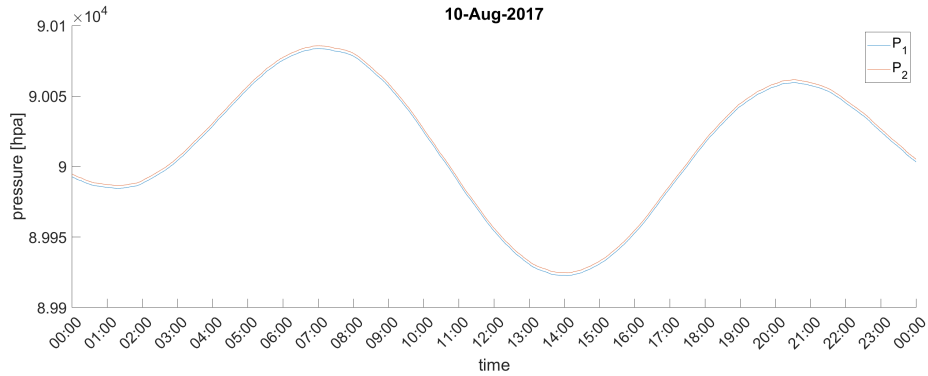
Where  $a, b, c, d$  and  $t_0$  are the parameters.  $t_0$  is a time constant since the deployment which gives the minimum RMS. After a iteration,  $t_0 = 351$  day for gauge 1 and  $t_0 = 759$  day for gauge 2. A Least square will be implemented to get 2 parameters groups for gauge 1 and 2.  $T$  is the Temperature and  $T_{ref}$  the reference temperature which is here the median temperature since the deployment. A smoothing line is in Figure 1.3. Then, this calibration can be used for the whole datasets.



**Figure 1.3:** calibration model

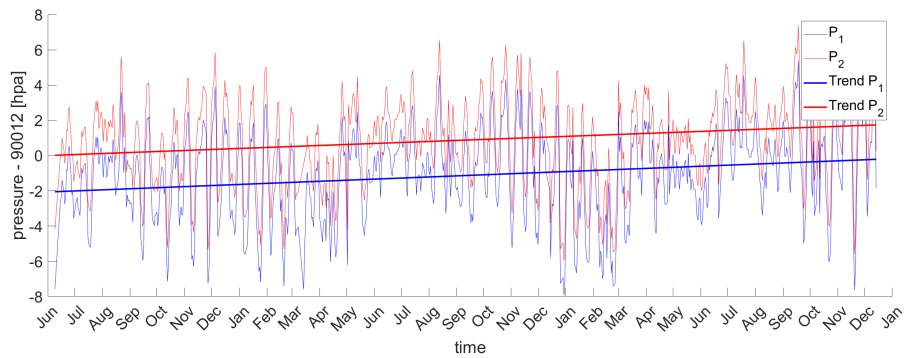
## 1.3 Data

The measurements took place every 1 minutes, Figure 1.4 shows a part of result in 10.Aug.2017. It is obvious that the pressure varies due to diurnal and semi-diurnal tides, but the difference between them remains stable.



*Figure 1.4: Pressure in 10.Aug.2017*

Therefor, the daily mean pressure will be calculated. We noticed that there are several gaps in the whole time series, if this gap is too big and can't be interpolated, the data of this day shall not be used. In the end, the results are shown in Figure 1.5.



*Figure 1.5: Pressure*

The Trend of two time series is 0.732 hpa/year for gauge 1 and 0.686 hpa/year. The difference between two time series are very stable after calibration.

## 1.4 Tilt

The platform tilt has been not taken into consideration till now, Figure 1.6 shows the change in accelerator measurements over deployment, and so the tilt can be calculated (Figure 1.7).

$$\text{tilt} = \frac{a_{\text{horizontal}}}{a_{\text{vertical}}} \quad (1.2)$$

$$= \frac{\sqrt{a_y^2 + a_z^2}}{a_x} \quad (1.3)$$

By Analysing the shape of the platform Figure 1.8, we can estimate a subsidence of one leg by nearly 1 cm of the observed tilt. To remove this effect, we remove external pressure over the experiment by 0.3 hpa.

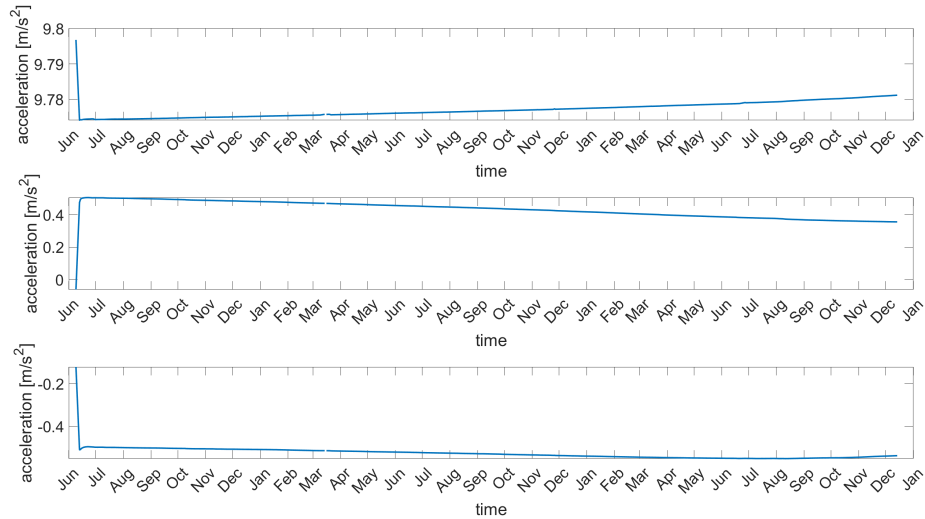


Figure 1.6: Acceleration

## 1.5 Another Effects

According to the paper, the tide influences should be excluded from the pressure measurements. This was done in the paper using low-pass Godin filter, which remains unclear. Therefore, this step was not performed in this work.

There are other error sources, which is the combined effects of changes in oil reservoir level, counting clock drift and the minimum subsidence. The correction is set to 0.1 hpa.

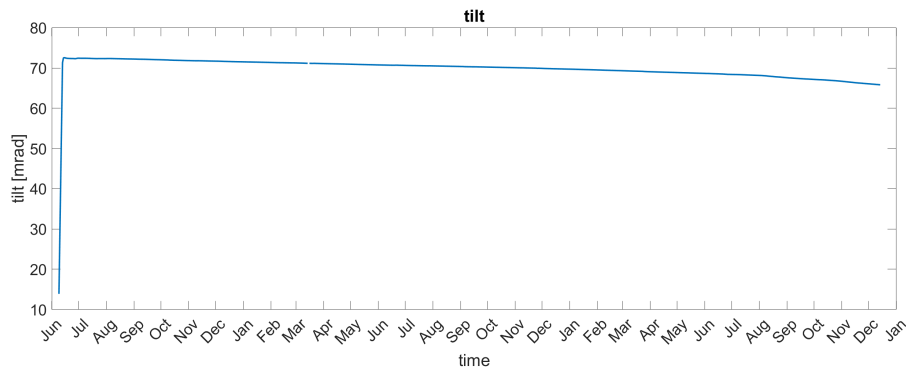


Figure 1.7: Tilt

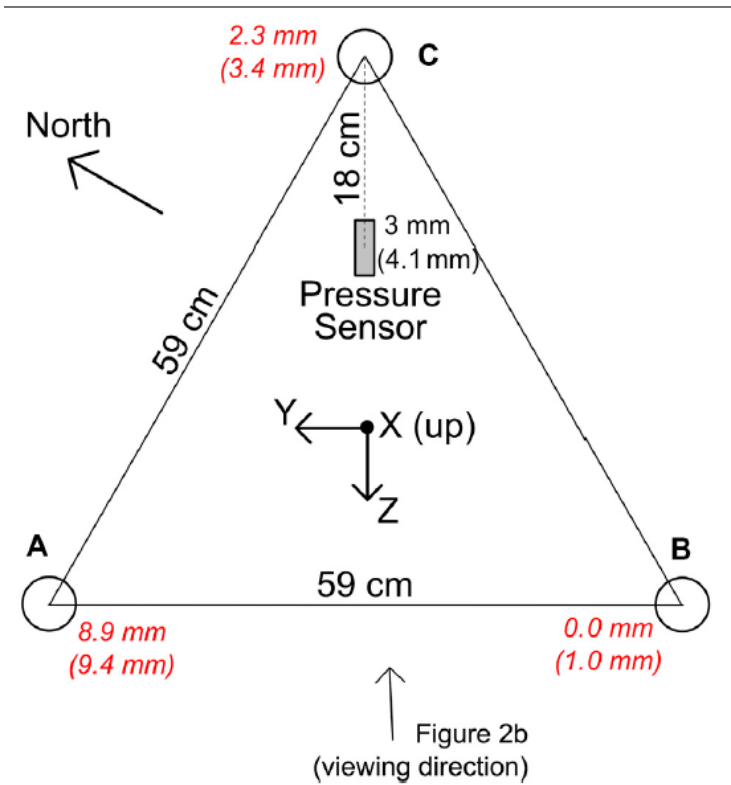


Figure 1.8: Platform

## 1.6 Discussion

This trend of increasing pressure are consistent with the trend with the trend expected from the El Niño Southern Oscillation. This effect is not a result of density structure but is related partly to variations in local wind stress moving water toward or away from the coast and partly due to the effects of coast trapped waves generated by wind stresses elsewhere. Ryan and Noble (2002). The first winter of the GSSM deployment coincided with weak La Niña event from October 2017-March 2018 while the latter parts of the deployment coincided with weak El Niño conditions from October 2018-June 2019 and November-December 2019.

## Bibliography

Ryan, H. and Noble, M. (2002), 'Sea level response to enso along the central california coast: how the 1997–1998 event compares with the historic record', (1-4), 149–169.