

# Estimating total magnetization direction using equivalent-layer technique

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## ABSTRACT

We have developed a new method for estimating the total magnetization direction of magnetic sources based on equivalent layer technique using total field anomaly data. In this approach, we do not have to impose a strong information about the shape and the depth of the sources, and do not require a regularly spaced data. Usually, this technique is used for processing potential data estimating a 2D magnetic moment distribution over a fictitious layer composed by dipoles below the observation plane. In certain conditions, when the magnetization direction of equivalent sources is almost the same of true body, the estimated magnetic property over the layer is all positive. The methodology uses a positivity constraint to estimate a set of magnetic moment over the layer and a magnetization direction of the layer through a iterative process. Mathematically, the algorithm solve a least squares problem in two steps: the first one solve a linear problem for estimating a magnetic moment and the second solve a non-linear problem for magnetization direction of the layer. We test the methodology applying to synthetic data for a complicated scenarios and

geometries of sources. Moreover, we applied the method to field data from Goias Alkaline Province (GAP), center of Brazil, over Montes Claros complex. The result suggests the area is composed by intrusions with remarkable strong remanent magnetization component, being in agreement with the current literature for this region.

## METHODOLOGY

### Fundamentals of magnetic equivalent layer and the positive magnetic-moment distribution

Considering a Cartesian coordinate system with  $x$ -,  $y$ - and  $z$ -axis being oriented to north, east and downward, respectively. Let  $\Delta T_i \equiv \Delta T(x_i, y_i, z_i)$  be the total-field anomaly, at the  $i$ th position  $(x_i, y_i, z_i)$ , produced by a continuous layer located below the observation plane at a depth equal to  $z_c$ , where  $z_c > z_i$ , and  $p(x', y', z_c)$  is the distribution of magnetic dipoles per unit area over the layer. The total-field anomaly produced by this continuous layer is given by

$$\Delta T_i = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} p(x', y', z_c) [\gamma_m \hat{\mathbf{F}}_0^T \mathbf{M}(x_i, y_i, z_i, x', y', z_c) \hat{\mathbf{m}}(\mathbf{q})] dx' dy', \quad (1)$$

where  $\gamma_m$  is a constant proportional to the vacuum permeability,  $\hat{\mathbf{F}}_0$  is a unit vector with the same direction of the main geomagnetic field given by

$$\hat{\mathbf{F}}_0 = \begin{bmatrix} \cos I \cos D \\ \cos I \sin D \\ \sin I \end{bmatrix}, \quad (2)$$

where  $I$  and  $D$  are, respectively, the inclination and declination and  $\mathbf{M}(x_i, y_i, z_i, x', y', z_c)$  is a  $3 \times 3$  dimensional matrix (Oliveira-Jr et al., 2015) equal to

$$\mathbf{M}(x_i, y_i, z_i, x', y', z_c) = \begin{bmatrix} \partial_{xx}\phi & \partial_{xy}\phi & \partial_{xz}\phi \\ \partial_{yx}\phi & \partial_{yy}\phi & \partial_{yz}\phi \\ \partial_{zx}\phi & \partial_{zy}\phi & \partial_{zz}\phi \end{bmatrix}, \quad (3)$$

where  $\partial_{\alpha\beta}\phi$ ,  $\alpha = x, y, z$  and  $\beta = x, y, z$ , is the second derivative of the scalar function

$$\phi(x_i, y_i, z_i, x', y', z_c) = \frac{1}{[(x_i - x')^2 + (y_i - y')^2 + (z_i - z_c)^2]^{\frac{1}{2}}}. \quad (4)$$

with respect to the Cartesian coordinates  $x_i$ ,  $y_i$  and  $z_i$  of the observation points. The  $\hat{\mathbf{m}}(\mathbf{q})$  is a unit vector with the magnetization direction of the dipoles over layer given by

$$\hat{\mathbf{m}}(\mathbf{q}) = \begin{bmatrix} \cos \tilde{\mathbf{i}} \cos \tilde{d} \\ \cos \tilde{\mathbf{i}} \sin \tilde{d} \\ \sin \tilde{\mathbf{i}} \end{bmatrix} \quad (5)$$

and  $\mathbf{q}$  is a  $2 \times 1$  vector with components given by

$$\mathbf{q} = \begin{bmatrix} \tilde{\mathbf{i}} \\ \tilde{d} \end{bmatrix}, \quad (6)$$

where  $\tilde{\mathbf{i}}$  and  $\tilde{d}$  are the inclination and declination of the magnetization direction of the dipoles on the layer, respectively. We can also notice that the vector defined in equation 5 has a single and uniform magnetization direction of all dipoles on the layer. For convenience, this unit vector can be rewritten as follows

$$\hat{\mathbf{m}}(\mathbf{q}) = \mathbf{R} \hat{\mathbf{h}}, \quad (7)$$

where  $\hat{\mathbf{h}}$  defines the uniform magnetization direction of an arbitrary magnetic source and  $\mathbf{R}$  is a  $3 \times 3$  matrix obtained from Euler's rotation theorem. This theorem states that any rotation can be parametrized by using three parameters called Euler angles (H. Goldstein and Safko, 1980). That is, if all dipoles that set up the equivalent layer have the same magnetization direction  $\hat{\mathbf{m}}(\mathbf{q})$  and this direction is the same as the true magnetic source  $\hat{\mathbf{h}}$ ,

then the matrix  $\mathbf{R}$  (equation 7) is equal to identity. For this reason, the total-field anomaly produced by equivalent layer at the  $i$ th position  $(x_i, y_i, z_i)$  (equation 1) can be rewritten as

$$\Delta T_i = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} p(x', y', z_c) [\gamma_m \hat{\mathbf{F}}_0^T \mathbf{M}(x_i, y_i, z_i, x', y', z_c) \hat{\mathbf{h}}] dx' dy', \quad (8)$$

which represents the total-field anomaly produced by continuous layer with the same direction of the arbitrary magnetic source. Thus, the RTP field  $\Delta T_i^{PL}$  produced by equivalent layer at the point  $(x_i, y_i, z_i)$  is equal to

$$\Delta T_i^{PL} = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} p(x', y', z_c) [\gamma_m \partial_{zz} \phi(x_i, y_i, z_i, x', y', z_c)] dx' dy', \quad (9)$$

where  $\partial_{zz} \phi(x_i, y_i, z_i, x', y', z_c)$  is the second derivative of the inverse of distance (equation 4) with respect of  $z_i$ , evaluated at the point  $(x_i, y_i, z_i)$ . However, by considering the RTP field  $\Delta T_i^{PS}$  produced by an arbitrary uniformly magnetized source, we have

$$\Delta T_i^{PS} = \gamma_m \partial_{zz} \Gamma(x_i, y_i, z_i) m, \quad (10)$$

which represents the total-field anomaly produced at the pole, where  $m$  is the magnetization intensity of the magnetic source. The  $\partial_{zz} \Gamma(x_i, y_i, z_i)$  is the second derivative in relation to  $z_i$  of a scalar function  $\Gamma(x_i, y_i, z_i)$  that depends on source geometry and the observation point  $(x_i, y_i, z_i)$ . From equation 9 and 10, we obtain

$$m \partial_{zz} \Gamma(x_i, y_i, z_i) = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} p(x', y', z_c) \partial_{zz} \phi(x_i, y_i, z_i, x', y', z_c) dx' dy'. \quad (11)$$

We can notice that equation 11 can be calculated differentiating the following equation

$$m \partial_z \Gamma(x_i, y_i, z_i) = \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} \frac{p(x', y', z_c)(z_c - z_i)}{[(x_i - x')^2 + (y_i - y')^2 + (z_i - z_c)^2]^{\frac{3}{2}}} dx' dy', \quad (12)$$

where  $z_c > z_i$ , with respect to the vertical component  $z_i$ . From potential-field theory, we can highlight the classical upward continuation integral

$$U(x_i, y_i, z_i) = \frac{(z_c - z_i)}{2\pi} \int_{-\infty}^{+\infty} \int_{-\infty}^{+\infty} \frac{U(x', y', z_c)}{[(x_i - x')^2 + (y_i - y')^2 + (z_i - z_c)^2]^{\frac{3}{2}}} dx' dy', \quad (13)$$

where the function  $U(x_i, y_i, z_i)$  is an hamornic function at all  $(x_i, y_i, z_i)$  (Blakely, 1996). In this case, this function represents the total-field anomaly at the point  $(x_i, y_i, z_i)$  which can be mathematically interpreted as the convolution between its values  $U(x', y', z_c)$  and the vertical derivative in relation to  $z_i$  of the equation 4, evaluated on the horizontal plane  $z_i = z_c$ . Therefore, according the classical upward continuation function (equation 13), the magnetic-moment distribution  $p(x', y', z_c)$  in equation 12 assumes the form

$$p(x', y', z_c) = \frac{m}{2\pi} \partial_z \Gamma(x', y', z_c), \quad (14)$$

where  $\partial_z \Gamma(x', y', z_c)$  is the derivative of the scalar function  $\partial_z \Gamma(x_i, y_i, z_i)$  in relation to  $z_i$  evaluated over the equivalent layer. The most interesting aspect of equation 14 is that the magnetic-moment distribution is defined as the product of a positive constant  $\frac{m}{2\pi}$  and the function  $\partial_z \Gamma(x', y', z_c)$ , which is all positive at all points  $(x', y', z_c)$  over the equivalent layer. This relation is simular to that presented by Pedersen (1991) and Li et al. (2014). In the wavenumber domain, these authors determined the magnetic-moment distribution within a continuous equivalent layer with the same magnetization direction as the local-geomagnetic field at the pole. They also considered a planar equivalent layer located below and parallel to a horizontal plane containing the observed total-field anomaly. They assume a magnetic

source having a purely induced magnetization. Under these assumptions, Pedersen (1991) and Li et al. (2014) concluded that the magnetic moment distribution within the continuous equivalent layer is proportional to the pseudogravity anomaly produced by the source on the plane of the equivalent layer. By following different approaches using magnetic microscopy data, Baratchart et al. (2013) and Lima and Weiss (2016) pointed out by imposing a non-negativity constraint, the solution to the inverse problem is a unique magnetic moment distribution. Here, we do not follow the same wavenumber-domain reasoning used by all these authors. Equation 14, however, generalizes this positivity condition because it (1) holds true for all cases in which the magnetization of the equivalent layer has the same direction as the true magnetization of the sources, whenever it is purely induced or not, (2) does not require that the observed total-field anomaly data be on a plane and, (3) does not require a planar equivalent layer.

## Parametrization and forward problem

In practical situations, it's not possible to determine a continuous magnetic-moment distribution  $p(x', y', z_c)$  over the layer as shown in equation 1. For this reason, the layer has to be approximated by a discrete set of dipoles with unit volume located at a constant depth  $z = z_c$ . By discretizing the integrand of equation 1, the predicted total-field anomaly at the point  $(x_i, y_i, z_i)$  is given by

$$\Delta T_i = f_i(\mathbf{s}), \quad (15)$$

where  $\Delta T_i$  is the  $i$ th element of the  $N$ -dimensional predicted data vector  $\Delta \mathbf{T}(\mathbf{s})$ . The function  $f_i$  maps the parameter onto the data, in which parameter vector  $\mathbf{s}$  is formed by

an  $M$ -dimensional vector  $\mathbf{p}$  vector whose  $j$ th element  $p_j$  is the magnetic moment of the  $j$ th dipole with magnetization direction  $\mathbf{q}$  (equation 6) to all dipoles (figure 1). Explicitly, the function  $f_i$  is described as

$$f_i(\mathbf{s}) = \sum_{j=1}^M p_j g_{ij}(\mathbf{q}) = \mathbf{p}^T \mathbf{g}_i(\mathbf{q}) \quad (16)$$

where

$$g_{ij}(\mathbf{q}) = \gamma_m \hat{\mathbf{F}}_0^T \mathbf{M}_{ij} \hat{\mathbf{m}}(\mathbf{q}) \quad (17)$$

is an harmonic function representing the total-field anomaly produced at the  $i$ th position  $(x_i, y_i, z_i)$  by a dipole located at  $(x_j, y_j, z_c)$  with unitary magnetic-moment intensity. The matrix  $\mathbf{M}_{ij}$  is formed by the second derivatives of a function  $\phi_{ij}$  that depends on the inverse of the scalar function  $r_{ij} = [(x_i - x_j)^2 + (y_i - y_j)^2 + (z_i - z_c)^2]^{1/2}$ , analogously to equation 3 and 4. From equation 15-17, we can notice that the predicted total-field anomaly  $\Delta \mathbf{T}(\mathbf{s})$  has a linear relation with the magnetic moment  $\mathbf{p}$  and a nonlinear relation with the magnetizaion direction  $\mathbf{q}$ .

## Inverse problem

Let  $\Delta \mathbf{T}^o$  be an  $N$ -dimensional vector whose  $i$ th element  $\Delta T_i^o$  is the total-field anomaly observation produced by a magnetic source at the point  $(x_i, y_i, z_i)$ ,  $i = 1, \dots, N$ . In order to estimate the magnetization direction, we have to formulate an inverse problem by imposing a positivity constraint on the magnetic-moment distribution. It can be performed minimizing the difference between the observed data  $\Delta \mathbf{T}^o$  and the predicted data  $\Delta \mathbf{T}(\mathbf{s})$  by solving the following constrained problem of



$$\text{minimizing} \quad \Psi(\mathbf{s}) = \|\Delta \mathbf{T}^o - \Delta \mathbf{T}(\mathbf{s})\|_2^2 + \mu f_0 \|\mathbf{p}\|_2^2 \quad (18a)$$

$$\text{subject to} \quad \mathbf{p} \geq 0 \quad (18b)$$

where in equation 18a the first term is the data misfit function, the second term of equation 18a is a zeroth-order Tikhonov regularization function,  $\mu$  is the regularizing parameter,  $\|\cdot\|_2^2$  represents the squared Euclidean norm and  $f_0$  is a normalizing factor. In equation 18b,  $\mathbf{0}$  is a null vector and the inequality sign is applied element by element. This inequality imposes positivity constraints on the estimated magnetic moments of all dipoles, which is solved by using the nonnegative least squares (NNLS) proposed by Lawson and Hanson (1974).

The process of minimizing the goal function shown in equation 18a starts with an initial approximation to the parameter vector  $\mathbf{s}_0$  and then solving a sequence of linear problem of estimating a correction  $\Delta \mathbf{s}$ . The procedure is repeated until a minimum of goal function (equation 18a) is reached. This procedure is determined by using a gradient-based iterative optimization method like Gauss-Newton (Aster et al., 2005). Mathematically, the correction  $\Delta \mathbf{s}$  is expressed as a second-order expansion of the goal function given by

$$\Psi(\mathbf{s}_0 + \Delta \mathbf{s}) \approx \mathbf{J}(\mathbf{s}_0) + \mathbf{J}^T(\mathbf{s}_0)\Delta \mathbf{s} + \frac{1}{2}\Delta \mathbf{s}^T \mathbf{H}(\mathbf{s}_0)\Delta \mathbf{s} \quad (19)$$

in which  $\mathbf{J}(\mathbf{s}_0)$  and  $\mathbf{H}(\mathbf{s}_0)$  are, respectively, the gradient vector and the Hessian matrix of equation 18a. Thus, taking the gradient of the function 19, the parameter perturbation vector  $\Delta \mathbf{s}^k$  at the  $k$ th iteration is obtained by solving the linear system

$$\mathbf{H}^k(\mathbf{s}_0)\Delta \mathbf{s}^k = -\mathbf{J}^k(\mathbf{s}_0). \quad (20)$$

However, looking in more detail at the equation 20, the linear system can be given by equation

$$\left[ \begin{array}{c|c} \mathbf{H}_{pp}^k & \mathbf{H}_{pq}^k \\ \hline \mathbf{H}_{qp}^k & \mathbf{H}_{qq}^k \end{array} \right] \begin{bmatrix} \Delta \mathbf{p}^k \\ \Delta \mathbf{q}^k \end{bmatrix} = - \begin{bmatrix} \mathbf{J}_p^k \\ \mathbf{J}_q^k \end{bmatrix}, \quad (21)$$

in which  $\mathbf{J}_\alpha^k$  and  $\mathbf{H}_{\alpha\beta}^k$ , where  $\alpha = p, q$  and  $\beta = p, q$ , are the gradient vector and the Hessian matrix calculated in relation to each component of the magnetic-moment vector  $\mathbf{p}$  and the magnetization direction vector  $\mathbf{q}$ , respectively. Besides, in order to simplify the linear system 21, we consider null cross-derivatives. As we shown earlier, the predicted total-field anomaly has a linear relation with magnetic moments  $\mathbf{p}^k$ . For this reason, the gradient vector and the Gauss-Newton approximation of the goal function (equation 18a) relative to the part of magnetic moments are, respectively,

$$\mathbf{J}_p^k = -2\mathbf{G}_p^{kT} \Delta \mathbf{T}^o \quad (22)$$

and

$$\mathbf{H}_{pp}^k = 2\mathbf{G}_p^{kT} \mathbf{G}_p^k + 2\mu f_0^k \mathbf{I} \quad (23)$$

where  $\mathbf{G}_p^k$  is the  $N \times M$  sensitivity matrix at the  $k$ th iteration, whose elements are composed by the derivative of equation 18a in relation of  $j$ th element of the vector  $\mathbf{p}^k$ ,  $\mathbf{I}$  is an identity matrix and the normalizing factor  $f_0^k$  is equal to

$$f_0^k = \frac{tr(\mathbf{G}_p^{kT} \mathbf{G}_p^k)}{M}, \quad (24)$$

where  $tr$  is denotated as the trace of the matrix  $\mathbf{G}_p^{kT} \mathbf{G}_p^k$  and  $M$  is the total number of

dipoles composing the layer. Thus, at the  $k$ th iteration, we have to solve the linear system given by

$$\left[ \mathbf{G}_p^{kT} \mathbf{G}_p^k + \mu f_0^k \mathbf{I} \right] \mathbf{p}^k = \mathbf{G}_p^{kT} \Delta \mathbf{T}^o. \quad (25)$$

Owing to nonlinear relation of the magnetization direction  $\mathbf{q}^k$  with the predicted total-field anomaly, the gradient vector and the Hessian matrix for this case are, respectively,

$$\mathbf{J}_q^k = -2\mathbf{G}_q^{kT} [\Delta \mathbf{T}^o - \Delta \mathbf{T}(\mathbf{s}^k)] \quad (26)$$

and

$$\mathbf{H}_{qq}^k = 2\mathbf{G}_q^{kT} \mathbf{G}_q^k \quad (27)$$

in which  $\mathbf{G}_q^k$  is a  $N \times 2$  sensitivity matrix, whose elements are composed by derivative of equation 18 in relation of each component of the vector  $\mathbf{q}^k$ , that are the inclination and declination, respectively. Nevertheless, in order to calculate the correction  $\Delta \mathbf{q}^k$  at the  $k$ th iteration, we use the Levenberg-Marquardt method (Aster et al., 2005) by solving the linear system

$$\left[ \mathbf{G}_q^{kT} \mathbf{G}_q^k + \lambda^k \mathbf{I} \right] \Delta \mathbf{q}^k = \mathbf{G}_q^{kT} [\Delta \mathbf{T}^o - \Delta \mathbf{T}(\mathbf{s}^k)], \quad (28)$$

where  $\lambda$  is the Marquardt parameter that is updated along the iterative process and  $\mathbf{I}$  is an identity matrix.

## Iterative process for magnetization estimation

The estimation of the magnetic moments  $\mathbf{p}^\sharp$  and the magnetization direction  $\mathbf{q}^\sharp$  is formulated as an inverse problem of minimizing the difference between the observed data  $\Delta\mathbf{T}^o$  and the predicted data  $\Delta\mathbf{T}(\mathbf{s})$ . As we can notice, the solution of the inversion is given by solving two linear systems, one shown in equation 25 and other in equation 28. For this reason, we propose a nested algorithm for solving the inverse problem in two steps.

Our iterative algorithm starts with an initial guess for the magnetization direction  $\mathbf{q}_0$ . At the  $k$ th iteration, we calculated a set of magnetic moment  $\mathbf{p}^k$  by imposing a positivity constraint on equation 25. After estimating the magnetic-moment distribution  $\mathbf{p}^k$  at the  $k$ th iteration using the previous estimate  $\mathbf{q}^k$  for the magnetization direction, we estimate a new vector  $\mathbf{q}^{k+1}$  by solving an unconstrained nonlinear inverse problem and applying the correction  $\Delta\mathbf{q}^k$  using the equation 28. The iterative process stops when the minimum of the goal function (equation 18a) is reached. An overview of the algorithm is shown in figure 2 and in the algorithm 1.

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**Algorithm 1:** Nested NNLS and Levenberg-Marquardt method

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**Input** :  $\Delta \mathbf{T}^o, \mathbf{q}_0$

**Output:**  $\mathbf{p}^\sharp, \mathbf{q}^\sharp$

```
while (not converge) or ( $i < i_{max}$ ) do
    step 1: Solve equation 25 using NNLS ;
    step 2: Compute goal function (18) ;
    while (not converge) or ( $j < j_{max}$ ) do
        step 3: Initialize the Levenberg-Marquardt algorithm ;
        step 4: Compute goal function (18);
        while  $k < k_{marq}$  do
            step 5: Solve equation 28 ;
            step 6: Apply the correction for magnetization direction ;
            step 7: Compute goal function (18);
        end
        step 8: Analysis of convergence for inner loop.
    end
    step 9: Analysis of convergence for outer loop.
end
```

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### The choice of layer depth $z_c$ and regularization parameter $\mu$

The procedure for the use of our methodology for estimating the total magnetization require the choice of two main parameters. The first one is the layer depth  $z_c$  as shown in figure 1 and the second is the regularization parameter  $\mu$  shown in equation 25.

The method of the choice of layer is based on a classical approach proposed by Dampney

(1969). The author pointed out that the layer depth should satisfy an interval from 2.5 to 6 times the grid spacing. It should be notice that this rule was applied on an evenly spaced data. However, the choice for applying our method should correspond to an interval from 2 to 3 times the greater grid spacing. It is necessarily to point out that this range of values was found empirically due to the application on a irregular grid data.

To solve the equation 25 we have to choose a reliable regularization parameter  $\mu$ . For this purpose, we use the L-curve method (Hansen and OLeary, 1993). This approach is widely used in the literature to find a regularizing parameter, which filtering out enough noise whithout loosing to much information in the final solution. The procedure of finding the parameter is basically to plot a curve of optimal values between the solution norm and residual norm. The corner of the curve is the final result which gives a threshold between the regularization function and the data misfit.

## APPLICATION TO SYNTHETIC DATA

We test the proposed method by applying it to three different synthetic data sets simulating complex geological scenarios. The first one is generated by a model containing a set of sources with different geometries, all of them with the same magnetization direction. The second data set is generated by a set of magnetized bodies, but one of them is a shallow interfering source with the same magnetization direction. In the third test, we violate the unidirectional approach by simulating an interfering shallow source with different magnetization direction from the other bodies.

For all tests, the simulated data were computed on an irregular grid of  $49 \times 25$  points (a total of  $N = 1225$  observations) at a constant height of  $100\text{ m}$ . We assume an observation area of extending  $12\text{ km}$  along the x- and y-axis, resulting in a grid spacing of  $250\text{ m}$  and  $500\text{ m}$  on x- and y-axis, respectively. The data were contaminated with pseudorandom Gaussian noise with zero mean and  $10\text{ nT}$  standard deviation. The main field direction simulated was  $I = -40^\circ$  and  $D = -22^\circ$  for the inclination and declination, respectively. For the inversion, we use an equivalent layer composed by a grid of  $49 \times 25$  dipoles (a total of  $M = 1225$  equivalent sources) positioned at a depth of  $z_c = 1150\text{ m}$  below the observation plane (2.5 times the greater grid spacing). We use the L-curve to choose the regularizing parameter ( $\mu$ ). The algorithm 1 starts with an initial guess  $\mathbf{q}_0 = (-10^\circ, -10^\circ)$  for inclination and declination, respectively.

### Unidirectional magnetization direction sources

In order to test the methodology, we generate a 3D prism with polygonal cross-section whose top is positioned at a depth of  $450\text{ m}$  and the bottom  $3150\text{ m}$  with magnetization

intensity of  $4 \text{ A/m}$ . We also generate two spheres with radius equal to  $500 \text{ m}$ , one of them with the coordinates of center  $x_c = 1800 \text{ m}$ ,  $y_c = -1800 \text{ m}$  and  $z_c = 1000 \text{ m}$  and the other with  $x_c = 800 \text{ m}$ ,  $y_c = 800 \text{ m}$  and  $z_c = 1000 \text{ m}$ . The magnetization intensity for the two both spheres is equal to  $3 \text{ A/m}$ . We produce two rectangular prisms with the same  $2.5 \text{ A/m}$  of magnetization intensity. The smaller prism has the top at a depth of  $450 \text{ m}$  and side lengths of  $1000 \text{ m}$ ,  $700 \text{ m}$  and  $500 \text{ m}$  along x-,y- and z-axis, respectively. The greater prism has the top at a depth of  $500 \text{ m}$  and side lengths of  $1000 \text{ m}$ ,  $2000 \text{ m}$  and  $1550 \text{ m}$  along x-,y- and z-axis. All simulated sources have inclination  $-25^\circ$  and declination  $30^\circ$ . The noise-corrupted data is shown in figure 3a.

Figure 3b shows the predicted data produced by equivalent layer. Figure 3c shows the residuals defined as the difference between the simulated data (figure 3a) and the predicted data (figure 3b). The residuals appear normally distributed with a mean of  $-0.30 \text{ nT}$  and a standard deviation of  $9.67 \text{ nT}$  as shown in figure 3d. The estimated magnetization direction  $\mathbf{q}^\#$  has inclination  $-28.6^\circ$  and declination  $30.8^\circ$ . Figure 3e shows the estimated magnetic-moment distribution  $\mathbf{p}^\#$ . The convergence of the algorithm 1 is shown in figure 3f. These results show that the all-positive estimated magnetic-moment distribution and the estimated magnetization direction produce an acceptable data fitting.



## APPENDIX A

### CONSEQUENCES OF HIGH-LATITUDE ESTIMATION

## REFERENCES

- Aster, R. C., B. Borchers, and C. H. Thurber, 2005, Parameter estimation and inverse problems (international geophysics): Academic Press.
- Baratchart, L., D. P. Hardin, E. A. Lima, E. B. Saff, and B. P. Weiss, 2013, Characterizing kernels of operators related to thin-plate magnetizations via generalizations of Hodge decompositions: *Inverse Problems*, **29**, 015004(29pp).
- Blakely, R. J., 1996, Potential theory in gravity and magnetic applications: Cambridge University press.
- Dampney, C. N. G., 1969, The equivalent source technique: *GEOPHYSICS*, **34**, 39–53.
- H. Goldstein, C. P. P. J., and J. L. Safko, 1980, Classical mechanics: Addison-Wesley.
- Hansen, P., and D. OLeary, 1993, The use of the l-curve in the regularization of discrete ill-posed problems: *SIAM Journal on Scientific Computing*, **14**, 1487–1503.
- Lawson, C. L., and R. J. Hanson, 1974, Solving least squares problems: SIAM.
- Li, Y., M. Nabighian, and D. W. Oldenburg, 2014, Using an equivalent source with positivity for low-latitude reduction to the pole without striation: *GEOPHYSICS*, **79**, J81–J90.
- Lima, E. A., and B. P. Weiss, 2016, Ultra-high sensitivity moment magnetometry of geological samples using magnetic microscopy: *Geochemistry, Geophysics, Geosystems*, **17**, 3754–3774.
- Oliveira-Jr, V. C., D. P. Sales, V. C. F. Barbosa, and L. Uieda, 2015, Estimation of the total magnetization direction of approximately spherical bodies: *Nonlinear Processes in Geophysics*, **22**, 215–232.
- Pedersen, L. B., 1991, Relations between potential fields and some equivalent sources: *GEOPHYSICS*, **56**, 961–971.

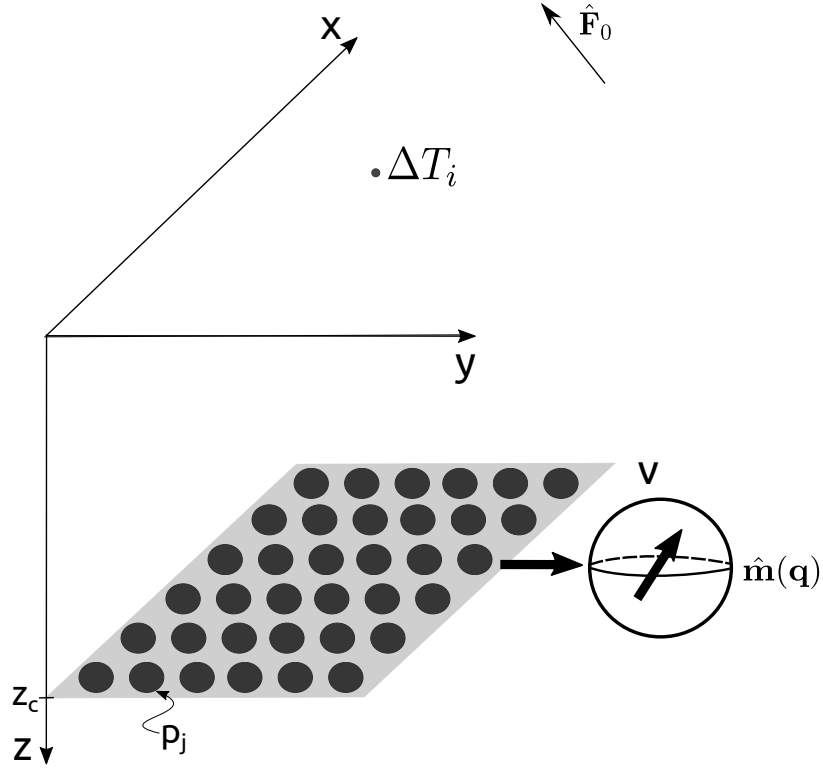


Figure 1: Schematic representation of an equivalent layer. The layer is positioned over the horizontal plane at a depth of  $z = z_c$ .  $\Delta T_i = f_i(\mathbf{s})$  is the predicted total-field anomaly at the point  $(x_i, y_i, z_i)$  produced by the set of  $M$  equivalent sources (black dots). Each source is located at the point  $(x_j, y_j, z_c)$ ,  $j = 1, \dots, M$ , and represented by a dipole with unity volume  $v$  with magnetization direction  $\hat{\mathbf{m}}(\mathbf{q})$  and magnetic moment  $p_j$ .

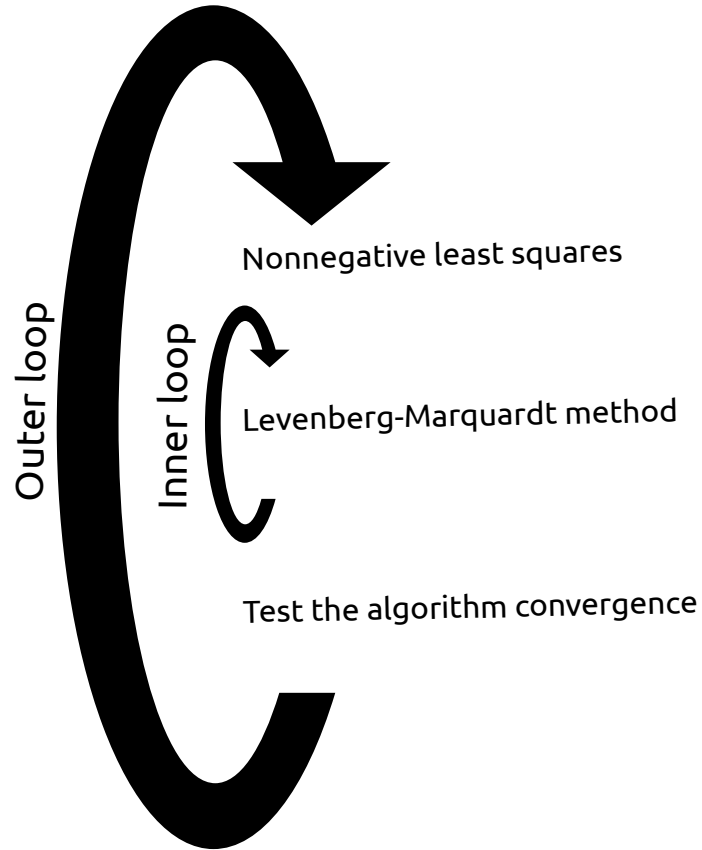


Figure 2: Iterative scheme overview for NNLS and Levenberg-Marquardt method for estimating magnetization direction. The outer loop is the nonnegative solution for magnetic-moment distribution and the inner loop calculates the magnetization direction correction using Levenberg-Marquardt method.

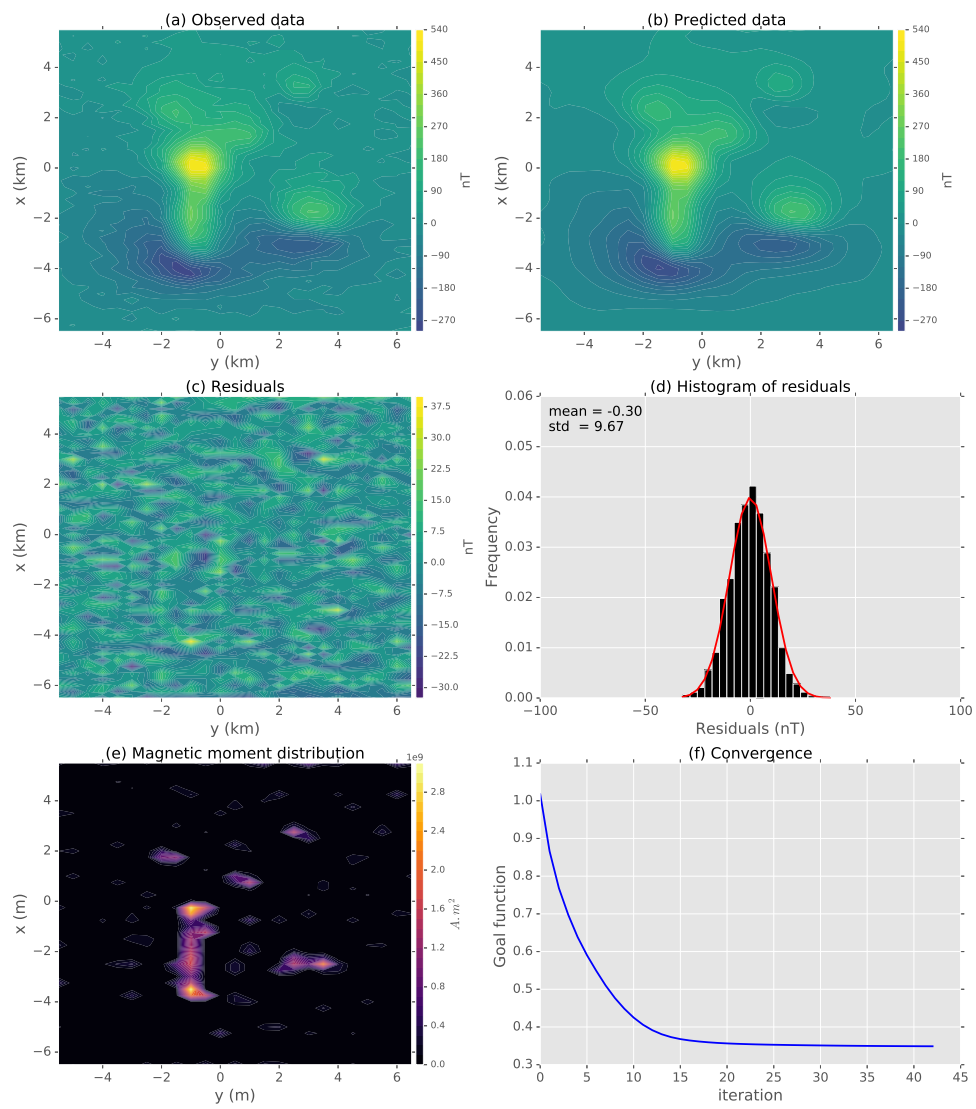


Figure 3: