Using the Sediment Transport and Vegetation Operators in ANUGA-Sed

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This document describes the implementation of operators for sediment transport and vegetation drag in ANUGA-Sed, a derivative of the hydrodynamic model ANUGA. The sediment transport operator can currently only be used with a single processor. Making it parallel safe would require modifying, at minimum, the method used for sediment flux. The vegetation drag operator should be parallel safe but has not been tested.

These operators use the equations presented in *Simpson and Castelltort* [2006] and *Davy and Lague* [2009] for calculating sediment transport and momentum sinks, and *Nepf* [1999] and *Kean and Smith* [2006] for vegetation drag.

1 Files

The two operators are located in the directory operators, this documentation is in the direction doc, and example files for using these two operators are in the directory examples/operators/sed.

2 Using the Sediment Transport Operator

2.1 Creating the domain

To use the sedimene transport operator, the quantity *concentration* must be defined as an evolved quantity when the domain is created. Exceptions are raised if this quantity does not exist when the operator is used.

The existing function for creating a rectangular domain already accepts a list of evolved quantities as an argument:

Creating a rectangular domain

Creating a domain that uses a mesh as input for *elevation* is usually done with the function create_domain_from_regions, which does not accept a list of evolved quantities. To accomplish the same functionality, use create_mesh_from_regions to generate the mesh and then create the domain:

Creating a domain from a mesh file

Because it updates the elevation of the bed, the sediment transport operator can only be used with one of ANUGA's Discontinuos Elevation ('DE') flow algorithms. Changes in the elevation of the bed and the distribution of sediment in transport during the simulation can be recorded in the SWW file by including elevation and concentration in domain.set_quantities_to_be_stored with a value of 2 (to save them at every yieldstep):

Setting the flow algorithm and recording the output

2.2 Creating the operator

The Sediment Transport operator is initialized in the run file with:

```
from anuga.operators.sed_transport_operator import Sed_transport_operator
sed_op = Sed_transport_operator(domain)
```

Initializing the sediment transport operator

Variable name	Default value	Units	Quantity
rho_w	1000	${\rm kg}~{\rm m}^{-3}$	Density of fluid
rho_s	2650	${\rm kg~m^{-3}}$	Density of sediment
nu	$1 \text{x} 10^{-6}$	$\mathrm{m}^2~\mathrm{s}^{-1}$	Kinematic viscosity of water
porosity	0.3	-	Porosity of the bed
criticalshear_star	0.06	-	Dimensionless critical shear stress
<pre>grain_size</pre>	0.00013	m	Grain size

Table 1: Default values for equation parameters in Sediment transport operator

Table 1 summarizes the default values that various equation parameters take when Sed_transport_operator is initialized. These values can be modified from the run file after the operator is initialized.

The initial concentration of sediment in the flow within the domain (as a fraction of the volume of the water column) can be set at any point after the domain is created with domain.set_quantity(). The concentration of any flow entering the domain through a Dirichlet boundary can be set after initializing the operator (for all boundaries, only once before the start of the run) as the operator variable inflow_concentration. If not set, inflow_concentration takes the maximum value of the quantity concentration at the first timestep.

```
domain.set_quantity('concentration', 0.01) # 1 percent
sed_op.inflow_concentration = 0.02 # 2 percent
sed_op.grain_size = 0.0002 # 0.2 mm grains
sed_op.criticalshear_star = 0.03
```

Modifying parameter values

Creating a vegetation drag operator

The function arguments turbulence, momentum sinks, and verbose, False by default, will be set to True if they are True for either of the operators.

2.3 Operator call

The operators act only on vertices with a flow depth greater than a minimum depth (set in sed_transport_config. py as min_depth=0.005m as default) in order to avoid extremely high erosion rates due to abnormally high velocities at low flow depths and to minimize the likelihood of thicknesses of aggradation that exceed the water depth, and permits the use of the operators without having to reduce the length of the timesteps.

The order of operations when both Sed_transport_operator and Vegetation_operator are called is:

- 1. Calculate drag on vegetation, compute the decreased flow velocity
- 2. Calculate erosion, deposition, and change in concentration
- 3. Compute sediment flux, update quantity concentration
- 4. Update quantity elevation
- 5. Compute sources and sinks of momentum, update quantities xmomentum and ymomentum
- 6. Compute turbulence, update quantity diffusivity

2.3.1 Drag on vegetation

The drag force F_D that vegetation imparts on the flow is given by:

$$F_D = \frac{1}{2} C_D a U_{ref}^2 \tag{1}$$

where $a = d_s/\Delta S^2$ is the projected plant area per unit volume (vegetation density per meter). This assumes that the vegetation can be approximated as a regular array of cylindrical stems that emerge from the flow. The drag force is allowed to reduce the flow velocity to zero but not to change its direction.

2.3.2 Sediment transport processes

The rate of change of elevation of the bed is a mass balance equation written as:

$$\frac{\partial z}{\partial t} = \frac{\dot{D} - \dot{E}}{1 - \phi} \tag{2}$$

where z is the elevation of the bed above some datum, ϕ is the porosity of the bed material, and \dot{D} and \dot{E} are the rates of sediment deposition and entrainment. The entrainment rate is given by:

$$\dot{E} = K_e \left(\tau^* - \tau_c^* \right) \tag{3}$$

where K_e is an erosion coefficient and τ_c^* is the dimensionless critical shear stress. τ^* is the dimensionless shear stress that the flow applies on the bed:

$$\tau^* = u^* \frac{1}{R g d_q} \tag{4}$$

where $u^* = \sqrt{f/8}$ is the shear velocity, f is the Darcy friction factor (f = 0.05), d_g is the grain diameter, $R = (\rho_s - \rho_w)/\rho_w$ is the submerged specific gravity of sediment, ρ_w and ρ_s are the densities of water and sediment, and g is the gravitational acceleration. K_e is given by:

$$K_e = K_e^* d_q \sqrt{R g d_q} \tag{5}$$

where K_e^* is a dimensionless erosion coefficient ($K_e^* = 12$).

The rate of aggradation D can be written as:

$$\dot{D} = D^* C v_s \tag{6}$$

where D^* is a dimensionless number equal to 1 if the sediment flux is uniformly distributed throughout the depth of the flow, and v_s is the settling velocity of grains in the fluid [Ferguson and Church, 2004]:

$$v_s = \frac{R g \, d_g^2}{C_1 \, \nu + (0.75 \, C_2 \, R \, g \, d_g^3)^2} \tag{7}$$

where C_1 and C_2 are 18 and 0.4 for smooth spheres, and ν is the kinematic viscosity of the fluid.

The time rate of change of concentration of sediment in the flow can be written as:

$$\frac{\partial Ch}{\partial t} = \dot{E} - \dot{D} - \frac{\partial q_{sx}}{\partial x} - \frac{\partial q_{sy}}{\partial y} \tag{8}$$

where q_{s_x} and q_{s_y} are sediment discharges per unit width in the x and y directions.

2.3.3 Sources and sinks of momentum

Four sources or sinks of momentum can be included by these operators into the equations for conservation of momentum: (1) friction loss, (2) loss due to spatial variations in sediment concentration, (3) loss of momentum due to the exchange of mass between the flow and the bed, and (4) loss due to fluid drag on vegetation. Changes to the momentum of the flow due to turbulence are calculated when the function argument *turbulence* is True.

- 1. Loss due to friction is computed using the existing manning_friction_implicit function
- 2. Loss due to spatial variations in concentration is given by the equation:

$$P_C = \frac{(\rho_s - \rho_w)gh^2}{2(\rho_w(1 - C) + \rho_s C)} \nabla \cdot C \tag{9}$$

3. Loss due to the exchange of mass between the flow and the bed is given by:

$$P_e = \frac{\dot{D} - \dot{E}}{1 - \phi} \left(\frac{\rho_w \phi + \rho_s (1 - \phi)}{\rho_w (1 - C) + \rho_s C} - 1 \right) \vec{v}$$
 (10)

where \vec{v} is the flow velocity.

4. Loss of momentum due to fluid drag on vegetation is given by:

$$P_d = -\vec{v}\vec{F_D}h\tag{11}$$

where $\vec{F_D}$ is the drag force of vegetation on the flow. The quantities *xmomentum* and *ymomentum* are modified by this equation whenever the Vegetation_operator is used, regardless of the momentum_sinks argument.

2.3.4 Updating quantities

All operators listed in fractional_step_operators are called at every timestep by domain.evolve after the flow calculations have been performed and the conserved quantities updated.

Both the Sed_transport_operator and Vegetation_operator directly modify the values of several quantities through functions in sed_transport_mesh.py.

- Momentum is modified by using the explicit_update (re-set to zero) and update methods of the quantities xmomentum and ymomentum.
- The shape of the bed is altered by using the set_values function at the vertices for the quantity elevation, followed by smooth_vertex_values() to eliminate discontinuities in the topography and distribute the updated elevations to the centroids and edges.
- The flux of sediment across the cell edges is computed by (1) obtaining the total volume of sediment in the flow within each cell from the values of the quantity concentration at the centroids, (2) finding the normal component of the flow velocities at the edges of the cells, (3) calculating, from the values of concentration at the edges, the volume of sediment that crosses each edge within the timestep, and (4) integrating the flux in and out of each cell across all edges to (5) change the volume of sediment present in the flow within each cell. This volume is then (6) converted back to a sediment concentration (of the form Ch) at the centroid of each cell, and (7) the updated value at each vertex is calculated by averaging the value at the centroids around it. (8) The quantity concentration is updated using the function set_values at the vertices.

3 Utilities

The file sed_transport_utils.py contains a set of functions, similar to existing ones, that are tailored to the needs of these specific operators.

```
create_domain_from_regions_sed(...)
```

This function is equivalent to __init__.create_domain_from_regions but accepts lists of quantities to be created within the domain. See example in section 2.1.

Reflective_boundary_Sed(domain)

This boundary type is equivalent to Reflective_bondary but manages the boundary value for concentration.

Dirichlet_boundary_Sed([stagexmomymomC])

This boundary is equivalent to the existing $Dirichlet_boundary$ but accepts a fourth entry for fractional concentration (C, not Ch) at the boundary. This boundary type should only be used for inflow boundaries. Flow outlets should be declared using the regular Dirichlet boundary type, which will allow sediment to be carried across the boundary by the flow instead of fixing the concentration at the cell edges. An exception is raised if a Dirichlet_boundary_Sed is has a fixed stage that is lower than the lowest elevation in the domain.

```
'side2' : Br,
'top' : Bd,
'side3' : Br,
'side4' : Br,
'exterior' : Br})
```

Listing 1: Using operator-specific boundary types

set_quantity_NNeigh(quantity_name,filename)

Equivalent to the function set_quantity, it assigns values from the point file filename to quantity quantity name. In contrast to set_quantity, which interpolates between the values in the point file to obtain the value at the centroids, this function uses a Nearest Neighbour algorithm to assign each centroid the value of the nearest point in the file. This is necessary when assigning values from an ASCII file to the quantity vegetation.

Two files in the directory operators/sed_transport/file_conversion can be used to transform ASCII files into point files that represent quantities other than *elevation*.

```
generic_asc2dem(name_in,quantity_name,...)
```

Equivalent to anuga.asc2dem, this function transforms an .asc file into a .dem file. While the original function assumes that the values in the files will be assigned to the quantity *elevation* (and codes this into the files), this function allows the user to define the name of that quantity in the argument *quantity name*.

```
generic_dem2pts(name_in,quantity_name,...)
```

Equivalent to anuga.dem2pts, this function transforms a .dem file into a .pts file for future use with the quantity quantity name. This function does not create that quantity or set its values.

```
from anuga.operators.sed_transport.file_conversion.generic_asc2dem
   import generic_asc2dem
from anuga.operators.sed_transport.file_conversion.generic_dem2pts
   import generic_dem2pts
from anuga.operators.sed_transport.sed_transport_utils
   import set_quantity_NNeigh
generic_asc2dem('veg.asc',
                quantity_name = 'vegetation',
                use_cache = False,
                verbose = True)
generic_dem2pts('veg.dem',
                quantity_name = 'vegetation',
                use_cache = False.
                verbose = True)
set_quantity_NNeigh(domain,
                     vegetation',
                    filename='veg.pts')
```

Listing 2: Using file conversion and set quantity utilities

4 Tests

A suite of unit tests and analytical solutions will be added to the operator package in the immediate future.

4.0.5 Process-specific quantities

concentration

Following Simpson and Castelltort [2006], concentration is given by:

Concentration
$$= \frac{Q_s}{Q}h = Ch$$
 (12)

where Q_s is the sediment discharge, Q is the total discharge, h is the flow depth, and C is the fractional concentration of sediment (by volume) in the flow. The operators store Ch as the quantity concentration. It is important to note that the sediment transport-specific Dirichlet boundary takes C, not Ch, as the input value.

vegetation

The values that are stored as the quantity *vegetation* are numeric codes (positive integers) for different types of vegetation, each with corresponding values of stem diameter and stem spacing in a look-up table. The default value for *vegetation* is zero, which is the code for no vegetation. The values of this quantity can be set using the existing functions in ANUGA or ones included in the directory /file_conversion and the file sed_transport_utils.py (see description in section 3).

If any entry of *vegetation* is greater than zero, the operators will attempt to import the csv file specified in the function argument *vegfile* of Vegetation_operator.

An exception is raised if the file is not found or if vegfile is not specified. In the case of a domain with two different types of vegetation (vegetation > 1) as well as bare areas (vegetation = 0), this file would contain a column vegcode with the corresponding code in the quantity vegetation, and two columns for the values of $stem\ diameter$ and $stem\ spacing$ (in meters).

```
vegcode, stem_diameter, stem_spacing
1, 0.01, 0.1
2, 0.01, 0.3
```

Listing 3: Example of vegfile

diffusivity

The effects of turbulence on the momentum of the flow can be calculated using these operators by harnessing the existing kinematic viscosity operator (domain.set_use_kinematic_viscosity) and dynamically modifying the values of the quantity diffusivity to reflect the conditions of the flow. An exception is raised by the operators if the argument turbulence is True but the quantity diffusivity does not exist. A message is then displayed and the run continues without calculating turbulence.

The values for *diffusivity* are calculated using the equation:

$$K_d = \sqrt{k_e} \, l + ad \, U_{ref} \, d_s \tag{13}$$

where l is the mixing length, U_{ref} is the flow velocity felt by the stems, d_s is the stem diameter, $ad = d_s^2/\Delta S^2$ is the fractional volume of the flow domain occupied by plants, and ΔS is the mean stem spacing. The velocity scale $\sqrt{k_e}$ is given by:

$$k_e = C_b U_{ref} + (\overline{C_D} ad)^{2/3} U_{ref}^2$$
 (14)

where C_b is the bed drag coefficient ($C_b = 0.001$) and $\overline{C_D}$ is the bulk drag coefficient for the vegetation array. If the stems are approximated as emergent cylinders, $\overline{C_D} = 1.2$. The mixing length l is given by:

$$l = h - \frac{h - h(ad - 0.005)}{0.005} \tag{15}$$

where h is the flow depth. The second term of all three equations is equal to zero when vegetation is not present (ad = 0). This equation is an approximation of the data shown in Nepf [1999], figure 6.

References

- Davy, P., and D. Lague (2009), Fluvial erosion/transport equation of landscape evolution models revisited, *Journal of Geophysical Research-Earth Surface*, 114(F3), F03,007, doi:10.1029/2008JF001146.
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