



# Numerical Solution to the Barotropic Vorticity Equation on an f-plane

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Final Project, Physics 781

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# Vorticity

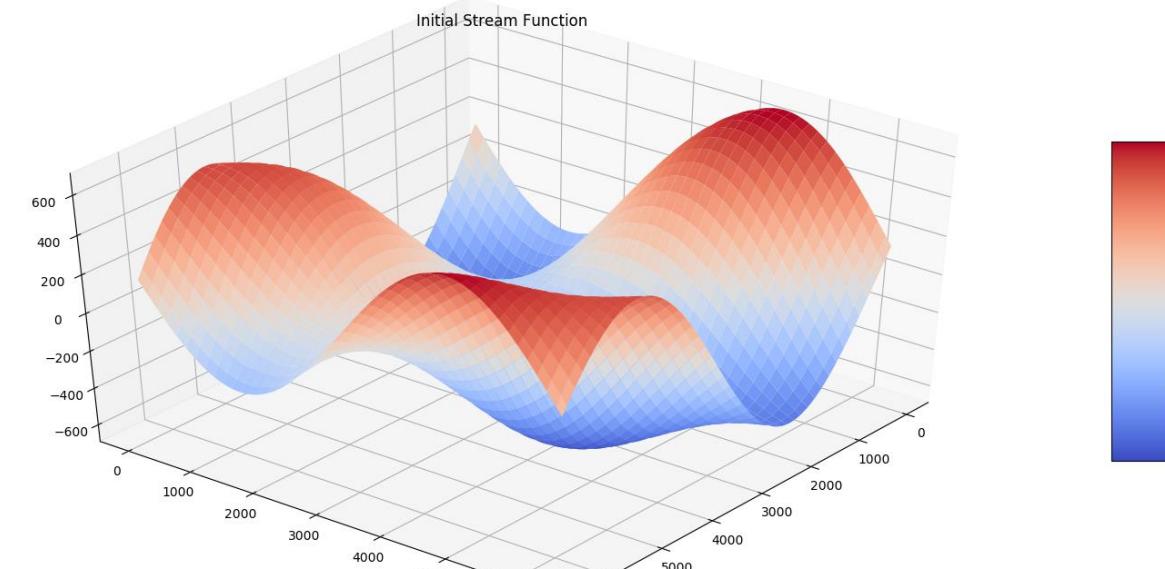
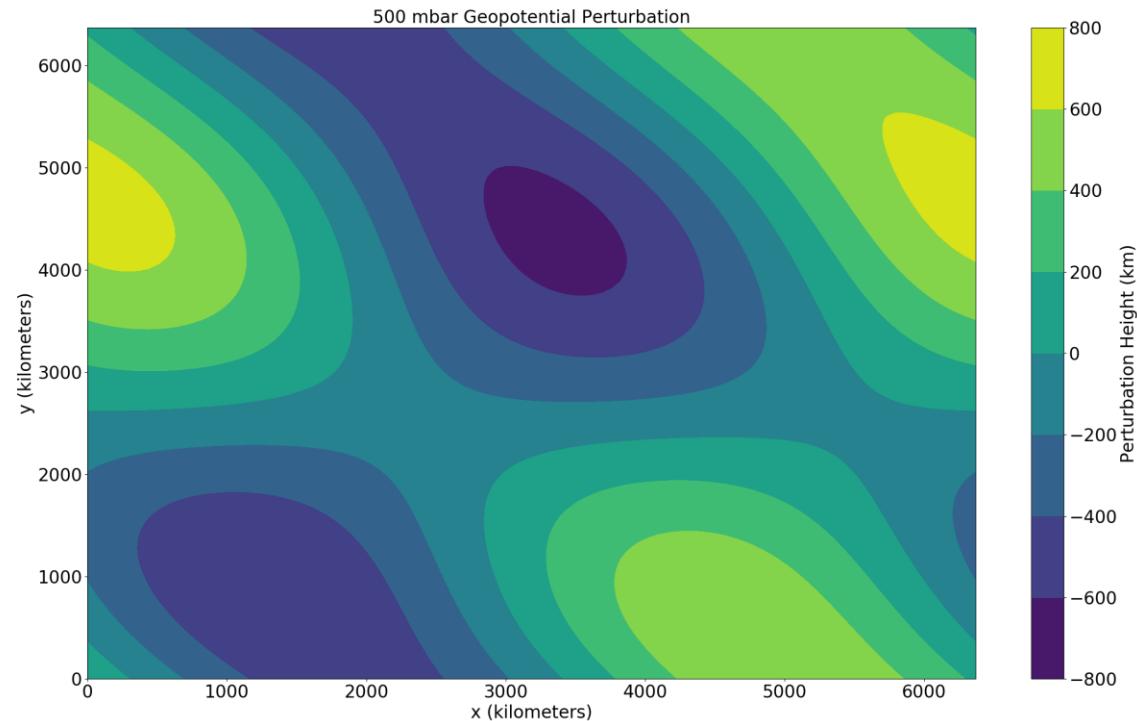
- Describes the local rotation of a fluid near some central point
- Analog to angular momentum for fluids
- Conserved!
- $\eta = \zeta + f$ , Absolute vorticity is equal to the relative vorticity plus contribution from Earth's rotation.
- $\frac{d}{dt} [\zeta + f] = 0$ . Vorticity Conservation
- Helpful to think of a “potential” function, the stream-function:  $\nabla^2 \Psi = \zeta$



# Barotropic Vorticity Equation

$$\frac{d}{dt} (\nabla^2 - F) w + \frac{g}{f} J(\psi, \nabla^2 \psi) + \beta \frac{\partial}{\partial x} w = 0$$

- $w$  is stream function,  $w = \frac{g}{f} \Phi$ ,  $\Phi$ =Geopotential Height
- $g$  is gravity,  $f$  is Coriolis parameter:  $2\Omega \sin(\text{latitude})$ ,  $\Omega$  is planetary rotation rate
- $F$  is small scale dissipation term
- $\beta = \frac{2\Omega \cos(\text{latitude})}{\text{Radius of Earth}}$
- $J(w, \nabla^2 w)$  is the Arakawa Jacobian,  $J(w, \nabla^2 w) = \frac{\partial w}{\partial x} \frac{\partial (\nabla^2 w)}{\partial y} - \frac{\partial w}{\partial y} \frac{\partial (\nabla^2 w)}{\partial x}$
- Consider constant latitude at 45 degrees North



# Solving the Problem

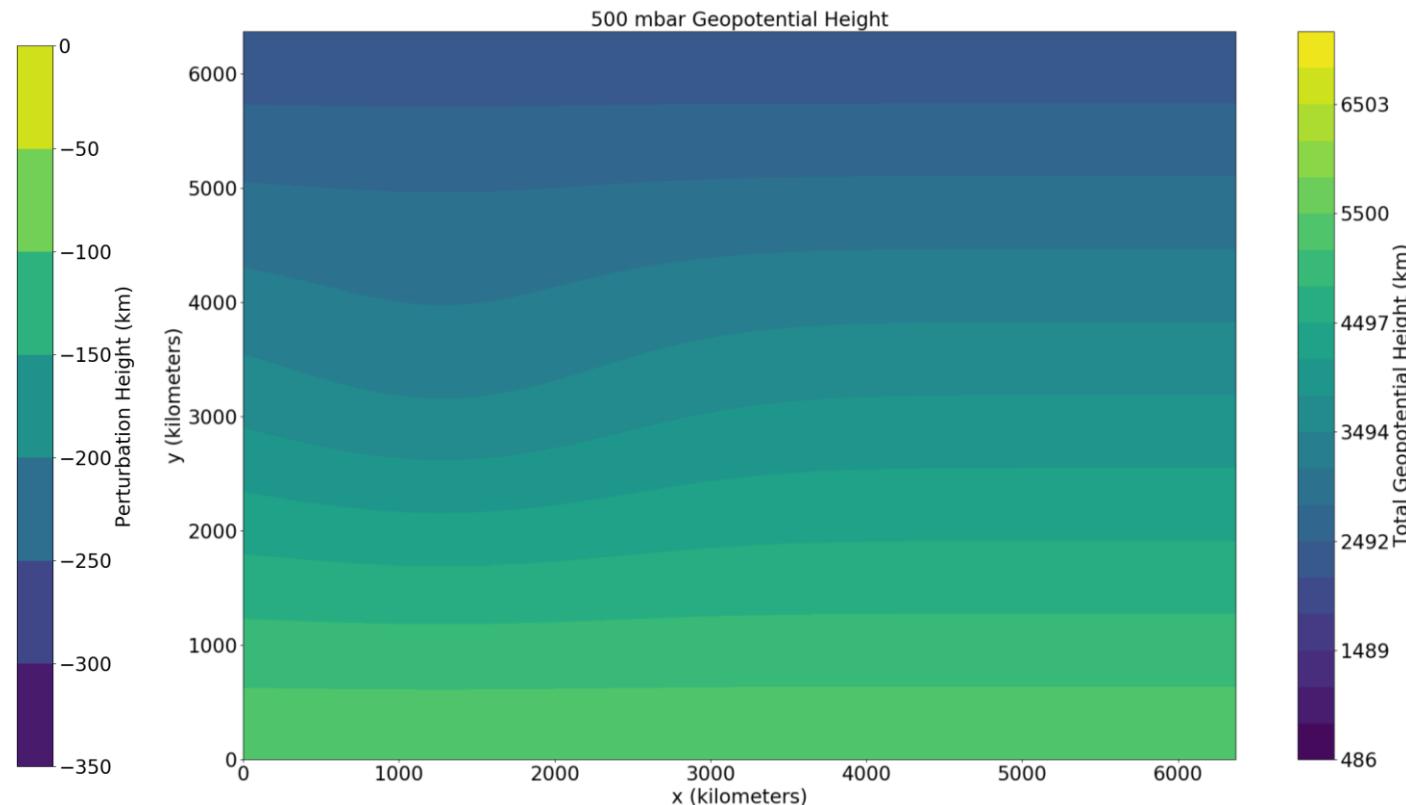
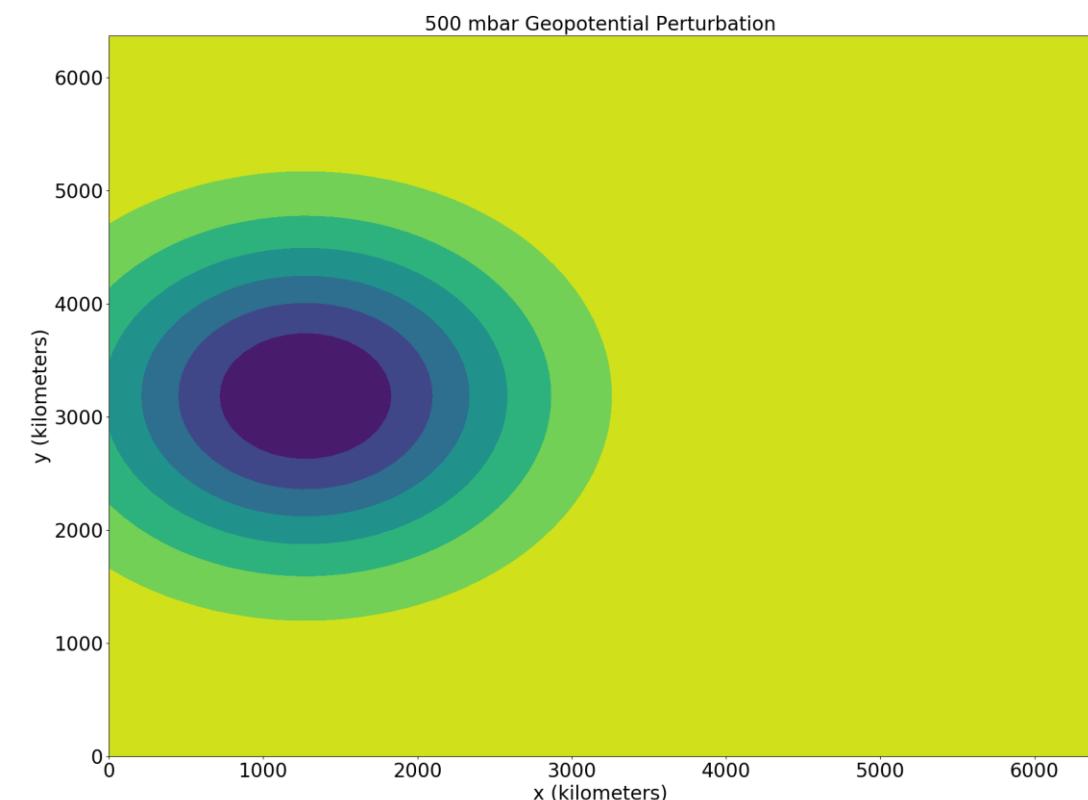
$$\frac{d}{dt}(\nabla^2 - F)w = -\frac{g}{f} J(w, \nabla^2 w) - \beta \frac{\partial}{\partial x} w$$
$$J(w, \nabla^2 w) = \frac{\partial w}{\partial x} \frac{\partial (\nabla^2 w)}{\partial y} - \frac{\partial w}{\partial y} \frac{\partial (\nabla^2 w)}{\partial x}$$

- Define grid
- Generate initial geopotential height perturbation
- Take derivatives using finite-difference method, appropriate to boundary conditions
- Leap-frog forward right side of equation to find left side of equation at new time step
- Use spectral method on left side of equation to find  $\psi$
- Add average horizontal flow
- Repeat

# Defining Grid

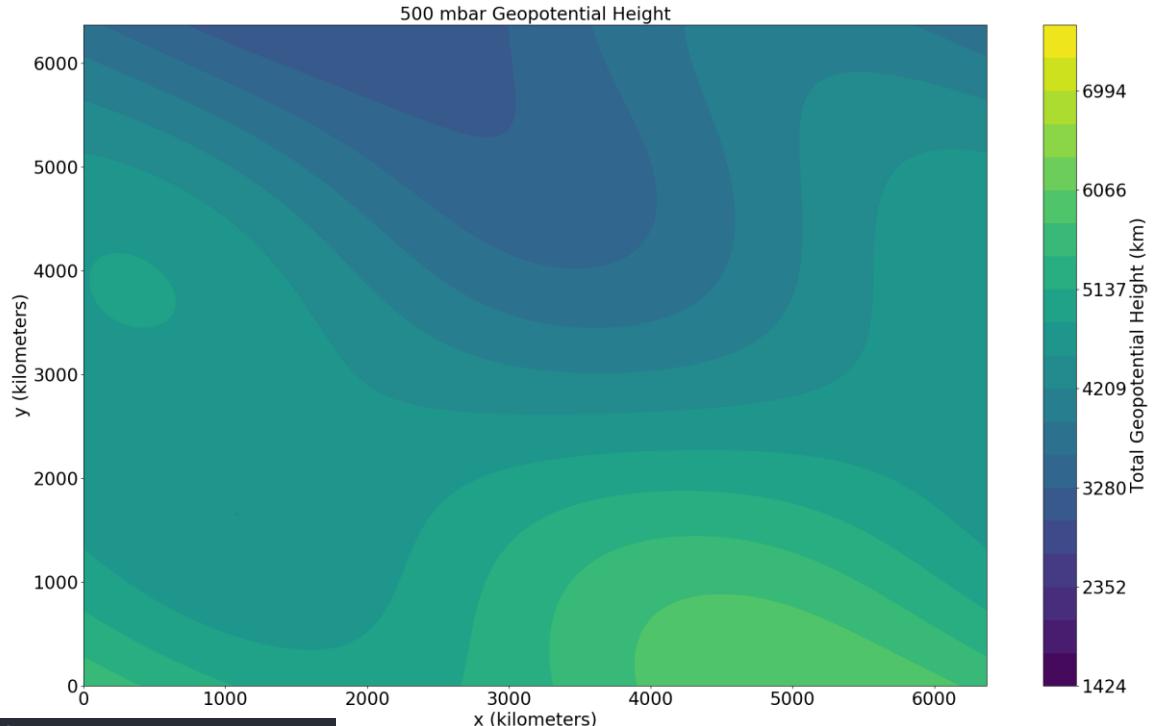
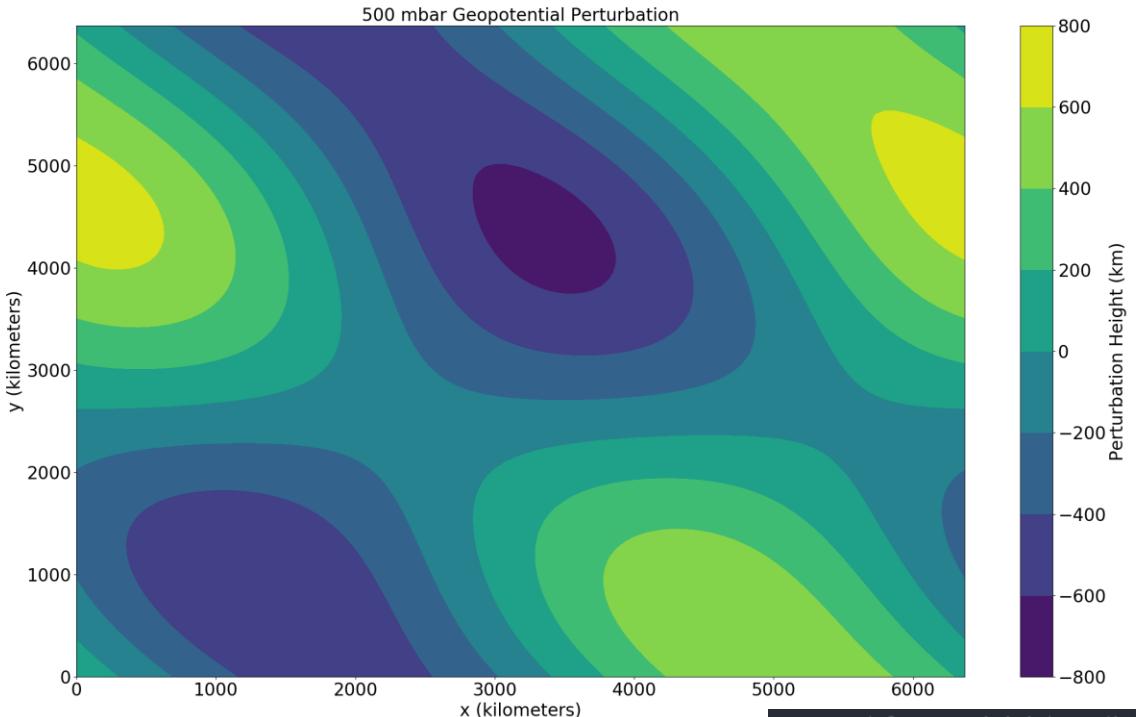
- 6000 x 6000 km grid cartesian grid
- Spatial resolution 50-150 km depending on what one is doing
- Temporal Resolution of 37 seconds to 5 minutes depending on what one is doing
- Coriolis parameter kept constant at 45 degrees for simplicity

# Generating Initial Conditions: Gaussian Perturbation



```
Z_θ = zeros((nx+1,ny+1));
term = -350*exp(((XX/xlen-.2)**2-(YY/ylen-.5)**2))/ .05
Z_θ = Z_θ + term
Z_θ[0:ny]=Z_θ[0:ny+1,ny-1]
Z_θ[0:ny+1,0]=Z_θ[0:ny+1,1]
```

# Generating Initial Conditions: 2-D waves



```
#set seed for same initial conditions every time
seed(a=2)
#randomize seed for different initial conditions
#seed()
Z_0 = zeros((nx+1,ny+1));

# Specify the number of waves in x and y

NwaveX = 1
NwaveY = 1
Nwaves = (2*NwaveX+1)*(2*NwaveY+1)

for kwave in range(-NwaveX,NwaveX+1):
    for lwave in range(-NwaveY,NwaveY+1):
        Amplitude = (2000.0*2*(random()-0.5)) / Nwaves
        phase= 2*pi*(random()-0.5)
        term = cos(2*pi*(kwave*(XX/xlen)+lwave*(YY/ylen)+phase))
        Z_0 = Z_0 + Amplitude*term
```

# Boundary Appropriate Finite Differencing

- We have an  $N \times N$  grid, when we take a derivative for the **bulk of the matrix** we use a central difference method (Order  $h^2$ )
- At edges we have to be careful and compute derivatives in accordance with our boundary conditions.

	$f(x-h)$	$f(x)$	$f(x+h)$
$f(0)$	$f(h)$		$f(N*h)$

- For a step size  $h$ , central difference:

$$f'(x) = \frac{f(x+h) - f(x-h)}{2h}, f''(x) = \frac{f(x+h) + f(x-h) - 2f(x)}{h^2}$$

- At **periodic boundary**  $f'(0) = \frac{f(h) - f(N*h)}{2h}, f''(0) = \frac{f(h) + f(N*h) - 2f(0)}{h^2}$

# Boundary Appropriate Finite Differencing continued...

- In this method one can use periodic or symmetric boundary conditions in  $y$
- For symmetric boundary conditions  $y$ , the derivative must be redefined using backward and forward differencing

- For a step size  $h$ , forward difference:

$$f'(y) = \frac{f(y+h) - f(y)}{h}, f''(y) = \frac{f(y+2h) + f(y) - 2f(y+h)}{h^2}$$

- At symmetric boundary  $f'(0) = \frac{f(h) - f(0)}{h}, f''(0) = \frac{f(0) + f(2h) - 2f(h)}{h^2}$

$f(0)$		
$f(0)$		
$f(h)$		
$f(2h)$		$f(N - 3h)$
		$f(N - 2h)$
		$f(N - h)$
		$f(N - h)$

# Leap-frog forward

$$\frac{d}{dt} (\nabla^2 - F) w = -\frac{g}{f} J(w, \nabla^2 w) - \beta \frac{\partial}{\partial x} w$$

$$R = -\frac{g}{f} J(w, \nabla^2 w) - \beta \frac{\partial}{\partial x} w$$

- $(\nabla^2 - F)w = Q \rightarrow \frac{d}{dt}Q = \frac{Q(n+1) - Q(n)}{\Delta t} \rightarrow Q(n+1) = Q(n) + R\Delta t$
- For first time step use  $\Delta t/2$

# Spectral Method to find $\psi$

$Q(n+1) = (\nabla^2 - F)w(n+1) \rightarrow$  Fourier Transform  $\rightarrow Q(n+1) = (-k^2 - F)w(n+1) \rightarrow$

$\rightarrow \frac{Q(n+1)}{(-k^2 - F)} = w(n+1) \rightarrow$  Inverse Fourier Transform  $\rightarrow w(n+1)$

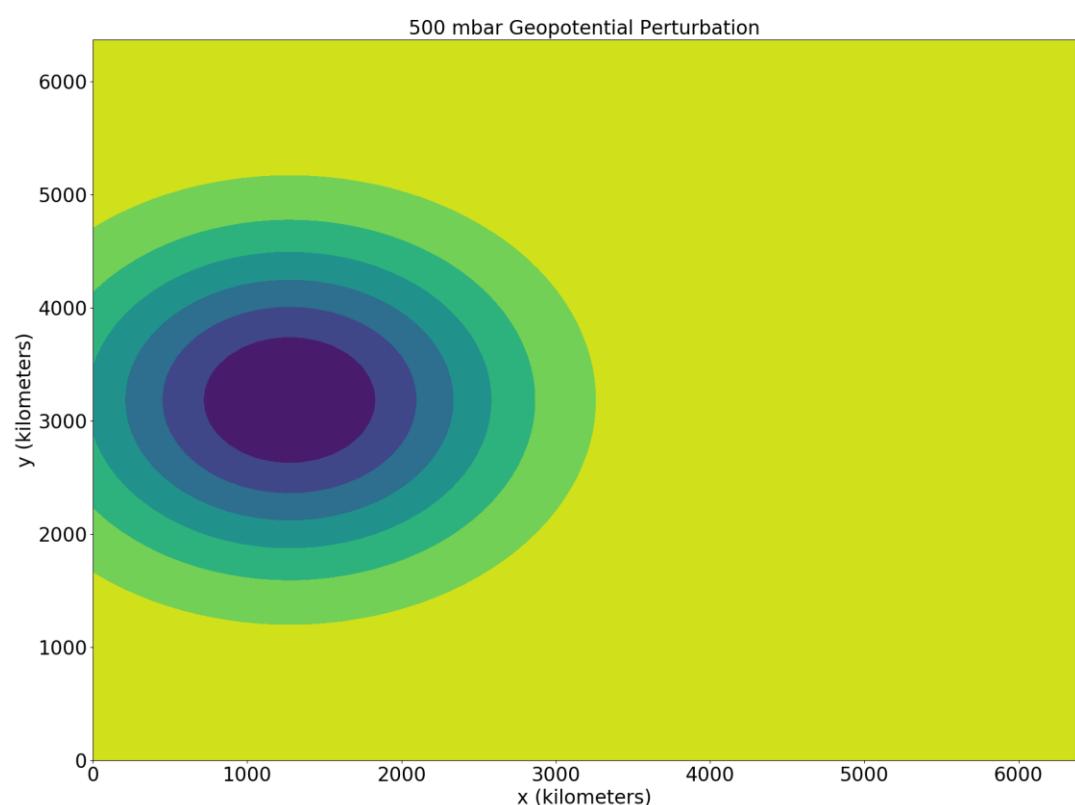
```
#Section 3.3: Solve the Helmholtz Equation  $(\nabla^2 - F)w = R$ . Compute the fft of the
#right hand side (strip off additional row and column).
R[0:nx,0:ny] = Q_np1[0:nx,0:ny]
R_hat = fft2(R)

#Compute the transform of the solution
W_hat = R_hat/C_rs
#Compute the inverse transform to get the solution at  $(n+1) * \Delta t$ .
w_new = real(ifft2(W_hat)) # We assume w is real
w[0:nx,0:ny] = w_new
w[nx,0:ny] = w[0,0:ny]      # Fill in additional column at east.
w[0:nx+1,ny]=w[0:nx+1,0]    # Fill in additional row at north.
```

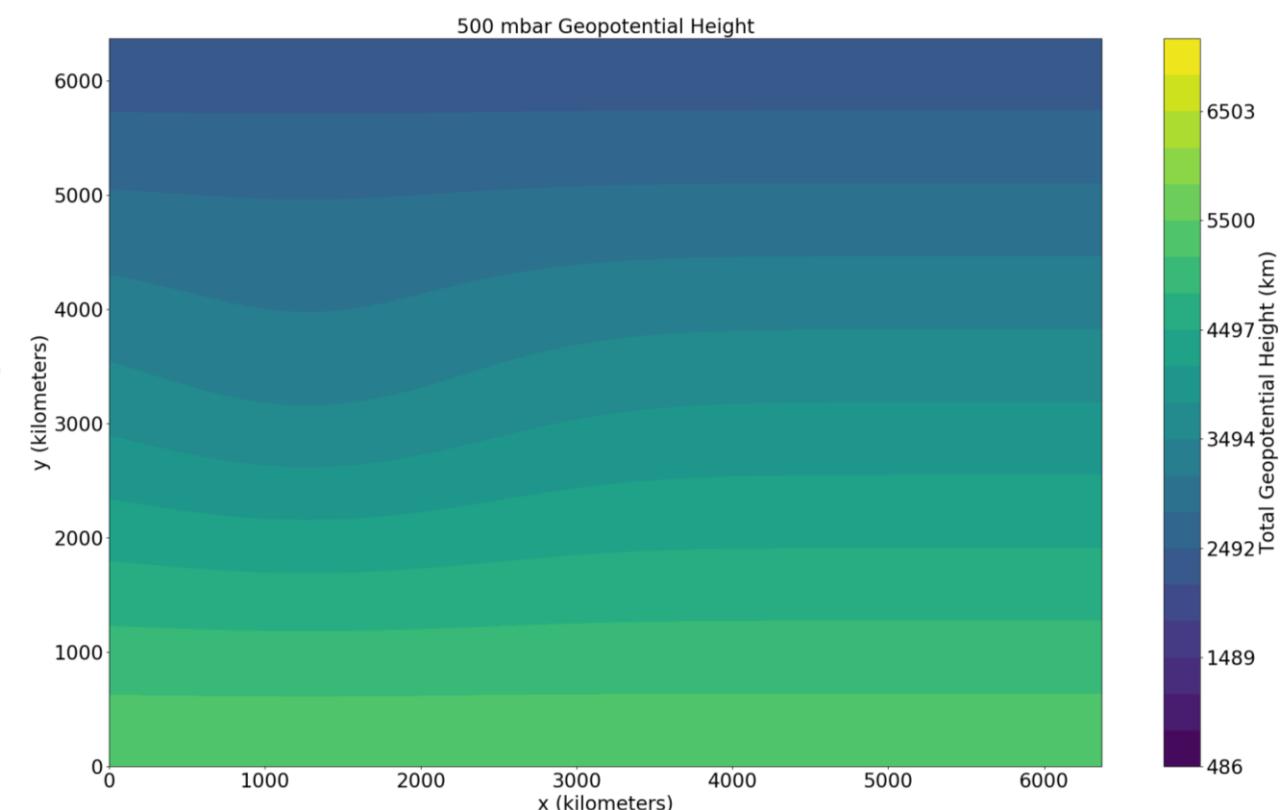
Comparisons and spatial and  
temporal resolution

$$C = u \frac{\Delta t}{\Delta x}$$

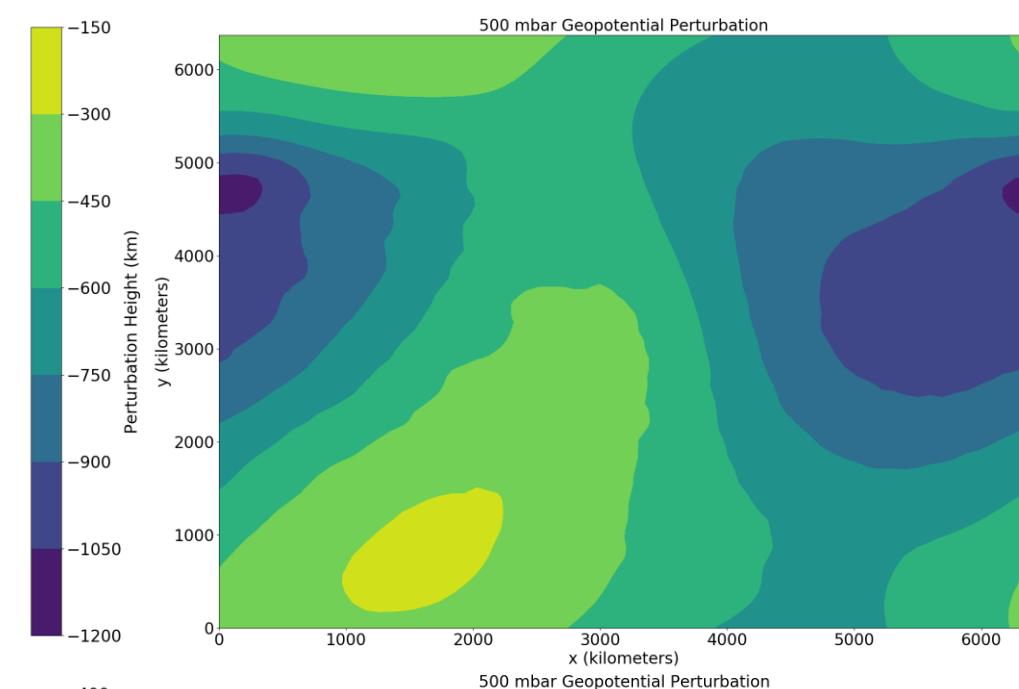
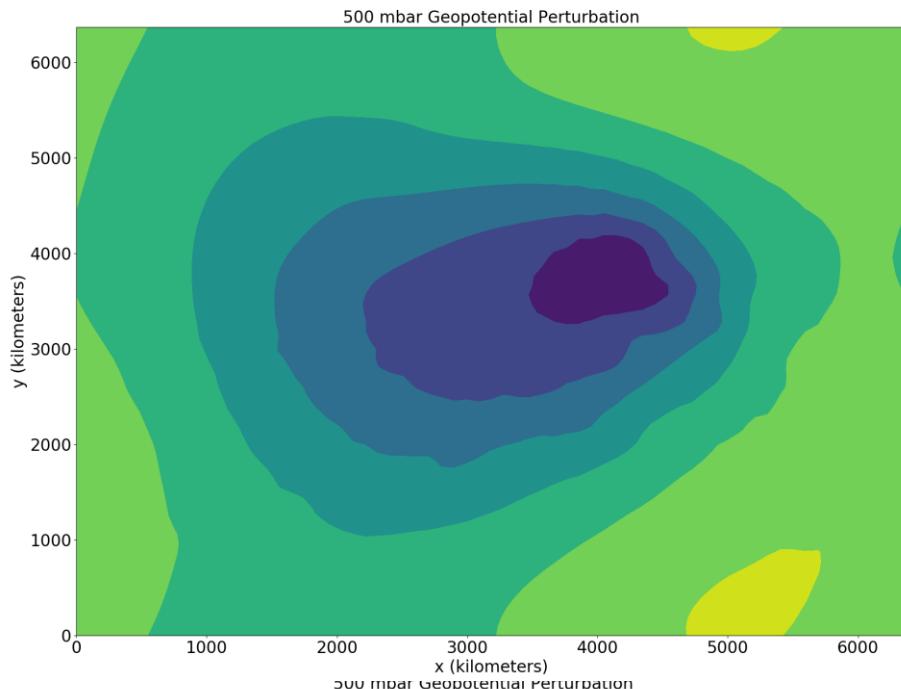
500 mb pressure anomaly



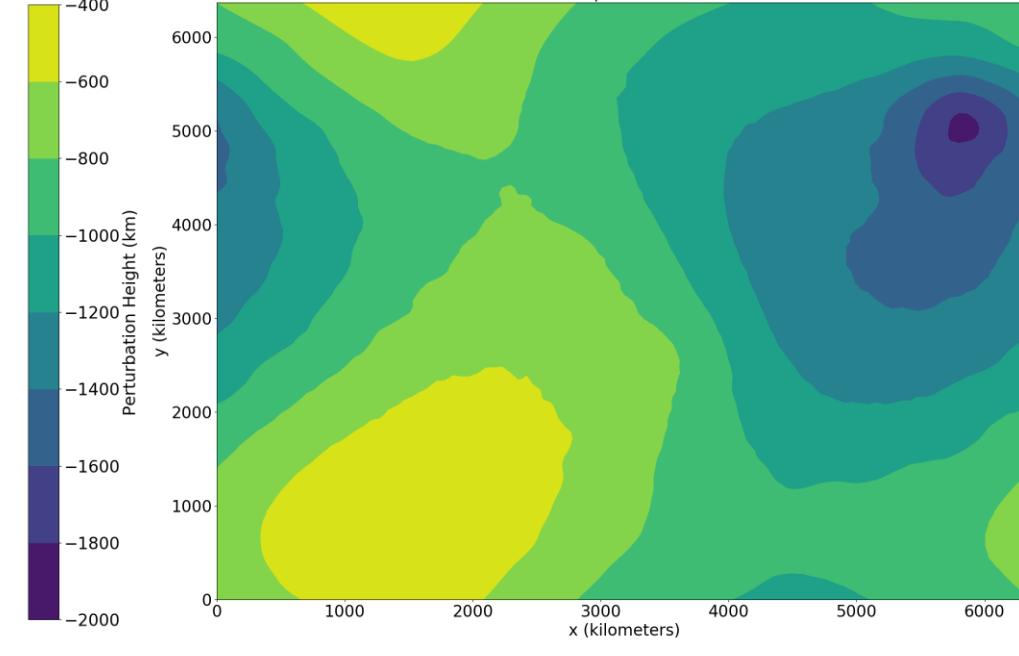
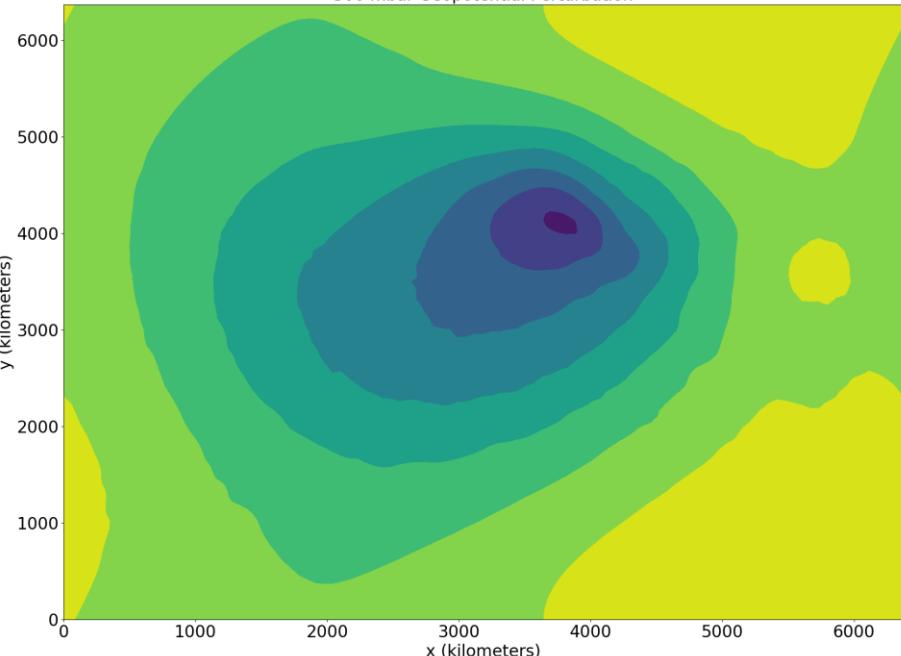
Initial geopotential height field



96.53 km  
Resolution,  
66x66 grid



63.71 km  
Resolution,  
100x100 grid

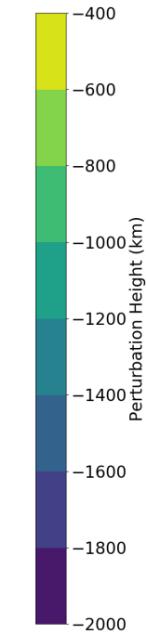
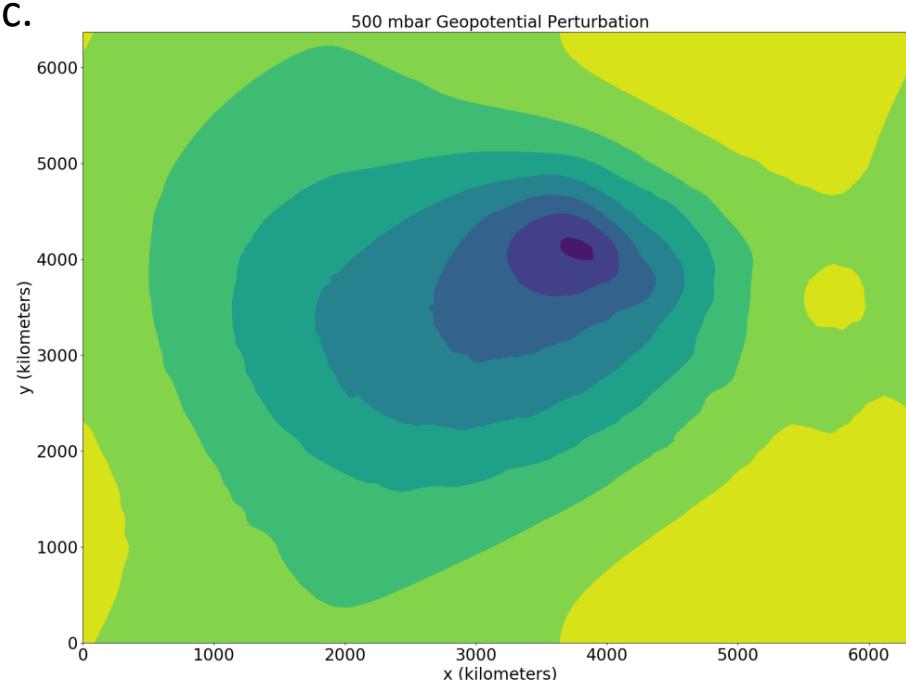


100x100 grid,  
50 m/s  
Mean zonal  
wind,  
Symmetric B.C.

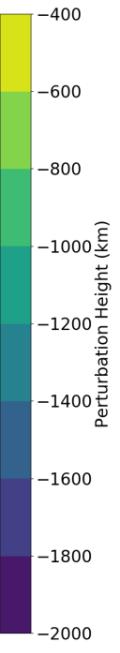
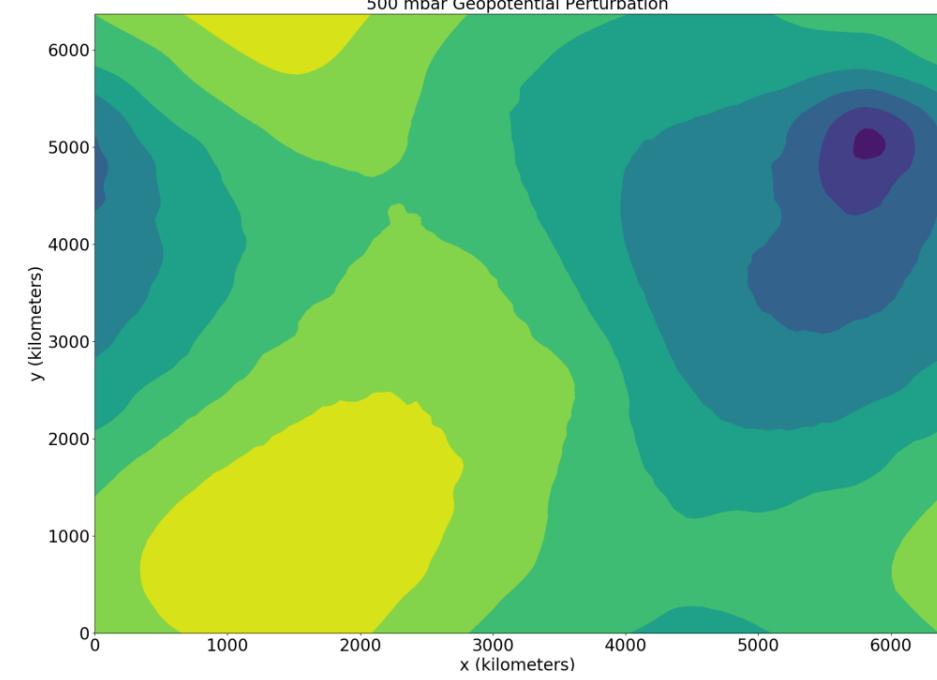
End of 1 Day

End of 2 Days

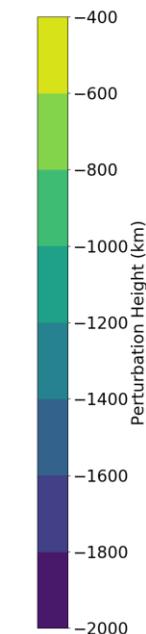
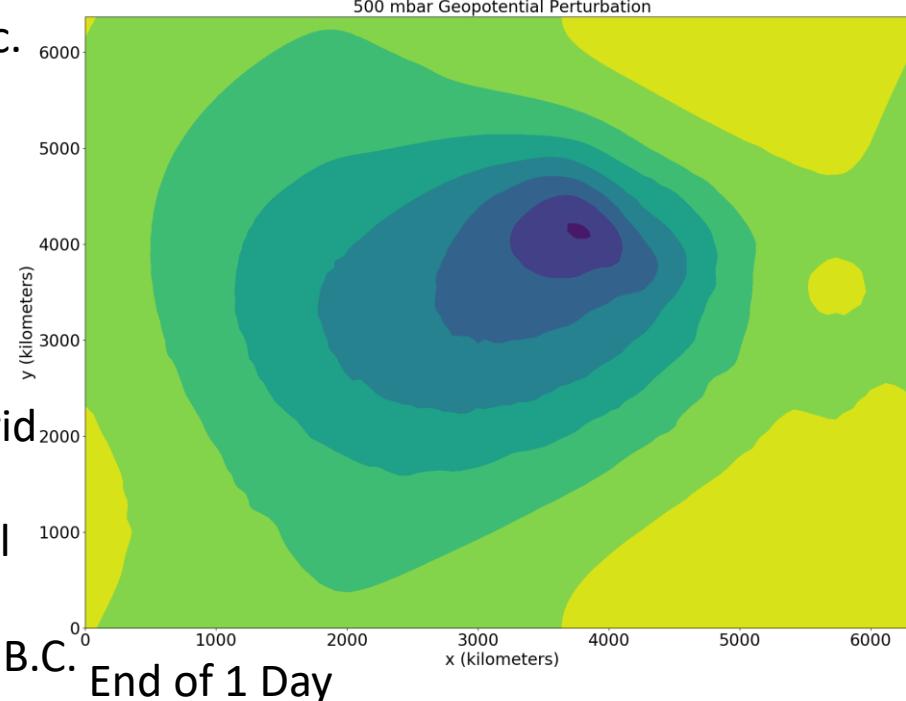
$\Delta t=100$  sec.



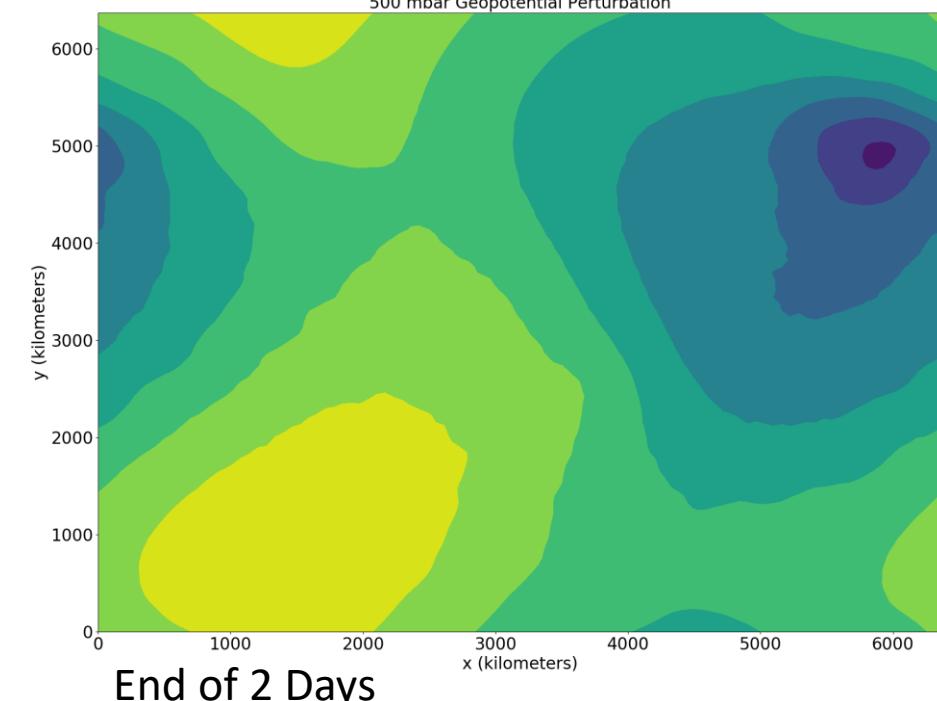
500 mbar Geopotential Perturbation



$\Delta t=37.5$  sec.



500 mbar Geopotential Perturbation



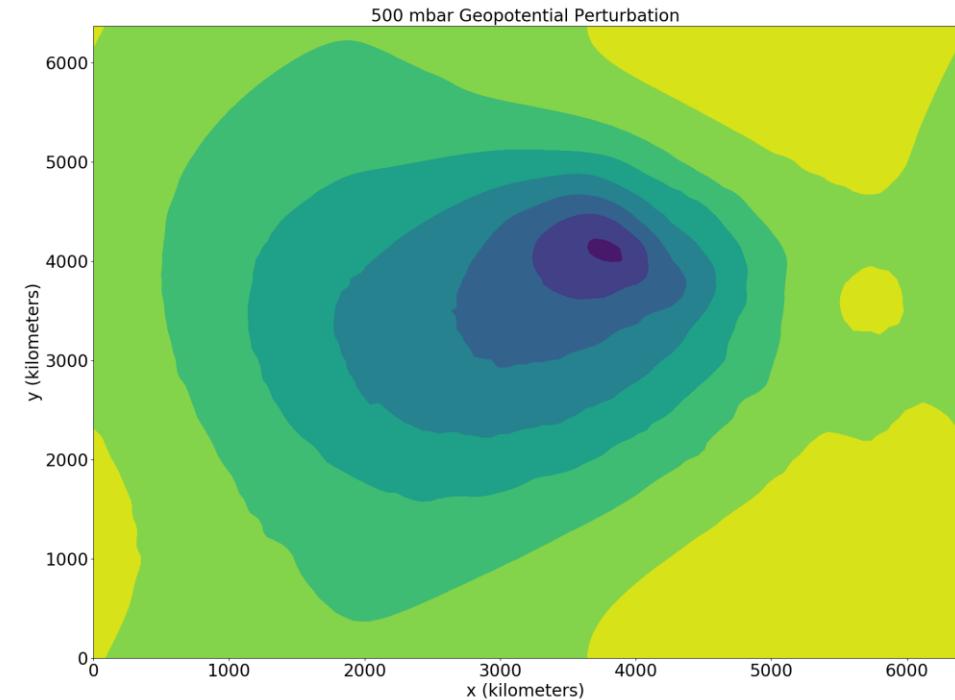
100x100 grid  
50 m/s  
Mean zonal  
wind,  
Symmetric B.C.

End of 1 Day

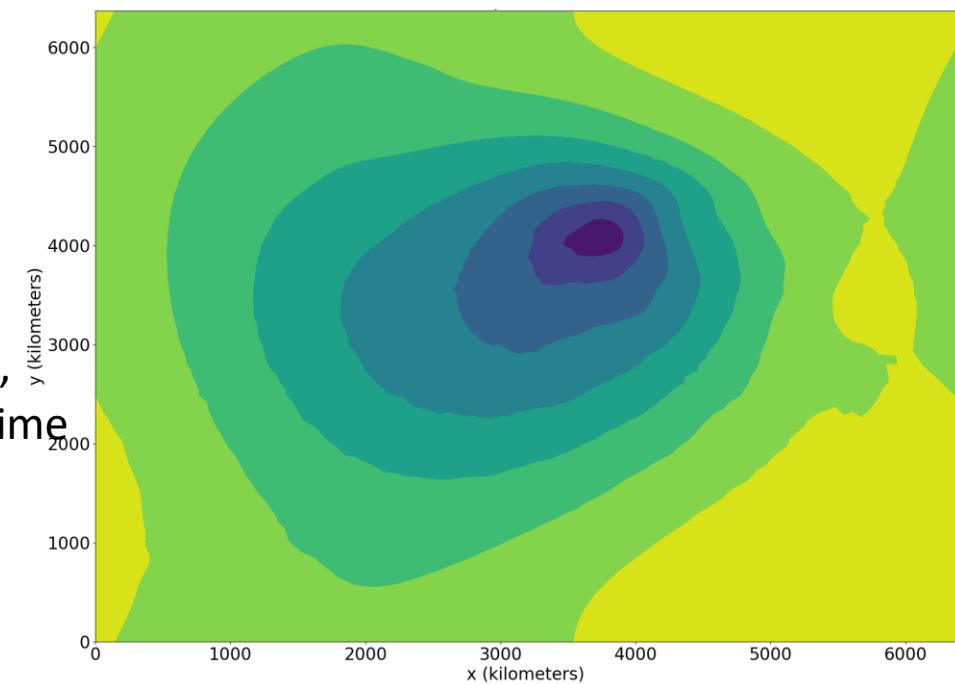
End of 2 Days

# Periodic vs Symmetric B.C.'s

Symmetric  
B.C.

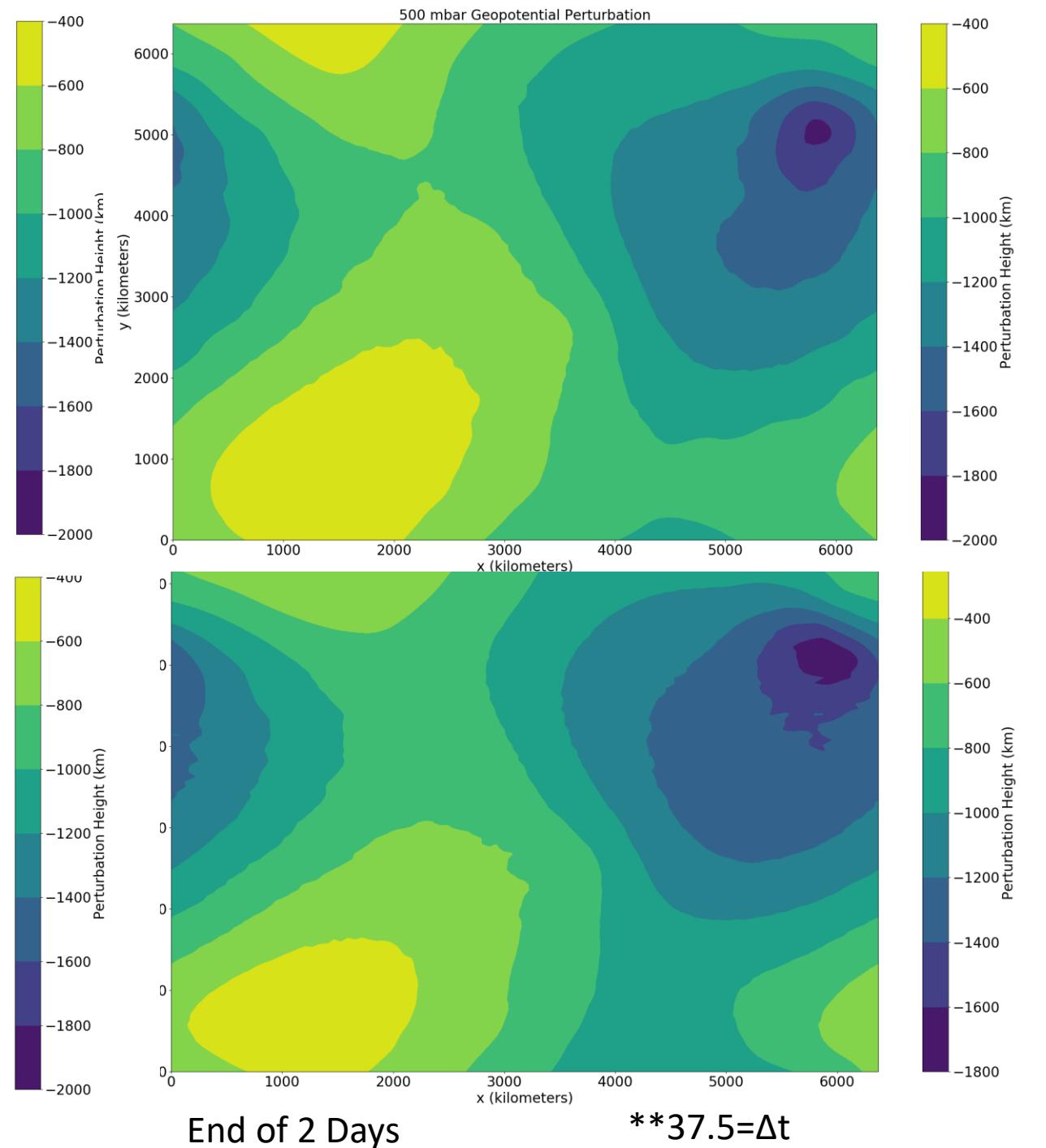


Periodic  
B.C.



100x100 grid,  
100 second time  
step,  
50 m/s  
Mean zonal  
wind

End of 1 Day



End of 2 Days

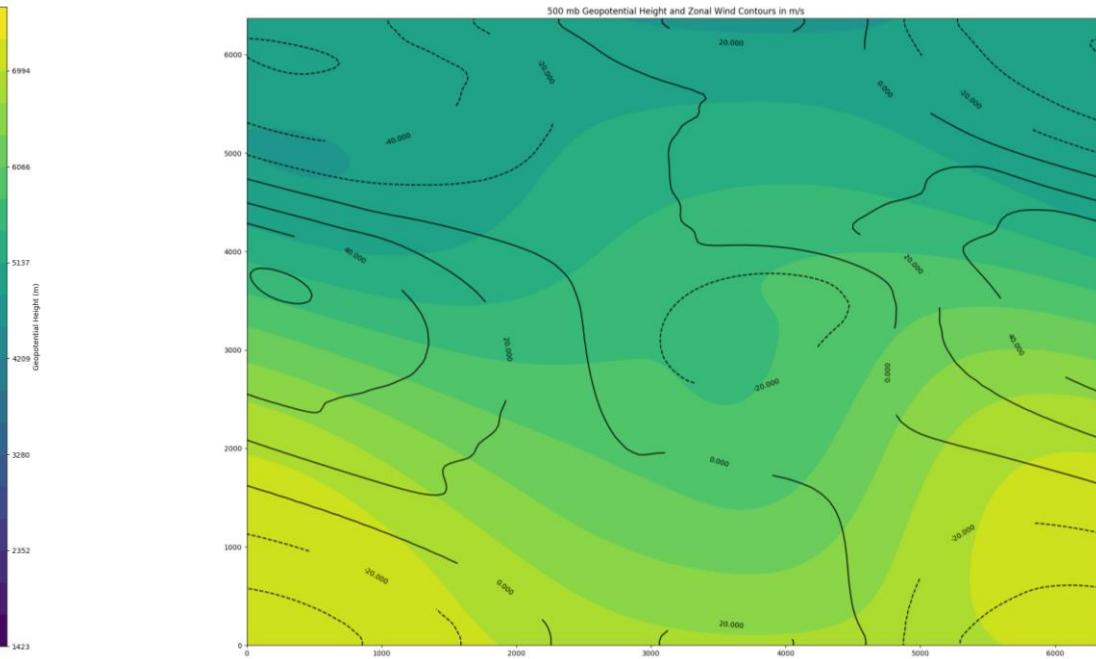
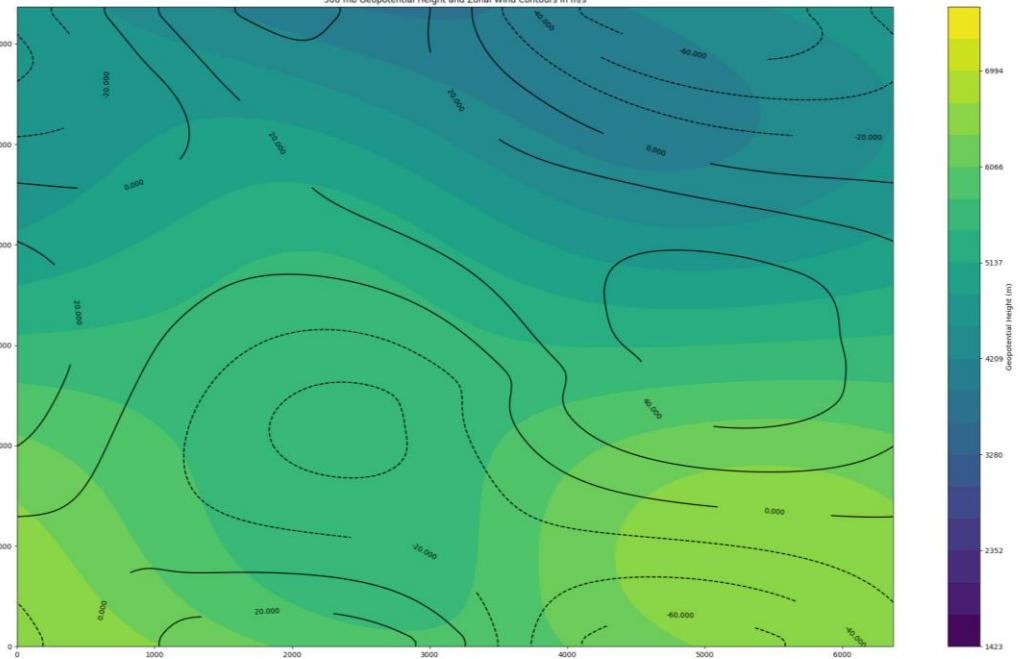
$\text{**}37.5 = \Delta t$

How does the zonal wind average change as the background wind state is modulated?

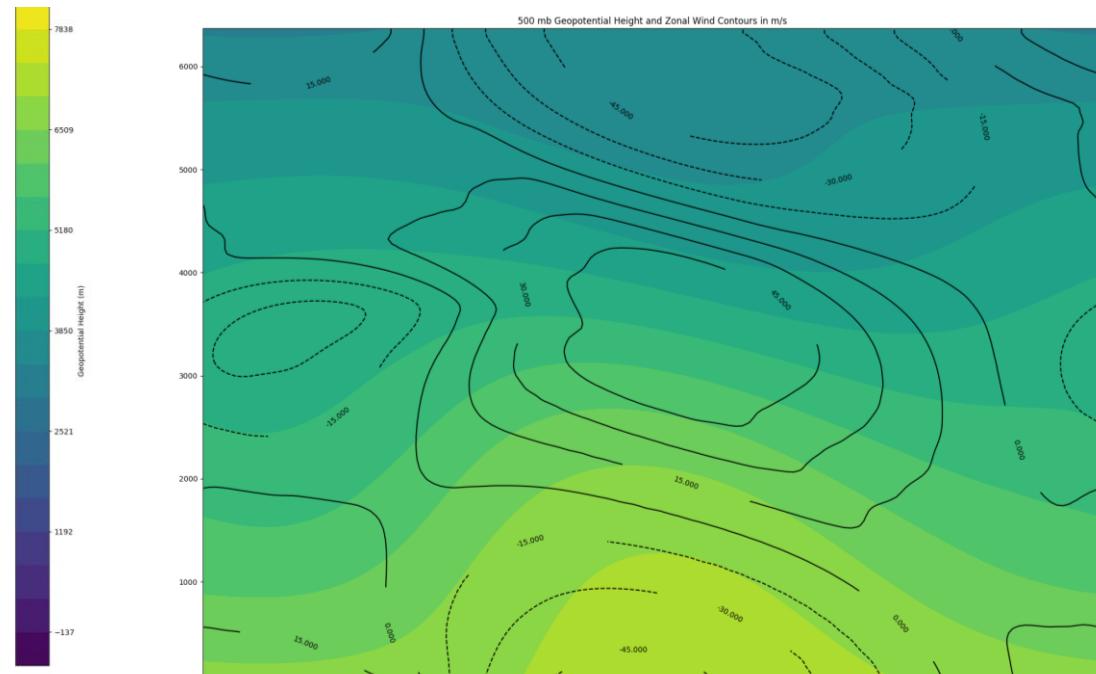
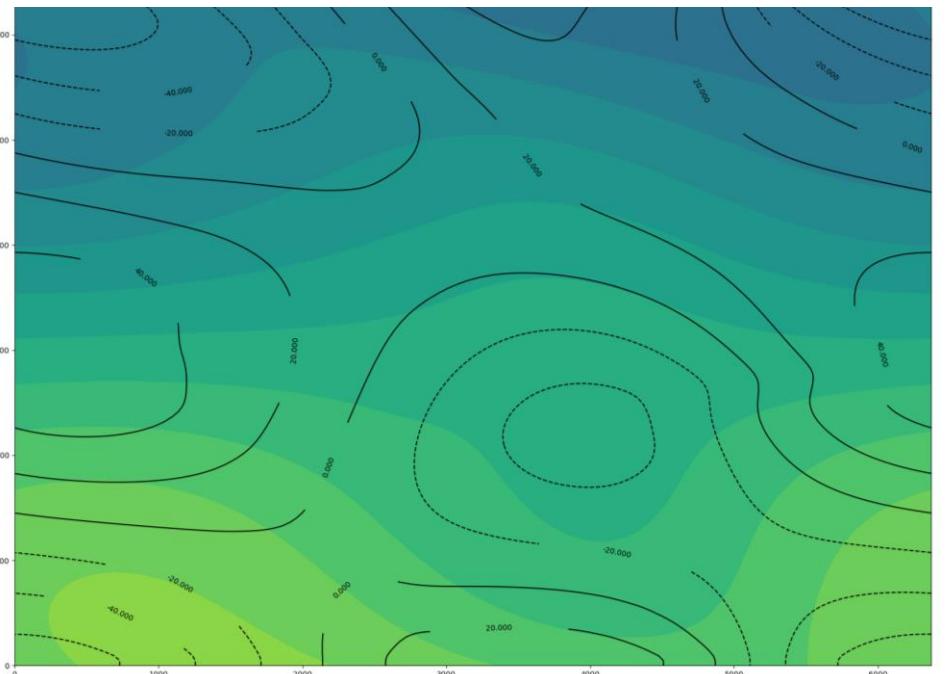
Wave Initial Condition - Symmetric B.C., 1/96<sup>th</sup>  
time step, 140x140 grid

Gaussian Initial Condition – Symmetric B.C, 1/12<sup>th</sup>  
time step, 66x66 grid

30 m/s



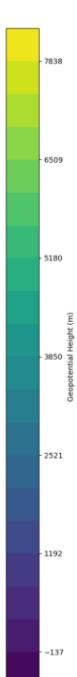
50 m/s



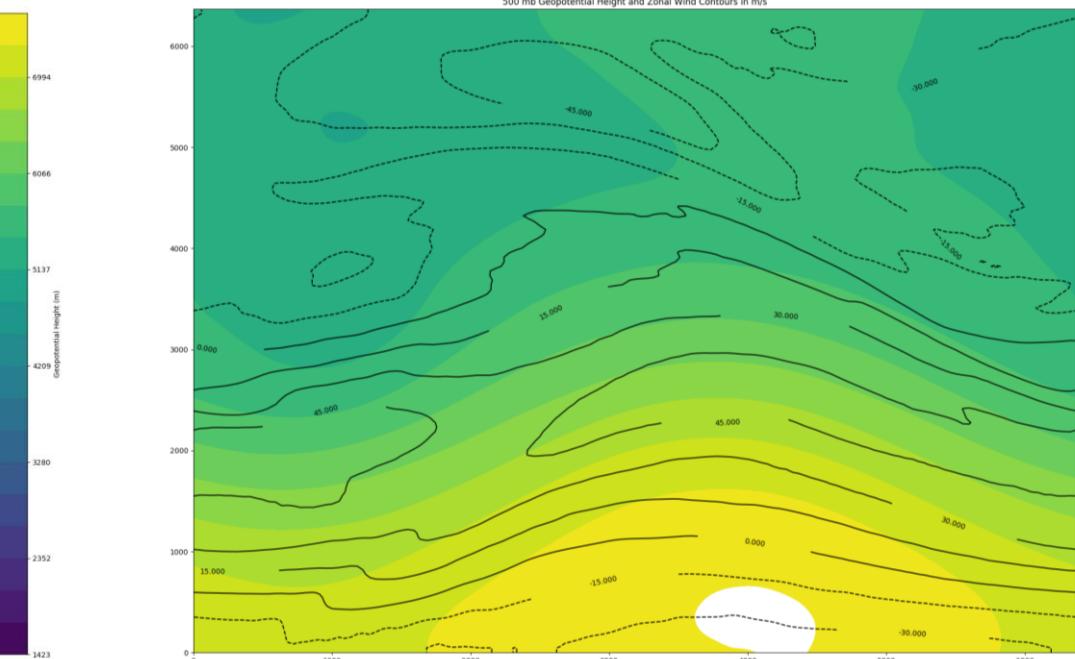
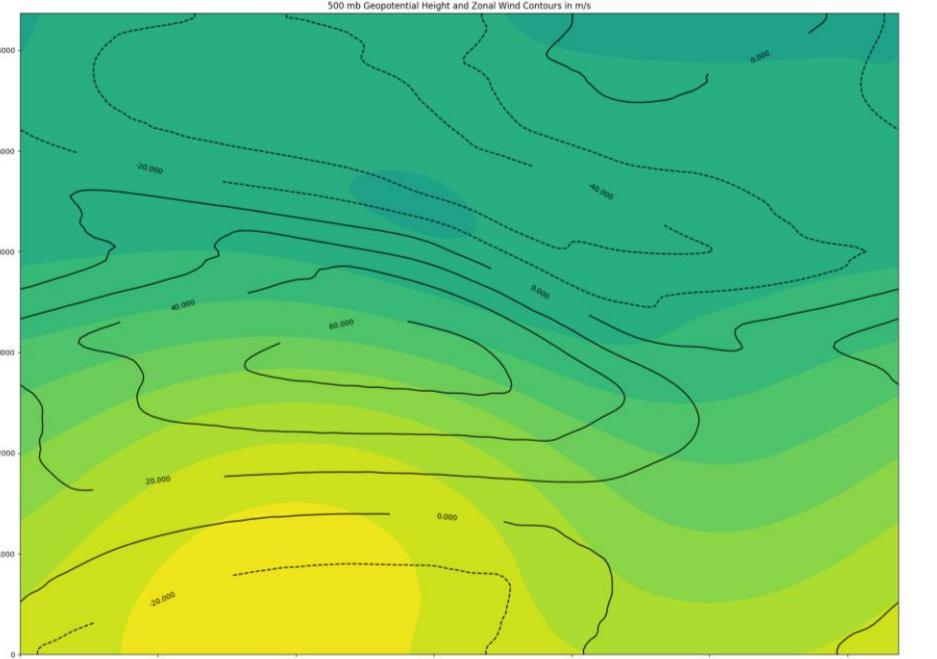
100x100 grid,  
50 m/s  
Mean zonal  
wind

End of 1 Day

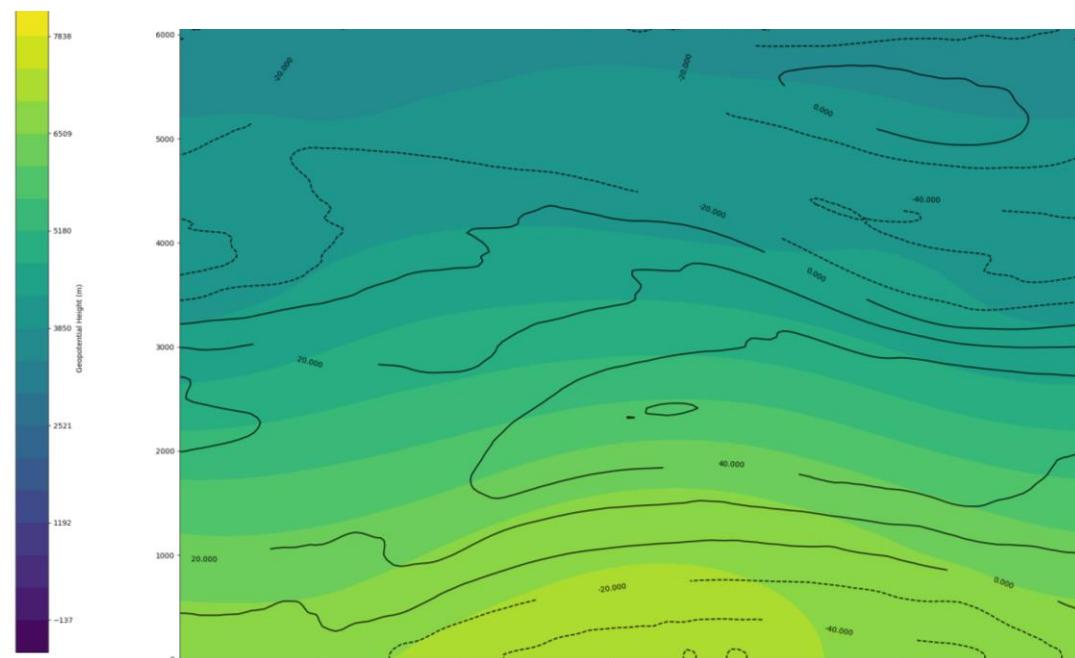
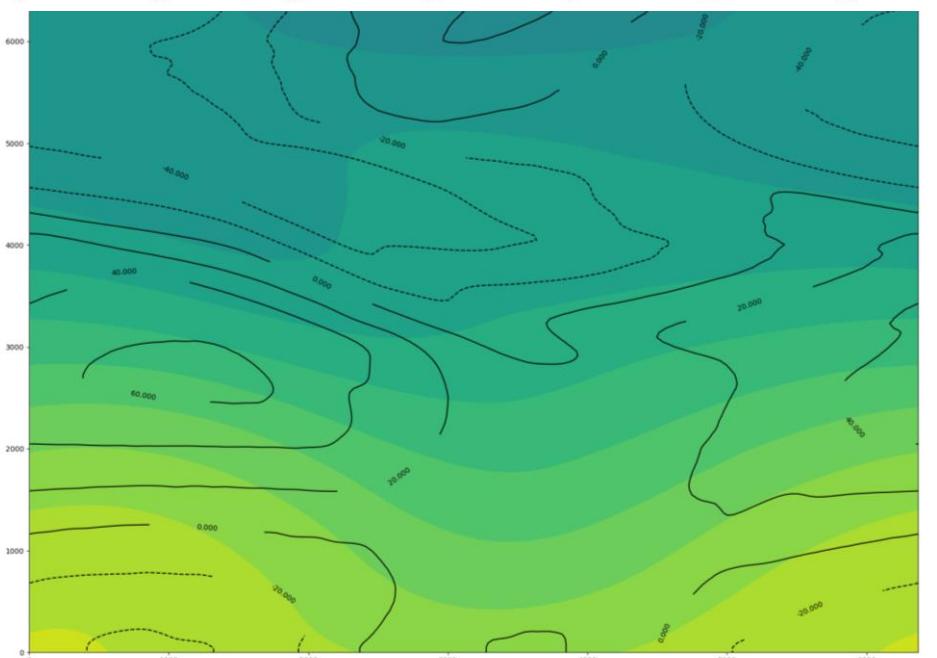
End of 2 Days



30 m/s



50 m/s

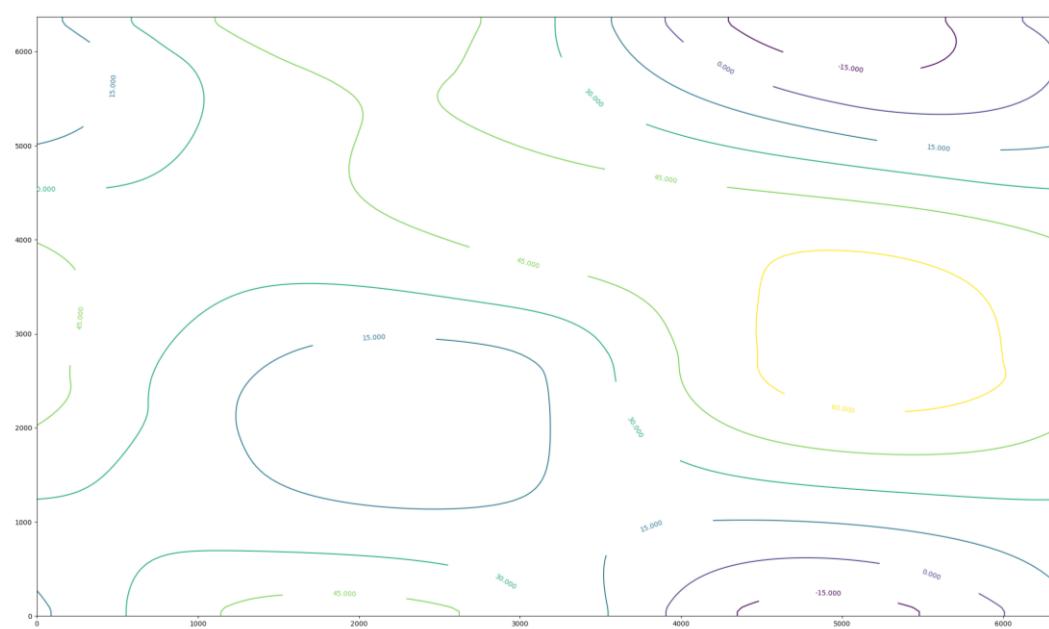
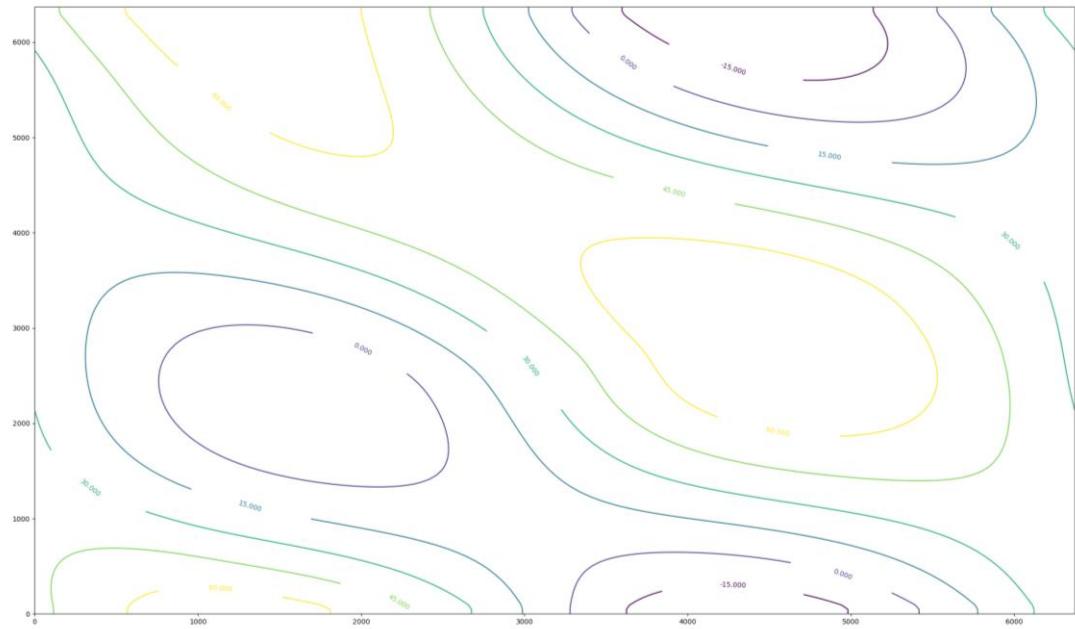


100x100 grid,  
50 m/s  
Mean zonal  
wind

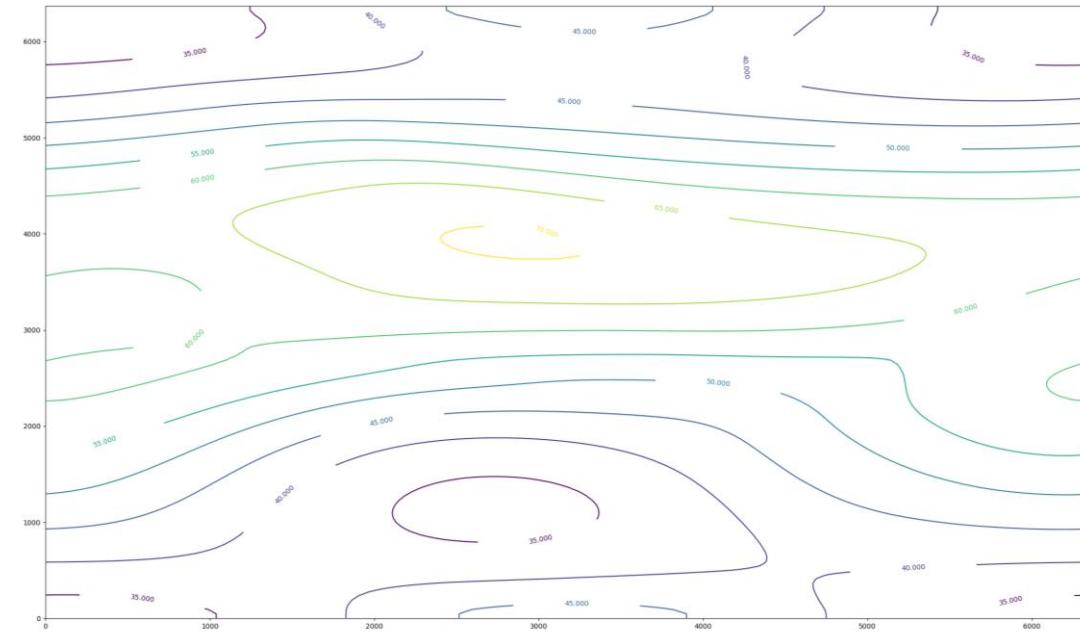
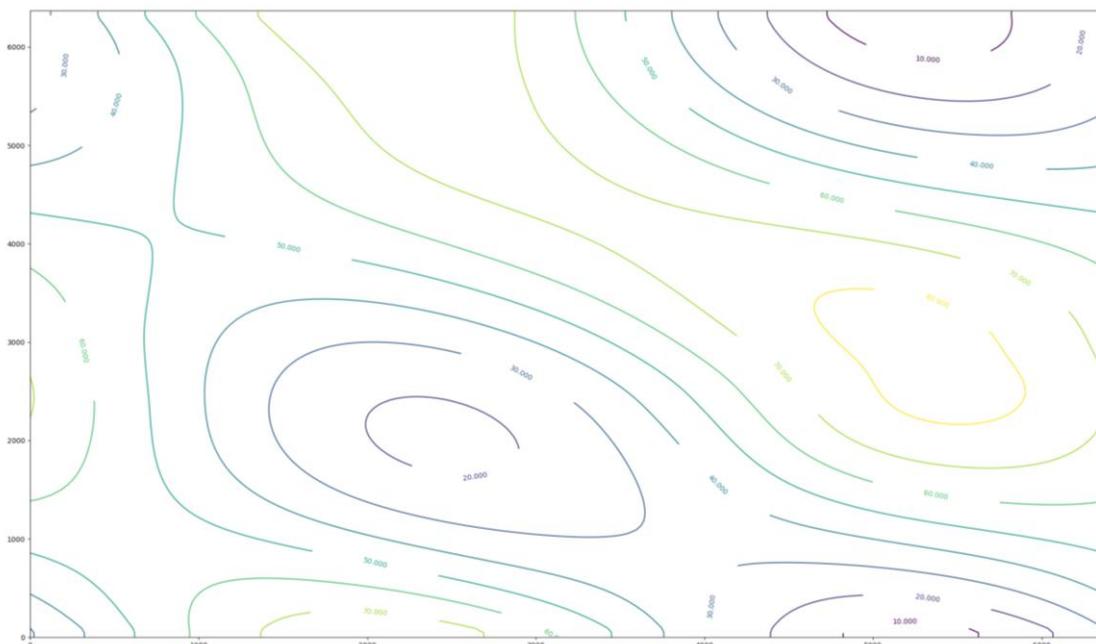
End of 3 Days

End of 4 Days

30 m/s



50 m/s

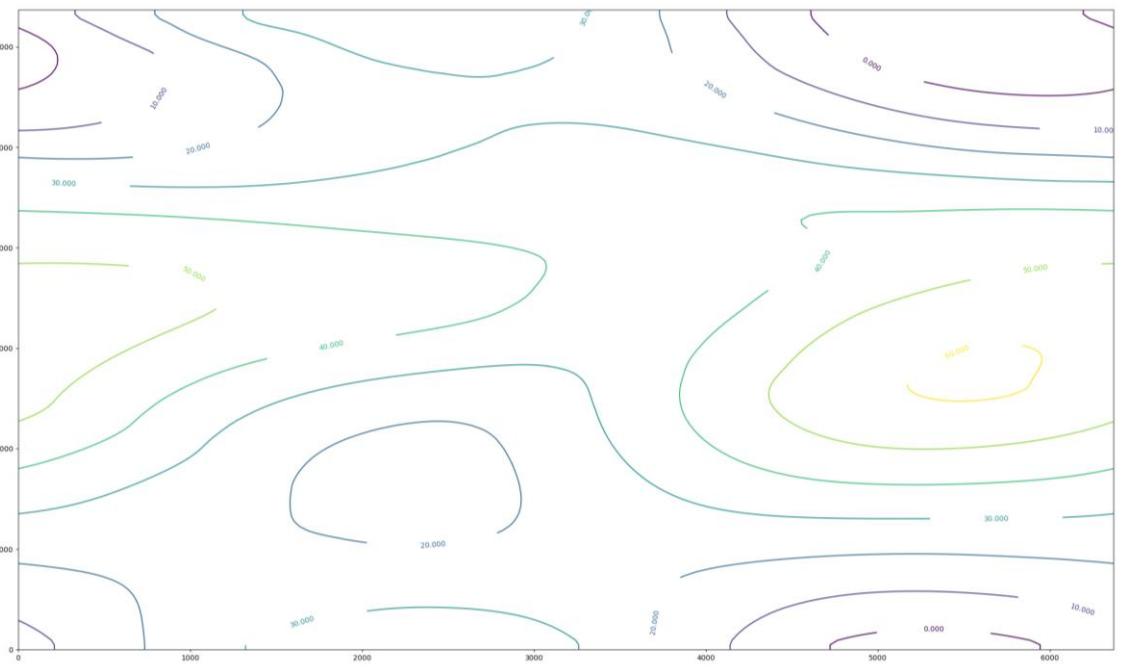


100x100 grid,  
50 m/s  
Mean zonal  
wind

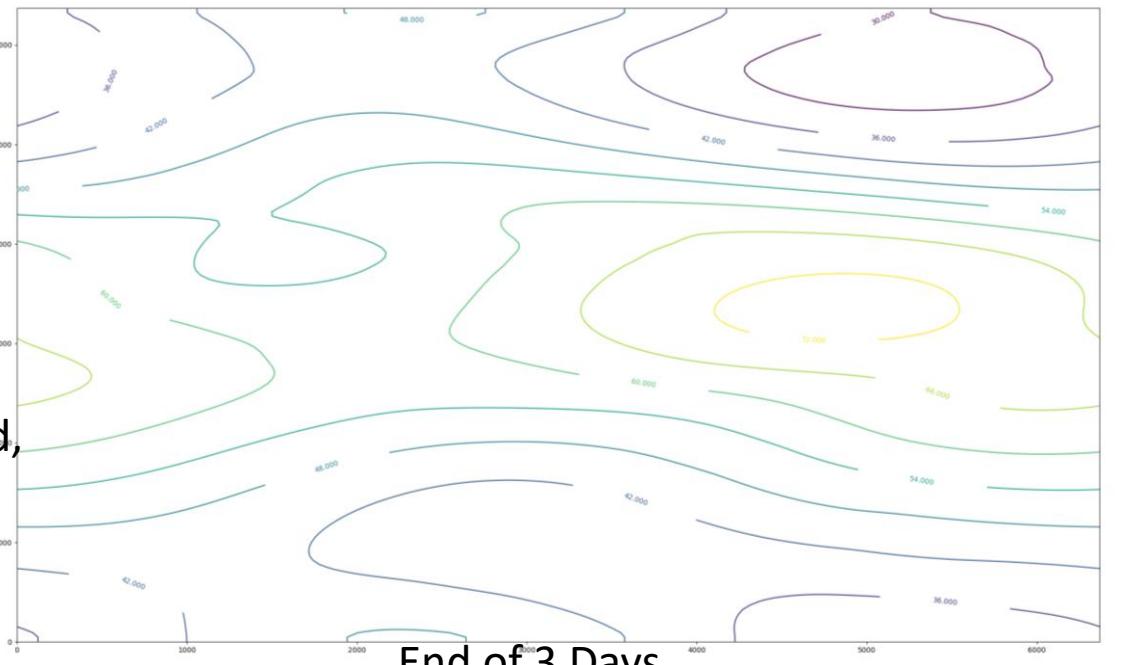
End of 1 Day

End of 2 Days

30 m/s

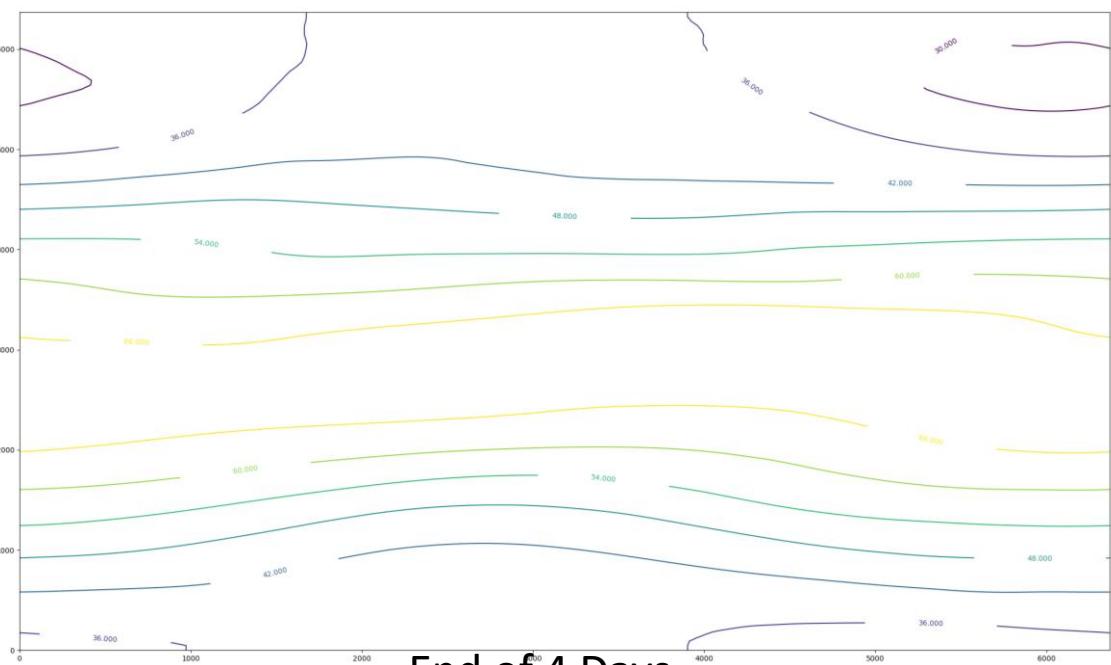
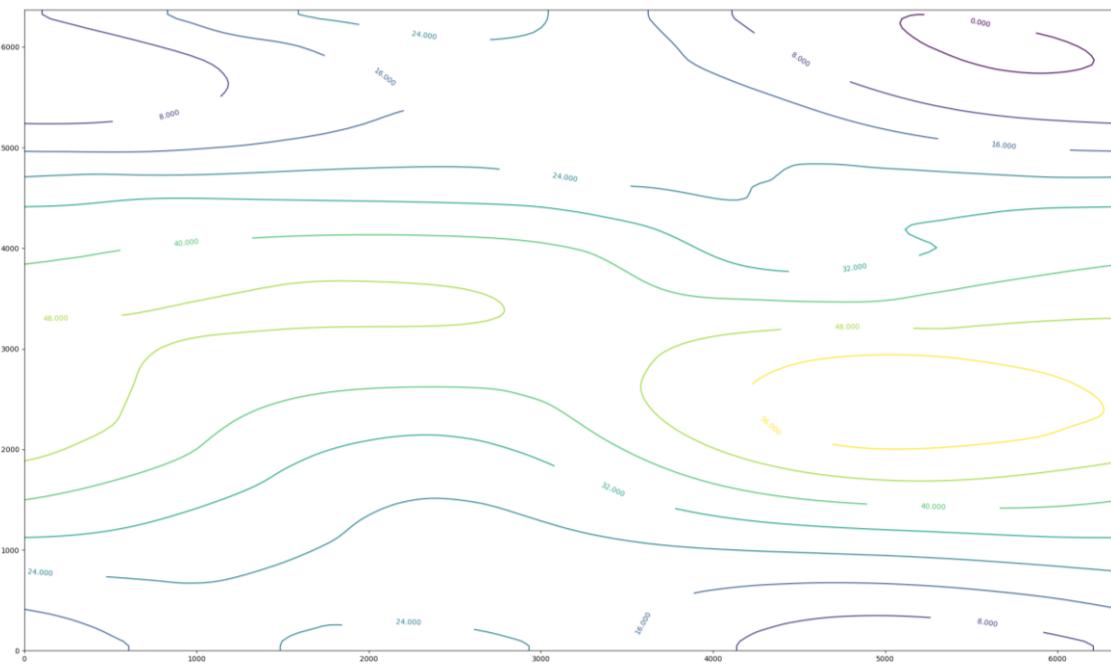


50 m/s



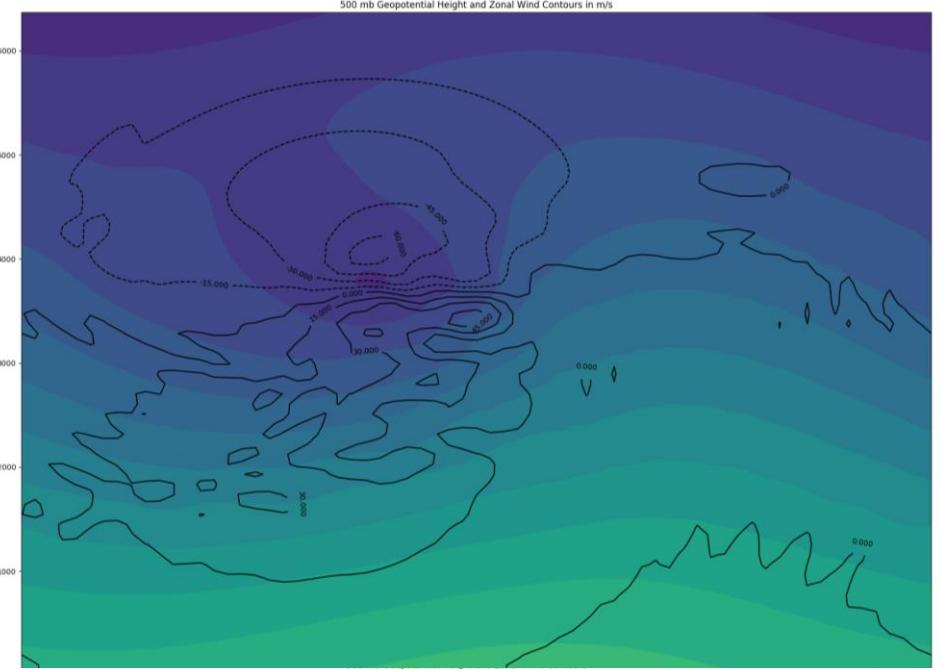
100x100 grid,  
50 m/s  
Mean zonal  
wind

End of 3 Days

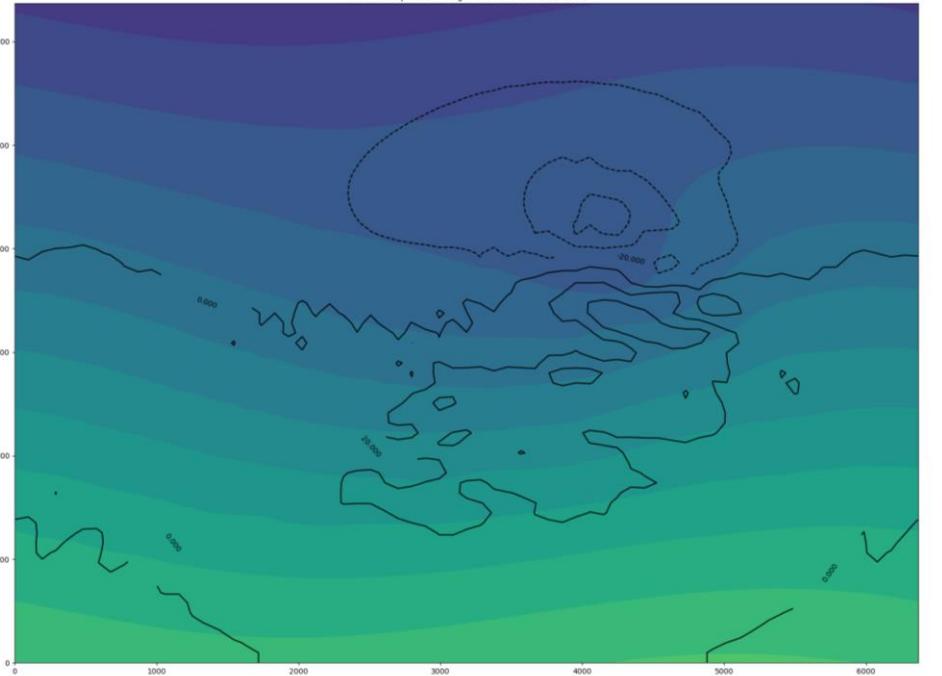


End of 4 Days

30 m/s

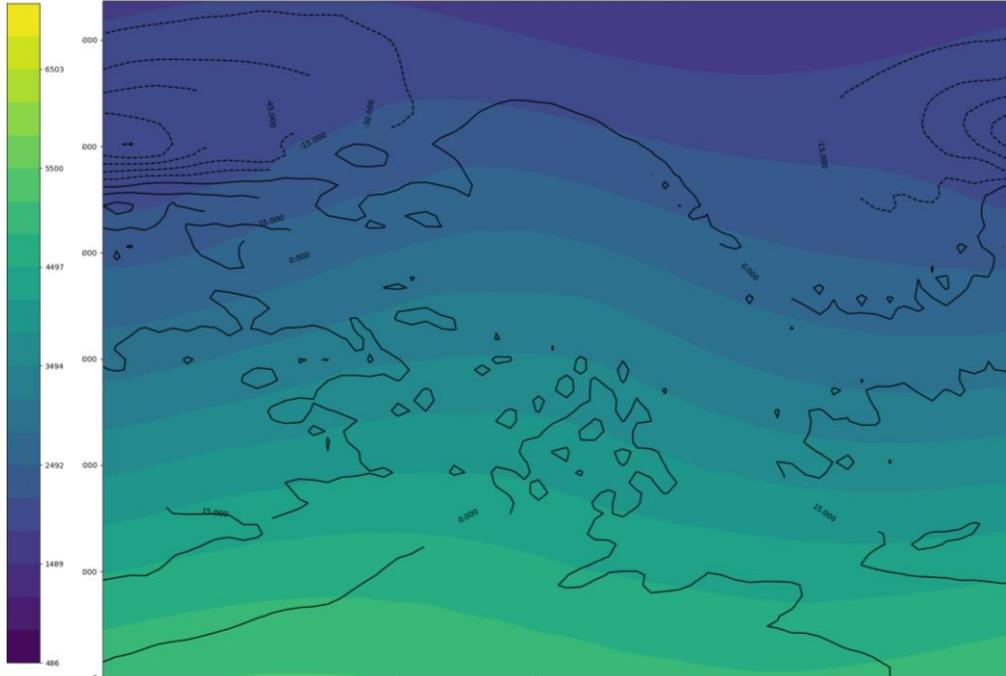
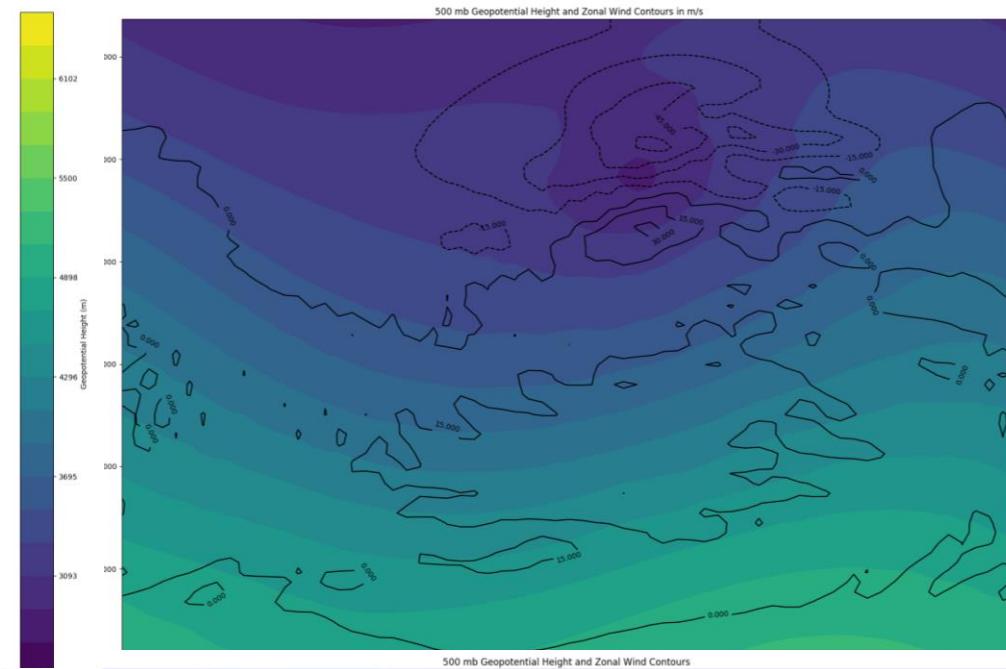


50 m/s



100x100 grid,  
50 m/s  
Mean zonal  
wind

End of 1 Day

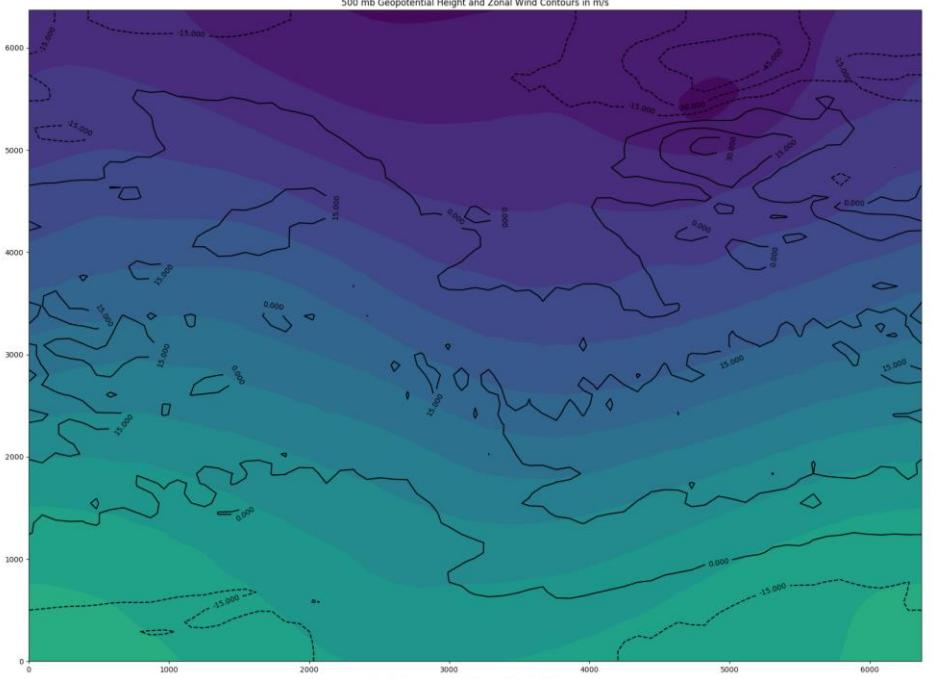


End of 2 Days

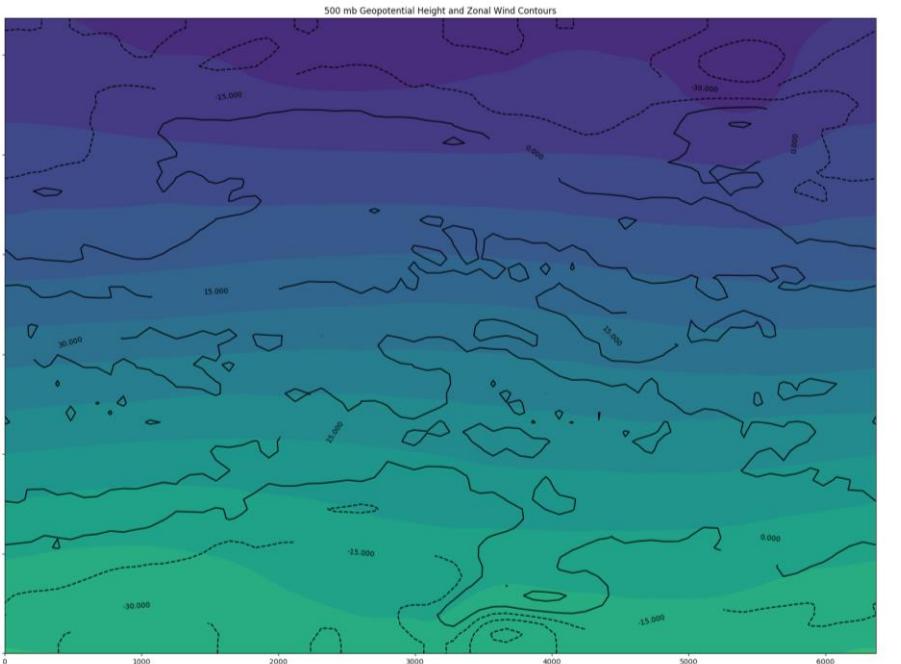
Geopotential Height (m)

Geopotential Height (m)

30 m/s

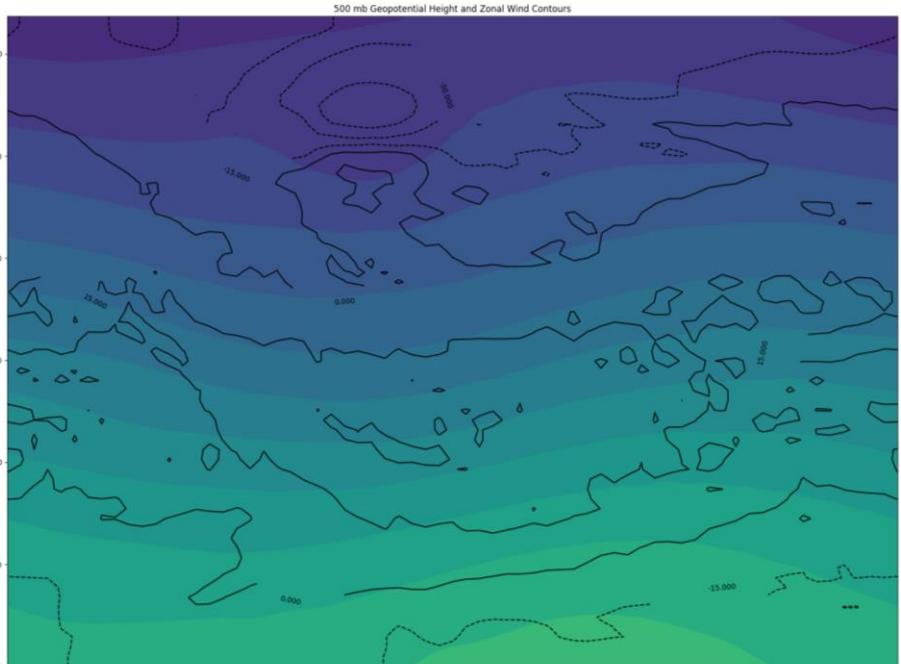
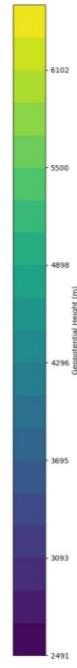
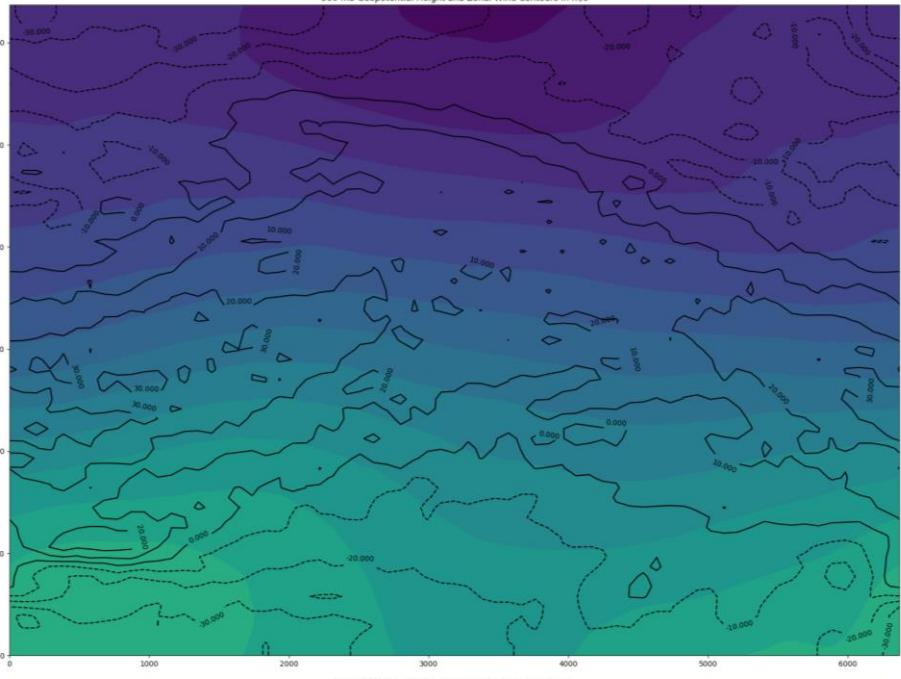


50 m/s



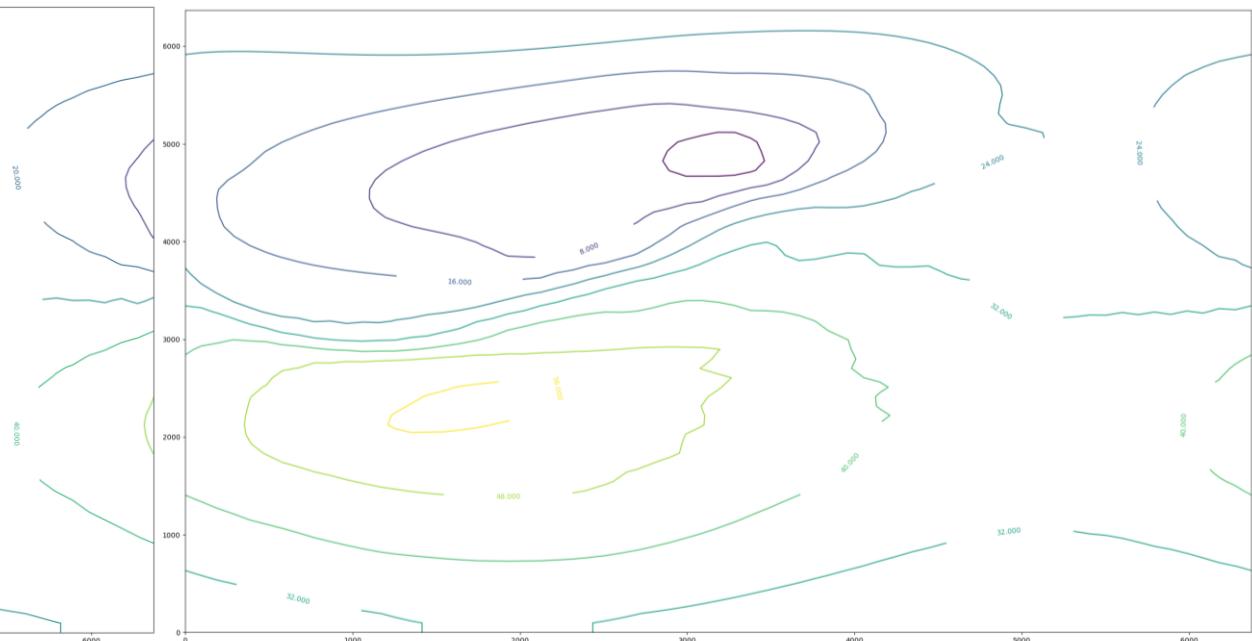
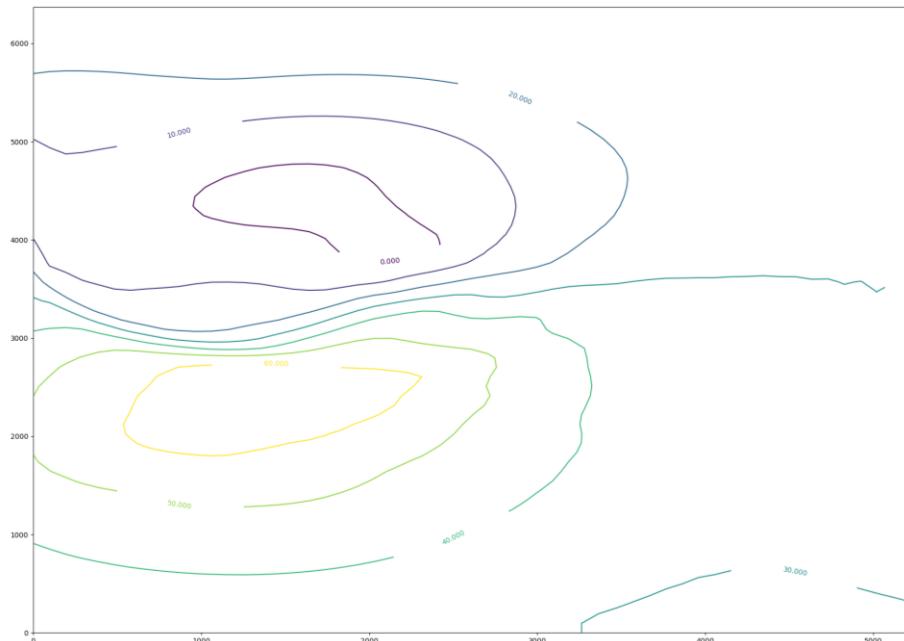
100x100 grid,  
50 m/s  
Mean zonal  
wind

End of 3 Days

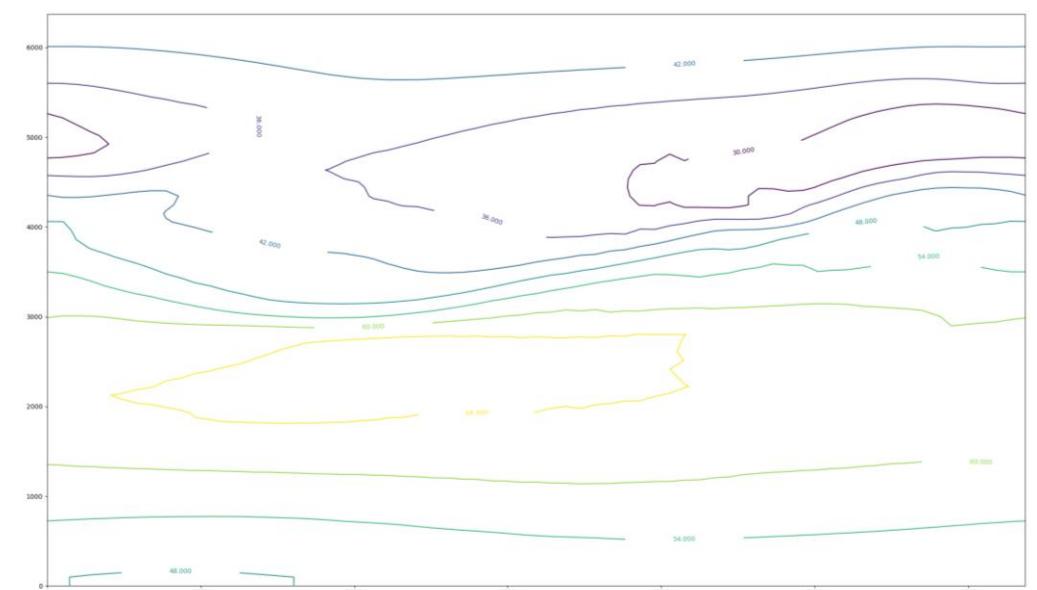
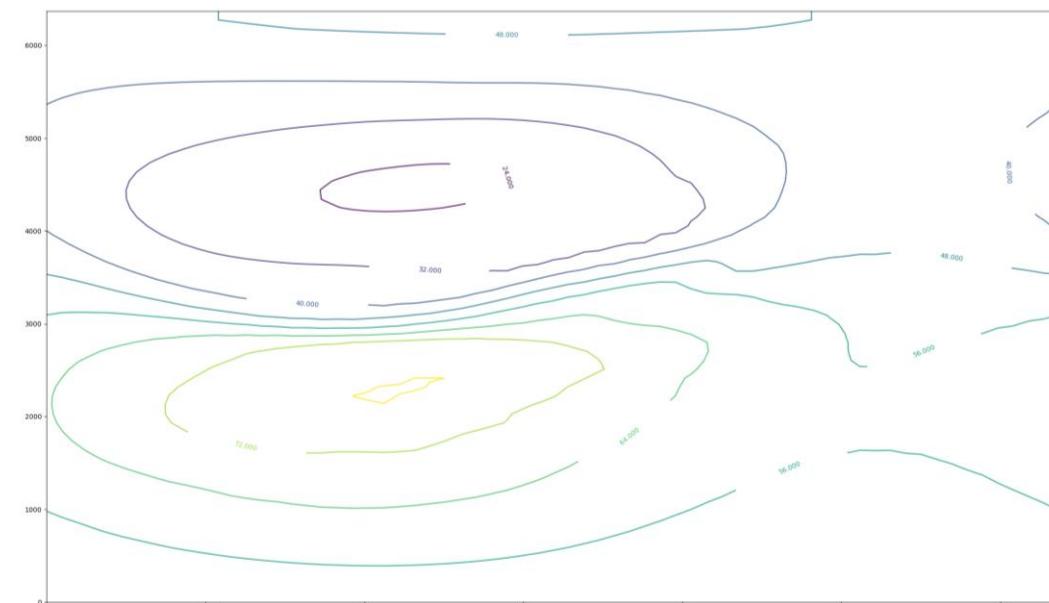


End of 4 Days

30 m/s



50 m/s

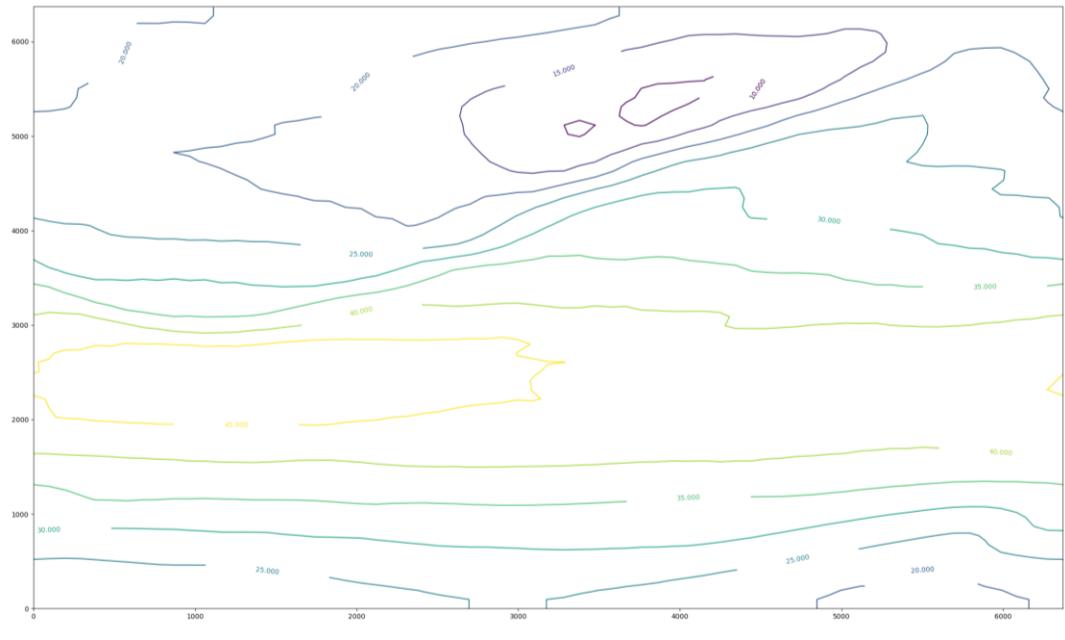


100x100 grid,  
50 m/s  
Mean zonal  
wind

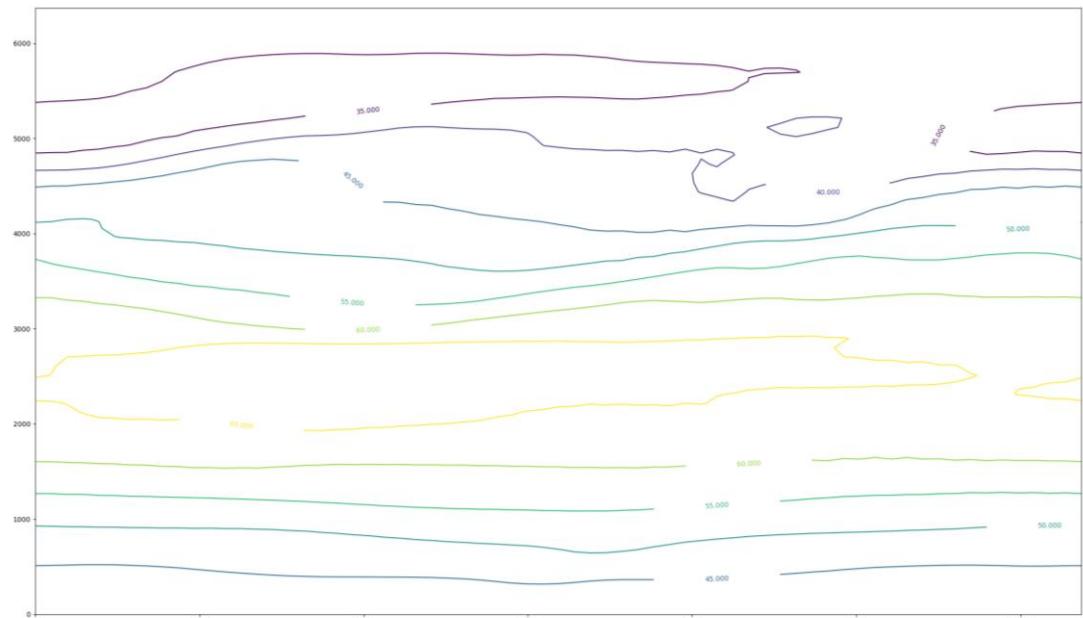
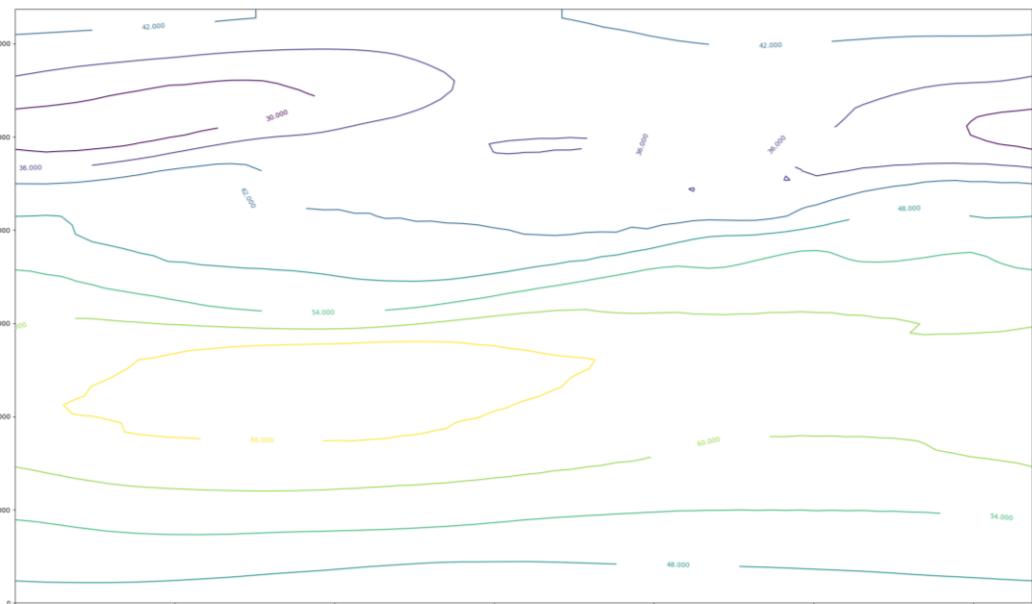
End of 1 Day

End of 2 Days

30 m/s



50 m/s



100x100 grid  
50 m/s  
Mean zonal  
wind

End of 3 Day

End of 4 Days