

METEOROLOGY

→ PA = ELV - 27 x (QNHlocal + 1013)

→ TA = PA + 27(QNHlocal - QNH)

- T.A. = I.A. + 27 x (QNH - 1013) + (0,004 x ISA dev x PA) →
- TA = IA + Xc → Xc = ISA correction x IA ISA correction = 4% each 10°C

→ ISAtem = 15 - [(Height/1000)x2]

→ QNH = QFE + (FE ÷ 27)

→ QFE = QNH - (FE ÷ 27)

→ IND ALT = TRUE ALT - (4 x thickness/1000 x ISA Dev)

→ FL (If I have HPA) = 1660 - 550 x LOG(hpa)

1)

Considering the North Atlantic route from the Azores to Bermuda, the mean height of the tropical tropopause during summer is approximately..

Tropopause Height = (16 x Cos(lat)) x 3280

- Summer +3000
- Winter -3000

→ IF QNH is higher than 1013 → TA(+) is HIGHER than IA(-)

→ IF QNH is lower than 1013 ← TA(-) is LOWER than IA(+)

→ Temperature COLDER than ISA = TA will be LOWER than IA

WIND

1) Ballot's.

→ With your back UPPER wind, the cold air is to your left (northern hemisphere)

High on the left = Head wind

Low on the left = tail wind

→ SH with your back upper wind the cold air is to your right....

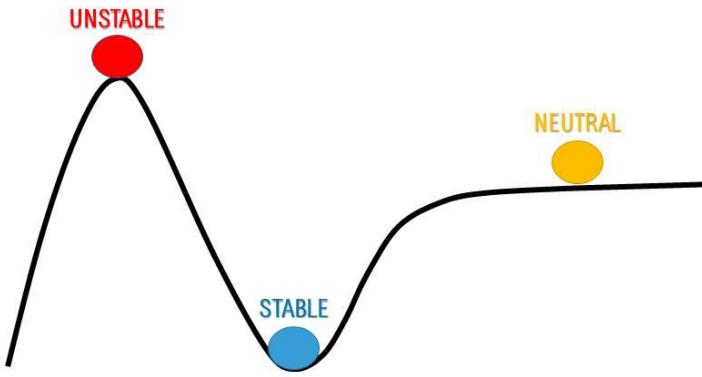
2) Geostrophic wind depends:

- Density
- Earth's Rotation
- Geographic latitude

UNSTABLE
T parcel > T air
The parcel is warmer than its surroundings, so it continues to rise and expand.

STABLE
T parcel < T air
The parcel is cooler than its surroundings, so it sinks and compresses

NEUTRAL
T parcel = T air
The parcel is the same temperature as its surroundings, no change



1) RH(Relative Humidity)=100-5(OAT-DW)

2) When the increasing the Air Temperature, the relative humidity will decrease

3) $T_d = T - (100 - RH)/5$ = Relative humidity given the Temperature and dewpoint