

ASYMPTOTIC WAVE-FRONT

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The Fourier transform

Transform and inversion (may be defined in slightly different ways)

$$\tilde{f} = \int_{-\infty}^{\infty} f(x) e^{-kx} dx, \quad f = \frac{1}{2\pi} \int_{-\infty}^{\infty} \tilde{f}(k) e^{kx} dk.$$

Initial surface elevation: $\eta(x, 0) = \eta_0(x)$; no initial velocities.

Fourier transform on linear inviscid set \Rightarrow The right-going wave system:

$$\eta(x, t) = \frac{1}{2\pi} \Re \int_0^{\infty} \tilde{\eta}_0 e^{i\chi} dk$$
$$\chi \equiv kx - \omega(k)t,$$

where η_0 is assumed symmetric around $x = 0$.

Stationary phase

Stationary point

$$\frac{d\chi(k_s)}{dk} = 0$$

corresponding to

$$c_g(k_s) = \frac{x}{t}$$

Neighbourhood of $k_s \Rightarrow$ dominant contribution to Fourier integral

Limitation in finite depth

Depth: H

$c_g \leq c_0 \equiv \sqrt{gH} \Rightarrow$ Stationary phase for $\frac{x}{t} \leq c_0$

Neighbourhood $\frac{x}{t} \sim c_0 \Rightarrow$ different approach

Wave front: close to $x = c_0 t$

Main contributions in Fourier integral from $k \rightarrow 0$.

$$\omega = \sqrt{gk \tanh(kH)} = c_0 k \left(1 - \frac{1}{6}(kH)^2 + O((kH)^4)\right),$$

where $c_0 = \sqrt{gh}$.

$$\chi \approx kx - c_0 \left(k - \frac{H^2}{6}k^3\right)t$$

and

$$\begin{aligned}\eta(x, t) &\sim \frac{1}{2\pi} \Re \int_0^\infty \tilde{\eta}_0(0) e^{i\left(kx - \left(c_0 k - \frac{H^2}{6}c_0 k^3\right)t\right)} dk \\ &= \frac{\hat{\eta}_0(0)}{2\pi} \int_0^\infty \cos\left(k\left(x - c_0 t\right) - \frac{H^2}{6}c_0 t k^3\right) dk\end{aligned}\tag{1}$$

Transformation to the Airy function

The substitution $k = s/m$ and $z = (x - c_0 t)/m$, where

$$m = \left(\frac{1}{2} H^2 c_0 t \right)^{\frac{1}{3}},$$

in (1) \Rightarrow

$$\eta(x, t) \sim \frac{\eta_0(0)}{2m} I(z),$$

with

$$I(z) = \frac{1}{\pi} \int_0^{\infty} \cos \left(\frac{1}{3} s^3 + zs \right) ds$$

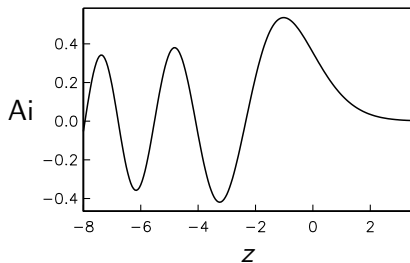
Integral formula for the Airy function, Ai.

The Airy function

$F = \text{Ai}$ is the solution of

$$\frac{d^2 F}{dz^2} - zF = 0, \quad (2)$$

which vanish as $z \rightarrow \infty$ and fulfills $\int_{-\infty}^{\infty} F dz = 1$.



Ai links an exponential behaviour to an oscillating one.

The wave front

close to $x = c_0 t$

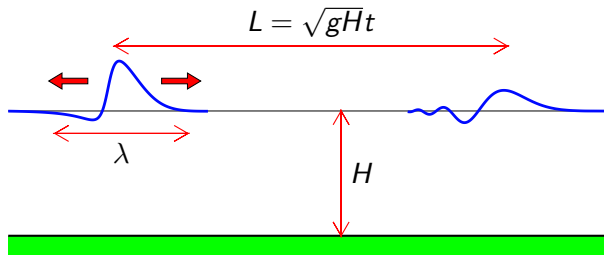
$$\eta \sim \frac{V}{2m} \text{Ai}(z) = \frac{\frac{1}{2}V}{(\frac{1}{2}c_0 H^2 t)^{\frac{1}{3}}} \text{Ai}\left(\frac{x - c_0 t}{(\frac{1}{2}c_0 H^2 t)^{\frac{1}{3}}}\right), \quad (3)$$

where $V = \tilde{\eta}(0)$ is the volume (per width) under $\eta(x, 0)$

- Leading crest close to $x = c_0 t$
- Shape independent of shape of $\eta(x, 0)$
- Maximum wave height decays as $\sim t^{-\frac{1}{3}}$
Will eventually dominate trailing waves ($\sim t^{-\frac{1}{2}}$).
- Length of leading crest increases as $\sim t^{\frac{1}{3}}$

Analysis may be extended to $V = 0$. Then front decays faster than trailing waves.

Example: Wave front and dispersion in tsunami propagation



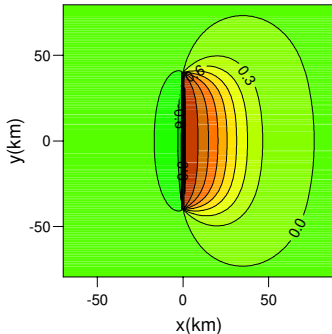
Dispersion often neglected for tsunamis due to large wavelength ($kH \ll 1 \Rightarrow c \approx c_g \approx \sqrt{gH}$).

Its significance depends on:

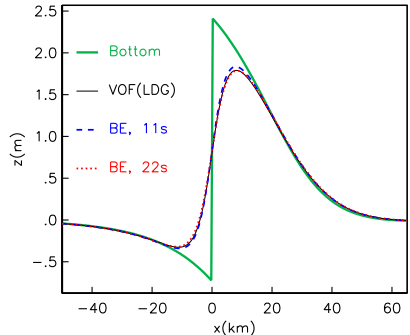
- 1 Extent of source relative to depth.
- 2 Propagation distance.

Example inspired by earthquake off Portugal (1969)

Bottom elevation



Surface response



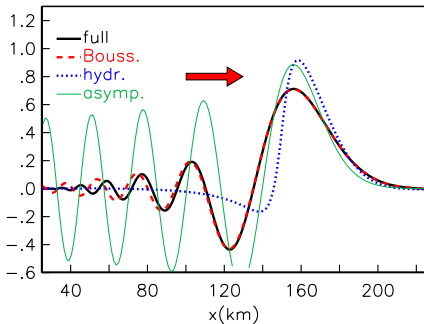
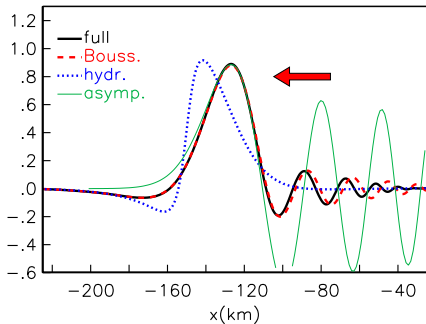
Magnitude: $M_s = 7.9$, $H = 5000$ m, inverse thrust fault, large dip angle $\approx 50^\circ$, fault length ≈ 70 km \Rightarrow rather confined bottom uplift

Left panel: co-seismic bottom-uplift from Okada's formula

Right panel: hydrodynamic response for center line from 2D theories

2D response used as initial condition.

After $t = 11.3 \Rightarrow L = c_0 t = 150 \text{ km}$



Curves:

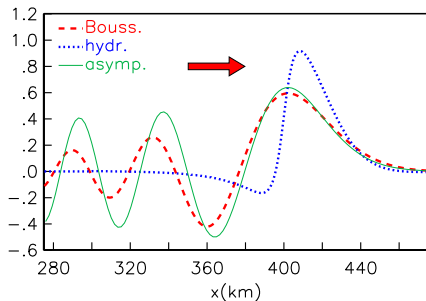
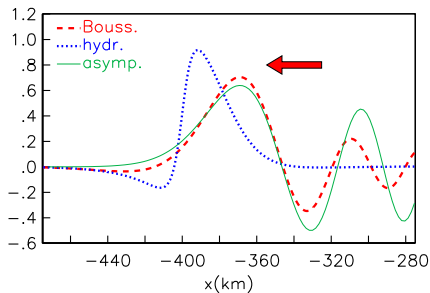
'full': full potential theory (numerical)

'Bouss.': Boussinesq equation (sec. 2.11 in Comp.) (numerical)

'hydr.': Shallow water solution (sec. 2.10); half the initial shape

'asyp.': Asymptotic solution for wave front

After $t = 30$ min $\Rightarrow L = 400\text{km}$



Now asymptote and numerical solution are becoming close.