

3D inversion for estimating total magnetization direction using equivalent layer technique

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ABSTRACT

We have developed a new method for estimating the total magnetization direction of magnetic sources based on equivalent layer technique using total field anomaly data. In this approach, we do not have to impose a strong information about the shape and the depth of the sources, and do not require a regularly spaced data. Usually, this technique is used for processing potential data estimating a 2D magnetic moment distribution over a fictitious layer composed by dipoles below the observation plane. In certain conditions, when the magnetization direction of equivalent sources is almost the same of true body, the estimated magnetic property over the layer is all positive. The methodology uses a positivity constraint to estimate a set of magnetic moment and a magnetization direction of the layer through an iterative process. Mathematically, the algorithm solves a least squares problem in two steps: the first one solves a linear problem for estimating a magnetic moment and the second solves a non-linear problem for magnetization direction of the layer. We test the methodology applying to synthetic data for different geometries and magnetization types

of sources. Moreover, we applied this method to field data from Goias Alkaline Province (GAP), center of Brazil, showing that the methodology can be a good tool for estimating the magnetization component of the alkaline intrusion complex in the Diorama region. The result for this complex suggests that this source has a remarkable strong remanent magnetization component. The magnetization direction estimated for this complex is -47° and -111° for inclination and declination, respectively.

METHODOLOGY

Considering a Cartesian coordinate system with x -, y - and z -axis being oriented to north, east and downward, respectively. Let $\Delta T(x_i, y_i, z_i)$ be the total field anomaly, at the i -th position (x_i, y_i, z_i) , produced by a continuous layer located below the observation plane on the depth z_c , where $z_c > z_i$, and $p(x', y', z_c)$ is the distribution of magnetic dipoles per unit area over the layer surface. In this case, the total field anomaly produced by this continuous layer is given by the equation

$$\Delta T(x_i, y_i, z_i) = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} p(x', y', z_c) [\gamma_m \hat{\mathbf{F}}_0^T \mathbf{H} \hat{\mathbf{h}}(\mathbf{q})] dx' dy', \quad (1)$$

where γ_m is a constant proportional to the vacuum permeability, $\hat{\mathbf{F}}_0$ is a unit vector with the same direction of the geomagnetic field \mathbf{F}_0 and \mathbf{H} is a 3×3 matrix equal to

$$\mathbf{H} = \begin{bmatrix} \partial_{xx}\phi & \partial_{xy}\phi & \partial_{xz}\phi \\ \partial_{yx}\phi & \partial_{yy}\phi & \partial_{yz}\phi \\ \partial_{zx}\phi & \partial_{zy}\phi & \partial_{zz}\phi \end{bmatrix}, \quad (2)$$

where $\partial_{\alpha\beta}\phi$, $\alpha = x, y, z$, $\beta = x, y, z$, is the second derivative of the function

$$\phi(x - x', y - y', z' - z_c) = \frac{1}{r}, \quad (3)$$

where $r = [(x - x')^2 + (y - y')^2 + (z_c - z')^2]^{1/2}$ and $\hat{\mathbf{h}}(\mathbf{q})$ is a unit vector with the magnetization direction of the layer that depends on the vector \mathbf{q} given by

$$\mathbf{q} = \begin{bmatrix} i \\ d \end{bmatrix}, \quad (4)$$

where i and d is the inclination and declination, respectively.