



Remote Sensing of Snow

Jakob Steiner

Date: 7 April 2022

overview

Importance of snow

Crucial variables of snow

Remote sensing of snow

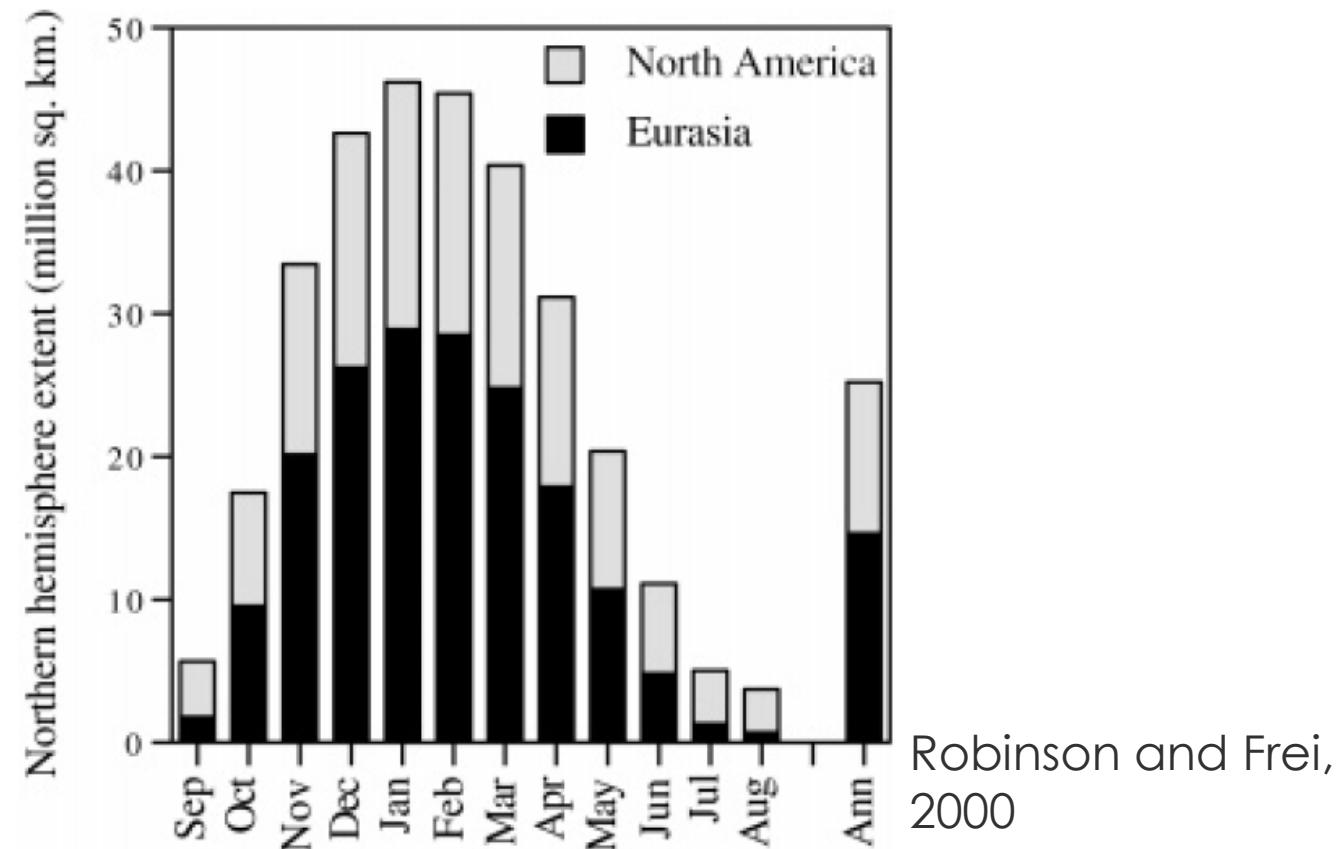
- Variables
- Active vs Passive sensors

What remote sensing can't see (yet)

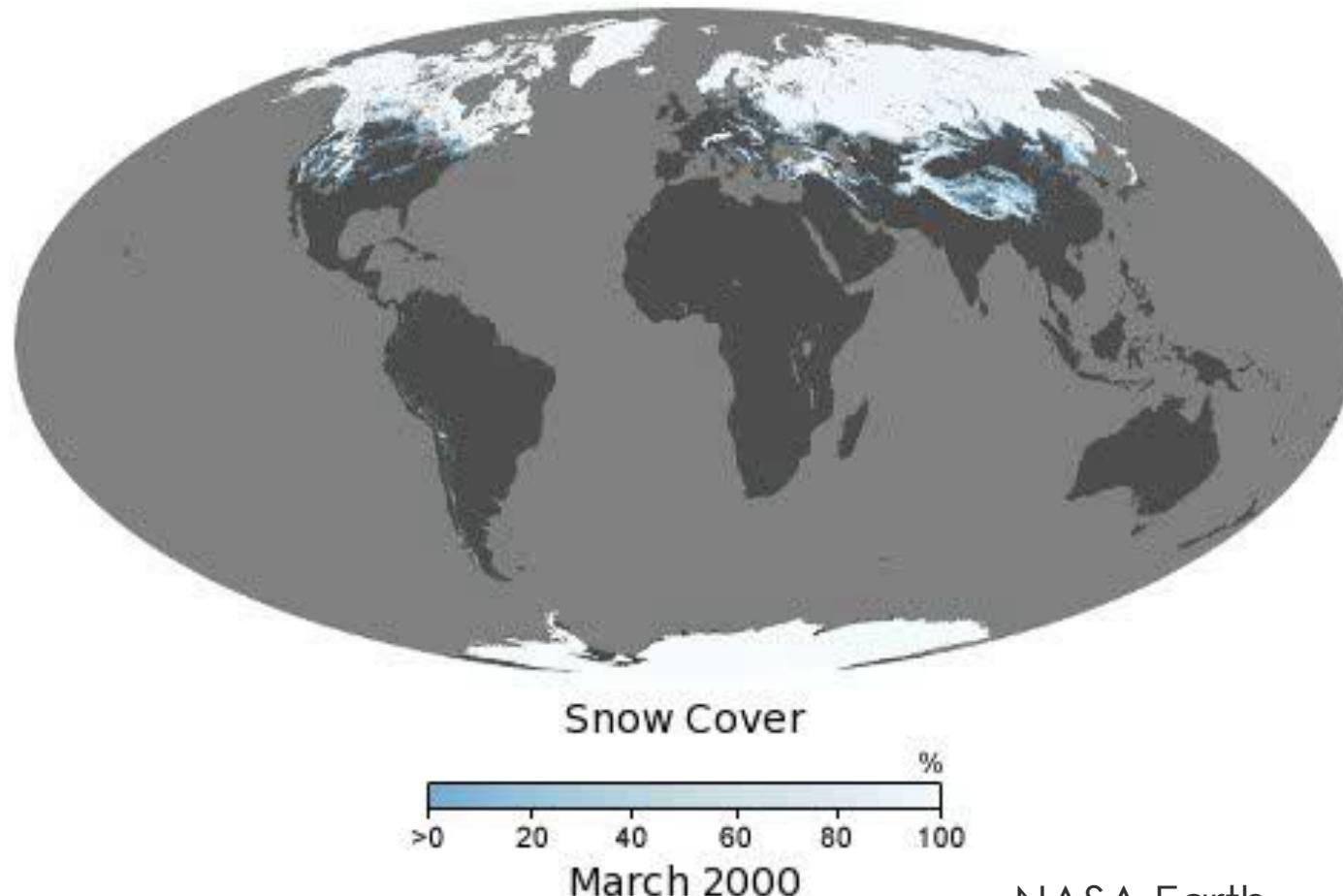


Global snow cover

Mean monthly SE (1972-1999)



Global snow cover



does it ever snow in Lahore?



DAWN, 2011

<https://www.dawn.com/news/609813/snow-day-in-lahore>

why does snow matter?

why does snow matter?

Hydrology

Climate

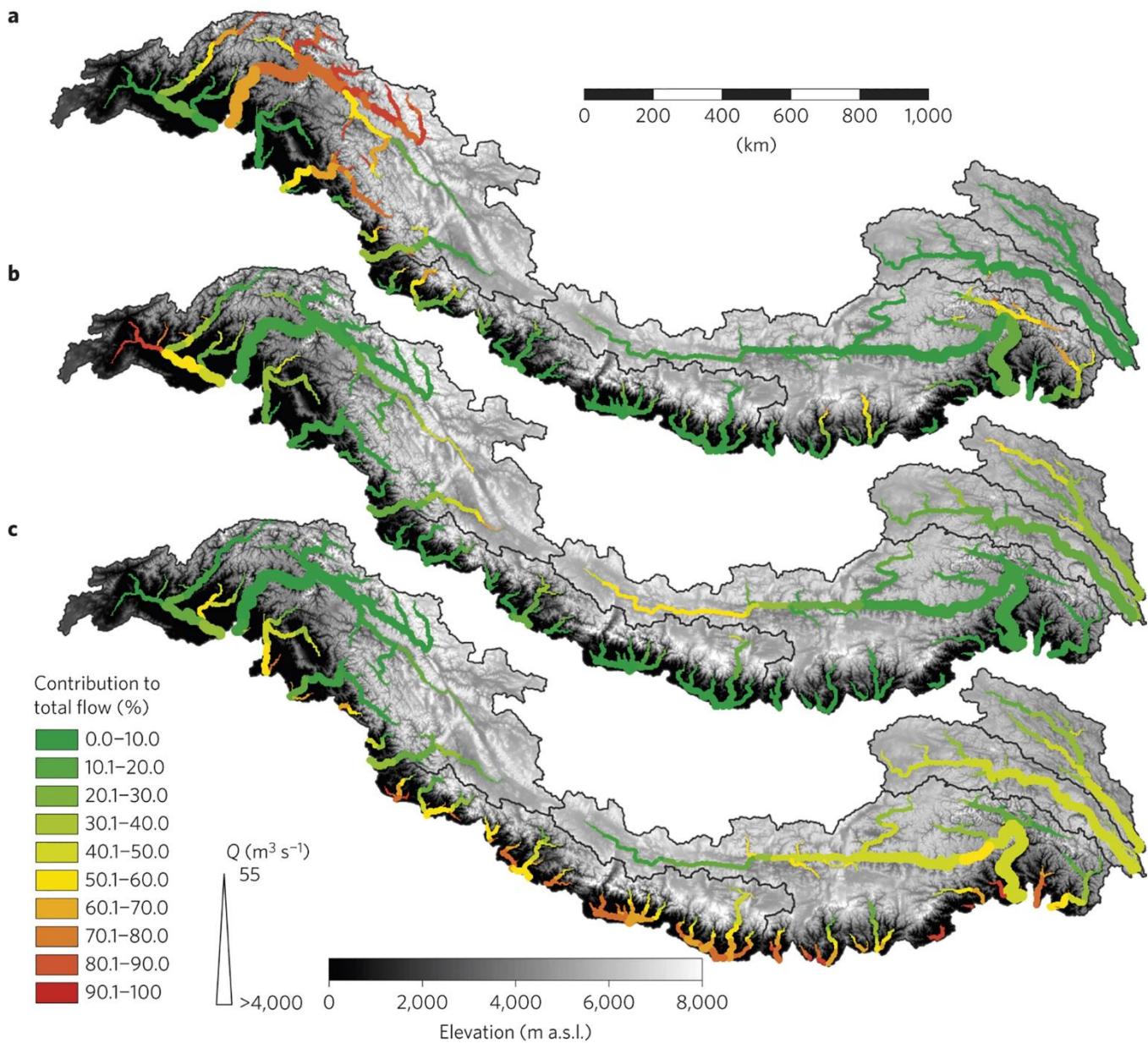
- albedo affects energy balance

Avalanches

Saturation of soils from melt

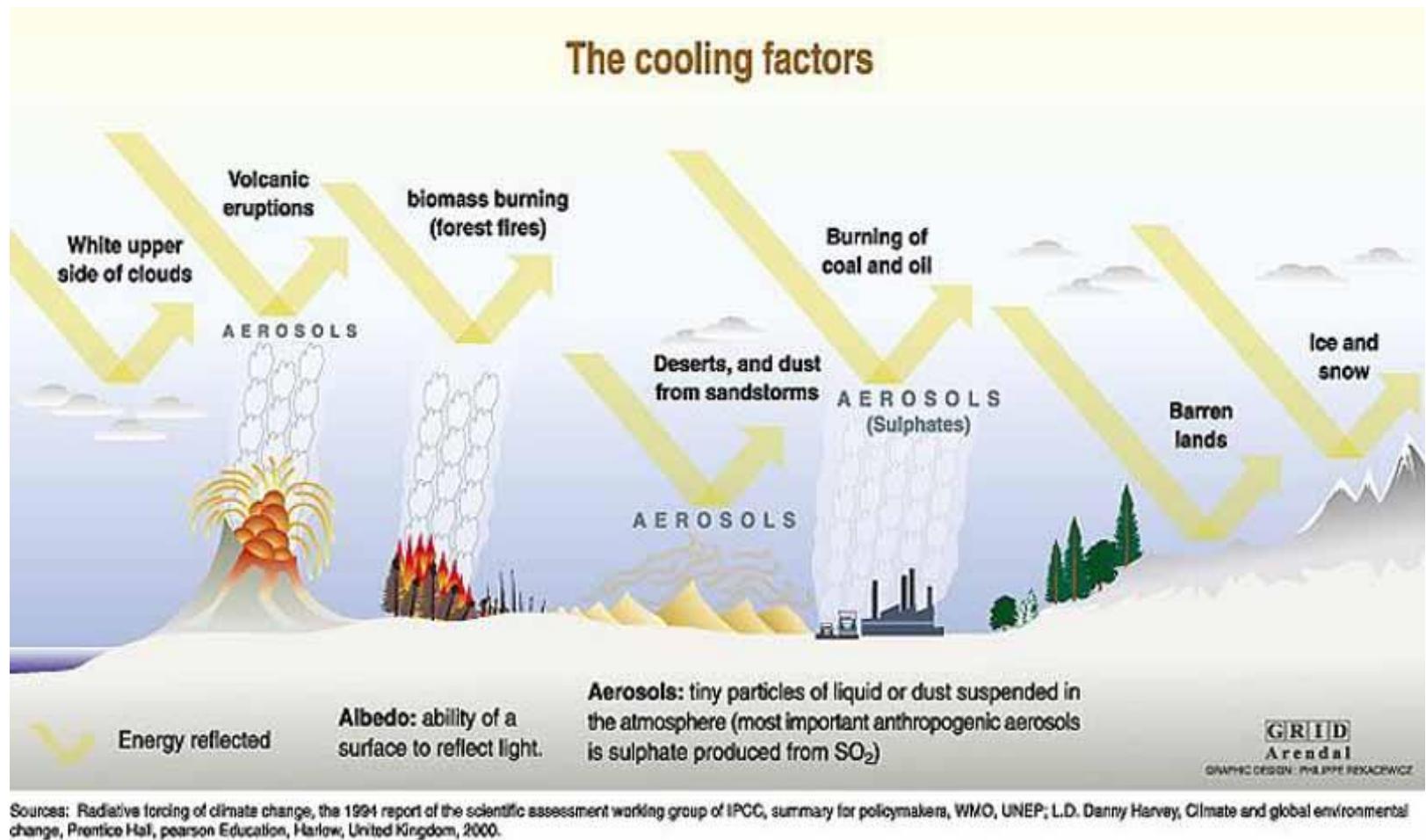
snow melt in discharge

- (a) glacier melt
- (b) snow melt
- (c) rainfall



albedo effect

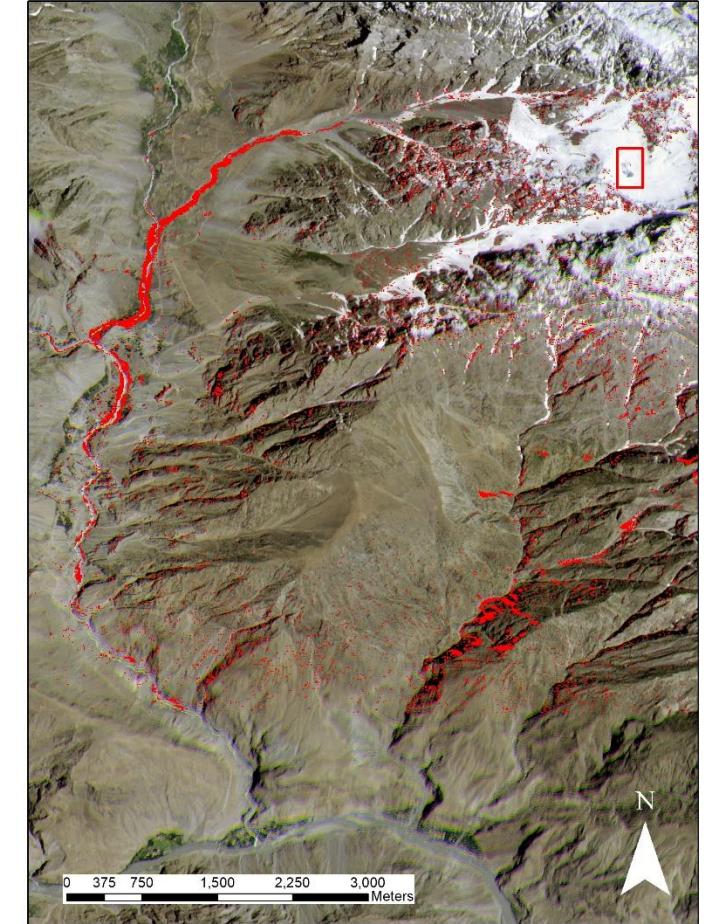
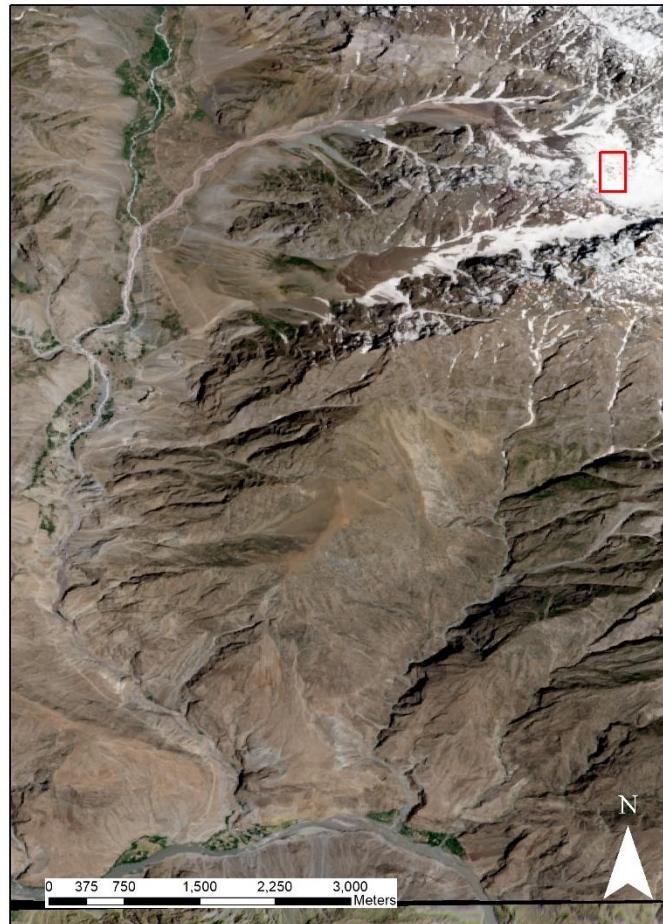
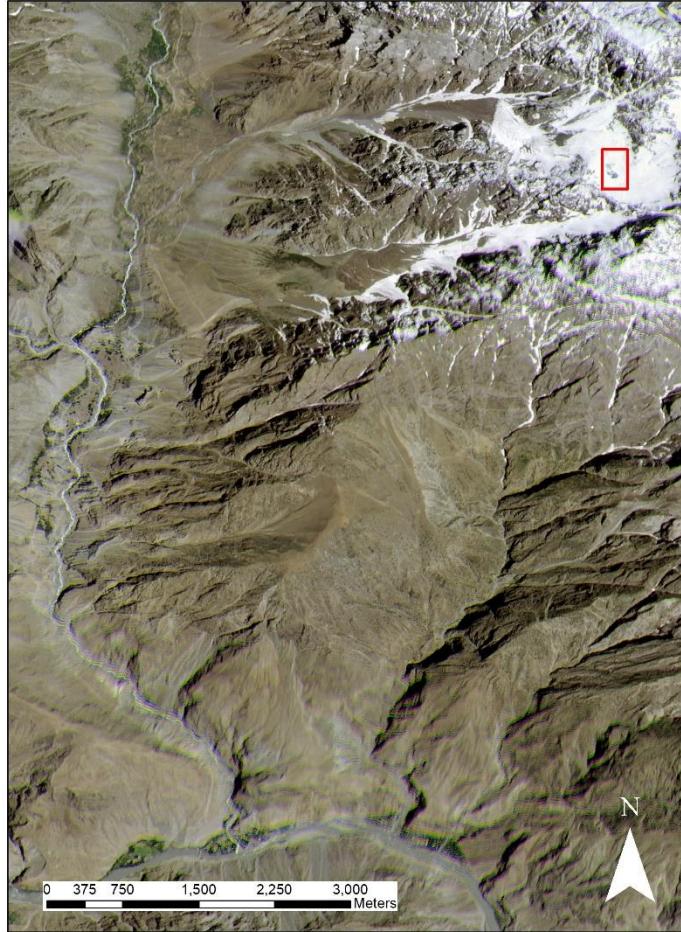
- Fresh snow results in ~90% of radiation to be reflected
- Feedback loop: less snow -> warmer earth -> even less snow



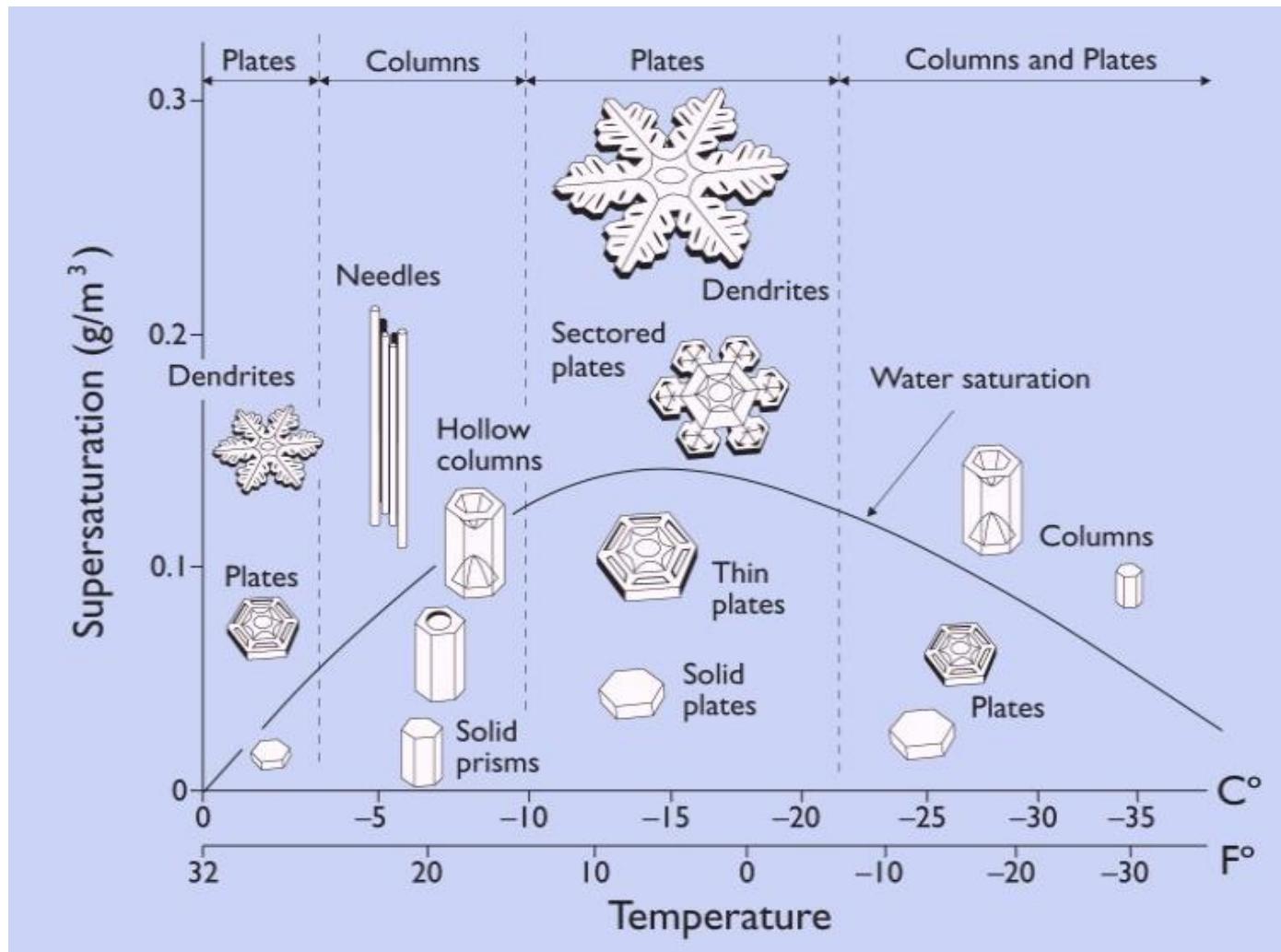
avalanches



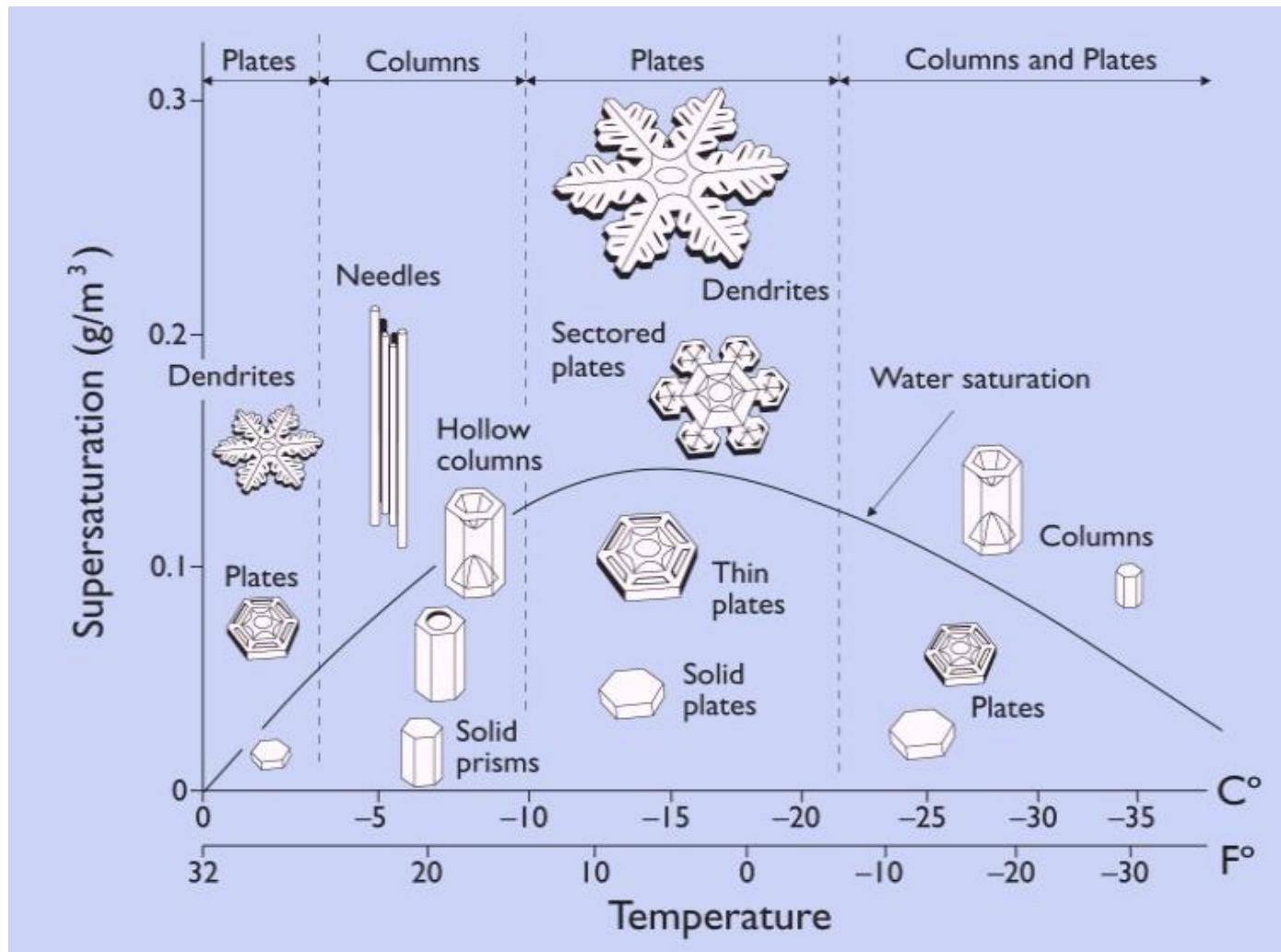
debris flows after soil saturation



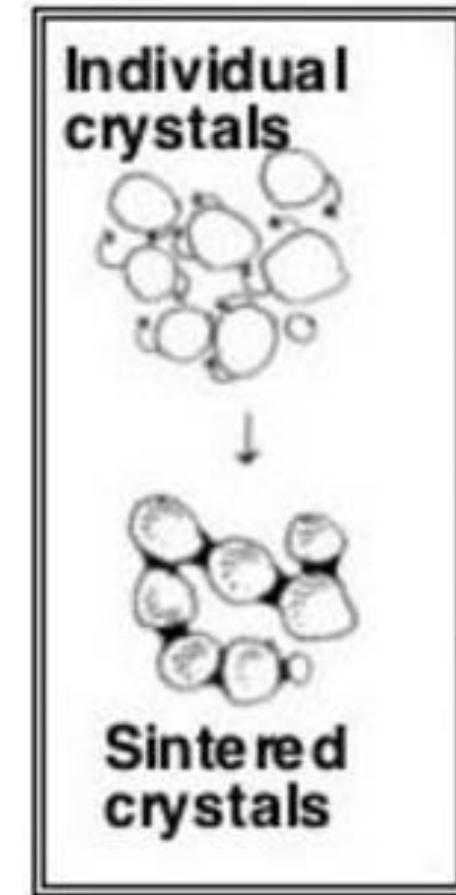
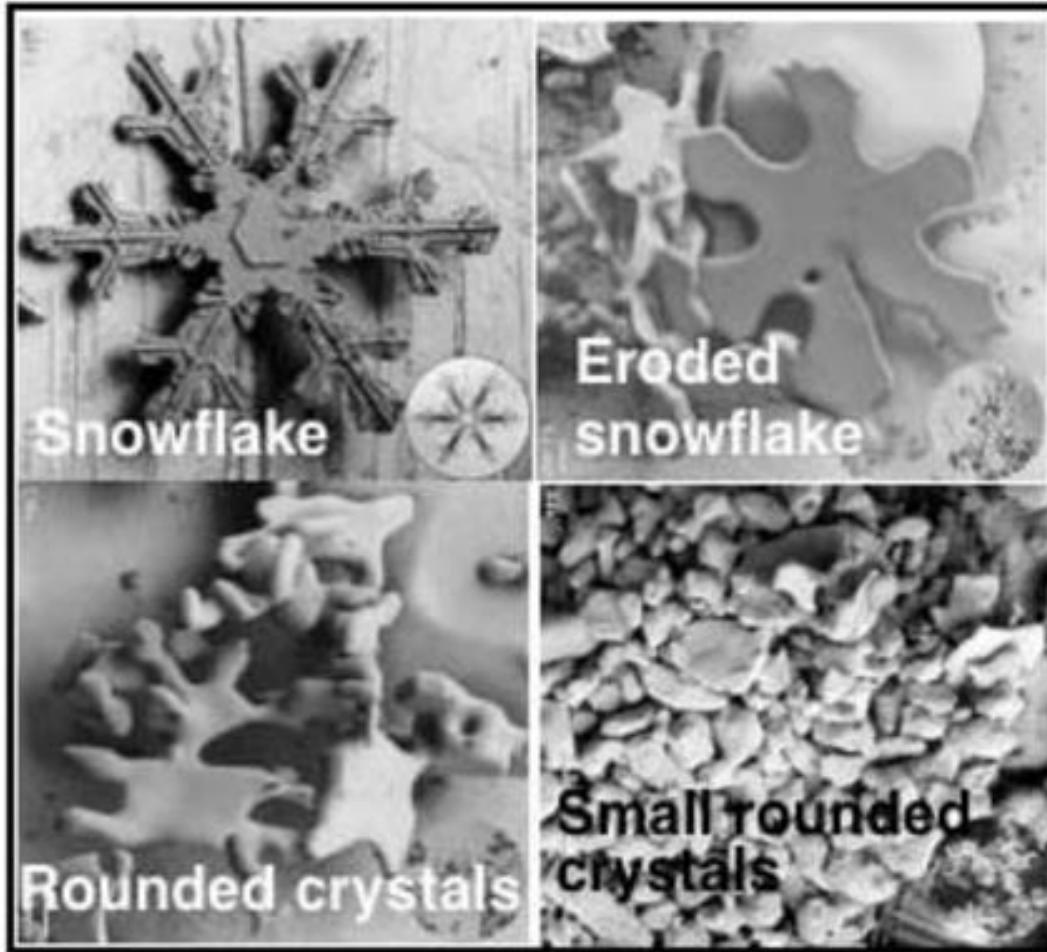
what is snow?



what is snow?



what is snow?



variables of snow

Property	Fresh snow	Settled snow
Snow cover extent or fraction (m ² or %)	-	-
Depth (m)	-	-
Density (kg m ⁻³)	20-150	250-550
Snow water equivalent (m w.e.)	-	-
Grain size (μm)	variable	variable
Albedo (-)	0.8-0.9	0.4-0.6
Thermal conductivity (W m ⁻¹ °C ⁻¹)	0.03-0.06	0.1-0.7

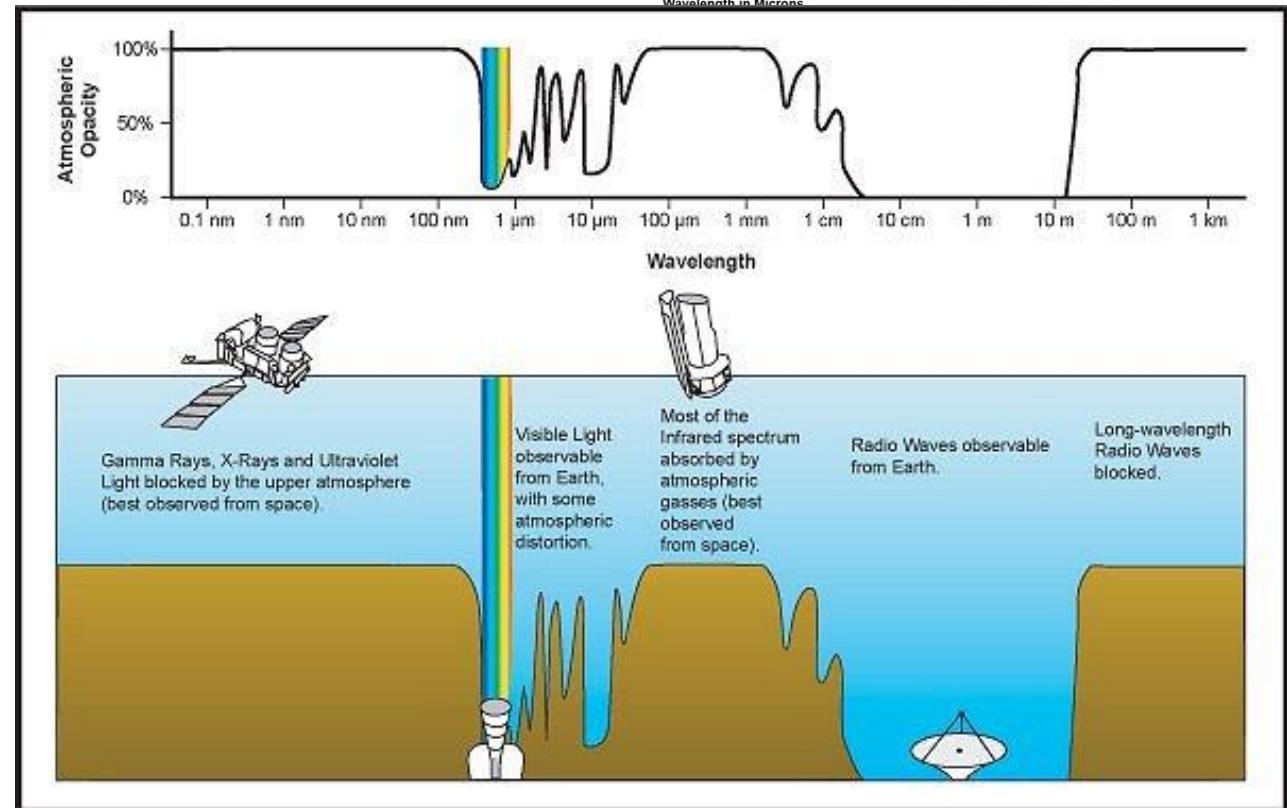
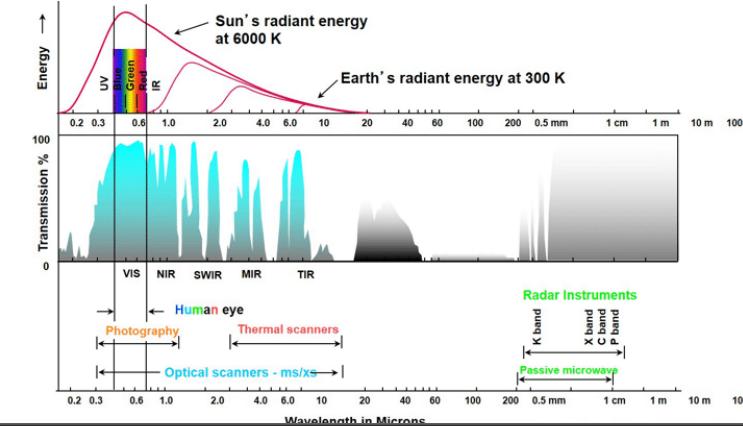


Remote sensing

Exploiting wavelengths

... of satellites

... of the earth surface



Active vs Passive Sensors

Passive sensors: record reflected or emitted electromagnetic signal

- visible light (reflected, optical imagery, daytime)
- microwave (emitted, AMSR, day and night)

Active sensors: send energy pulses and measure their return

- LIDAR (visible, SW)
- Radar (microwave)

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snow properties from optical imagery

Spectrometers

- Snow covered area/duration
- Snow grain size
- albedo/darkening (LAPs)

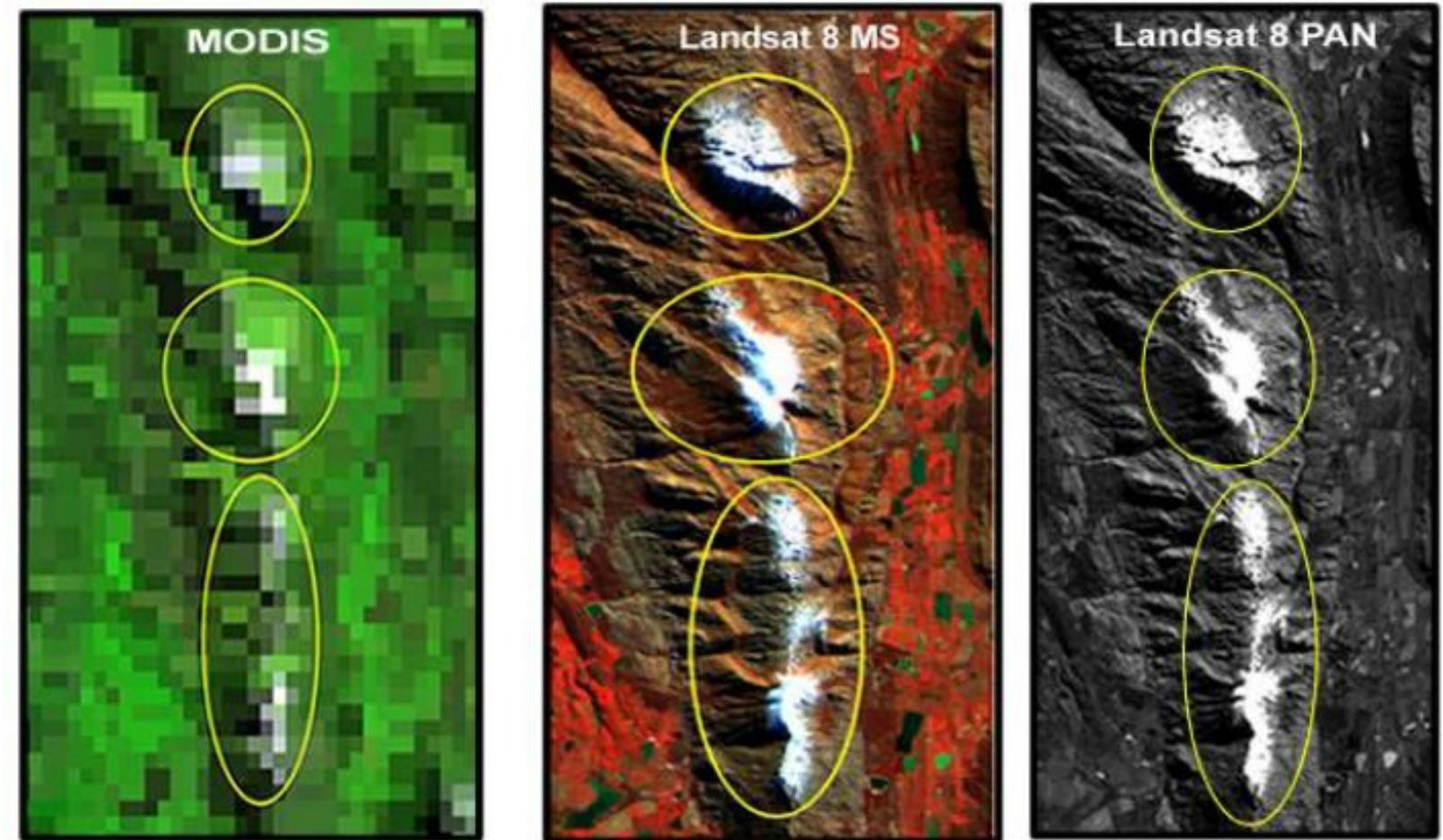
Cameras

- Repeat photogrammetry to produce snow depth



Resolution

- Spatial (0.5 – 5000 m)
- Temporal (revisit time, daily to bi-weekly)
- Spectral (number of bands, one to ~ten)
- Radiometric (variation in brightness)



Temporal Resolution

Landsat:

- + Spatial accuracy
- Long revisit time
- Impact of clouds

MODIS

- Spatially coarse
- + independent of clouds
- + near daily revisit

Satellite(s)—Instrument(s)	Operational since/until	Revisit Time	Spatial Resolution	Swath Width
Landsat MSS/TM/ETM+/OLI	1972/Today	16–18 days	30–100 m	185 km
Terra, Aqua-MODIS	2000/Today	Twice per day	250–1000 m	2330 km
TIROS/NOAA/Metop AVHRR	1978/Today	At least daily	1100 m	2400 km
Envisat/AATSR	2002/2012	2–3 days	1000 m	500 km
Envisat/MERIS	2002/2012	2–3 days	300 m	1150 km
ERS-2/ATSR-2	1995/2011	2–3 days	1000 m	512 km
Sentinel 2	to be launched in 2015	3–5 days	10–60 m	290 km
Sentinel 3-OLCI/SLSTR	to be launched in 2015	1–2 days	300–500 m	1270 km
Suomi-NPP-VIIRS	2011/Today	Daily	375–750 m	3040 km

Sources: [33–41].

snow properties from optical imagery

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Cameras

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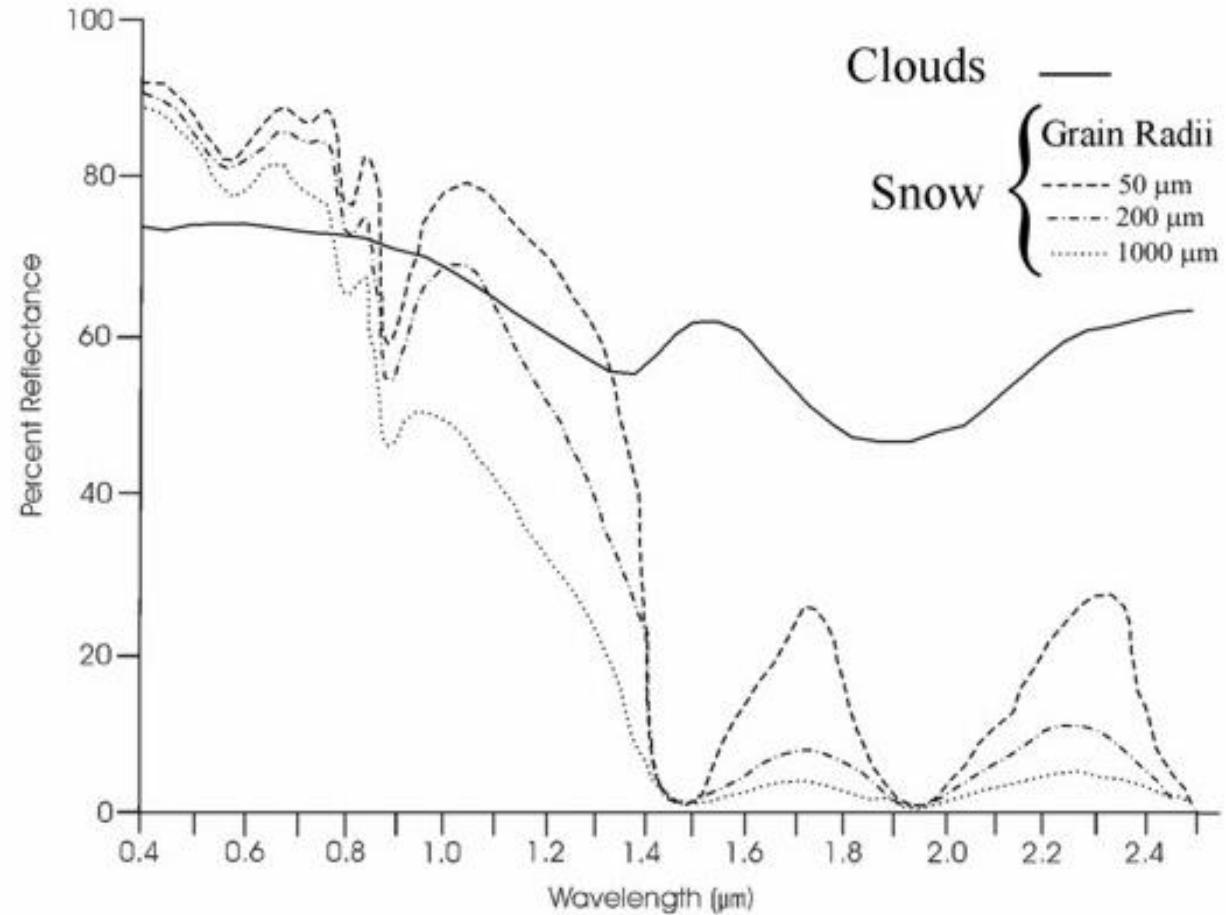
Spectral signature of snow

Reflectance in different parts of the electromagnetic spectrum

Snow is generally highly reflective as ice is transparent resulting in scattering

- For short wavelength absorption is controlled by light absorbing particles (LAPs)
- For longer wavelengths ice is absorptive and reflectance is controlled by grain size (radius, specific surface area)

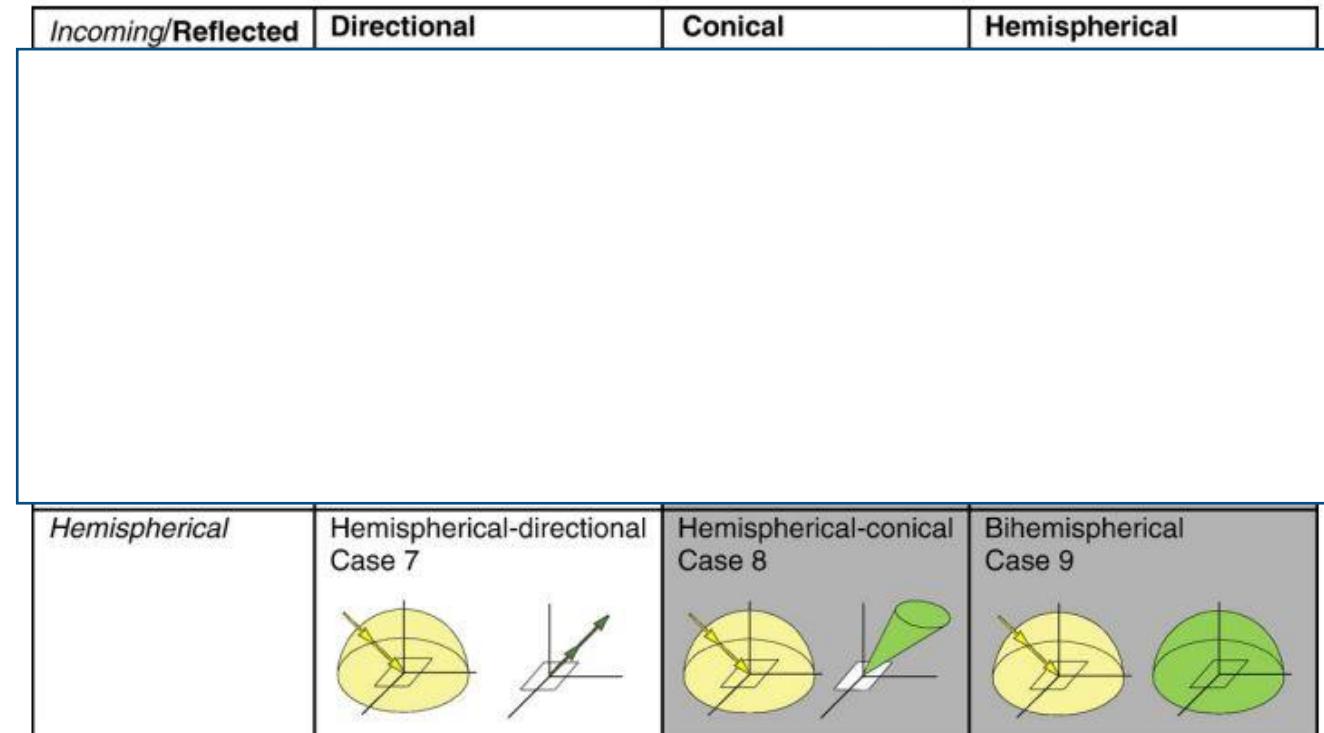
<http://snow.engin.umich.edu/>



Reflectance vs albedo

Reflected waves are picked up in a directional manner by eyes (case 7)/satellites (case 8)

Albedo is the hemispherical budget of reflected waves (case 9)



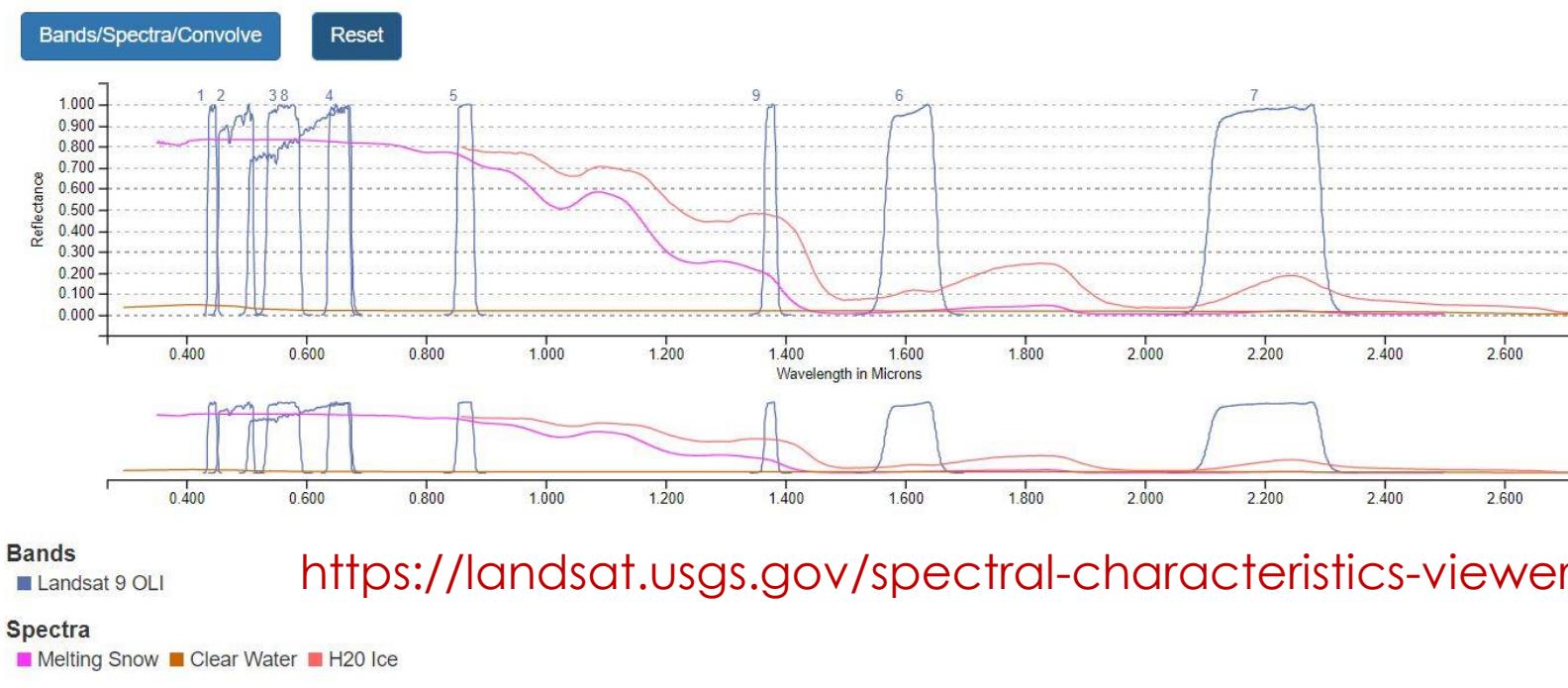
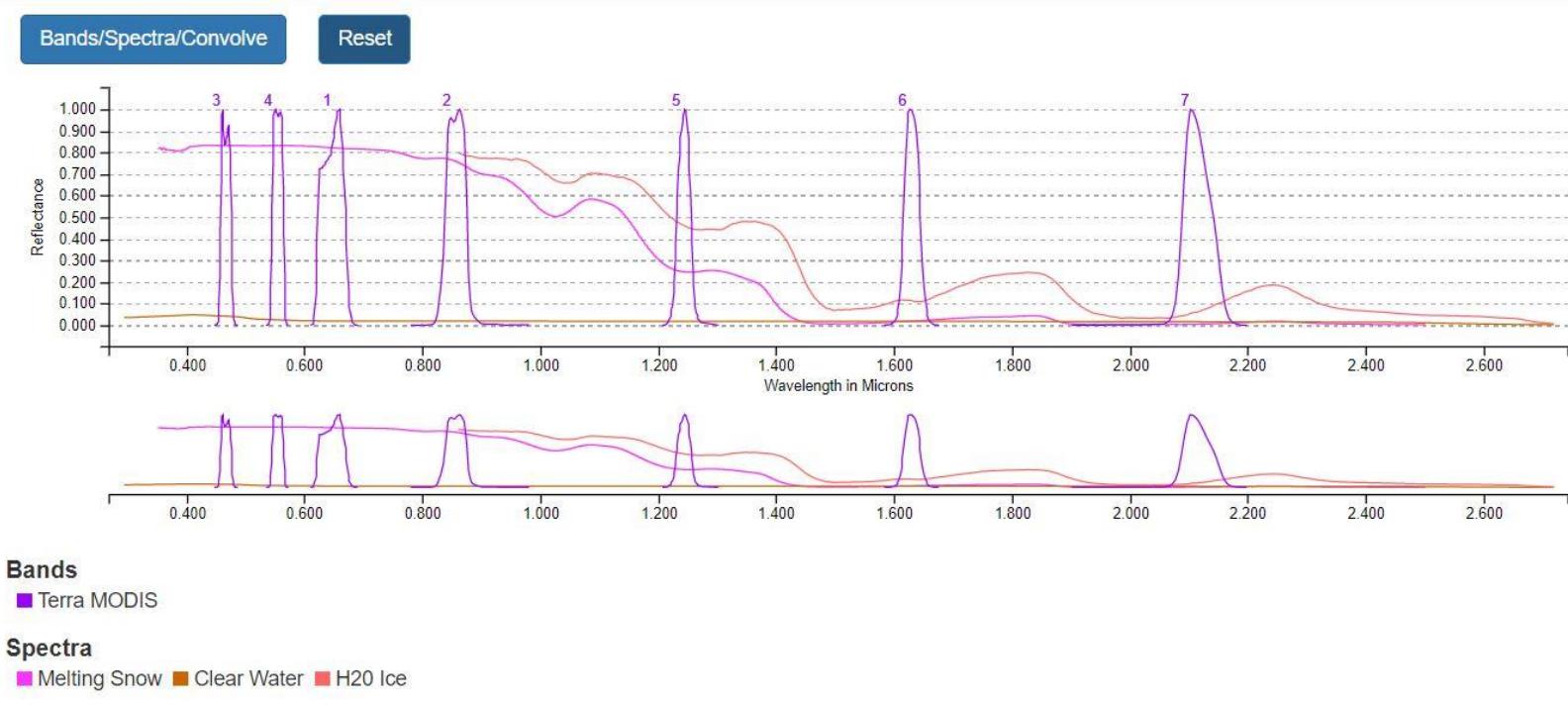
Multispectral bands

$$NDSI = \frac{VIS - SWIR}{VIS + SWIR}$$

MODIS: b 4 / b 6

Landsat: b 3 / b 6

NDSI > 0.4 → snow



snow covered area

NDSI (normalized differential snow index, binary)

- computationally efficient
- long practice and equivalent to other such indices (NDVI, NDWI etc)
- but...



Challenges ...

- Resolution becomes an issue:
 - MODIS 500 m
 - Landsat 30/90 m
- Fractional snow cover increasingly important for hydrology, permafrost and ecology



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snow covered area

NDSI (normalized difference snow index)

- computational
- long practice can take
- but...

fSCA (fractional snow covered area)

- uses all bands to calculate

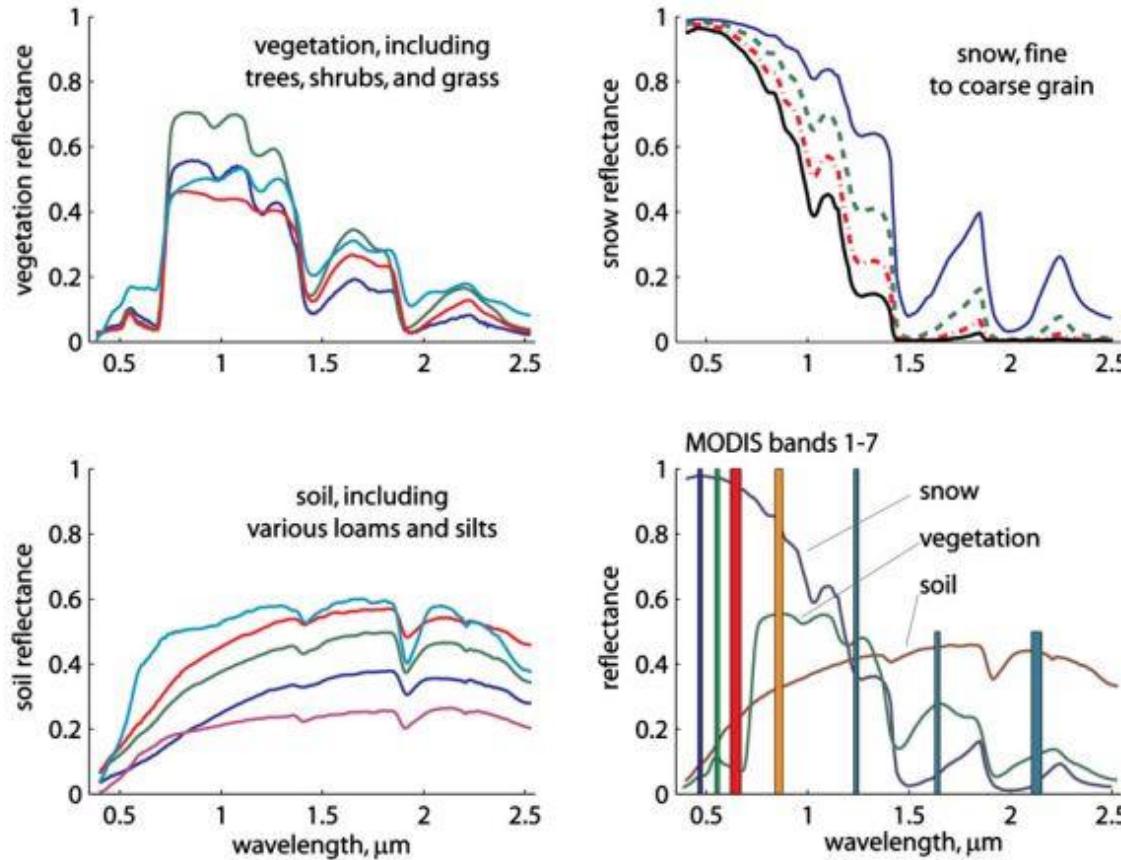


Fig. 1. Spectral reflectance of snow, vegetation, and soils. The vegetation and soils data are from laboratory or field measurements. The snow reflectance values are generated from a radiative transfer model. In the lower right graph, the MODIS land bands are shown, along with one sample spectrum each from vegetation, snow, and soil.

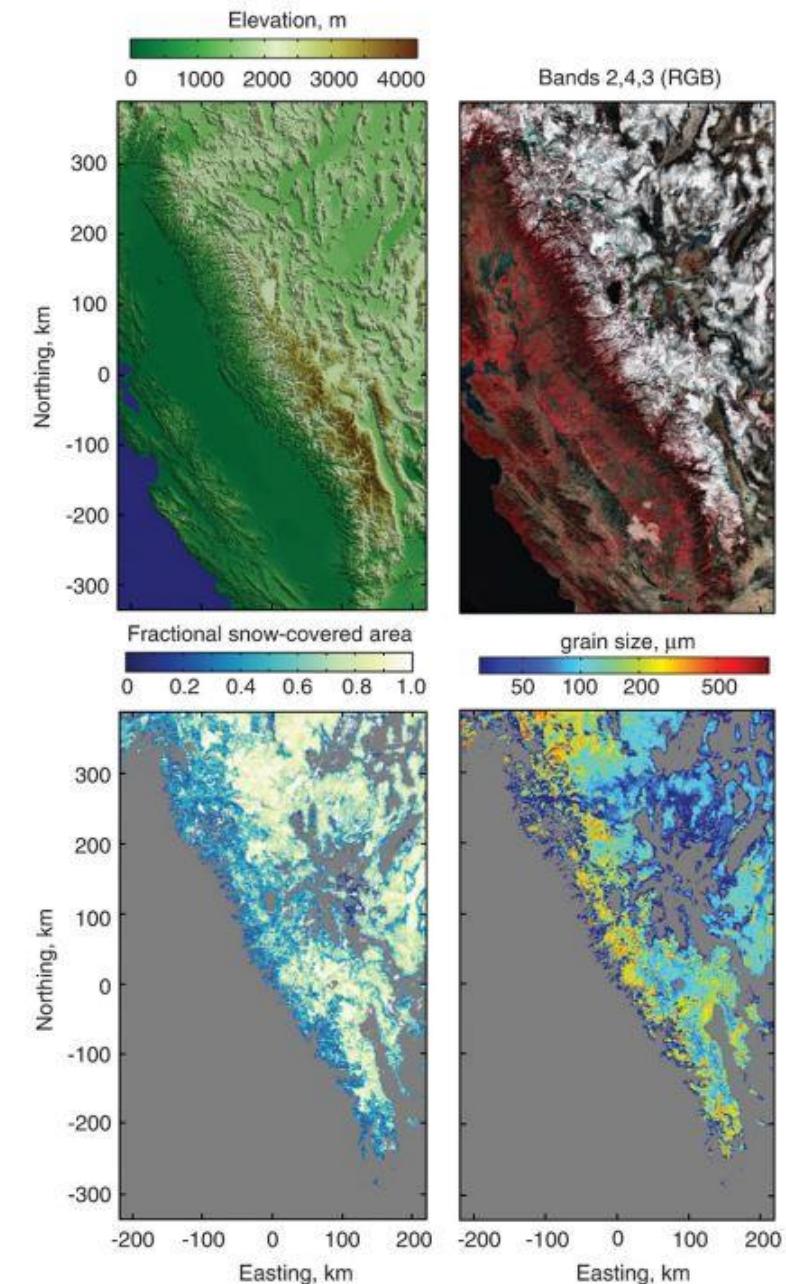
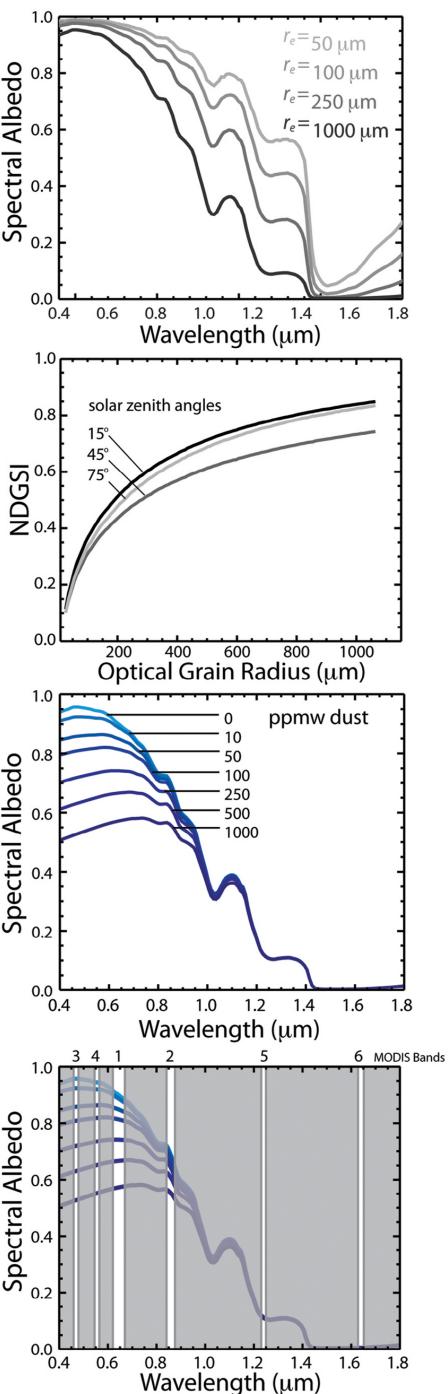


snow grain size

NDGSI (normalized differential grain size index)

$$\text{NDGSI} = \frac{\text{NIR} - \text{SWIR}}{\text{NIR} + \text{SWIR}}$$

MODIS: b 2 / b 5

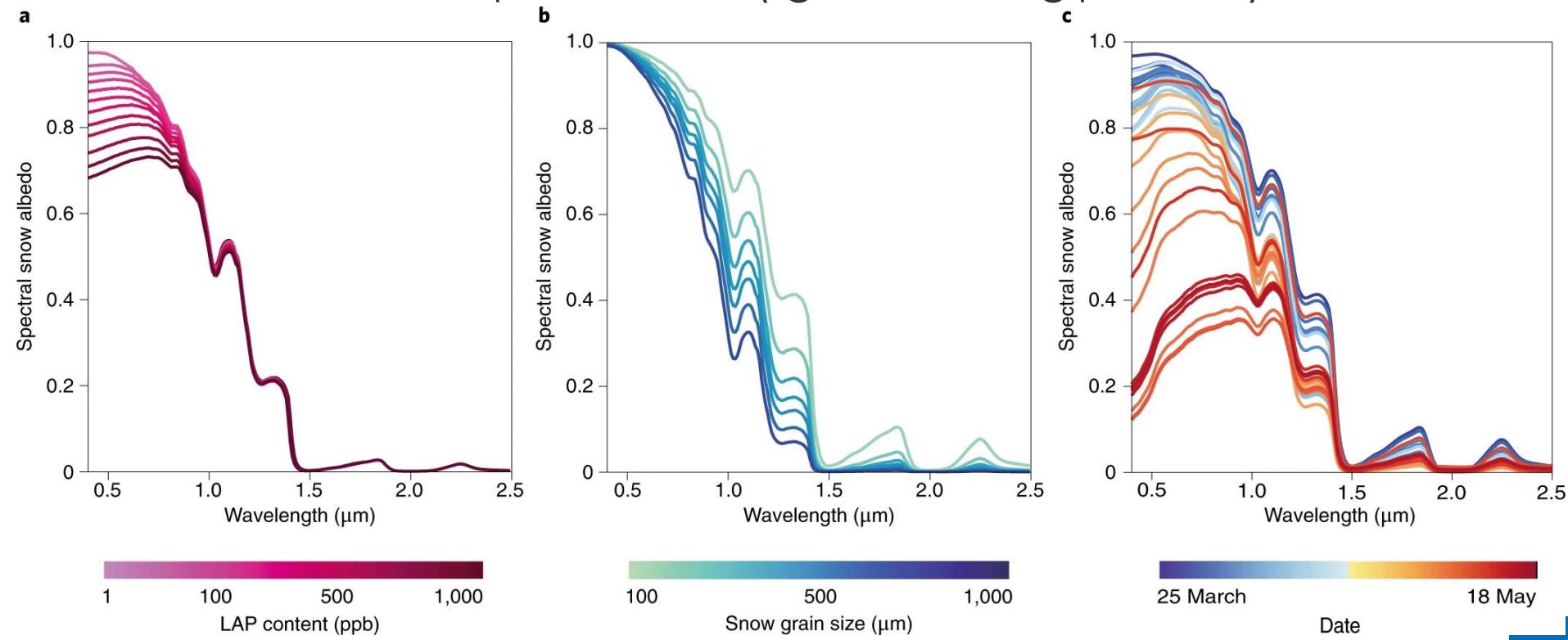


Painter, 2009



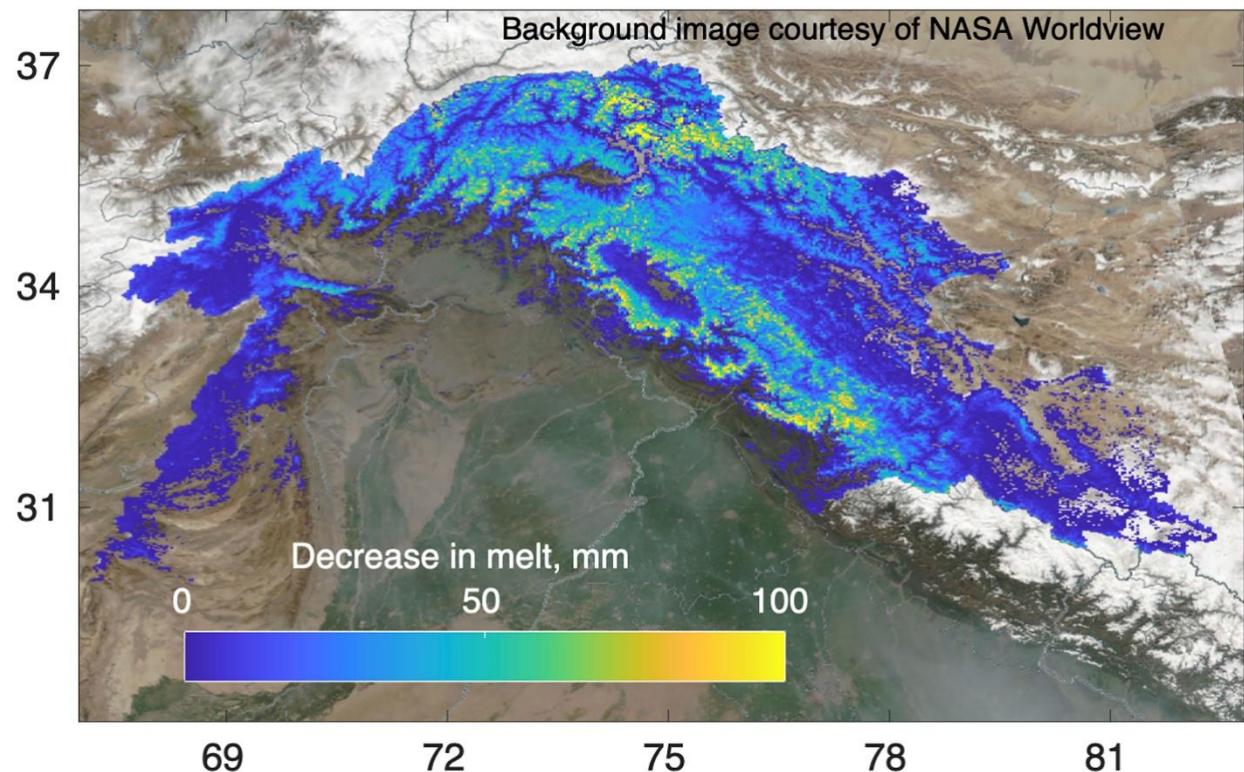
From grain size to albedo reduction

- Compute a theoretical spectral curve based on observed grain size
- Retrieve the actual snow spectra
- Calculate the difference to assess the impact of LAPs (light absorbing particles)



From grain size to albedo reduction

- Compute a theoretical spectral curve based on observed grain size
- Retrieve the actual snow spectra
- Calculate the difference to assess the impact of LAPs (light absorbing particles)
- combine LAPs with actual irradiance to calculate increased/decreased melt
- application for impact of lockdown measures (reduction in LAPs)



snow properties from optical imagery

Spectrometers

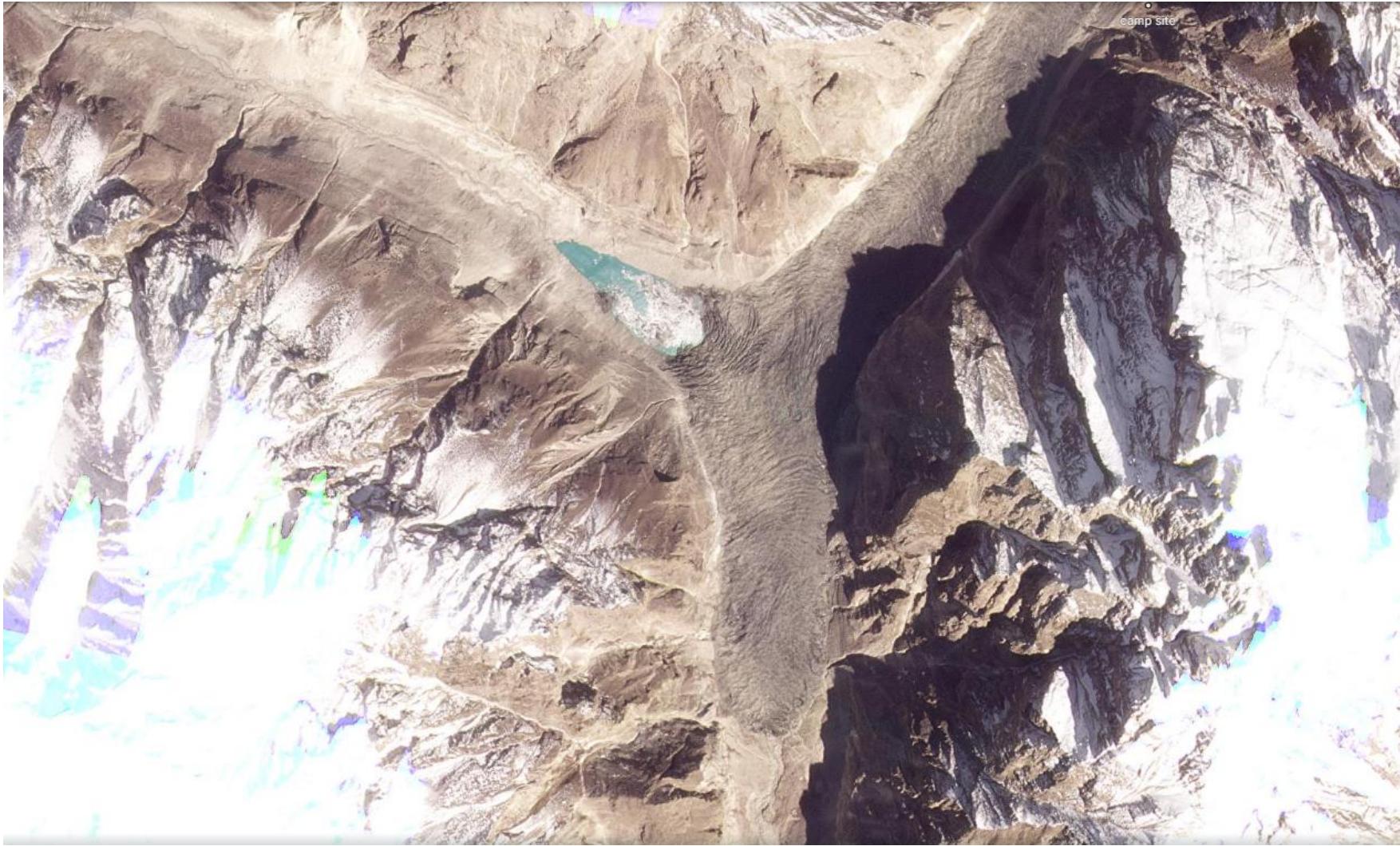
- Snow covered area/duration
- Snow grain size
- albedo/darkening (LAPs)

Cameras

- Repeat photogrammetry
- Stereo photogrammetry



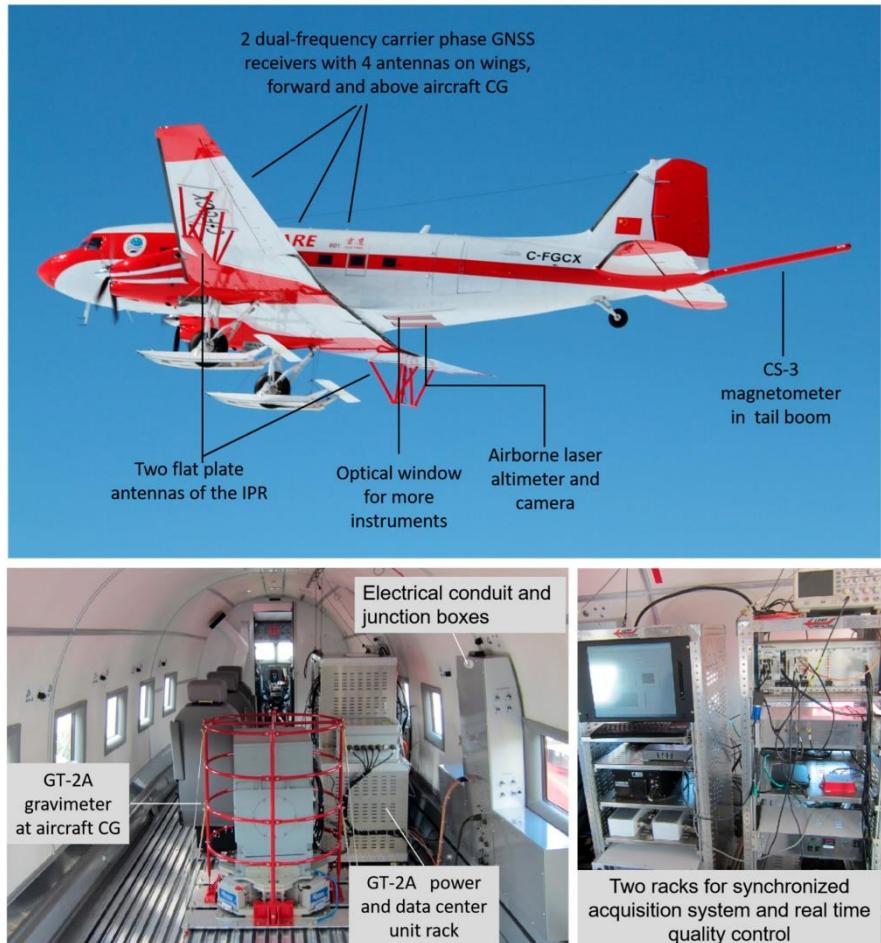
new suite of commercial satellites



- satellite data is virtually accessible to anyone with ~10h time lag



Mapping from planes or UAVs



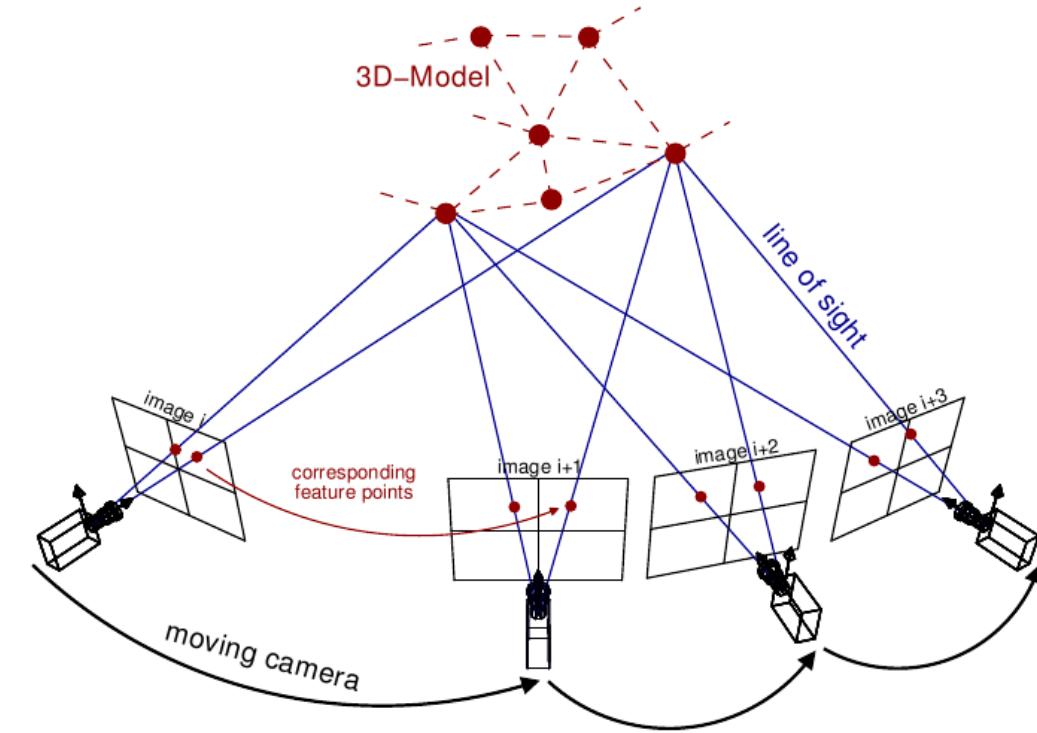
Cui et al., 2020



Tiggelen et al., 2022

structure from motion

- use simple oblique imagery to create 3D models
- wide range of applications and software
- generally done from handheld and UAV imagery



Riel et al., 2016



structure from motion

- surface roughness

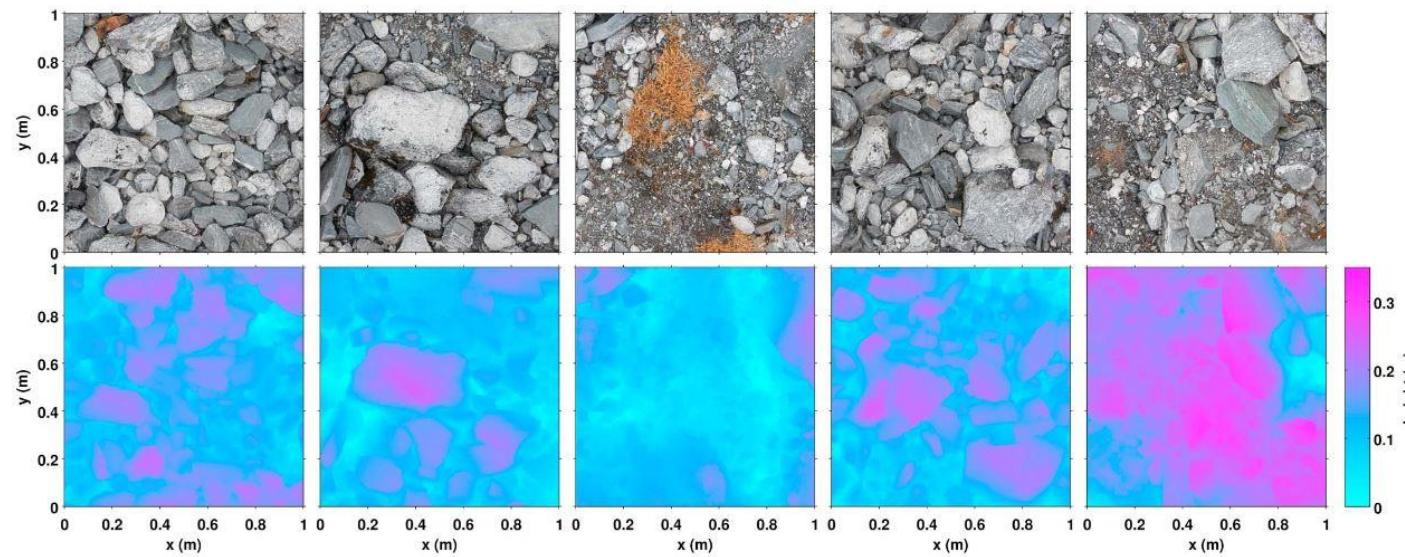
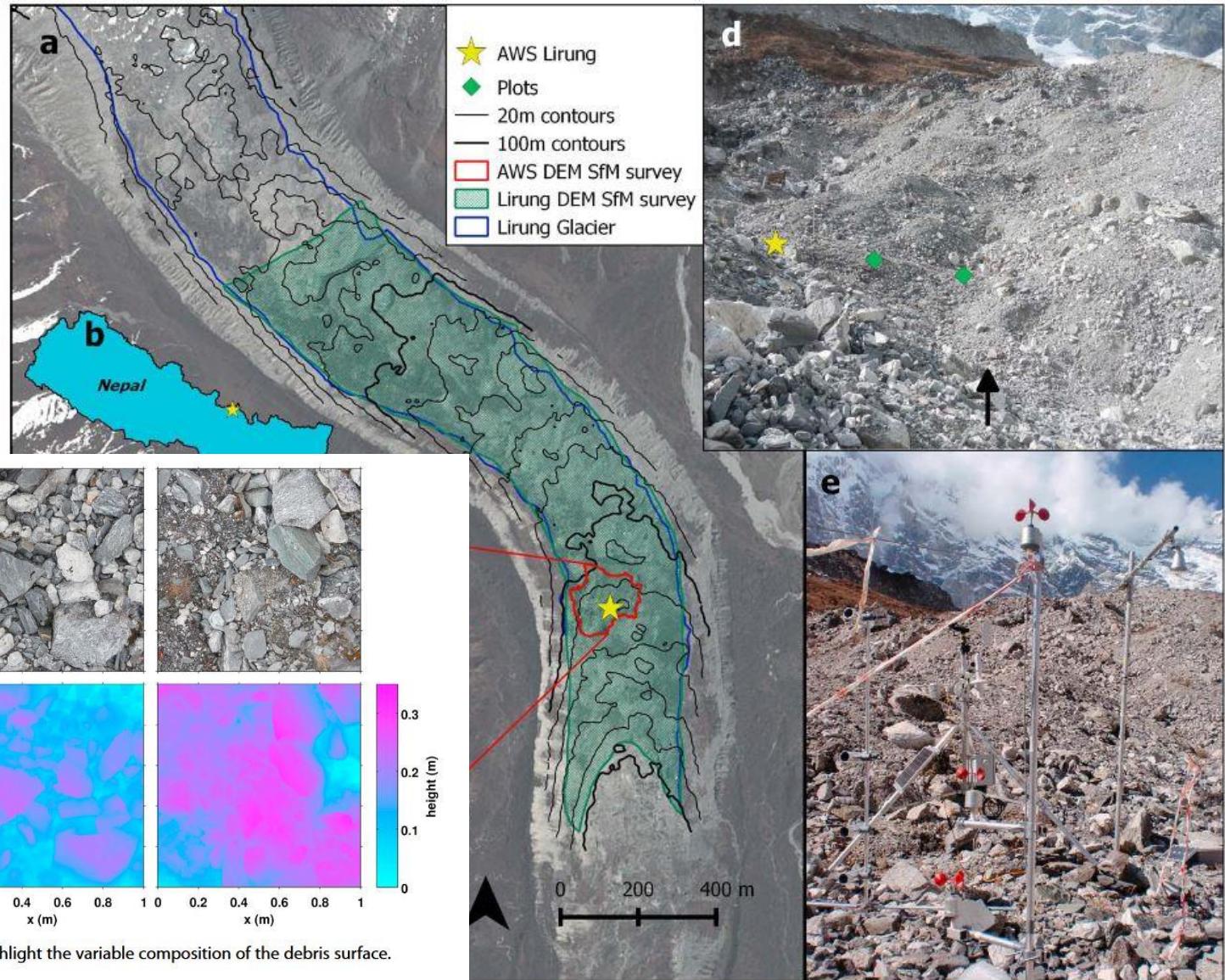
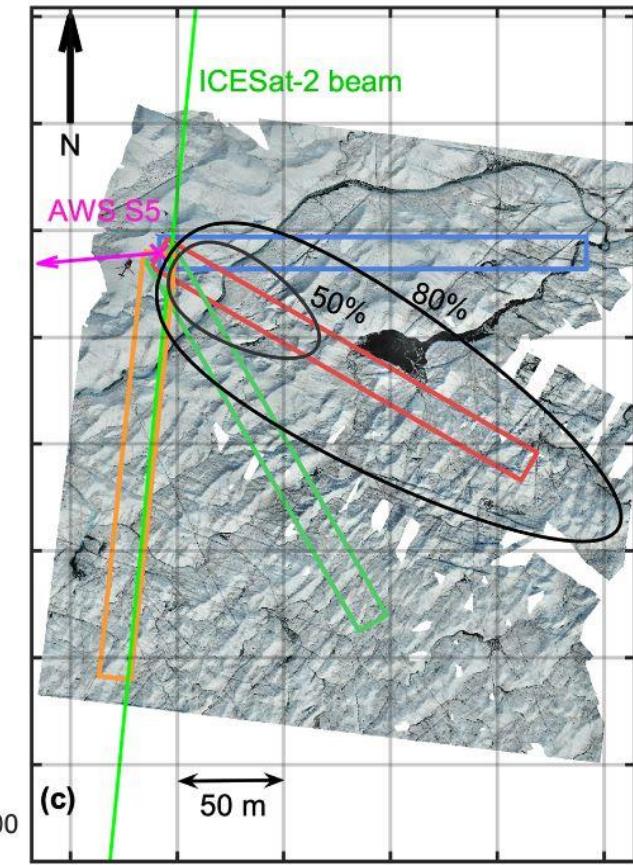
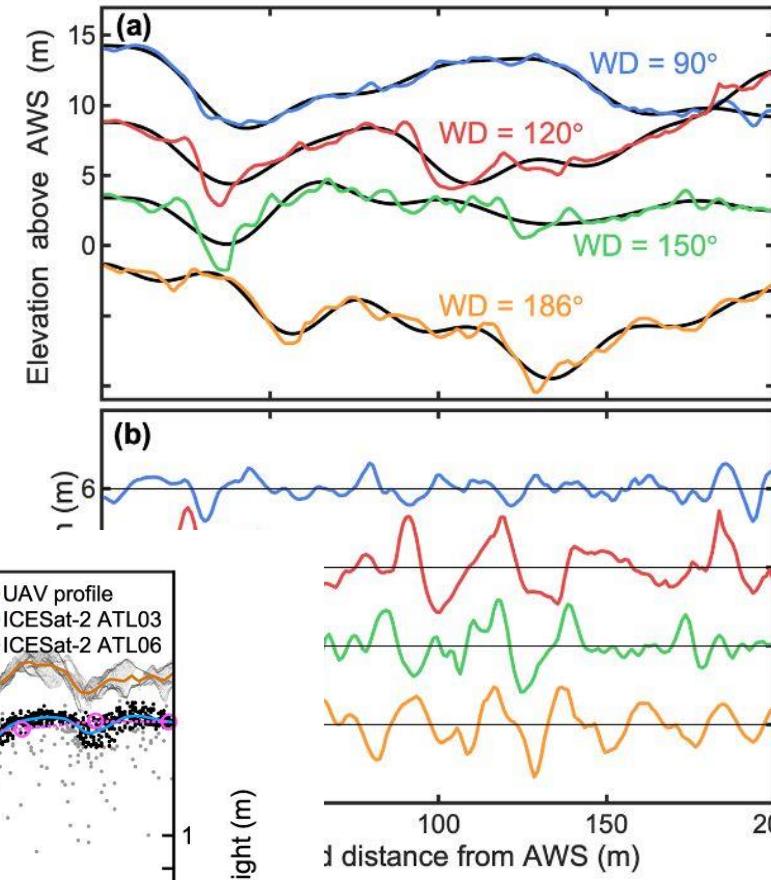
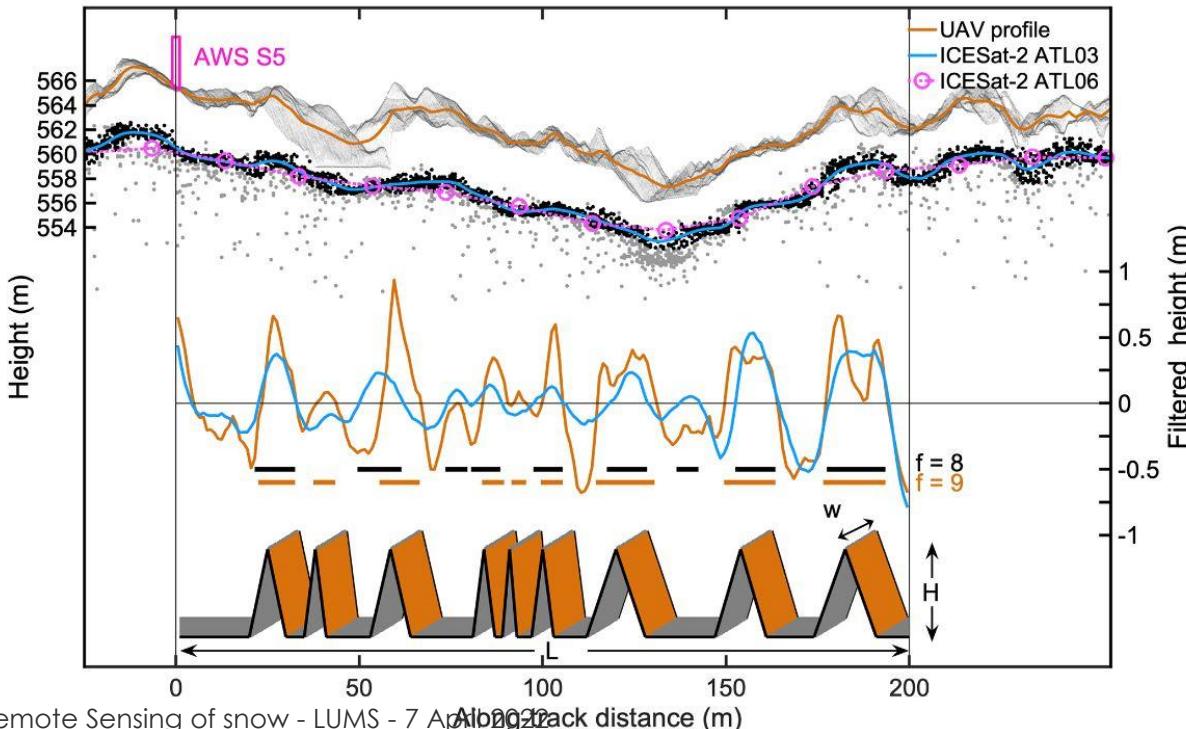


Figure 2. (top) SfM-generated orthophotos and (bottom) detrended DEMs for the five study plots highlight the variable composition of the debris surface. Up-glacier is oriented upward for each plot.



structure from motion

- surface roughness
- combined with RS laser (IceSat)



SfM for snow depth

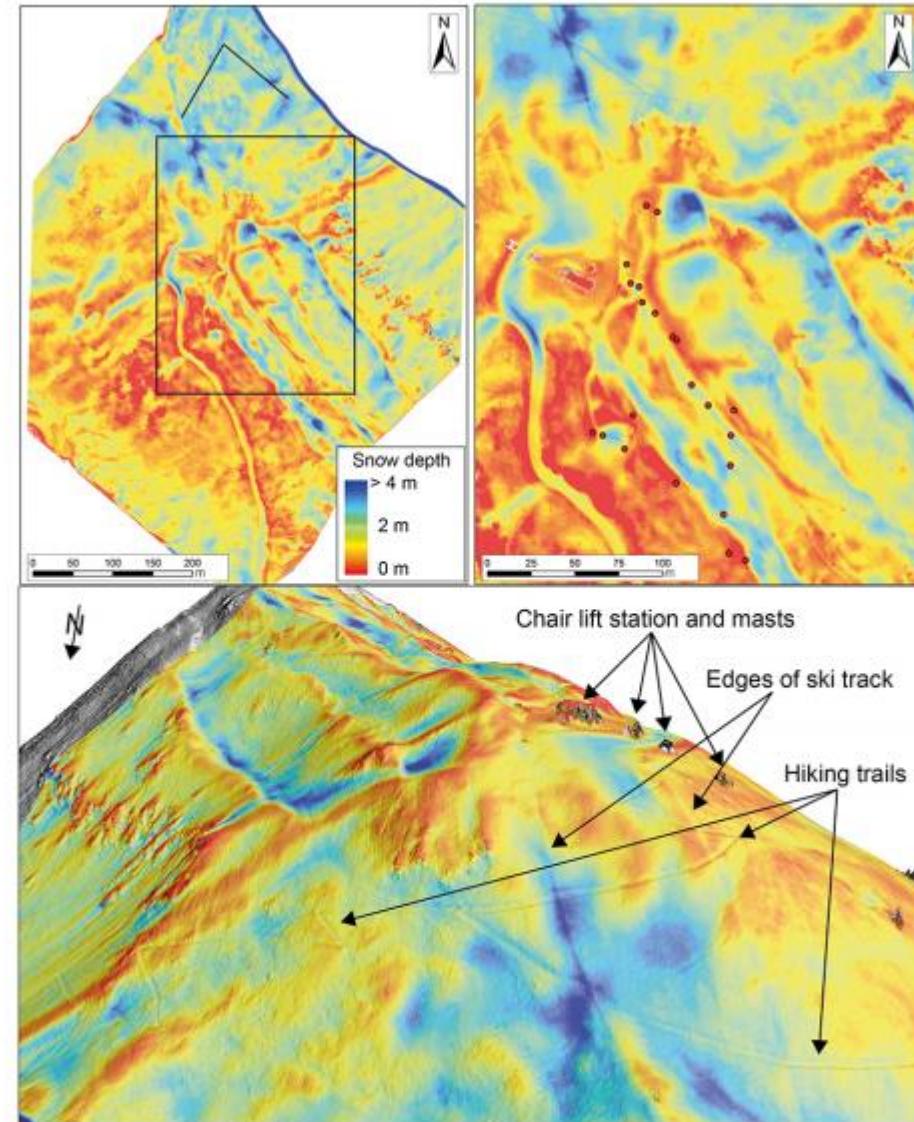
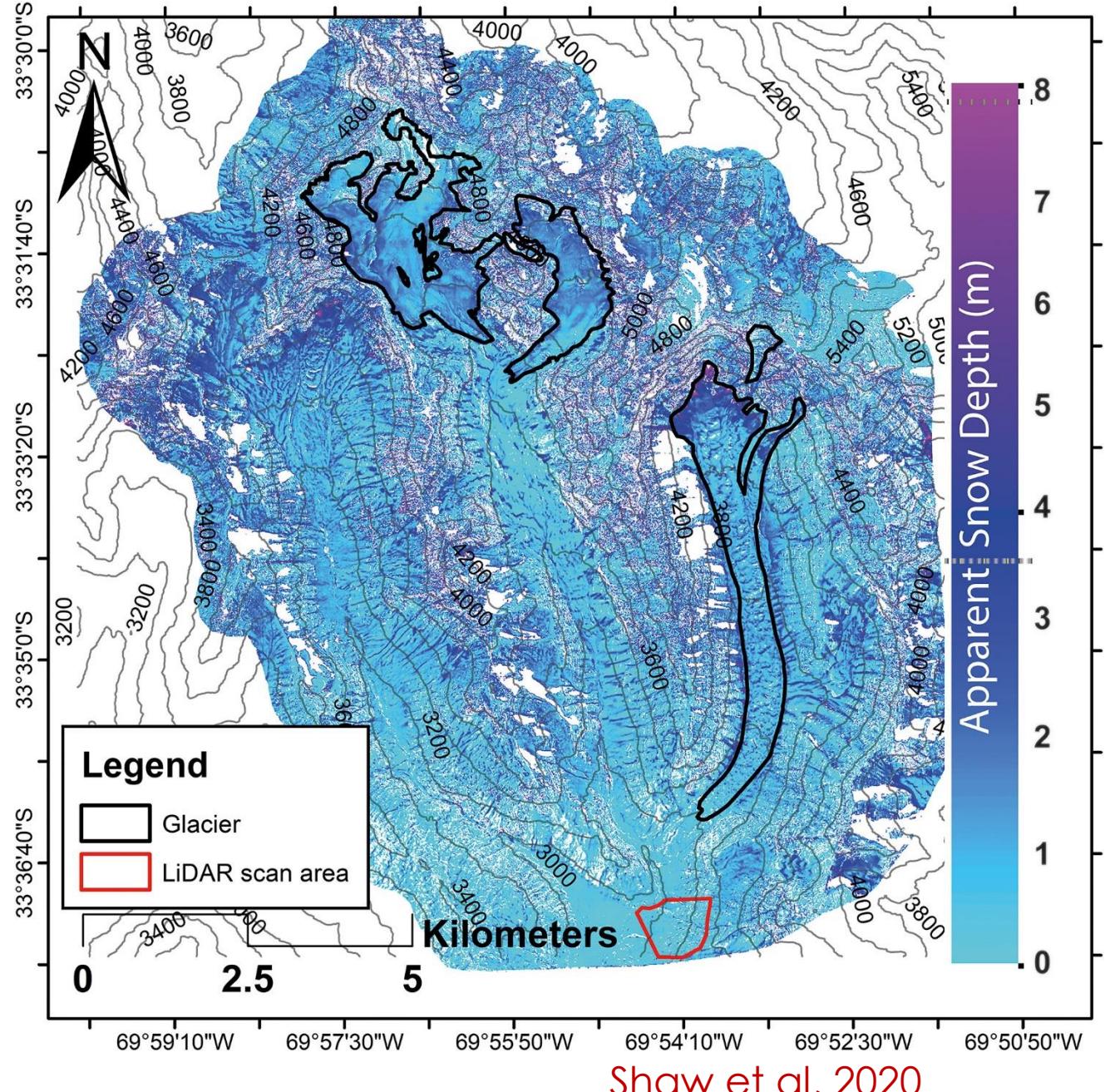


Figure 6. Overall HS map of the Brämabühl test site (top left panel) and close-up of the central part (top right panel). The locations of the reference plots are displayed as red circles. 3-D view of the HS draped over the hillshade of the snow-free DSM facing from north to south (bottom panel).

Tri-Stereo photogrammetry

- high resolution optical imagery
(Pleiades, CubeSats, SPOT etc)
- generally commercial



Active vs Passive Sensors

Passive sensors: record reflected or emitted electromagnetic signal

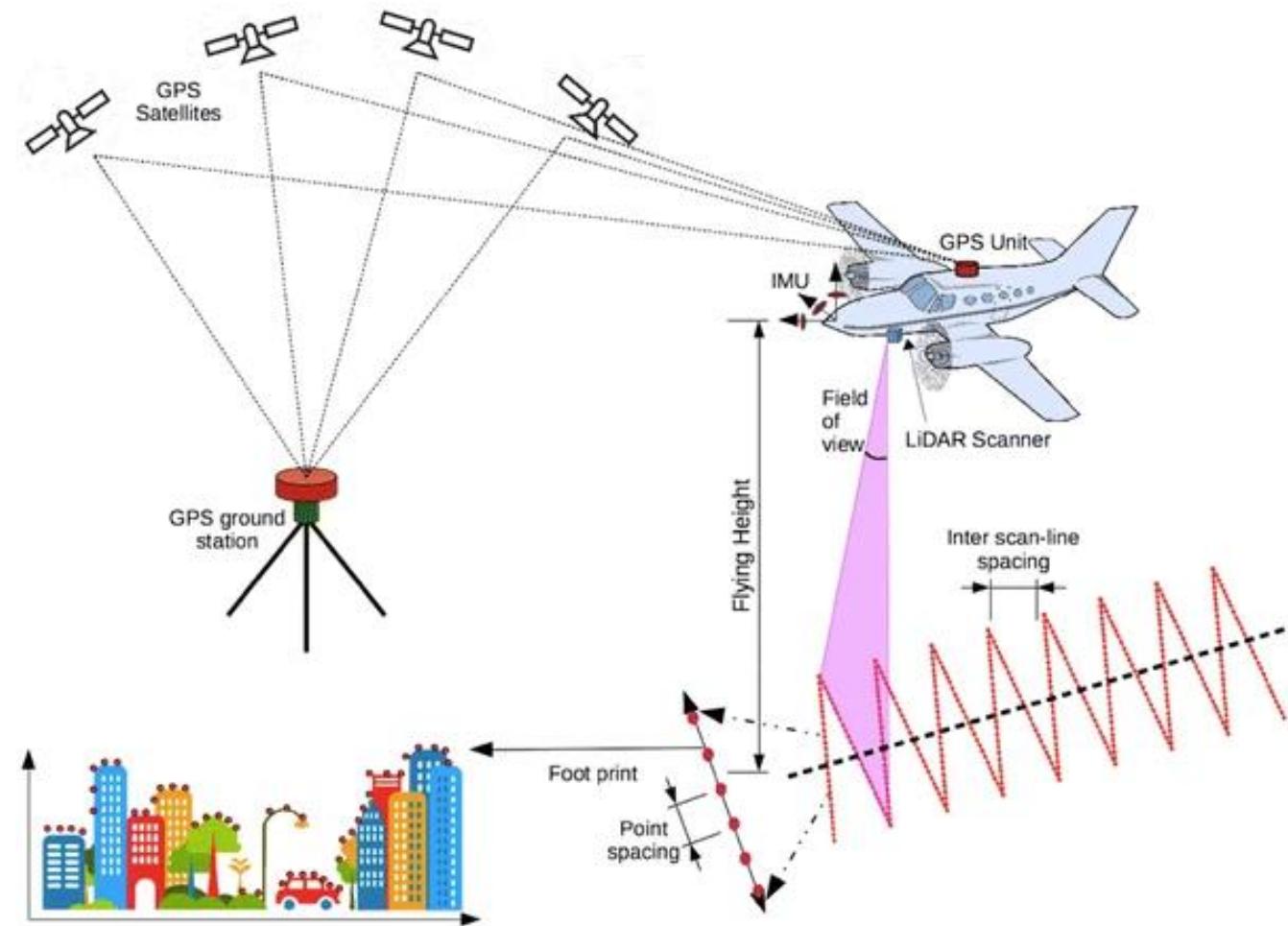
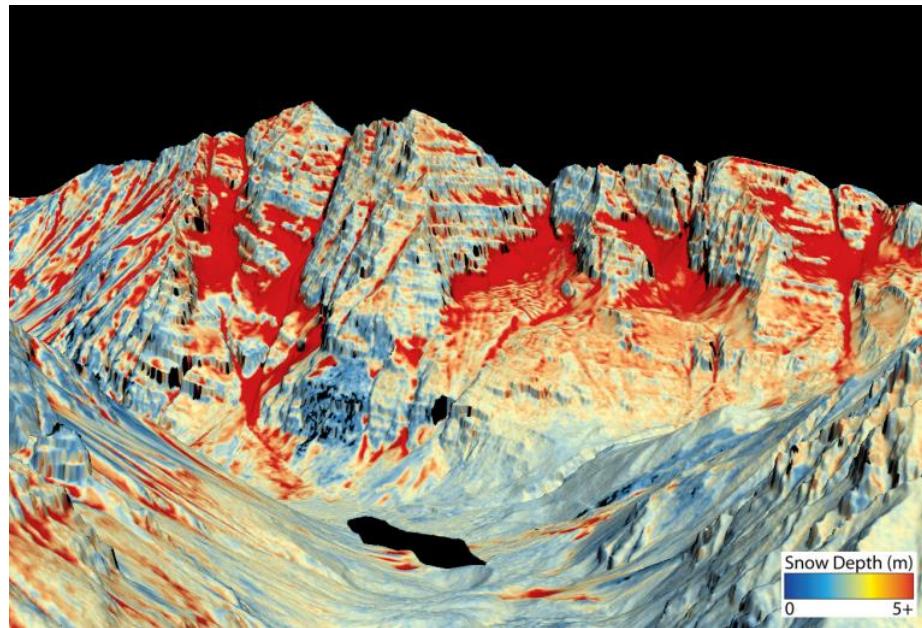
- visible light (reflected, optical imagery, daytime)
- microwave (emitted, AMSR, day and night)

Active sensors: send energy pulses and measure their return

- LIDAR (visible, SW)
- Radar (microwave)

LIDAR (light detection and ranging, active)

- Emitted laser, tracking the time/distance between sensor and surface
- application on planes and UAVs



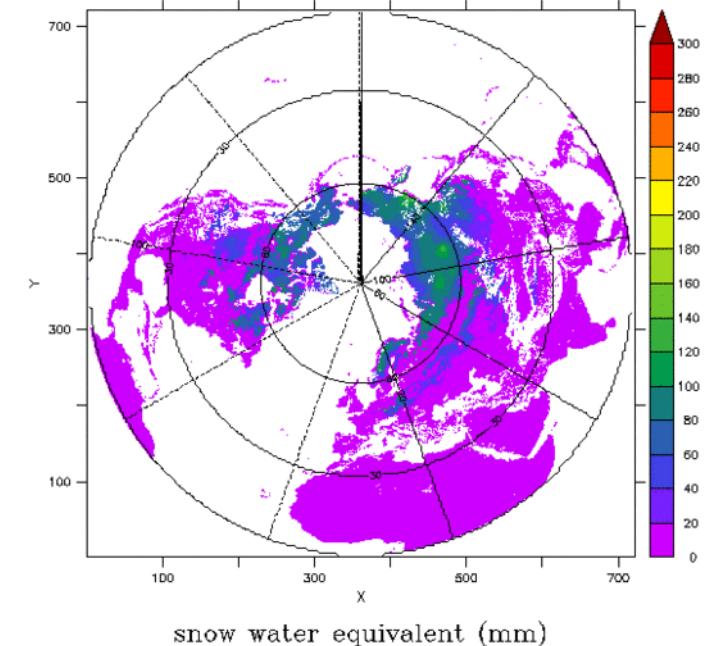
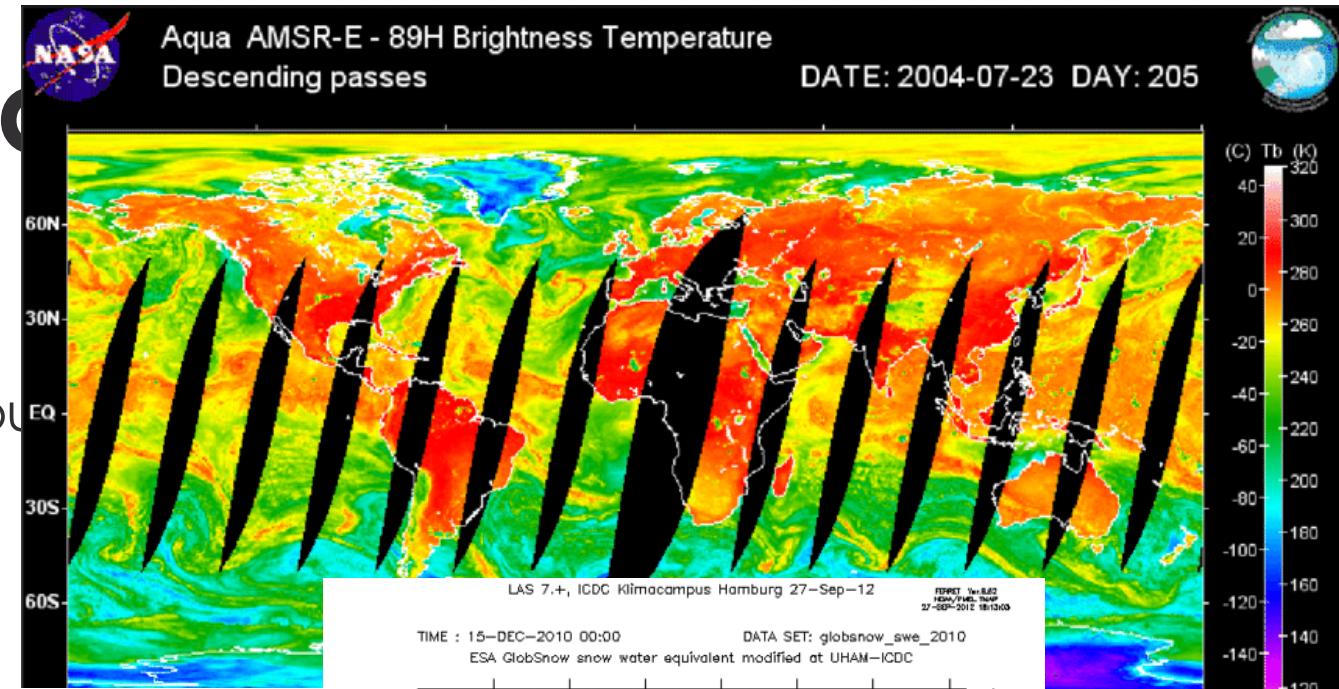
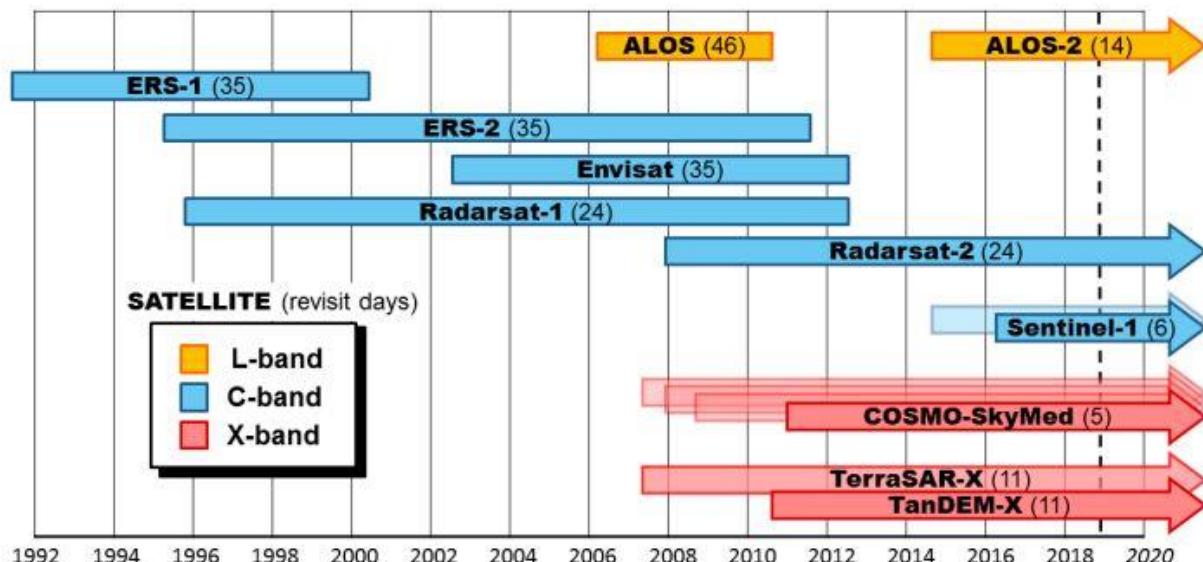
Satellite microwave remote sensing

Passive:

AMSR/SMMR: wide swath, near daily data but

- crucial is emissivity ϵ
- $\epsilon = 1$ for perfect black body
- $\epsilon = 0.7$ to 0.9 for dry snow

Active
Radar

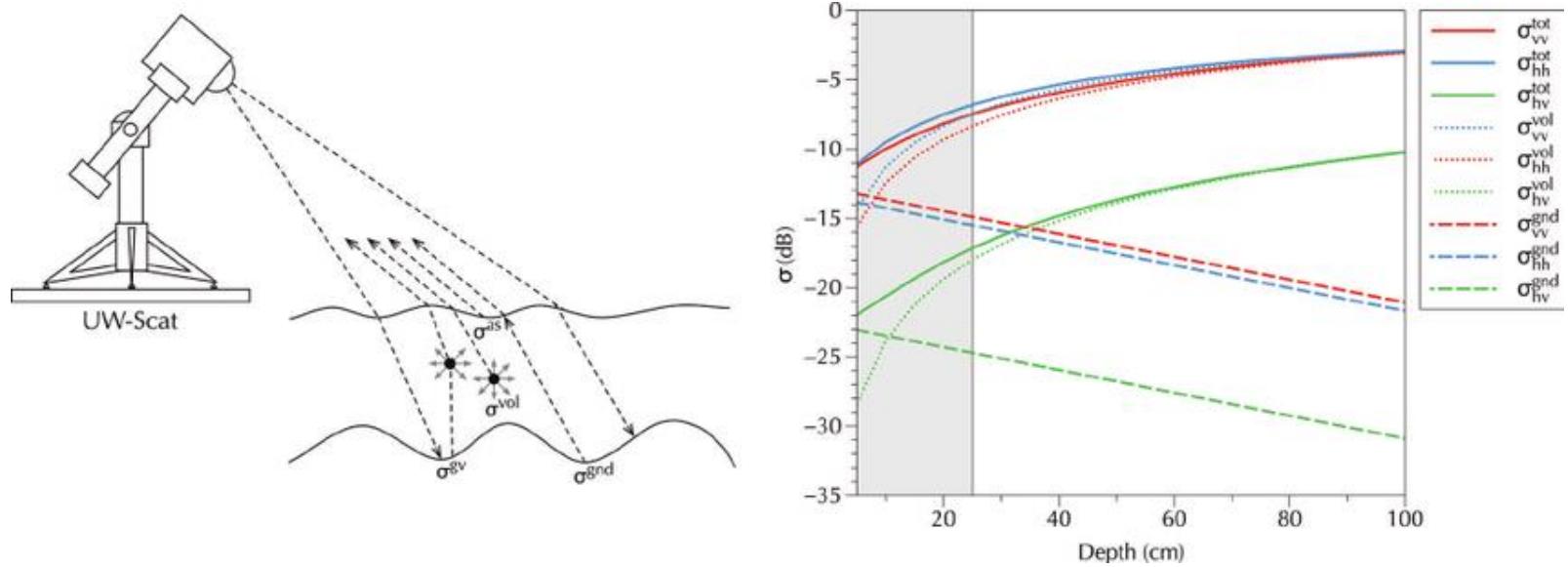


Radar

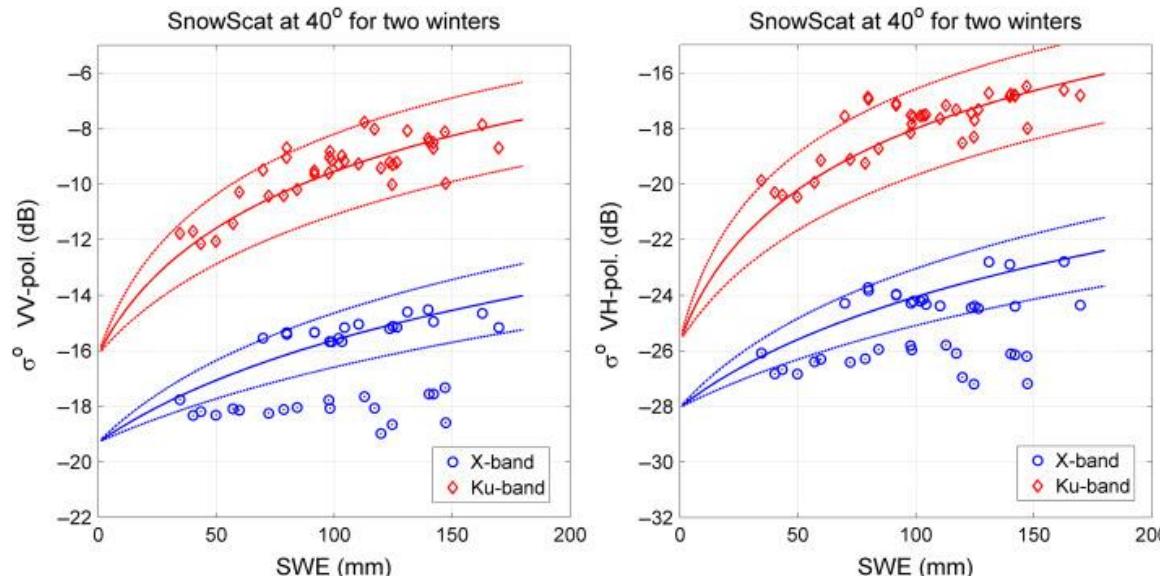
Snow Depth

Snow Water Equivalent

Grain Size (hoar etc)

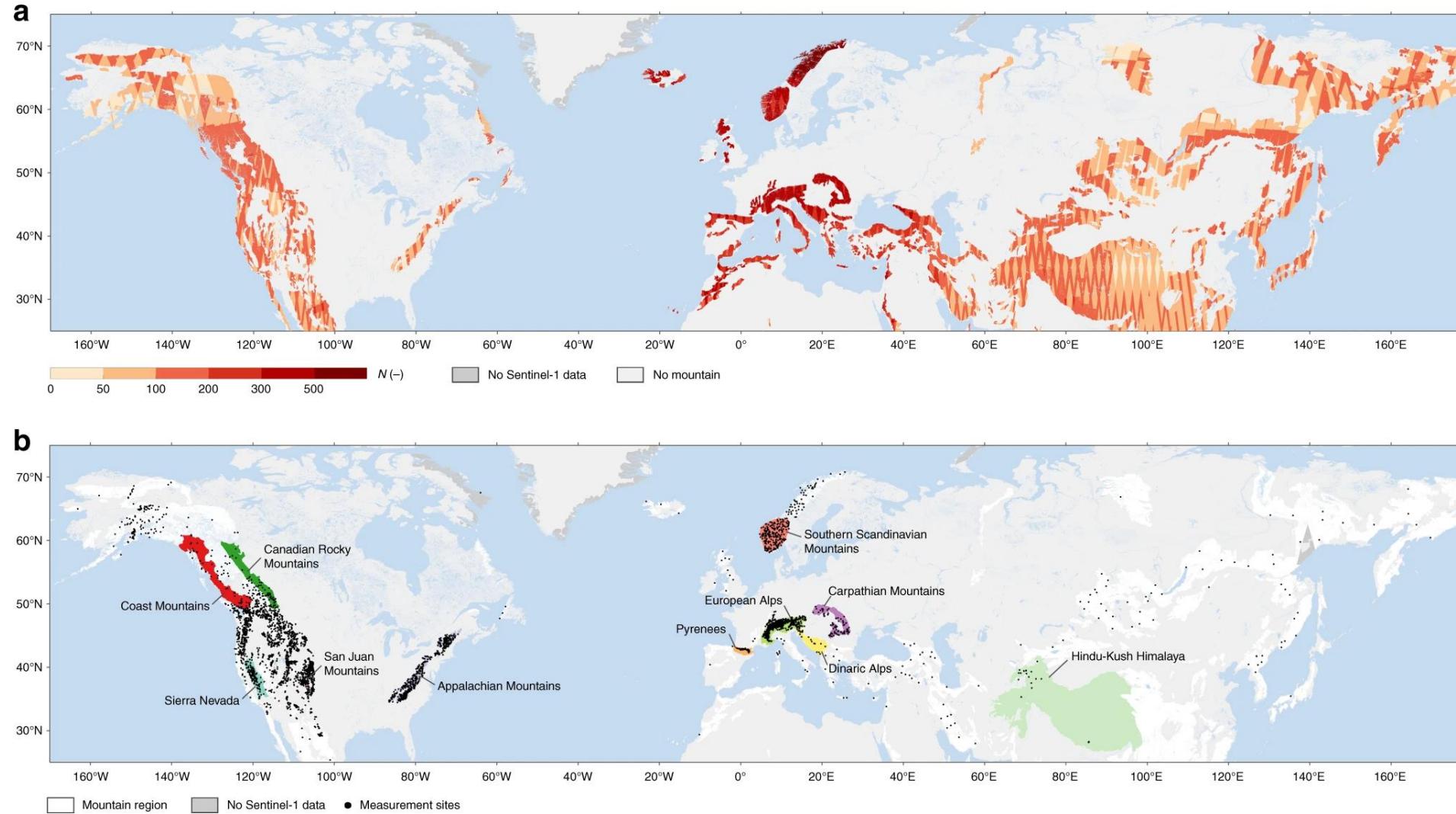


King et al., 2014

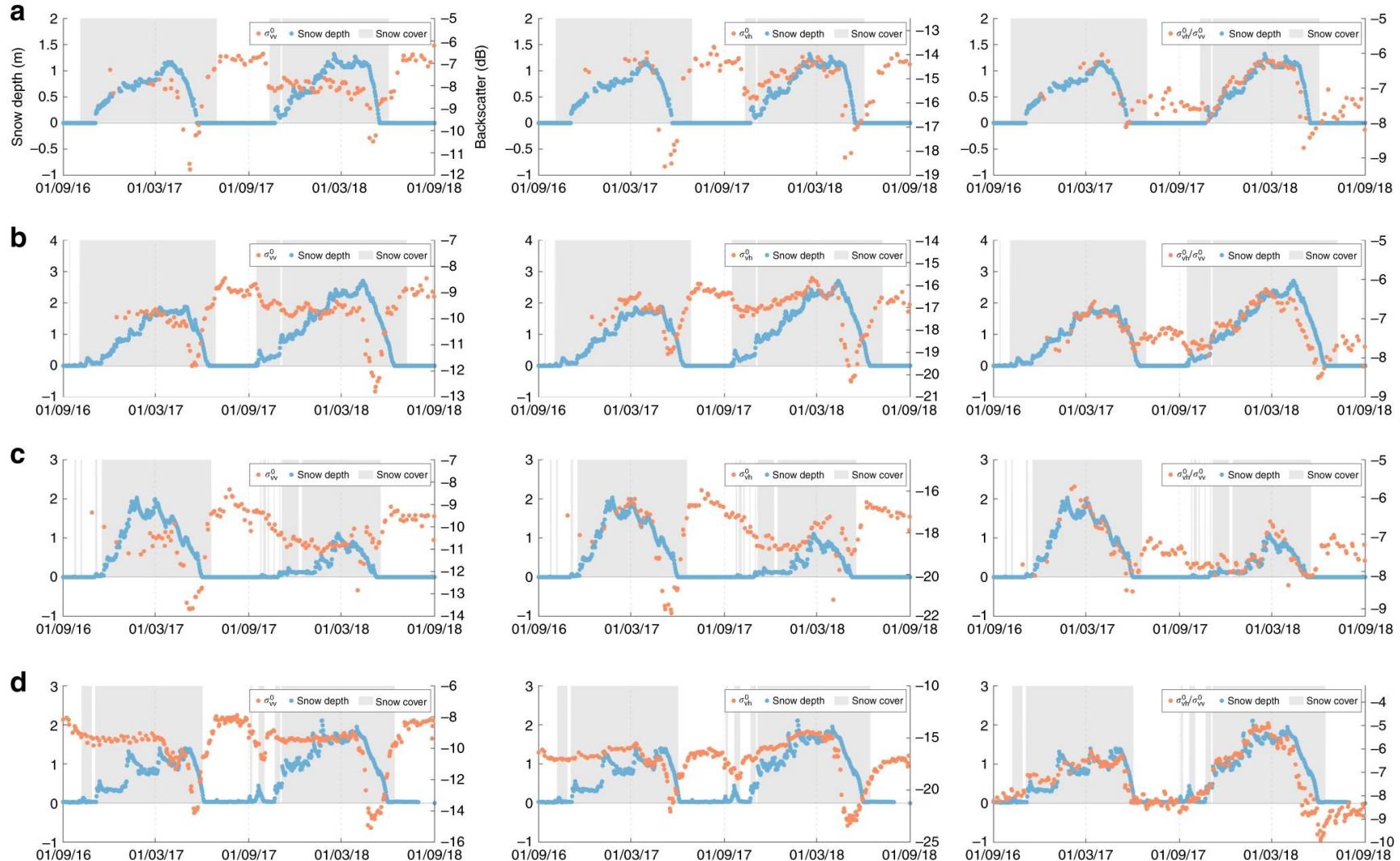


Rott et al., 2018

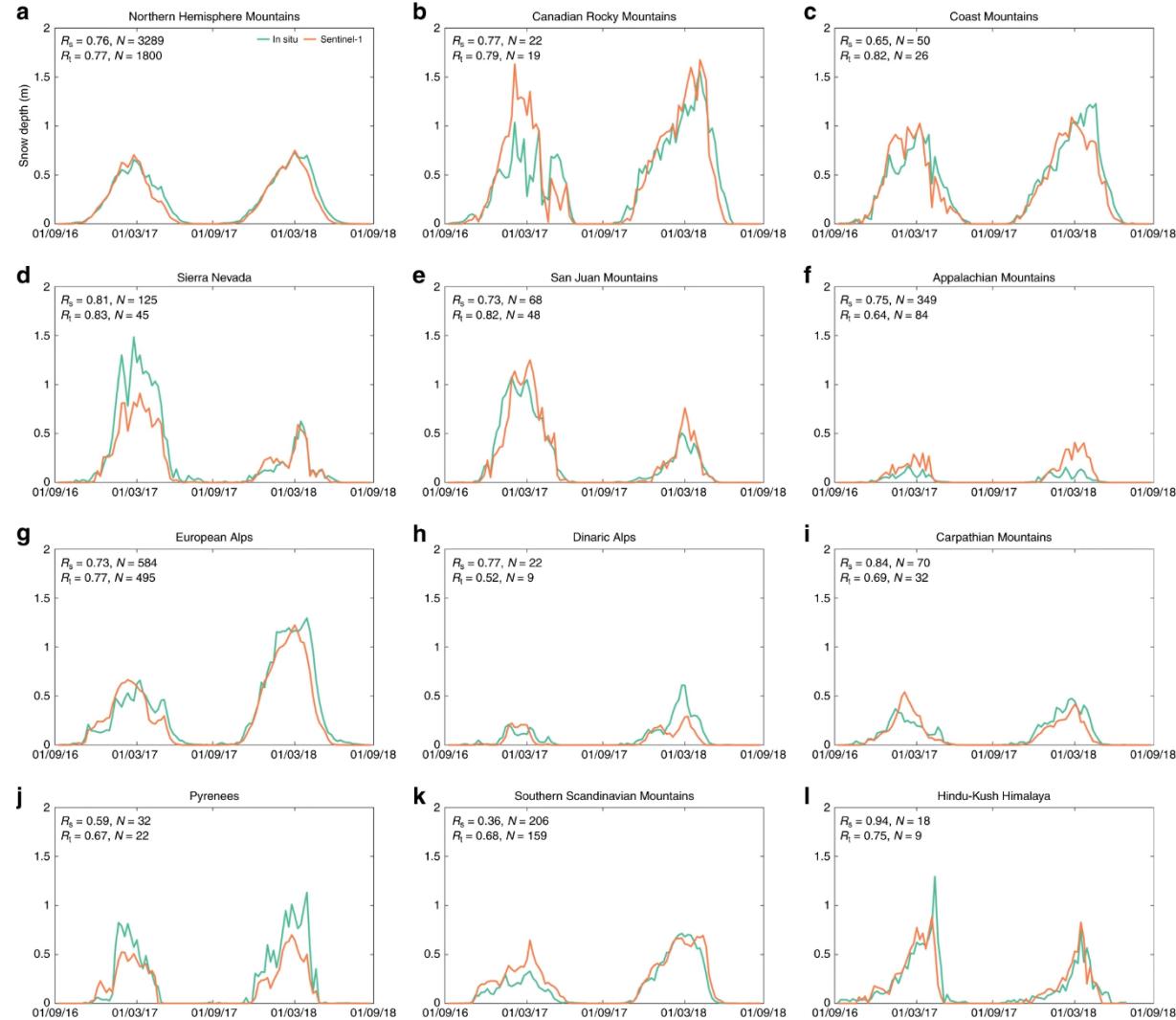
Radar (active) – Snow Depth



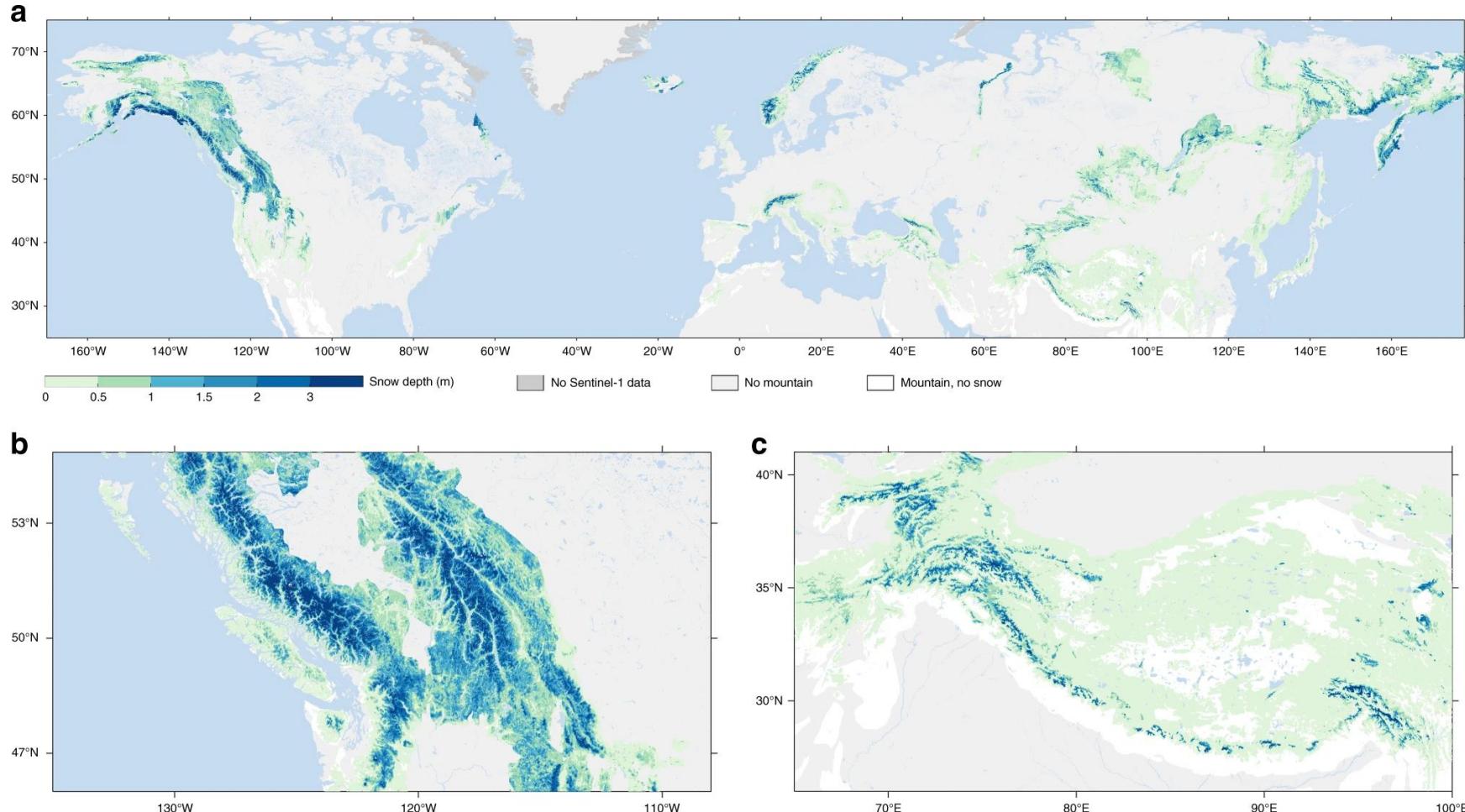
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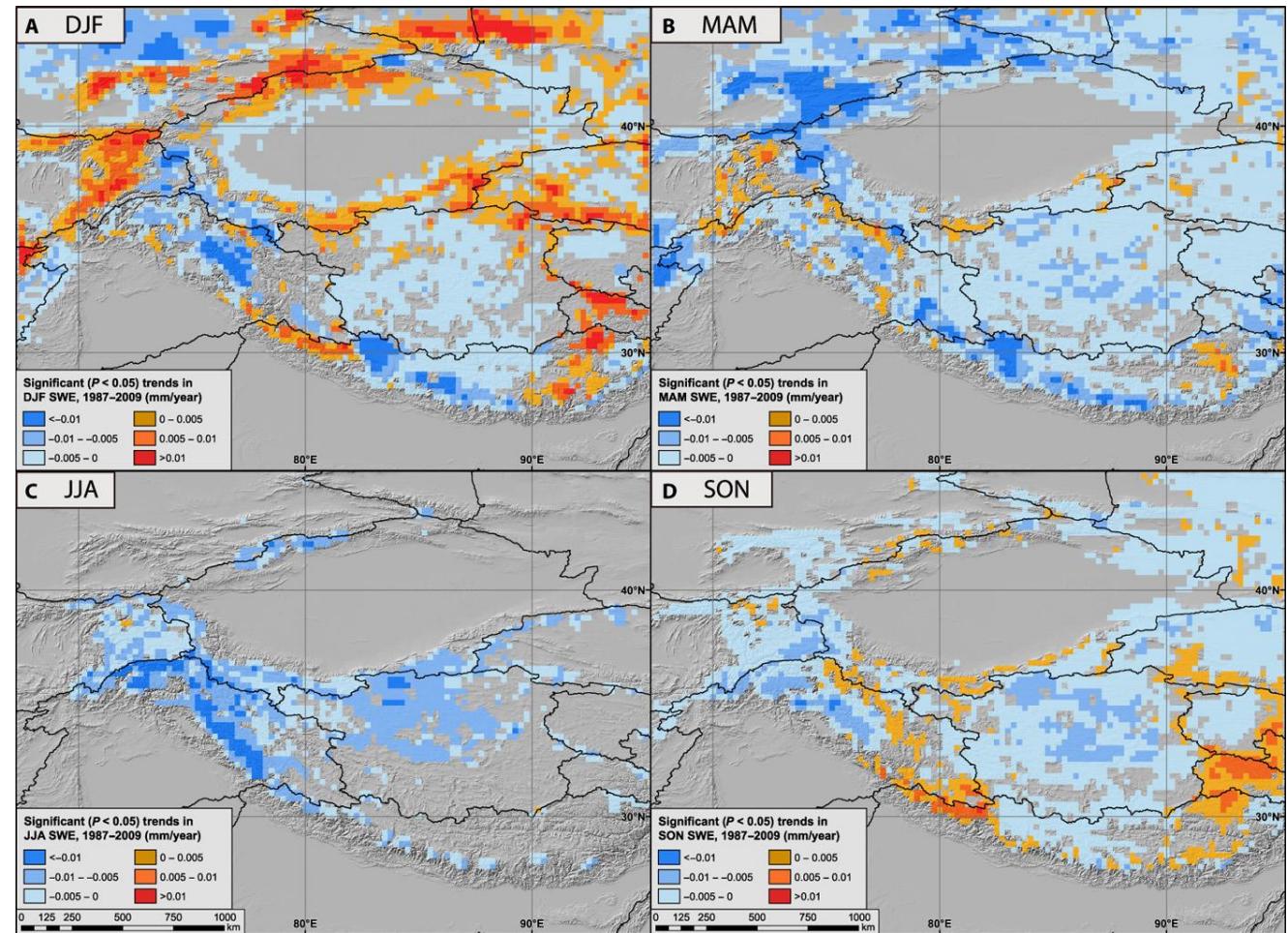


Radar (active) – Snow Depth



Radar (passive) – Snow Water Equivalent

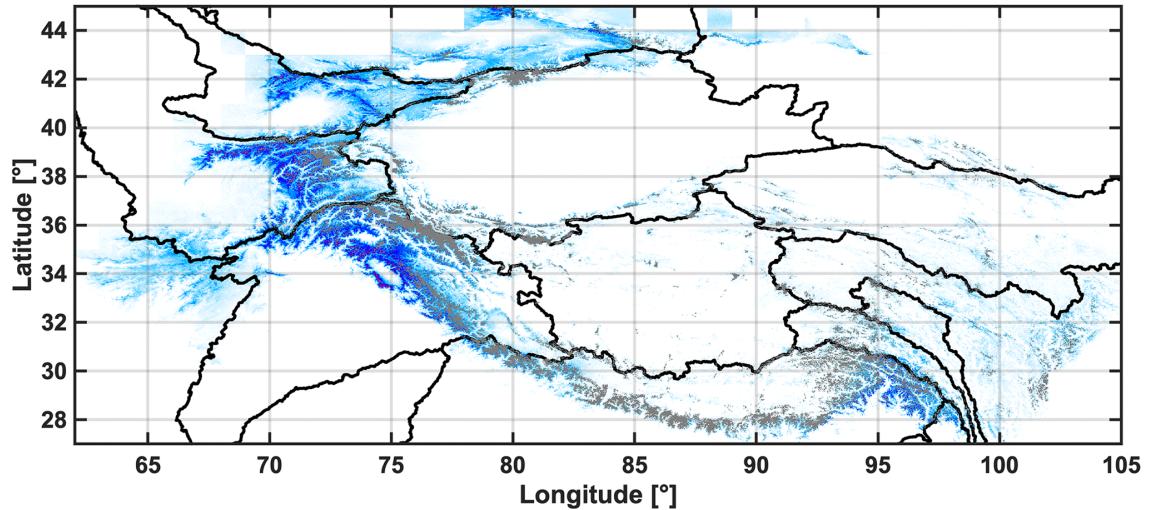
- SSMI (passive, 1979 – 2009)
- 27km resolution



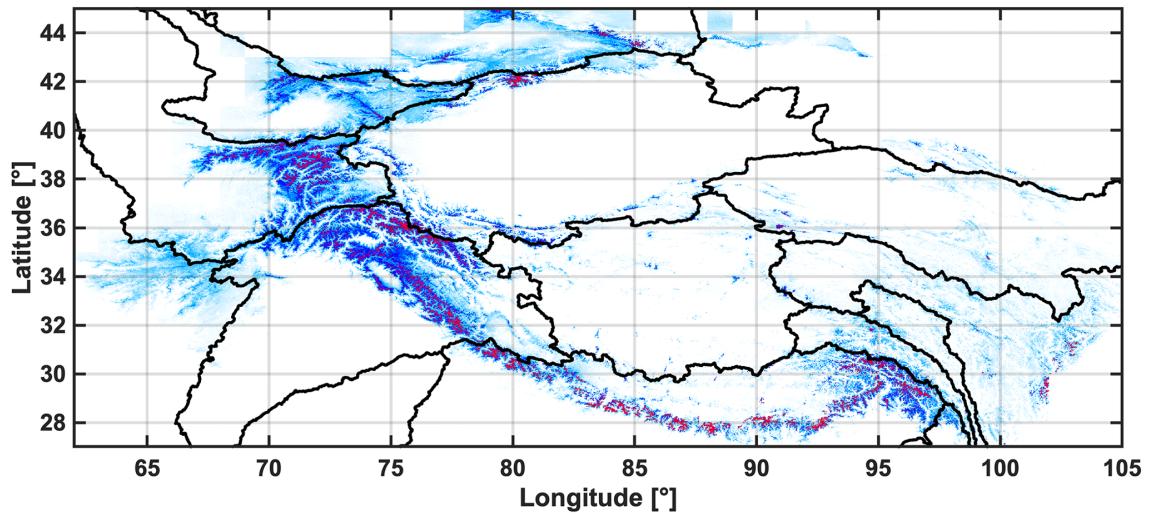
Modelled SWE

- Using fSCA
- Combine that with a simple model, based on climatic forcings
- 500 m resolution

(a)

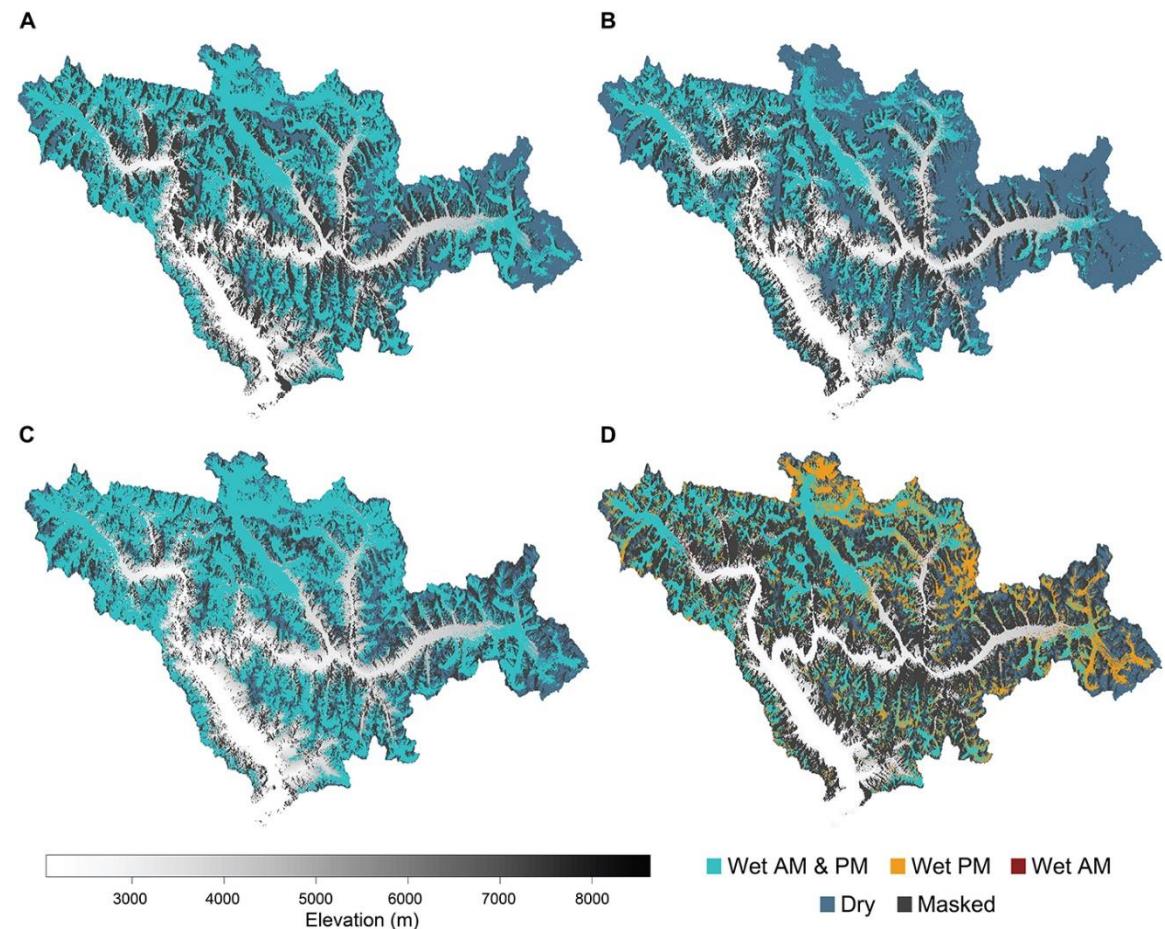


(b)



What can't we see (or remains difficult)

- Melt getting there
- Sublimation
- Refreezing
- Snow redistribution (wind drift/erosion)



Refreezing

- 20% - 30% of snowmelt is refrozen in Central Himalaya
- ~15% of positive energy is needed to “remelt” frozen meltwater

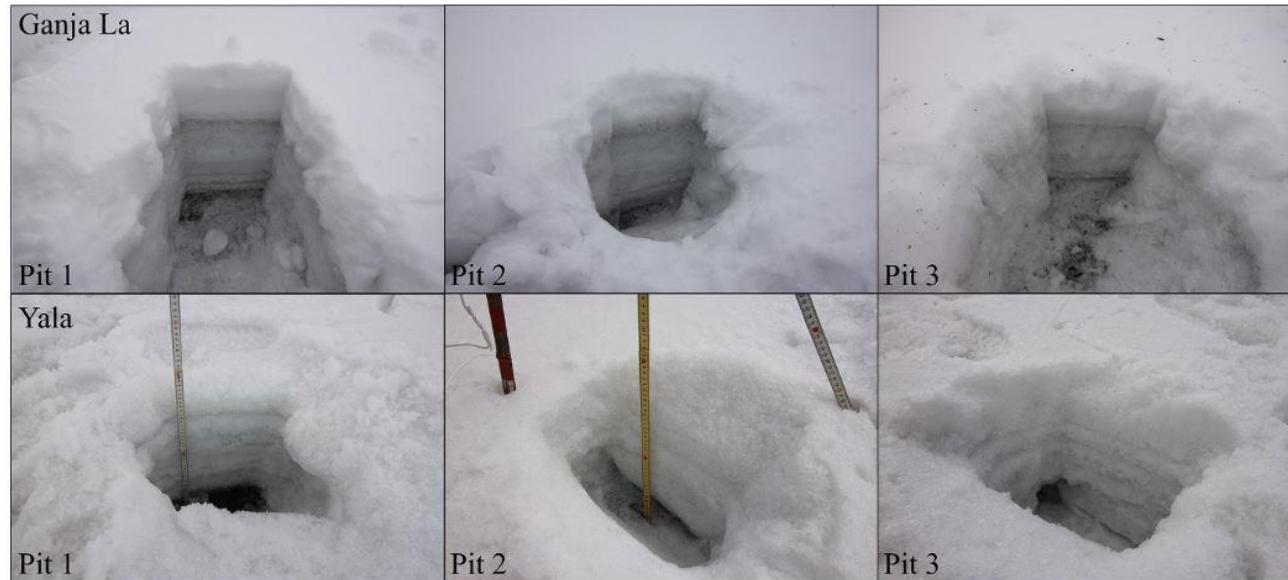


Fig. 8. Pictures of three snow pits at Ganja La on April 30th 2018 and Yala on April 25th 2018. The snow pits at Ganja La have a depth of 41–42 cm and contain four or five ice layers. The snow pits at Yala have a depth of 38–44 cm and contain five ice layers.

Refreezing

- catchment scale research allows us to understand and model processes
- acts as point scale validation of remote sensing products

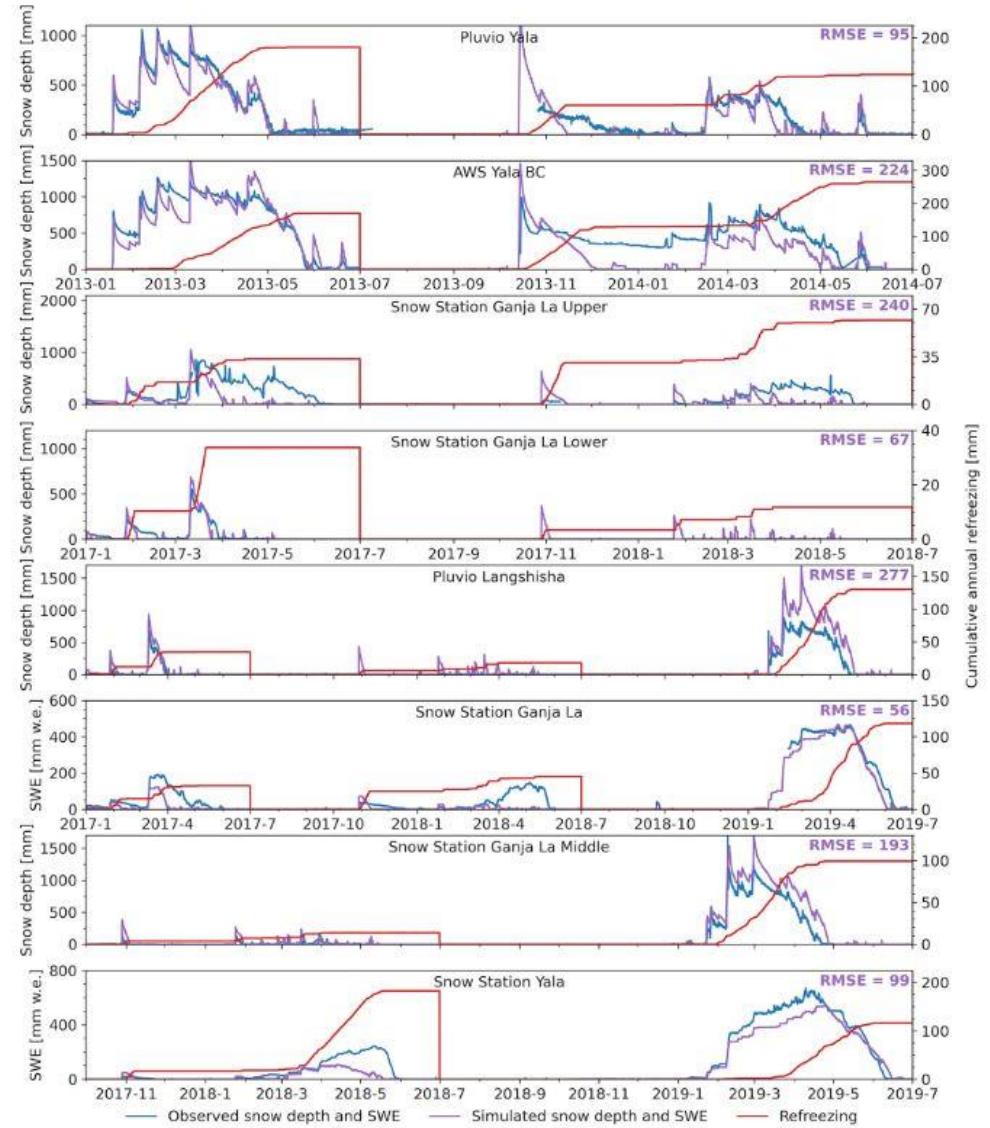


Fig. 6. Simulated and observed snow depth and SWE at stations for periods with measurements, including the simulated cumulative annual refreezing at the stations.

Measuring snow

Snow water equivalent

- snow pits
- snow pillows
- gamma-ray sensor



Measuring snow

Snow water equivalent





Thank you

Let's protect
the pulse.

