

物理海洋学笔记

2017 级海洋科学专业崔英哲

<https://github.com/Cuiyingzhe/UUC-Physical-Oceanography-Notes>

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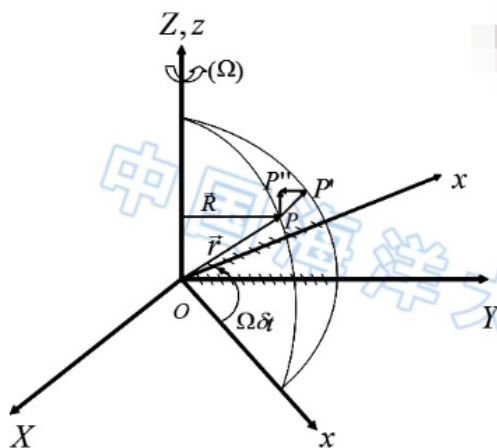
1 基本方程

1.1 旋转坐标系的速度和加速度

惯性坐标系: 静止的或是匀速直线运动的坐标系, 固定在恒星上的坐标系可以被看成惯性坐标系.

固定在地球上的坐标系: 地球对恒星的加速度主要是由地球自转引起的, 于是可以把地球当作一个对惯性坐标系作纯粹地转运动的物体.

1.1.1 旋转坐标系和惯性坐标系中的速度



惯性坐标系 (XYZ) 绝对位移: $p\vec{p}'' = \vec{V}_a \delta t$, \vec{V}_a 为绝对速度

旋转坐标系 (xyz) 相对位移: $p'\vec{p}'' = \vec{V} \delta t$, \vec{V} 为相对速度

$$\because p\vec{p}'' = p'\vec{p}'' + p\vec{p}'$$

$$\therefore \vec{V}_a \delta t = \vec{V} \delta t + \vec{V}_e \delta t \Rightarrow \vec{V}_a = \vec{V} + \vec{V}_e \text{ (绝对速度等于相对速度与牵连速度的向量和)}$$

$$\text{其中, } \vec{V}_e = \vec{\Omega} \times \vec{r} \Rightarrow \vec{V}_a = \vec{V} + \vec{\Omega} \times \vec{r}$$

$$\Rightarrow \frac{d_a \vec{r}}{dt} = \frac{d\vec{r}}{dt} + \vec{\Omega} \times \vec{r}$$

$$\frac{d_a \vec{A}}{dt} = \frac{d\vec{A}}{dt} + \vec{\Omega} \times \vec{A}$$

1.1.2 旋转坐标系和惯性坐标系中的加速度

$$\text{令 } \vec{A} = \vec{V}_a = \vec{V}_e + \vec{V} = \vec{V} + \vec{\Omega} \times \vec{r}$$

$$\begin{aligned} \frac{d\vec{V}_a}{dt} &= \frac{d_a}{dt} (\vec{V} + \vec{V}_e) = \frac{d_a}{dt} (\vec{V} + \vec{\Omega} \times \vec{r}) \\ &= \frac{d}{dt} (\vec{V} + \vec{\Omega} \times \vec{r}) + \vec{\Omega} \times (\vec{V} + \vec{\Omega} \times \vec{r}) \\ &= \frac{d\vec{V}}{dt} + \vec{\Omega} \times \vec{V} + \vec{\Omega} \times \vec{V} + \vec{\Omega} \times (\vec{\Omega} \times \vec{r}) \\ &= \frac{d\vec{V}}{dt} + 2\vec{\Omega} \times \vec{V} - \Omega^2 \vec{R} \end{aligned}$$

1.2 作用在海水微团上的外力运动方程的向量形式

压强梯度力: $\frac{1}{\rho} \nabla p$

分子粘性力 (摩擦力):

$$\begin{cases} F_x = \frac{1}{\rho} \mu \left(\frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2} + \frac{\partial^2 u}{\partial z^2} \right) = \frac{\mu}{\rho} \Delta u \\ F_y = \frac{1}{\rho} \mu \left(\frac{\partial^2 v}{\partial x^2} + \frac{\partial^2 v}{\partial y^2} + \frac{\partial^2 v}{\partial z^2} \right) = \frac{\mu}{\rho} \Delta v \\ F_z = \frac{1}{\rho} \mu \left(\frac{\partial^2 w}{\partial x^2} + \frac{\partial^2 w}{\partial y^2} + \frac{\partial^2 w}{\partial z^2} \right) = \frac{\mu}{\rho} \Delta w \end{cases} \Rightarrow \vec{F} = \frac{\mu}{\rho} \Delta \vec{V} = \gamma \Delta \vec{v}$$

重力 (地球引力与地球自转产生的惯性离心力的合力): $\vec{g} = -G \frac{M_g}{r^2} \cdot \left(\frac{\vec{r}}{r} \right)$

科氏力: $-2\vec{\Omega} \times \vec{V}$

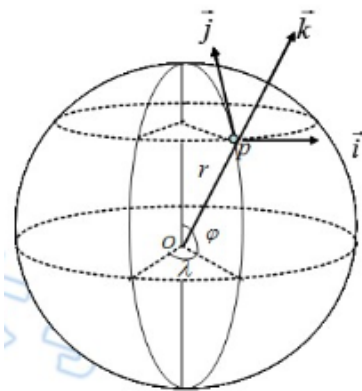
天体引潮力 (受其他天体万有引力与惯性力离心力的合力): $\vec{F}_M = -G \frac{M_M}{L^2} + G \frac{M_M}{D^2} \cdot \left(\frac{\vec{D}}{D} \right)$

由牛顿第二定律和坐标系转换关系:

$$\begin{cases} \frac{d_a \vec{V}_a}{dt} = \sum_i \vec{F}_i \\ \frac{d_a \vec{A}_a}{dt} = \frac{d\vec{V}}{dt} + 2\vec{\Omega} \times \vec{V} - \Omega^2 \vec{R} \end{cases}$$

$$\Rightarrow \frac{d\vec{V}}{dt} = -\frac{1}{\rho} \nabla P - 2\vec{\Omega} \times \vec{V} + \vec{g} + \nu \Delta \vec{V} + \vec{F}_T$$

1.3 运动方程在球坐标系的标量形式



速度:

$$\vec{V} = u\vec{i} + v\vec{j} + w\vec{k}$$

$$\Rightarrow \begin{cases} u = r \cos \varphi \frac{d\lambda}{dt} \\ v = r \frac{d\varphi}{dt} \\ w = \frac{dr}{dt} \end{cases}$$

加速度:

$$\begin{aligned} \frac{d\vec{A}}{dt} &= \frac{\frac{\partial \vec{A}}{\partial t} dt + \frac{\partial \vec{A}}{\partial \lambda} d\lambda + \frac{\partial \vec{A}}{\partial \varphi} d\varphi + \frac{\partial \vec{A}}{\partial r} dr}{dt} \\ &= \frac{\partial \vec{A}}{\partial t} + \frac{\partial \vec{A}}{\partial \lambda} \frac{d\lambda}{dt} + \frac{\partial \vec{A}}{\partial \varphi} \frac{d\varphi}{dt} + \frac{\partial \vec{A}}{\partial r} \frac{dr}{dt} \\ &= \frac{\partial \vec{A}}{\partial t} + \frac{u}{r \cos \varphi} \frac{\partial \vec{A}}{\partial \lambda} + \frac{v}{r} \frac{\partial \vec{A}}{\partial \varphi} + w \frac{\partial \vec{A}}{\partial r} \\ \Rightarrow \frac{d}{dt} &= \frac{\partial}{\partial t} + u \frac{\partial}{r \cos \varphi \partial \lambda} + v \frac{\partial}{r \partial \varphi} + w \frac{\partial}{\partial r} \\ \Rightarrow \frac{d}{dt} &= \frac{\partial}{\partial t} + (\vec{V} \cdot \nabla) \\ \Rightarrow \nabla &= \frac{\partial}{r \cos \varphi \partial \lambda} \vec{i} + \frac{\partial}{r \partial \varphi} \vec{j} + \frac{\partial}{\partial r} \vec{k} \\ \Rightarrow \begin{cases} \frac{d\vec{i}}{dt} = \frac{\partial \vec{i}}{\partial t} + u \frac{\partial \vec{i}}{r \cos \varphi \partial \lambda} + v \frac{\partial \vec{i}}{r \partial \varphi} + w \frac{\partial \vec{i}}{\partial r} \\ \frac{d\vec{j}}{dt} = \frac{\partial \vec{j}}{\partial t} + u \frac{\partial \vec{j}}{r \cos \varphi \partial \lambda} + v \frac{\partial \vec{j}}{r \partial \varphi} + w \frac{\partial \vec{j}}{\partial r} \\ \frac{d\vec{k}}{dt} = \frac{\partial \vec{k}}{\partial t} + u \frac{\partial \vec{k}}{r \cos \varphi \partial \lambda} + v \frac{\partial \vec{k}}{r \partial \varphi} + w \frac{\partial \vec{k}}{\partial r} \end{cases} \end{aligned}$$

$$\frac{d\vec{V}}{dt} = \frac{du}{dt}\vec{i} + \frac{dv}{dt}\vec{j} + \frac{dw}{dt}\vec{k} + u\frac{d\vec{i}}{dt} + v\frac{d\vec{j}}{dt} + w\frac{d\vec{k}}{dt}$$

$$\Rightarrow \frac{d\vec{V}}{dt} = \left(\frac{du}{dt} - \frac{uv \tan \varphi}{r} + \frac{uw}{r}\right)\vec{i} + \left(\frac{dv}{dt} + \frac{u^2 \tan \varphi}{r} + \frac{vw}{r}\right)\vec{j} + \left(\frac{dw}{dt} - \frac{u^2 + v^2}{r}\right)\vec{k}$$

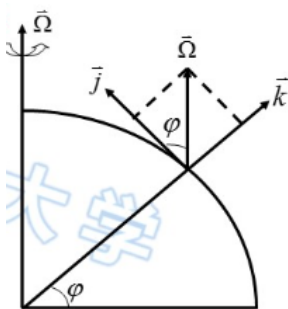
压强梯度力:

$$\frac{1}{\rho} \nabla p = -\frac{1}{\rho} \left(\frac{1}{r \cos \varphi} \frac{\partial p}{\partial \lambda} \vec{i} + \frac{1}{r} \frac{\partial p}{\partial \varphi} \vec{j} + \frac{\partial p}{\partial r} \vec{k} \right)$$

重力:

$$\vec{g} = -g\vec{k}$$

科氏力:



$$\vec{\Omega} = \Omega \sin \varphi \vec{k} + \Omega \cos \varphi \vec{j}$$

$$-2\vec{\Omega} \times \vec{V} = -2 \begin{vmatrix} \vec{i} & \vec{j} & \vec{k} \\ 0 & \Omega \cos \varphi & \Omega \sin \varphi \\ u & v & w \end{vmatrix}$$

$$= -2[(w\Omega \cos \varphi - v\Omega \sin \varphi)\vec{i} + (u\Omega \sin \varphi)\vec{j} + (-u\Omega \cos \varphi)\vec{k}]$$

$$\Rightarrow -2\vec{\Omega} \times \vec{V} = (fv - \tilde{f}w)\vec{i} - (fu)\vec{j} + (\tilde{f}u)\vec{k}$$

其中, $\begin{cases} f = 2\Omega \sin \varphi \\ \tilde{f} = 2\Omega \cos \varphi \end{cases}$

$$\Rightarrow \begin{cases} \frac{du}{dt} = -\frac{1}{\rho} \frac{\partial p}{r \cos \varphi \partial \lambda} + fv - \tilde{f}w + \frac{uv \tan \varphi}{r} - \frac{uw}{r} + \gamma(\Delta \vec{v})_\lambda - \frac{1}{r \cos \varphi} \frac{\partial \phi_T}{\partial \lambda} \\ \frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial p}{r \partial \varphi} - fu - \frac{u^2 \tan \varphi}{r} - \frac{vw}{r} + \gamma(\Delta \vec{v})_\varphi - \frac{1}{r} \frac{\partial \phi_T}{\partial \varphi} \\ \frac{dw}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial r} + \tilde{f}u = g + \frac{u^2 + v^2}{r} + \gamma(\Delta \vec{v})_r - \frac{\partial \phi_T}{\partial r} \end{cases}$$

1.4 直角坐标系的运动方程

略去地球曲率的影响

$$\Rightarrow \begin{cases} \frac{du}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial x} + fv + \tilde{f}w + F_{N\lambda} + F_{T\lambda} \\ \frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial y} - fu + F_{Ny} + F_{Ty} \\ \frac{dw}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial z} + \tilde{f}u - g + F_{Nz} + F_{Tz} \end{cases}$$

1.5 海水层流运动的基本方程组

1.5.1 连续方程

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \vec{V}) = 0 \Leftrightarrow \frac{d\rho}{dt} + \rho \nabla \cdot \vec{V} = 0$$

特别地, 对于不可压缩流体:

$$\nabla \cdot \vec{V} = 0$$

1.5.2 盐量扩散方程

$$\begin{aligned}
& \begin{array}{ccc} \text{盐量增加量} & \text{平流作用} & \text{分子扩散作用} \\ \frac{\partial}{\partial t} \iiint_{\tau} \rho s d\tau & = - \oiint_{\sigma} \rho s V_n d\sigma + & - \oiint_{\sigma} S_n d\sigma \end{array} \\
& \iiint_{\tau} \frac{\partial(\rho s)}{\partial t} d\tau = \iiint_{\tau} \nabla \cdot (\rho s \vec{V}) d\tau - \iiint_{\tau} \nabla \cdot \vec{S} d\tau \\
& \Rightarrow \frac{\partial(\rho s)}{\partial t} + \nabla \cdot (\rho s \vec{V}) + \nabla \cdot \vec{S} = 0 \\
& \Rightarrow \rho \frac{\partial s}{\partial t} + s \frac{\partial \rho}{\partial t} + s \nabla \cdot (\rho \vec{V}) + \rho \vec{V} \cdot \nabla s + \nabla \cdot \vec{S} = 0 \\
& \Rightarrow \left(\frac{\partial s}{\partial t} + \vec{V} \cdot \nabla s \right) + \frac{s}{\rho} \left[\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \vec{V}) \right] = -\frac{1}{\rho} \nabla \cdot \vec{S} \\
& \Rightarrow \frac{\partial s}{\partial t} + \vec{V} \cdot \nabla s = \frac{k}{\rho} \Delta s = k_D \Delta s
\end{aligned}$$

其中, $k_D = \frac{k}{\rho} \sim 1.1 \times 10^{-9} \text{ (m}^2/\text{s)}$

1.5.3 热传导方程

与上面类似:

$$\frac{\partial \theta}{\partial t} + \vec{V} \cdot \nabla \theta = \frac{\kappa}{\rho c_p} \Delta \theta = k_{\theta} \Delta \theta$$

其中, $k_{\theta} = \frac{\kappa}{\rho c_p} \sim 1.4 \times 10^{-7} \text{ (m}^2/\text{s)}$

1.5.4 热膨胀方程-状态方程

热膨胀方程:

$$\rho = \rho_0 \left(1 - \frac{\theta}{k} \right)$$

0°C 时的海水密度 海水的膨胀系数

EOS80 国际海水状态方程:

$$\rho(s, t, p) = \rho(s, t, 0) \left[1 - \frac{np}{k(s, t, p)} \right]^{-1}$$

1.6 基本方程的矢量形式和标量形式

矢量形式:

$$\left\{ \begin{array}{l} \frac{d\vec{V}}{dt} = -\frac{1}{\rho} \nabla p - 2\Omega \times \vec{V} + \vec{g} + \gamma \Delta \vec{V} - \nabla \phi_T \\ \frac{\partial \rho}{\partial t} + \rho \nabla \cdot \vec{V} = 0 \\ \frac{\partial s}{\partial t} + \vec{V} \cdot \nabla s = k_D \Delta s \\ \frac{\partial \theta}{\partial t} + \vec{V} \cdot \nabla \theta = k_{\theta} \Delta \theta \\ \rho = \rho(\theta, s, p) \end{array} \right.$$

标量形式 (直角坐标系):

$$\left\{ \begin{array}{l} \frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial y} - fu + \gamma \Delta v - \frac{\partial \phi_T}{\partial y} \\ \frac{dw}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial z} + \tilde{f}u - g + \gamma \Delta w - \frac{\partial \phi_T}{\partial z} \\ \frac{d\rho}{dt} + \rho \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} \right) = 0 \\ \frac{\partial s}{\partial t} + u \frac{\partial s}{\partial x} + v \frac{\partial s}{\partial y} + w \frac{\partial s}{\partial z} = k_D \left(\frac{\partial^2 s}{\partial x^2} + \frac{\partial^2 s}{\partial y^2} + \frac{\partial^2 s}{\partial z^2} \right) \\ \frac{\partial \theta}{\partial t} + u \frac{\partial \theta}{\partial x} + v \frac{\partial \theta}{\partial y} + w \frac{\partial \theta}{\partial z} = k_{\theta} \left(\frac{\partial^2 \theta}{\partial x^2} + \frac{\partial^2 \theta}{\partial y^2} + \frac{\partial^2 \theta}{\partial z^2} \right) \\ \rho = \rho(\theta, s, p) \end{array} \right.$$

1.7 边界条件

无质量交换的运动学边界条件：

$$\frac{\partial F}{\partial t} + \vec{c} \cdot \nabla F = 0$$

例：

$$(1) \text{ 海面 } (z = \zeta(x, y, t)): \frac{\partial \zeta}{\partial t} + \vec{V}_H \cdot \nabla_H \zeta - w = 0$$

$$(2) \text{ 海底 } (z = -h(x, y)): \vec{V}_H \cdot \nabla_H h + w = 0$$

动力学边界条件：

由牛顿第三定律，在界面法线方向有：

$$(\vec{p}_n)_1 = (\vec{p}_n)_2$$

1.8 * 时间平均的基本方程和边界条件 (直角坐标系)

连续方程：

$$\frac{\partial \bar{u}}{\partial x} + \frac{\partial \bar{v}}{\partial y} + \frac{\partial \bar{w}}{\partial z} = 0$$

运动方程：

$$\begin{cases} \frac{\partial \bar{u}}{\partial t} + \bar{u} \frac{\partial \bar{u}}{\partial x} + \bar{v} \frac{\partial \bar{u}}{\partial y} + \bar{w} \frac{\partial \bar{u}}{\partial z} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial x} + f\bar{v} - \bar{f}w + \gamma \Delta \bar{u} - \frac{\partial \bar{\phi}_T}{\partial x} + \frac{\partial}{\partial x} \left(A_{x\alpha} \frac{\partial \bar{u}}{\partial x} \right) + \frac{\partial}{\partial y} \left(A_{xy} \frac{\partial \bar{u}}{\partial y} \right) + \frac{\partial}{\partial z} \left(A_{xz} \frac{\partial \bar{u}}{\partial z} \right) \\ \frac{\partial \bar{v}}{\partial t} + \bar{u} \frac{\partial \bar{v}}{\partial x} + \bar{v} \frac{\partial \bar{v}}{\partial y} + \bar{w} \frac{\partial \bar{v}}{\partial z} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial y} - f\bar{u} + \gamma \Delta \bar{v} - \frac{\partial \bar{\phi}_T}{\partial y} + \frac{\partial}{\partial x} \left(A_{yx} \frac{\partial \bar{v}}{\partial x} \right) + \frac{\partial}{\partial y} \left(A_{yy} \frac{\partial \bar{v}}{\partial y} \right) + \frac{\partial}{\partial z} \left(A_{yz} \frac{\partial \bar{v}}{\partial z} \right) \\ \frac{\partial \bar{w}}{\partial t} + \bar{u} \frac{\partial \bar{w}}{\partial x} + \bar{v} \frac{\partial \bar{w}}{\partial y} + \bar{w} \frac{\partial \bar{w}}{\partial z} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial z} + \bar{f}u - g + \gamma \Delta \bar{w} - \frac{\partial \bar{\phi}_T}{\partial z} + \frac{\partial}{\partial x} \left(A_{2x} \frac{\partial \bar{w}}{\partial x} \right) + \frac{\partial}{\partial y} \left(A_{zy} \frac{\partial \bar{w}}{\partial y} \right) + \frac{\partial}{\partial z} \left(A_{zz} \frac{\partial \bar{w}}{\partial z} \right) \end{cases}$$

盐量扩散方程：

$$\frac{\partial \bar{s}}{\partial t} + \bar{u} \frac{\partial \bar{s}}{\partial x} + \bar{v} \frac{\partial \bar{s}}{\partial y} + \bar{w} \frac{\partial \bar{s}}{\partial z}$$

热传导方程：

$$\frac{\partial \bar{\theta}}{\partial t} + \bar{u} \frac{\partial \bar{\theta}}{\partial x} + \bar{v} \frac{\partial \bar{\theta}}{\partial y} + \bar{w} \frac{\partial \bar{\theta}}{\partial z} = k_\theta \Delta \bar{\theta} + \frac{\partial}{\partial x} \left(K_{\theta x} \frac{\partial \bar{\theta}}{\partial x} \right) + \frac{\partial}{\partial y} \left(K_{\theta y} \frac{\partial \bar{\theta}}{\partial y} \right) + \frac{\partial}{\partial z} \left(K_{\theta z} \frac{\partial \bar{\theta}}{\partial z} \right)$$

状态方程：

$$\bar{\rho} = \bar{\rho}(\bar{s}, \bar{\theta}, \bar{p})$$

1.9 铅直向平均的基本方程

$$\begin{aligned} \frac{\partial}{\partial x} [(h + \zeta) \langle u \rangle] + \frac{\partial}{\partial y} [(h + \zeta) \langle v \rangle] - \left[u|_\zeta \frac{\partial \zeta}{\partial x} + v|_\zeta \frac{\partial \zeta}{\partial y} - w|_\zeta \right] - \left[u|_{-h} \frac{\partial h}{\partial x} + v|_{-h} \frac{\partial h}{\partial y} + w|_{-h} \right] = 0 \\ \frac{\partial \zeta}{\partial t} + \frac{\partial [(h + \zeta) \langle u \rangle]}{\partial x} + \frac{\partial [(h + \zeta) \langle v \rangle]}{\partial y} = 0 \end{aligned}$$

1.10 尺度分析

$$\text{Rossby 数 } \text{Ro} = \frac{U}{FL} \begin{cases} \gg 1: \text{平流非线性项比 Coriolis 力重要, 大尺度运动} \\ = 1: \text{平流非线性项与 Coriolis 力同等重要} \\ \ll 1: \text{平流非线性项可以忽略, 小尺度运动} \end{cases}$$

$$\text{水平 Ekman 数 } E_l = \frac{A_l}{FL^2} \text{ 水平湍流摩擦项与 Coriolis 力比值}$$

$$\text{垂直 Ekman 数 } E_z = \frac{A_z}{FD^2} \text{ 垂直湍流摩擦项与 Coriolis 力比值}$$

准静力近似 f 平面近似 β 平面近似 Boussinesq 近似

2 海流

2.1 地转流

地转流：不考虑海面风的作用，远离沿岸的大洋中部大尺度、准水平、定常的海水流动。

产生原因：海水受热力和动力因素导致压力 (和密度) 在水平方向分布不均匀。

$$p = p_a + \rho gh \quad \rho \begin{cases} \neq \rho_0 \Rightarrow \text{梯度流} \\ = \rho_0 \Rightarrow \text{倾斜流} \end{cases}$$

2.1.1 梯度流

假定和方程

(1) 在相当长一段时间里海面温度变化和降水蒸发变化都不大，于是可以认为已形成的海水密度场、温度场和盐度场近似于定常，从而相应的海水运动也近似于定常： $\frac{\partial}{\partial t} = 0$ 。

(2) 海洋深而宽广，在远离海岸及海底的大洋中部海区，大尺度运动： $Ro \ll 1$ 。

(3) 不考虑海底摩擦及边界摩擦的影响，且海面无风力作用，则流动属一种无摩擦流动： $E_l, E_z \ll 1$ 。

(4) β 平面近似准静力近似

x 方向基本方程：

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial x} + fv + \frac{\partial}{\partial x} \left(A_{xx} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(A_{xy} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(A_{xz} \frac{\partial u}{\partial z} \right)$$

$$\text{假定 (1)} \Rightarrow \frac{\partial u}{\partial t} = 0$$

$$\text{假定 (2)} \Rightarrow u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = 0$$

$$\text{假定 (3)} \Rightarrow \frac{\partial}{\partial x} \left(A_{xx} \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left(A_{xy} \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left(A_{xz} \frac{\partial u}{\partial z} \right) = 0$$

可得梯度流的控制方程：

$$\begin{cases} 0 = -\frac{1}{\rho} \frac{\partial p}{\partial x} + fv \\ 0 = -\frac{1}{\rho} \frac{\partial p}{\partial y} - fu \\ 0 = -\frac{1}{\rho} \frac{\partial p}{\partial z} - g \\ \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0 \\ u \frac{\partial \theta}{\partial x} + v \frac{\partial \theta}{\partial y} + w \frac{\partial \theta}{\partial z} = 0 \\ u \frac{\partial s}{\partial x} + v \frac{\partial s}{\partial y} + w \frac{\partial s}{\partial z} = 0 \\ \rho = \rho(s, \theta) \end{cases}$$

特征

水平速度：

$$\begin{cases} 0 = -\frac{1}{\rho} \frac{\partial p}{\partial x} + fv \\ 0 = -\frac{1}{\rho} \frac{\partial p}{\partial y} - fu \end{cases} \quad (1)$$

$$\Rightarrow \begin{cases} v = \frac{1}{\rho f} \frac{\partial p}{\partial x} \\ u = -\frac{1}{\rho f} \frac{\partial p}{\partial y} \end{cases} \quad (2)$$

- (1) 水平速度和压强梯度成正比；
- (2) 与密度和科氏参数成反比；
- (3) 地转关系在赤道不成立 ($f = 0$).

垂向速度：

$$\begin{aligned}
 \frac{\partial(1) \times \rho}{\partial y} - \frac{\partial(2) \times \rho}{\partial x} &\Leftrightarrow \frac{\partial(\rho f v)}{\partial v} + \frac{\partial(\rho f u)}{\partial x} = 0 \\
 &\Leftrightarrow f \rho \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} \right) + f \left(u \frac{\partial \rho}{\partial x} + v \frac{\partial \rho}{\partial y} + w \frac{\partial \rho}{\partial z} \right) - f \rho \frac{\partial w}{\partial z} - f w \frac{\partial \rho}{\partial z} + \beta \rho v = 0 \\
 &\Leftrightarrow f \rho \stackrel{=0}{\boxed{\nabla \vec{V}}} + f \vec{V} \cdot \nabla \rho - f \rho \frac{\partial w}{\partial z} - f w \frac{\partial \rho}{\partial z} + \beta \rho v = 0
 \end{aligned} \tag{3}$$

$$\vec{V} \cdot \nabla \rho = \vec{V} \cdot \nabla \rho(s, \theta) = \vec{V} \cdot \left(\nabla_s \frac{\partial \rho}{\partial s} + \nabla_\theta \frac{\partial \rho}{\partial \theta} \right) = 0$$

$$(3) \Leftrightarrow f \left(\rho \frac{\partial w}{\partial z} + w \frac{\partial \rho}{\partial z} \right) = \beta \rho v \stackrel{\text{尺度分析}}{\Rightarrow} f \frac{\partial w}{\partial z} = \beta v \stackrel{\text{尺度分析}}{\Rightarrow} W = \frac{\beta D}{F} U \sim 2 \times 10^{-4} U$$

垂向流速比水平流速小得多，地转流为准水平运动.

运动特性

$$(1) \times u + (2) \times v \Leftrightarrow u \frac{\partial p}{\partial x} + v \frac{\partial p}{\partial y} = 0 \Leftrightarrow \vec{V}_H \cdot \nabla_H p = 0$$

- (1) 梯度流平行于等压线；
- (2) 北半球，流向右侧为高压，南半球相反；

密度特性

$$\frac{\partial(1) \times \rho}{\partial y} - \frac{\partial(2) \times \rho}{\partial x} \Leftrightarrow f \rho \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) + f \left(u \frac{\partial \rho}{\partial x} + v \frac{\partial \rho}{\partial y} \right) + \rho u \frac{\partial f}{\partial x} + \rho v \frac{\partial f}{\partial y} = 0 \Leftrightarrow \vec{V}_H \cdot \nabla_H \rho = 0$$

- (1) 梯度流近似平行于等密线；
- (2) 在北半球，流向右侧密度小；
- (3) 等压面倾斜与等密面倾斜方向相反.

温盐特性

忽略垂向运动：

$$\vec{V}_H \cdot \nabla_H \theta = 0$$

$$\vec{V}_H \cdot \nabla_H s = 0$$

- (1) 梯度流平行于等温线和等盐线；
- (2) 在北半球，流向右侧温度高，盐度低.

2.1.2 倾斜流

假定和方程 (1) 海水密度为常数；

(2) 水平方向的压强梯度是由海面倾斜引起的.

$$\Rightarrow p = p_a + \int_z^\zeta \rho g dz = p_a + \rho g(\zeta - z)$$

倾斜流的控制方程：

$$\begin{cases} f v = g \frac{\partial \zeta}{\partial x} \\ f u = -g \frac{\partial \zeta}{\partial y} \end{cases} \tag{4}$$

$$\tag{5}$$

性质：

$$(4) \times u - (5) \times v \Leftrightarrow u \frac{\partial \zeta}{\partial x} + v \frac{\partial \zeta}{\partial y} = 0 \Leftrightarrow \vec{V}_H \cdot \nabla \zeta = 0$$

- (1) 倾斜流平行于等水位线；
- (2) 在北半球，流向右侧水位高；
- (3) 倾斜流从表至底流速流向相同，压强梯度相同。

2.2 Ekman 漂流

由恒速定常的风长时间驱动大尺度、均匀密度的海洋，所产生的处于稳定状态的海流。

2.2.1 无限深海漂流

物理背景

Ekman 的老师 Nansen 在海洋调查时发现，冰山不是顺风漂移，而是沿着风向右方 $20^\circ \sim 40^\circ$ 的方向移动。Ekman 在 1905 年研究了这种现象并提出风海流理论 [Ekman, 1905]。

假定

无限深海 Ekman 漂流中用到了以下假定：

- 1) 海洋无限广阔，海洋无限深。

即无侧边界效应，仅有垂直湍流所生水平湍流摩擦力，并假定垂直湍流粘滞系数 A_z 为常量。由于海洋无限深， $z \rightarrow \infty, \vec{V} = 0$

- 2) 定常均匀风场长时间作用。

即运动的基本参量与时间和水平坐标无关且海面无升降、无水平压强梯度。

- 3) 密度分布均匀， ρ 为常数，不考虑热盐性质。

- 4) 采用 f 平面近似。

方程推导

控制方程和边界条件

首先给出一般的控制方程：

$$\begin{cases} \frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial x} + fv + A_l \left(\frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2} \right) + A_z \frac{\partial^2 u}{\partial z^2} \\ \frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial y} - fu + A_l \left(\frac{\partial^2 v}{\partial x^2} + \frac{\partial^2 v}{\partial y^2} \right) + A_z \frac{\partial^2 v}{\partial z^2} \\ \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0 \end{cases} \quad (6)$$

由假定 1)， $A_l \left(\frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2} \right) = 0, A_l \left(\frac{\partial^2 v}{\partial x^2} + \frac{\partial^2 v}{\partial y^2} \right) = 0$ ；

由假定 2)， $\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = 0, \frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} = 0$ ；

由假定 3)， $-\frac{1}{\rho} \frac{\partial p}{\partial x} = 0, -\frac{1}{\rho} \frac{\partial p}{\partial y} = 0$

则 (6) 化为：

$$\begin{cases} 0 = fv + A_z \frac{\partial^2 u}{\partial z^2} \end{cases} \quad (7a)$$

$$\begin{cases} 0 = -fu + A_z \frac{\partial^2 v}{\partial z^2} \end{cases} \quad (7b)$$

$$\begin{cases} \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0 \end{cases} \quad (7c)$$

不失一般性地，假定风力仅沿 y 方向作用，即 $\tau_x = 0, \tau_y = \text{const.}$ 再结合假定 1)，控制方程的边界条件为：

$$\begin{cases} z = 0, \rho A_z \frac{\partial u}{\partial z} = 0 & (8a) \end{cases}$$

$$\begin{cases} z = 0, \rho A_z \frac{\partial v}{\partial z} = -\tau_y & (8b) \end{cases}$$

$$\begin{cases} z = \infty, u = v = 0 & (8c) \end{cases}$$

方程求解

$$(7a) + (7b) \times i \Leftrightarrow A_z \frac{\partial^2(u + iv)}{\partial z^2} = if(u + iv)$$

令 $W = u + iv$ ，得：

$$A_z \frac{\partial^2 W}{\partial z^2} = ifW \Rightarrow \frac{\partial^2 W}{\partial z^2} = \frac{(1+i)^2 \Omega \sin \varphi}{A_z} W$$

令 $a = \sqrt{\Omega \sin \varphi / A_z}$ ， $j^2 = (1+i)^2 a^2$ ，得：

$$\frac{d^2 W}{dz^2} - j^2 W = 0 \quad (9)$$

(9) 式通解为： $W = Ae^{jz} + Be^{-jz}$

结合边界条件： $(8a) + (8b) \times i \Rightarrow z = 0, \rho A_z \frac{\partial W}{\partial z} = -\tau_y, z \rightarrow \infty, W = 0$

$$z \rightarrow \infty \Rightarrow A = 0, W = Be^{-jz}; z = 0, \rho A_z \frac{\partial W}{\partial z} \Big|_{z=0} = \rho A_z \frac{\partial (Be^{-jz})}{\partial z} \Big|_{z=0} = -\tau_y, \Rightarrow B = \tau_y / (j\rho A_z)$$

因此，方程的解为：

$$W = \frac{\tau_y}{j\rho A_z} e^{-jz} = \frac{i\tau_y}{(1+i)a\rho A_z} e^{-(1+i)az} = \frac{e^{i\frac{\pi}{2}}\tau_y}{\sqrt{2}e^{i\frac{\pi}{4}}a\rho A_z} e^{-(1+i)az}$$

令 $D_0 = \pi/a$ ，得到最终解的形式为：

$$W = \frac{\tau_y}{\sqrt{2}a\rho A_z} e^{-\frac{\pi}{D_0}z + i(\frac{\pi}{4} - \frac{\pi}{D_0}z)} \quad (10)$$

物理性质

运动速度

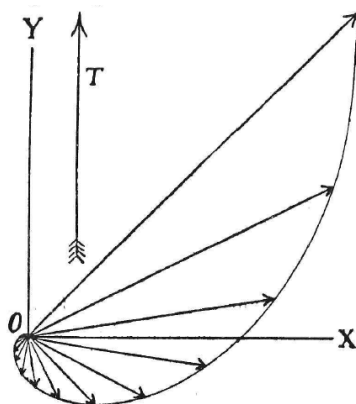
在海面 ($z = 0$) 处， $W_0 = \frac{\tau_y}{\sqrt{2}a\rho A_z} e^{i\frac{\pi}{4}}$ 。大小为 $|W_0| = \frac{\tau_y}{\sqrt{2}a\rho A_z}$ ，方向与 x 轴成 45° ，即与风向向右偏 45° 。

在任意深度处， $|W_z| = \frac{\tau_y}{\sqrt{2}a\rho A_z} e^{-\frac{\pi}{D_0}z}$ ，方向为 $\frac{\pi}{4} - \frac{\pi}{D_0}z$ ，即流速随深度增加呈指数形式减小，流向随深度的增加而逐渐向右偏。

在摩擦深度 $z = D_0$ 处， $|W_{D_0}| = \frac{\tau_y e^{-\pi}}{\sqrt{2}a\rho A_z} = e^{-\pi}|W_0| = 0.043|W_0|$ ，方向 $-\frac{3}{4}\pi$ ，即与 x 轴成 -135° ，与表面流向正好相反。

Ekman 螺旋和 Ekman 螺线

根据速度的垂向分布，表层流速最大，流向偏向风向的右方 45° ；随深度增加，流速逐渐减小，流向逐渐右偏；到摩擦深度，流速是表面流速的 4.3%，流向与表面流向相反，运动可以忽略。连连接各层流速的矢量端点，构成 Ekman 螺旋；Ekman 螺旋在平面上的投影，称为 Ekman 螺线 [Ekman, 1905]。



水平体积输运

体积输运：

$$\begin{aligned}
 S &= \int_0^\infty W dz \\
 &= \frac{\tau_y}{\sqrt{2}a\rho A_z} e^{i\frac{\pi}{4}} \int_0^\infty e^{-\frac{\pi}{D_0}(1+i)z} dz \\
 &= \frac{\tau_y}{\sqrt{2}a\rho A_z} e^{i\frac{\pi}{4}} \left[-\frac{D_0}{\pi} \frac{1}{(1+i)} \right] e^{-\frac{\pi}{D_0}(1+i)z} \Big|_0^\infty \\
 &= \frac{\tau_y}{2\Omega \sin \varphi \rho} = \frac{\tau_y}{f\rho}
 \end{aligned}$$

可以发现，得到的输运结果只有实部，没有虚部，说明体积输运方向为 x 轴正向，即在北半球水体向风向右侧 90° 输运.

2.2.2 有限深海漂流

假定

有限深海 Ekman 漂流中用到了以下假定：

1) 海区无限广阔、有限深，远离海岸.

即无侧边界效应，仅有垂直湍流所生水平湍流摩擦力，并假定垂直湍流粘滞系数 A_z 为常量. 由于海洋有限深， $z \rightarrow h, \vec{V} = 0$

2) 定常均匀风场长时间作用.

即运动的基本参量与时间和水平坐标无关且海面无升降、无水平压强梯度.

3) 密度分布均匀， ρ 为常数，不考虑热盐性质.

4) 采用 f 平面近似.

控制方程和边界条件：

$$\begin{cases}
 0 = fv + A_z \frac{\partial^2 u}{\partial z^2} \\
 0 = -fu + A_z \frac{\partial^2 v}{\partial z^2} \\
 \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0 \\
 z = 0 : \rho A_z \frac{\partial u}{\partial z} = 0, \rho A_z \frac{\partial v}{\partial z} = -\tau_y \\
 z \rightarrow h : u = v = 0
 \end{cases}$$

方程求解 令 $\xi = h - z$ ，定解问题化为：

$$\begin{cases}
 -fv = A_z \frac{\partial^2 u}{\partial \xi^2} \\
 fu = A_z \frac{\partial^2 v}{\partial \xi^2}
 \end{cases} \quad (11)$$

$$\begin{cases}
 \xi = h : \rho A_z \frac{\partial u}{\partial \xi} = 0, \rho A_z \frac{\partial v}{\partial \xi} = \tau_y \\
 \xi \rightarrow 0 : u = v = 0
 \end{cases} \quad (12)$$

令 $W = u + iv, \tau = \tau_x + i\tau_y$ ，控制方程：

$$(11) + (12) \times i \Leftrightarrow \frac{d^2 W}{d\xi^2} - j^2 W = 0$$

边界条件：

$$\begin{aligned}
 \xi = h : \rho A_z \frac{\partial W}{\partial \xi} &= \tau \\
 \xi = 0 : W &= 0
 \end{aligned}$$

与无限深海漂流解法类似，解得：

$$W = \frac{(1+i)\tau_y}{2a\rho A_z} \frac{sh(1+i)a\xi}{ch(1+i)ah}$$

物理性质

与水深的关系

(1) $h \geq 2D_0$ 时, 有限深海漂流流速流向与无限深海相同; (2) 水深越浅, 流向随深度增加右偏 (北半球) 越缓慢; (3) 从上层到下层的流速矢量越是趋近风矢量的方向.

体积输运

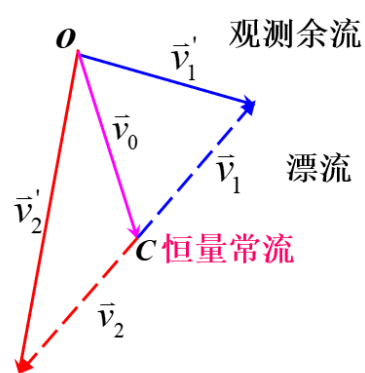
- (1) 在 x, y 方向 (平行和垂直风向) 都有输送;
- (2) 运输方向为风向右端, $\pm 90^\circ$ 之间:

$$S_x > 0; 0 < h < D_0, ah < \pi, S_y > 0; D_0 < h < 2D_0, \pi < ah < 2\pi, S_y < 0; h > 2D_0, S_y = 0$$

2.2.3 漂流分离

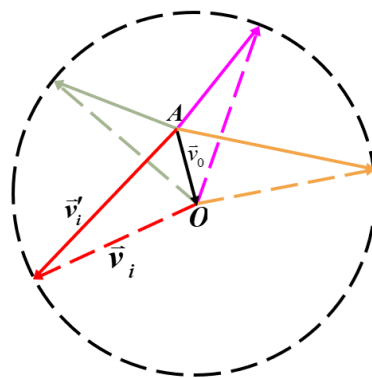
利用风速大小相等、方向相反的两组观测余流分离漂流

* 余流 = 漂流 + 恒量常流



利用一组风速大小相等、方向不同的实测余流分离漂流

* 漂流速度矢量端点落在同一圆周上

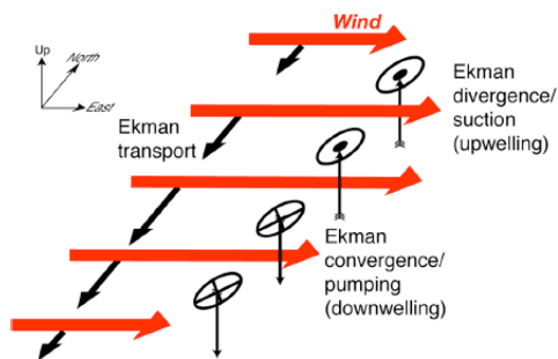


2.2.4 升降流

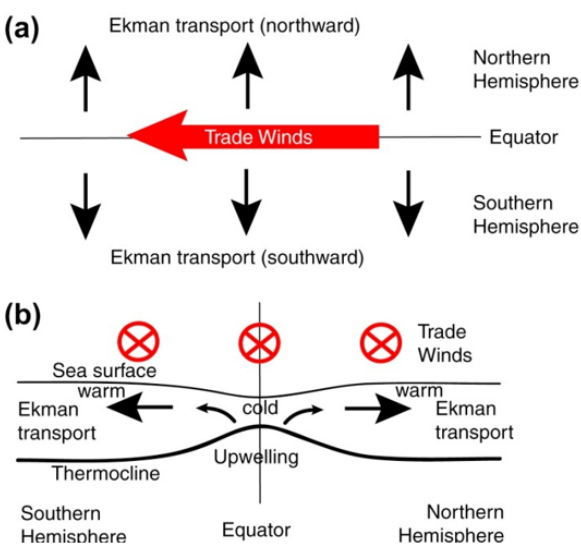
由不均匀风场或风场和地形配合产生的“较强烈”的铅直向流动。

物理背景

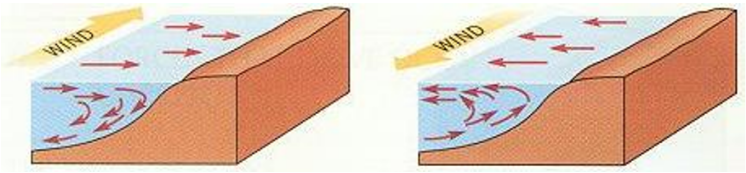
非均匀风场 \Rightarrow 非均匀 Ekman 漂流 \Rightarrow 非均匀体积输运 \Rightarrow 辐聚辐散 \Rightarrow 升降流



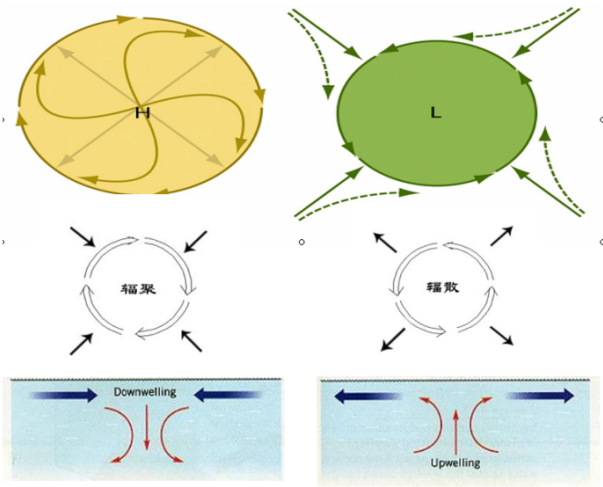
赤道附近的升降流



顺 (沿) 岸风产生的升降流



气旋和反气旋产生的海洋升降流



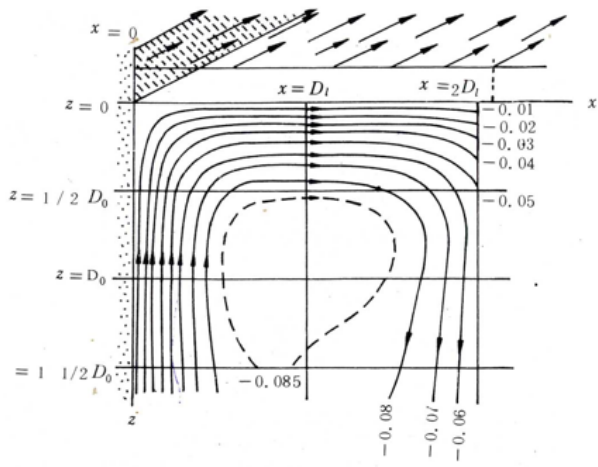
假定

- (1) ρ 为常数;
- (2) 直线风系, 风仅沿 x 方向有变化; 风区内为恒定的均匀风场; 风区外无风; $\frac{\partial}{\partial y} = 0$
- (3) 定常风场; $\frac{\partial}{\partial t} = 0$
- (4) 大尺度; $Ro \ll 1$
- (5) 有限深度. $h \geq 2D_0$

控制方程及边界条件

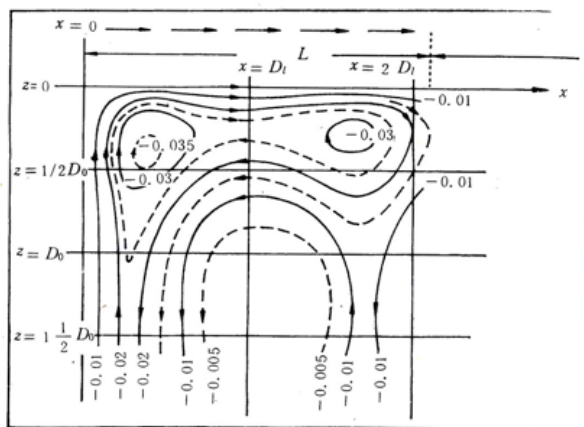
$$\begin{cases} A_l \frac{\partial^2 u}{\partial x^2} + A_z \frac{\partial^2 u}{\partial z^2} + f v + g \frac{\partial \zeta}{\partial x} = 0 \\ A_l \frac{\partial^2 v}{\partial x^2} + A_z \frac{\partial^2 v}{\partial z^2} - f u = 0 \\ \frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0 \end{cases}$$

结果讨论



- (1) 近岸产生上升流 $x \leq 0.5D_l$;
- (2) 风区外延附近下降流 $x = 2D_l$;
- (3) 上升流来自 $z = 1.5D_0$ 或更深;
- (4) 最大 w 出现在 $z = D_0$;
- (5) 上层离岸流, 下层向岸流, 构成一个循环.

若风向与海岸成 θ 角:



- (1) 三个升降流系统：两个顺时针，一个逆时针；
- (2) 大顺时针循环；
- (3) $\theta = 21.5^\circ$ 时，升降流达最大强度；
- (4) 纬度越低，升降流越强.

2.3 非定常运动

2.3.1 漂流的发展

假定

- (1) 远离海岸和海底的开阔大洋；
- (2) 风场均匀恒定；
- (3) ρ 为常数；
- (4) 海面无倾斜；
- (5) 运动非定常。

控制方程

$$\begin{cases} \frac{\partial u}{\partial t} - fv = A_z \frac{\partial^2 u}{\partial z^2} \\ \frac{\partial v}{\partial t} + fu = A_z \frac{\partial^2 v}{\partial z^2} \end{cases}$$

初边值条件

$$\begin{cases} z = 0 : A_z \frac{\partial u}{\partial z} = 0, \rho A_z \frac{\partial v}{\partial z} = -\tau_y (t > 0) \\ z \rightarrow \infty : u = v = 0 \\ t = 0 : u = v = 0 / u = C_1, v = C_2 \end{cases}$$

解的讨论

$$\begin{cases} u = \frac{2\pi\tau_y}{\rho f D_0} \int_0^{t'} \frac{\sin(2\pi\eta)}{\sqrt{\eta}} e^{\frac{\pi z^2}{4D_0^2}} d\eta \\ v = \frac{2\pi\tau_y}{\rho f D_0} \int_0^{t'} \frac{\cos(2\pi\eta)}{\sqrt{\eta}} e^{\frac{\pi z^2}{4D_0^2}} d\eta \end{cases}$$

根据下图 [Ekman, 1905]:

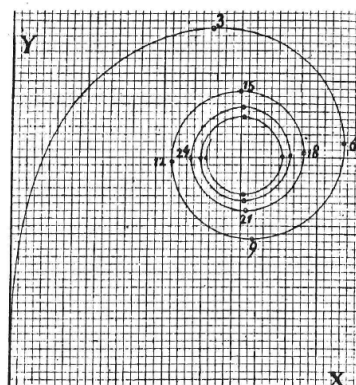


Fig. 3. $z=0$.

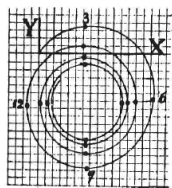


Fig. 4. $z=0,5 D$.

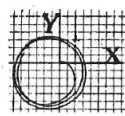


Fig. 5. $z=D$.



Fig. 6. $z=2D$.

随时间增加，空间某点流苏端点顺时针旋转（北半球），逐渐趋向一个极限值（即漂流）。

2.3.2 惯性流

假定和方程

- (1) 风场消失或者流动离开风区；
- (2) 外部驱动小时，湍摩擦失去作用；
- (3) 流动转为由惯性项维持平衡；

(4) 强制流动转变为自由流动.

$$\begin{cases} \frac{du}{dt} - fv = 0 \\ \frac{dv}{dt} + fu = 0 \end{cases} \quad (13)$$

$$\quad (14)$$

求解

$$\begin{aligned} (13) \times u + (14) \times v &\Leftrightarrow u \frac{du}{dt} + v \frac{dv}{dt} = \frac{1}{2} \frac{du^2}{dt} + \frac{1}{2} \frac{dv^2}{dt} = 0 \\ &\Leftrightarrow \frac{d}{dt} (u^2 + v^2) = 0 \\ &\Leftrightarrow u^2 + v^2 = c = V_0^2 \\ (13), (14) &\Rightarrow \begin{cases} v = \frac{dy}{dt} = \frac{1}{f} \frac{du}{dt} \\ u = \frac{dx}{dt} = \frac{1}{f} \frac{dv}{dt} \end{cases} \\ &\Rightarrow \begin{cases} y - y' = \frac{1}{f} (u - u') \\ x - x' = -\frac{1}{f} (v - v') \end{cases} \\ &\Rightarrow \begin{cases} y - \left(y' - \frac{u'}{f} \right) = \frac{u}{f} \\ x - \left(x' + \frac{v'}{f} \right) = -\frac{v}{f} \end{cases} \\ &\Rightarrow \boxed{(y - y_0)^2 + (x - x_0)^2 = \frac{1}{f^2} (u^2 + v^2) = \frac{V^2}{f^2} = r^2} \end{aligned}$$

流体质点沿半径为 r 的圆周作匀速运动，这个圆称之为惯性圆，对应的流动为惯性流.

惯性圆半径

$$\text{科氏力充当向心力: } \frac{V_0^2}{r} = fV_0 \Rightarrow V_0 = fr \Rightarrow r = \frac{V_0}{f} = \frac{V_0}{2\omega \sin \varphi}$$

随纬度增加而减小；赤道 $r \rightarrow \infty$ ，水质点作直线运动.

周期

$$T_i = \frac{2\pi r}{V_0} = \frac{2\pi r}{fr} = \frac{2\pi}{f} = \frac{\pi}{\omega \sin \varphi}$$

运动方向

北半球，顺时针；南半球，逆时针.

背景流

- (1) 当无其他外加流动时，所有惯性圆的圆心位于同一条铅直线上，因而海水就像以角速度 $2\omega \sin \varphi$ 旋转的刚体一样；
- (2) 当有其他外加流动时，除了在同一水平面上所有海水质点皆在同一时刻由同一流速速率外，还依外加流动速度方向移动.

2.4 风生大洋环流

2.4.1 Sverdrup 理论

假定

- (1) 远离海岸的大洋中部海区, $Ro \ll 1$ 大尺度、等深大洋 h 为常数;
- (2) 远离边界、无侧边界影响, 无水平湍摩擦应力 $E_l \ll 1$;
- (3) 定常风定常流动;
- (4) ρ 为常数;
- (5) β 平面近似.

控制方程

$$\begin{cases} -fv = -\frac{1}{\rho} \frac{\partial p}{\partial x} + A_z \frac{\partial^2 u}{\partial z^2} \\ fu = -\frac{1}{\rho} \frac{\partial p}{\partial y} + A_z \frac{\partial^2 v}{\partial z^2} \\ 0 = -\frac{1}{\rho} \frac{\partial p}{\partial z} - g \\ \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0 \end{cases}$$

边界条件

$$\begin{cases} z = \zeta(\text{海面}) : \rho A_z \frac{\partial u}{\partial z} = \tau_{x\zeta}, \rho A_z \frac{\partial v}{\partial z} = \tau_{y\zeta} \\ z = -h(\text{足够深}) : u = v = 0, \frac{\partial u}{\partial z} = \frac{\partial v}{\partial z} = 0, \frac{\partial p}{\partial x} = \frac{\partial p}{\partial y} = 0 \end{cases}$$

求解

对上式进行垂直积分:

$$\begin{cases} -fM_y = -\frac{\partial P}{\partial x} + \tau_{x\zeta} & (15) \\ fM_x = -\frac{\partial P}{\partial y} + \tau_{y\zeta} & (16) \\ \frac{\partial M_x}{\partial x} + \frac{\partial M_y}{\partial y} = 0 \end{cases}$$

$$\text{其中, } M_x = \int_{-h}^0 \rho u dz, M_y = \int_{-h}^0 \rho v dz, P = \int_{-h}^0 p dz, \tau_{x\zeta} = \int_{-h}^0 \rho A_z \frac{\partial^2 u}{\partial z^2} = \rho A_z \left(\frac{\partial u}{\partial z} \Big|_{z=0} - \frac{\partial u}{\partial z} \Big|_{z=-h} \right) \quad (\zeta \ll h)$$

$$\begin{aligned} \frac{\partial(1)}{\partial y} - \frac{\partial(2)}{\partial x} &\Leftrightarrow -M_y \frac{\partial f}{\partial y} \overset{0}{\boxed{-f \frac{\partial M_y}{\partial y} - f \frac{\partial M_x}{\partial x}}} = \frac{\partial \tau_{x\zeta}}{\partial y} - \frac{\partial \tau_{y\zeta}}{\partial x} \\ &\Leftrightarrow M_y \frac{\partial f}{\partial y} = \frac{\partial \tau_{y\zeta}}{\partial x} - \frac{\partial \tau_{x\zeta}}{\partial y} \\ &\Leftrightarrow \boxed{\beta M_y = \text{rot}_z \tau_\zeta} \\ &\text{Sverdrup 方程} \end{aligned}$$

Sverdrup 方程的物理意义 1: 海水南北向的输运由风应力旋度所驱动.

讨论

将质量输运分为 Ekman 漂流输运与地转流输运两部分:

$$\begin{aligned} M_x &= M_{xE}(\text{Ekman 漂流}) + M_{xg}(\text{地转流}) \\ M_y &= M_{yE}(\text{Ekman 漂流}) + M_{yg}(\text{地转流}) \end{aligned}$$

$$\begin{cases} -fM_{yE} = \tau_{x\zeta} & (17) \\ fM_{xE} = \tau_{y\zeta} & (18) \\ -fM_{yg} = -\frac{\partial P}{\partial x} & (19) \\ fM_{yg} = -\frac{\partial P}{\partial y} & (20) \end{cases}$$

$$\frac{\partial(17)}{\partial y} - \frac{\partial(18)}{\partial x} \Leftrightarrow \frac{\partial M_{xE}}{\partial x} + \frac{\partial M_{yE}}{\partial y} = \boxed{\nabla \cdot \vec{M}_E = (\text{rot}_z \vec{\tau}_\zeta - \beta M_{yE}) / f} \quad (21)$$

$$\frac{\partial(19)}{\partial y} - \frac{\partial(20)}{\partial x} \Leftrightarrow \frac{\partial M_{xg}}{\partial x} + \frac{\partial M_{yg}}{\partial y} = \boxed{\nabla \cdot \vec{M}_g = -\beta M_{yg} / f} \quad (22)$$

(1) Ekman 漂流质量输运的水平散度与 ① 风应力旋度 ② f ③ β 有关.

(2) 地转流质量输运的水平辐散引起南北向的地转运动.

$$(21) + (22) \Leftrightarrow \text{rot}_z \tau_{x\zeta} - \beta M_y = 0 \Leftrightarrow \boxed{\beta M_y = \text{rot}_z \vec{\tau}_\zeta}$$

Sverdrup 方程的物理意义 2: 地转流流量的散度和 Ekman 漂流流量的散度相平衡, 所以 Sverdrup 方程又称 Sverdrup 平衡.

(1) 所有南北向的地转运动, 必须显示水平散度;

(2) 虽然 Ekman 漂流流量的散度与地转流流量的散度本身不为 0, 但是它们的和, 即总流量的水平散度必须为 0, 说明 Ekman 漂流流量的散度和地转流流量的散度刚好取得平衡;

(3) $\text{rot}_z \vec{\tau}_\zeta = 0$: 只存在东西方向的输运, $\text{rot}_z \vec{\tau}_\zeta > 0$: 质量输运向北 (北半球), $\text{rot}_z \vec{\tau}_\zeta < 0$: 质量输运向南 (南半球);

(4) 地转流引起的南北质量运输量比 Ekman 漂流引起的大.

缺陷

设仅有纬向风, 且 $\tau_{x\zeta}$ 仅为 y 的函数:

$$M_y = \frac{1}{\beta} \text{rot}_z \vec{\tau}_\zeta = \frac{1}{\beta} \left(\cancel{\frac{\partial \tau_{x\zeta}}{\partial x}} - \frac{\partial \tau_{x\zeta}}{\partial y} \right) = -\frac{a}{2\omega \cos \varphi} \frac{\partial \tau_{x\zeta}}{\partial y}$$

$$\frac{\partial M_x}{\partial x} = -\frac{\partial M_y}{\partial y} \Rightarrow M_x = \frac{x}{2\omega \cos \varphi} \left(a \frac{\partial^2 \tau_{xf}}{\partial y^2} + \frac{\partial \tau_{xf}}{\partial y} \text{tg} \varphi \right) + c(y)$$

在东、西边界, $M_x = 0$ 而 Sverdrup 理论不能同时满足.

参考文献

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