


Observing Earth as an Exoplanet with *TESS*: Photometric Mapping of Surface Features in Reflected Light

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ABSTRACT

To date, a handful of exoplanets have been photometrically mapped using phase-modulated reflection or emission from their surfaces, but the small amplitudes of such signals have limited previous maps almost exclusively to coarse dipolar features on hot giant planets. Future instruments may enable more detailed mapping of terrestrial planets, including those in the habitable zone, but significant challenges will remain in analyzing and interpreting these data. In this work, we extract a rotationally-modulated reflected light curve of the Earth from the time-variable background flux in data from the Transiting Exoplanet Survey Satellite (*TESS*). We treat these data as that of an unknown exoplanet and reconstruct surface features using a spherical harmonic-based reflection model developed as an extension of the **starry** package. We recover a time-variable albedo map including persistent regions indicative of the Earth’s major oceans and cloud banks. Our results suggest that the recovery of surface features on a terrestrial exoplanet from reflected-light photometric monitoring will be possible, but will require a considerable investment of observing time to achieve the high cadence and long time baseline needed. 

Keywords: methods: data analysis, techniques: photometric, planets and satellites: surfaces, Earth

1. INTRODUCTION

Of the many challenges and opportunities facing the field of exoplanets in the future, the prospect of mapping the surface of a potentially habitable planet is one of the most exciting – and one of the most difficult. Present-day missions including the Transiting Exoplanet Survey Satellite (*TESS*; [Ricker et al. 2015](#)) aim to discover small terrestrial planets around bright, nearby stars. Once these targets are found,

more detailed observational characterization must follow to evaluate their suitability for harboring life. The discovery of surface features such as continents and oceans would be a critical step toward understanding the habitability of an alien world.

The most straightforward method of observing features on a terrestrial exoplanet would be to directly image the planet at sufficient resolution. However, such an observation would require an effective telescope aperture many orders of magnitude larger and equipped with far more advanced coronagraphic optics than what we have available today. Fortunately, surface mapping can also be done indirectly by observing rotational phase-dependent variations of the planetary signal in precise photometric time series. This is a complex data analysis problem: strong degeneracies are involved when transforming a 1-dimensional time series into a 2-dimensional surface map (Cowan et al. 2013). Interpretation is also challenging in the face of many potential explanations for a poorly resolved surface feature. Nevertheless, this technique has already proven successful in the mapping of large dipolar features on the surfaces of hot giant planets, most notably the hot Jupiter HD189733b (Knutson et al. 2007; Majeau et al. 2012; de Wit et al. 2012). More recently, this technique has also been applied to the potentially terrestrial planet 55 Cancri e, suggesting a large longitudinal offset in its peak surface emission due to either lava flows on the surface or strong atmospheric recirculation (Demory et al. 2016a,b; Hammond & Pierrehumbert 2017).

Obtaining a map of a temperate terrestrial exoplanet is still beyond the capabilities of current facilities, as both the thermal and reflected light signals from these planets are orders of magnitude below current measurement capabilities. Nevertheless, much work has been done on the theoretical front on mapping terrestrial planets in the habitable zone (e.g., Ford et al. 2001; Tinetti et al. 2006a,b; Williams & Gaidos 2008; Pallé et al. 2008; Kawahara & Fujii 2010; Fujii & Kawahara 2012; Berdyugina & Kuhn 2017; Haggard & Cowan 2018; Farr et al. 2018; Lustig-Yaeger et al. 2018); for a review, see Cowan & Fujii (2018). In particular, several works have harnessed space-based observations of the Earth in proof-of-concept studies aimed at shedding light on the opportunities and challenges future exoplanet mapping missions may bring. Perhaps the most well-known of these is the work of Sagan et al. (1993), who used observations from the Galileo probe to infer the existence of life on Earth. This study made use of both spectroscopic observations and spatially-resolved images of the planet at high resolution, neither of which will necessarily be available for an analogous exoplanet. For a more exoplanet-like approach, Goode et al. (2001) and Woolf et al. (2002) used unresolved measurements of Earthshine reflecting off the face of the Moon to infer the planet’s seasonally-variable albedo and its atmospheric transmission spectrum, respectively. Similarly, Hasinoff et al. (2011) presented a method to infer the two-dimensional map of the Earth from single-pixel measurements of the planet’s diffuse reflectance on the Moon. Cowan et al. (2009) used unresolved images of the Earth taken by the EPOXI spacecraft to construct low resolution albedo maps, recovering approximate longitudinal distributions of oceans continents on the planet. More

recently, [Jiang et al. \(2018\)](#) made use of daily images of the fully-illuminated disk of the Earth taken by the Deep Space Climate Observatory (DSCOVR) to recover the planet’s rotation rate and seasonal albedo variations.


In all of these studies, the resolution and fidelity of the inferred maps of the Earth have been strongly limited by the cadence and/or baseline of observations. Inferring high spatial resolution features necessitates observations at high temporal resolution, while long baselines spanning different seasons on the planet are required to break the degeneracies inherent to the mapping problem.

In this work we use scattered light present in the background of detectors aboard the *TESS* spacecraft to infer a map of the surface of the Earth. Although *TESS* was designed as an all-sky photometric exoplanet survey, significant levels of scattered light from the Earth (“Earthshine”) are present across the *TESS* detector during several of the mission’s observing sectors. Despite the presence of a sunshade and a lens hood to limit this kind of contamination, Earthshine is a dominant background signal particularly during sectors 1 and 2, when the Earth spent a significant fraction of the time above the edge of the *TESS* sunshade. Although the Earth is never directly in the field of view of the *TESS* cameras, Earthshine reflects off the inside of the lens hoods, illuminating each of the four cameras with a signal that is both spatially and temporally variable due to the complex optics at play. For more information, refer to the *TESS* Instrument Handbook¹.

Although the Earthshine signal in *TESS* is convolved with complicated systematics, this dataset offers an unprecedented opportunity to tackle the mapping problem in detail, given its extremely high cadence (observations recorded every 2 minutes in short cadence mode) and the long observing baseline of *TESS* (a 2 year primary mission, with the Earth in view for several months of the year). Moreover, recent advances on the modeling side have made the reconstruction of detailed maps more feasible than ever. [Luger et al. \(2019\)](#) developed an analytic algorithm for generating—and inverting—light curves of stars and planets whose surface emission is described in terms of spherical harmonics, dubbed **starry**. The analyticity of this method makes the computation of the model both fast and precise. In turn, [Haggard & Cowan \(2018\)](#) derived analytic expressions for the analogous case in reflected light, packaged into the **EARL** code. In this work, we derive similar expressions within the **starry** framework to perform fast, analytical inversion of the light curve of the Earth and thus construct a map of the planet in reflected light.

The paper is organized as follows: we give an overview of the *TESS* data used in §2, and in §3 we describe the methods employed to infer a map from these data, including the adaptation of the **starry** algorithm to reflection mapping, the modeling of spacecraft-related systematics, and the likelihood and priors used. We present results in §4. In §5, we assess the accuracy of our map and discuss the relevance of this

¹ https://archive.stsci.edu/missions/tess/doc/TESS_Instrument_Handbook_v0.1.pdf

analysis to future attempts at mapping terrestrial exoplanets. Finally, we conclude with a summary of our major findings in §6. The code used to generate all the figures in this paper is open source and available on GitHub². Clickable icons  next to each figure caption link to the source code used to generate them.

2. DATA

The *TESS* pipeline produces processed light curves for all short-cadence (2 minute) targets. Part of this processing is the determination of the localized background flux within each postage stamp, which is provided as an ancillary time series. As with all *TESS* time series data products, the time is given in units of *TESS* Julian Date (TJD), defined as $\text{BJD} - 2457000$, and the data are delivered in 27-day parcels corresponding to the varying pointing sectors of *TESS*. We use these background flux measurements from Sectors 1 and 2 as our primary data source in this work.

For each of the two sectors considered, we downloaded a random subsample of 1,000 short-cadence light curve files corresponding to postage stamps located across the full *TESS* field of view and extracted the background flux time series (`SAP_BKG`) and corresponding uncertainties (`SAP_BKG_ERR`) from each. We then interpolated all light curves in a given sector onto the same time grid, masked all cadences with nonzero `QUALITY` flags, and masked cadences during which Earth was below the edge of the sunshade.

Not all postage stamps are equally informative, since the desired signal from the Earth is inhomogeneously spread across the *TESS* detectors and manifests itself differently in different regions (see §3.2). We down-selected our pool of 1,000 targets to 75 maximally useful background light curves from Sector 1 and 20 from Sector 2. This down-selection was done with several criteria in mind: each light curve should contain a significant signal from Earth; they should be representative of the general background flux modulation seen across the *TESS* cameras, rather than containing any truly localized signals from astrophysical background sources; and the sample as a whole should contain the full range of typical behaviors of background across the entire *TESS* field of view. To accomplish this, we applied a categorization and outlier rejection scheme to the ensemble of light curves. During this selection step only, we normalized each light curve \mathbf{f}_n to have flux in the range 0–1:

$$\mathbf{f}'_n = \frac{\mathbf{f}_n - f_{n,0.05}}{f_{n,0.95} - f_{n,0.05}}, \quad (1)$$

where $f_{n,0.05}$ and $f_{n,0.95}$ represent the 5th and 95th percentile values of the flux in the n^{th} light curve, respectively. We then computed the median value of \mathbf{f}'_n at each cadence across all light curves, along with an estimate of the standard deviation based

² <https://github.com/rodluger/earthshine>

on the median absolute deviation across all light curves:

$$\begin{aligned}\mu &= \text{median}(\{\mathbf{f}'_n\}) \\ \sigma &= 1.4826 \times \text{median}(\{|\mathbf{f}'_n - \mu|\})\end{aligned}\tag{2}$$

For each light curve, we then calculated a score indicative of its similarity to the median light curve:

$$\chi_n^2 = \frac{1}{K} \sum_{k=0}^K \frac{(f'_{n,k} - \mu_k)^2}{\sigma^2},\tag{3}$$

where K is the number of cadences in the light curve. We removed all light curves with $\chi_n^2 < N$ from the pool, placing them in a separate group. We repeated these steps several times, classifying light curves into four groups for Sector 1 and two groups for Sector 2. We found that this procedure effectively grouped together light curves with similar features, increasing in variation within each group from the first group (most homogeneous) to the last group (most heterogeneous). In practice, we found light curves in all but the last group to be dominated by the signal of the Earth, while light curves in the last group were dominated by signals other than the periodic modulation of the planet. To obtain our final set of light curves, we sorted the light curves in each group by the amplitude of the signal and kept the first 25 in each of the first three groups in Sector 1 and the first 20 in the first group in Sector 2 for a total of 75 and 20 light curves, respectively. We then median-normalized all light curves (as opposed to the normalization used for the classification discussed above). Finally, we computed an estimate of the non-Earth background flux from the cadences during which the planet was below the edge of the sunshade and removed this baseline from the flux. Our final dataset consisted of 912,605 data points in total. A representative sample of these signals is shown in Figure 1.

3. METHODS

3.1. *Orbital Modeling*

We downloaded³ the Solar System and *TESS* ephemeris data for Sectors 1 and 2 and used `spiceypy` (Acton 1996; Acton et al. 2017; Annex 2017) to convert the TJD timestamps to ephemeris time (ET). We then used the `spekzr` utility of `spiceypy` to compute the relative positions of the Sun, the Earth, and *TESS* for every cadence in our dataset in the J2000 equatorial reference frame. In this right-handed coordinate frame, the Earth is centered at the origin, the y axis points along the planet’s spin axis, and the x axis points toward the vernal equinox. We do not apply any light travel correction, as the expected delay is on the order of a few seconds.

At any given cadence, we rotate the Earth about its spin axis to the correct phase, assuming a planetary rotation period of 0.9972696 days (the Earth’s sidereal rotation

³ <https://archive.stsci.edu/missions/tess/models/>

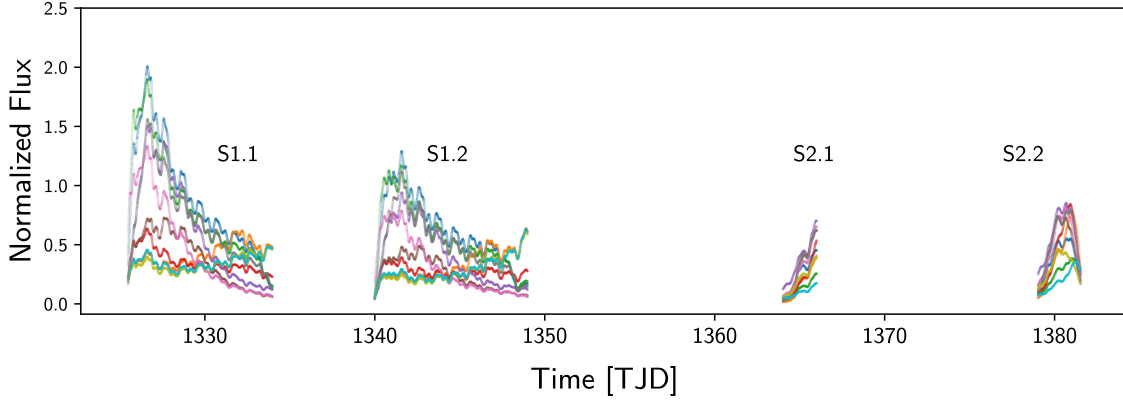


Figure 1. Normalized background flux in each of ten postage stamps for each of Sectors 1 and 2. Cadences where the Earth was below or very close to the edge of the sunshade were masked. While the planet is in view for the majority of both orbits in Sector 1, it is only visible in the data at the very end of each orbit in Sector 2. [☞](#)

period) and an obliquity of 23.437° . We determine the initial rotation phase of the Earth by computing the Earth Rotation Angle (ERA), given by (Urban & Seidelmann 2013)

$$\text{ERA} = 360^\circ(0.7790572732640 + 1.00273781191135448 t_U) \bmod 360^\circ \quad (4)$$

where $t_U = t - 2451545.0$. At $t = t_0 = 2458325.5$, the first timestamp in our dataset, we find $\text{ERA} = 303.4^\circ$, meaning the vernal equinox is aligned with the prime meridian $360^\circ - 303.4^\circ = 56.6^\circ = 3.77$ hours past t_0 . Finally, we rotate the frame to align *TESS* with the $+z$ axis, with the spin vector of the Earth along the $y-z$ plane. We verified our rotations by comparing our results on several days to the ephemerides provided by the JPL HORIZONS Web interface⁴.

Figure 2 shows the view *TESS* has of the Earth on each of the dates in our dataset; these images were produced by applying the sequences of rotations described above. The planet is illuminated by the Sun assuming it is a point source, and the night side is shaded black. Note that the map of the Earth shown in this figure is purely illustrative and does not correspond to the map inferred in this study. Note also that although the images are computed an integer number of rotation periods apart, the planet appears to slowly rotate over the course of the observation. This is due to the orbit of *TESS*, which changes its vantage point relative to the Earth over the course of its 13 day orbit.

3.2. Systematics Modeling

While the [signal of Earth](#) that we wish to analyze is present in all extracted light curves, its strength is modulated by time-variable systematics that are correlated but not identical across all postage stamps, since the illumination of the CCD by the

⁴ See <https://ssd.jpl.nasa.gov/horizons.cgi> and this [Jupyter notebook](#).

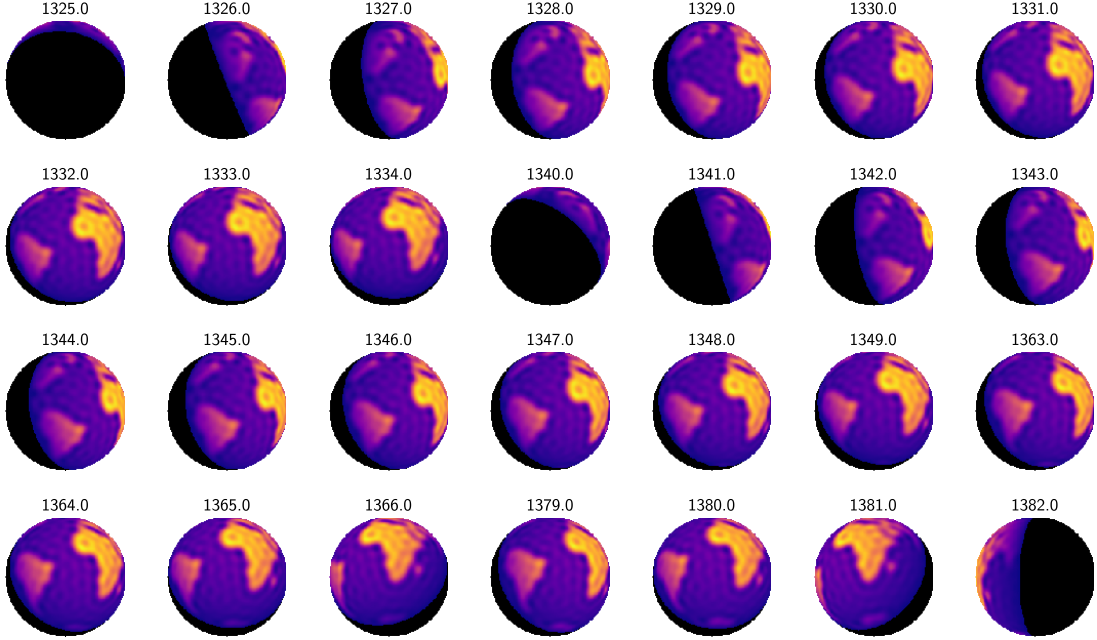


Figure 2. Synthetic images showing *TESS*’s view of the Earth on every date in Sectors 1 and 2 that was included in the regression. Each image is labeled with the corresponding timestamp in TJD. The dayside features shown here correspond to a simple continental model of the Earth, not to the map inferred in this study. [☞](#)

Earth is strongly inhomogeneous. A physical model of the reflections and scattering that gives rise to the signal of the Earth on the *TESS* detector is beyond the scope of this work. Instead, we treat this as a de-trending problem, where the signal we wish to fit is shared by all postage stamps, but multiplied by a systematic signal that is different for each target. This problem is precisely what the technique of pixel-level decorrelation (PLD; Deming et al. 2015; Luger et al. 2016, 2018) seeks to address. In PLD, one computes the sum over all N light curves at each point in time, and uses this quantity to normalize each of the individual light curves. Since (by assumption) each light curve is the product of the astrophysical signal and the systematics signal, this procedure divides out the astrophysical signal, producing a basis of N vectors that contain *only* functions of the systematics. This now serves as a “clean” basis one can then use to fit the systematics component of the data, with minimal risk of fitting out astrophysical signals.

We use PLD to determine a systematics basis set for each *TESS* sector, and from this basis we construct a separate design matrix \mathbf{B} for each sector. To increase the flexibility of the systematics model, we append to \mathbf{B} a basis of fifth order orthogonal polynomials in time for each of the two orbits in each sector. The systematics model ψ for the n^{th} signal is thus

$$\psi_n = \mathbf{B}\mathbf{w}_n \quad (5)$$

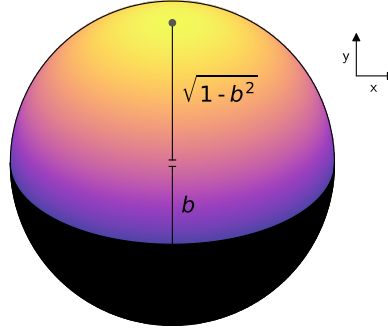


Figure 3. Geometry of the *starry* model in reflected light. In this frame, the y axis points up, the x axis points to the right, and the z axis points out of the page. The planet has unit radius and sits at the origin, while the illumination source is along the $y-z$ plane, with the sub-stellar point marked by a dot. The semi-major axis of the terminator is unity, and the semi-minor axis of the terminator is denoted b . The night side is colored black. [☞](#)

where \mathbf{w}_n are the weights of each regressor. Since we solve for different weights for each of the 75 light curves in Sector 1 and 20 light curves in Sector 2, our systematics model has $75 \times (75 + 2 \times 5) + 20 \times (20 + 2 \times 5) = 6,975$ free parameters. The total number of data points is 912,605, so overfitting is not particularly concerning; nevertheless, we impose an L2 prior on the weights to minimize this risk (see §3.4).

3.3. Reflected Light Mapping with *starry*

We model the astrophysical component of the signal—the rotational modulation of the Earth in reflected light—using the *starry* code package (Luger et al. 2019). *starry* computes phase curves and occultation light curves for bodies whose surfaces can be expressed as a sum of spherical harmonics. The algorithm works by first transforming from spherical harmonic coefficients to coefficients in the polynomial basis \tilde{p} , whose terms are of the form $x^i y^j z^k$ for (non-negative) integer i and j and $k = 0$ or 1. Luger et al. (2019) showed how to compute the visible flux by integrating each of the terms in \tilde{p} across the visible surface of the projected disk. For unocculted spheres observed in emitted light, the area of integration is the full disk, which makes the integration trivial: the flux contribution from each term in \tilde{p} is a constant easily computed via recursion relations involving factorials. The algorithm is thus extremely fast and very precise.

However, in the present work, we are interested in *reflected* light, so we must account for the illumination by the Sun. This illumination varies smoothly across the dayside of the planet, but its derivative discontinuously changes at the terminator, beyond which the illumination is zero everywhere. This requires two modifications to the *starry* algorithm: we must (1) weight all integrands by the illumination function on the dayside, and (2) modify the integration limits to truncate the integrals at the day/night terminator.

Let us consider a right-handed coordinate frame in which the planet has unit radius and is located at the origin, the illumination source is along the $y - z$ plane, and the sub-stellar point is on the $+y$ axis at $x = 0$ (see Figure 3). In this frame, the terminator [projected onto the \$x - y\$ plane](#) is a segment of an ellipse, with semi-major axis $a = 1$ aligned with the x axis and semi-minor axis b [aligned with the \$y\$ axis](#). It is straightforward to show that the sub-stellar point is located at $y = \sqrt{1 - b^2}$ and the illumination is given by

$$I(b; x, y) = \begin{cases} \sqrt{1 - b^2}y + bz & y \geq -b\sqrt{1 - x^2} \\ 0 & \text{otherwise} \end{cases}, \quad (6)$$

where $z = +\sqrt{1 - x^2 - y^2}$. Since the illumination function is just a polynomial in x , y , and z , weighting the terms in \tilde{p} by this function keeps all terms within the polynomial basis. Provided we always rotate the problem such that the sub-stellar point lies along the $+y$ axis as above, the limits of integration are $-1 \leq x \leq 1$ and $b\sqrt{1 - x^2} \leq y \leq \sqrt{1 - x^2}$. Computing the total reflected flux for any surface map is therefore a matter of solving integrals of the form

$$\begin{aligned} J_{ijk}(b) &= \int_{-1}^1 \int_{b\sqrt{1-x^2}}^{\sqrt{1-x^2}} x^i y^j z^k dy dx \\ &= \frac{\Gamma\left(\frac{i+1}{2}\right) \Gamma\left(\frac{j+1}{2}\right)}{2\Gamma\left(\frac{i+j+k+4}{2}\right)} \begin{cases} 0 & i \text{ odd} \\ 1 - b^{j+1} & k = 0 \\ \frac{\sqrt{\pi}}{2} - b^{j+1} \Gamma\left(\frac{j+4}{2}\right) {}_2\tilde{F}_1\left(-\frac{1}{2}, \frac{j+1}{2}; \frac{j+3}{2}; b^2\right) & k = 1 \end{cases}, \end{aligned} \quad (7)$$

where ${}_2\tilde{F}_1$ is the regularized hypergeometric function. For all values of i , j , and k in \tilde{p} , ${}_2\tilde{F}_1$ reduces to simple trigonometric functions involving b . Moreover, it is straightforward to derive recursion relations for $J_{ijk}(b)$, making its evaluation extremely fast. We implemented the algorithm for computing $J_{ijk}(b)$ in the development version of [starry](#)⁵; the details will be discussed [further](#) in upcoming work.

As with the systematics model (§3.2), the model for the astrophysical signal is *linear*. The signal \mathbf{s} is a linear operation on the vector of spherical harmonic coefficients \mathbf{y} that describes the body's surface:

$$\mathbf{s} = \mathbf{A}\mathbf{y} \quad (8)$$

where \mathbf{A} is the design matrix and \mathbf{y} is the vector of spherical harmonic coefficients. Note that in [Luger et al. \(2019\)](#), \mathbf{y} describes the emissivity of the surface, but here we will use it to represent the *albedo* of the planet.

⁵ <https://github.com/rodluger/starry/releases/tag/earthshine>

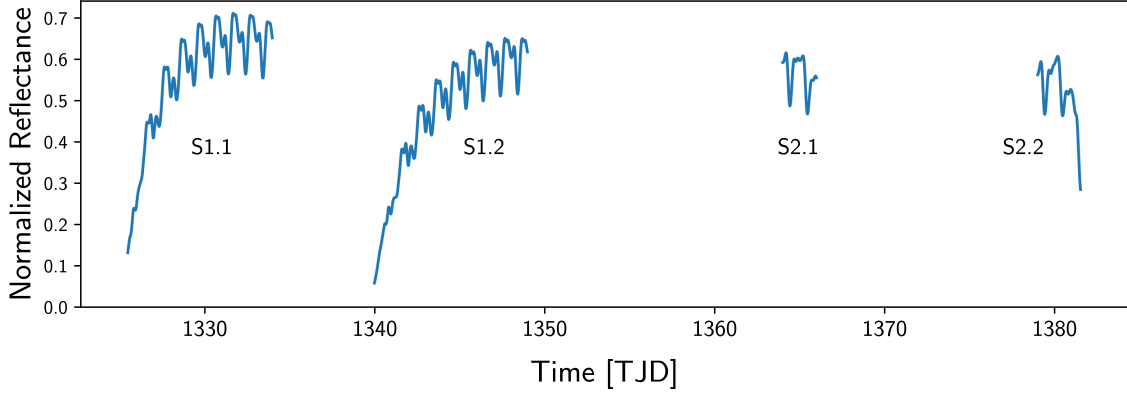


Figure 4. Normalized surface reflectance versus time for the maximum likelihood **starry** model in each of the four orbits of *TESS* in Sectors 1 and 2. [☐](#)

We chose a maximum spherical harmonic degree $l_{max} = 10$ for our fits. Increasing l_{max} above this value leads to a sharp rise in the amount of “ringing” without substantially increasing the quality of the fit. Note that $l = 10$ corresponds to a maximum surface resolution on the order of $180^\circ/10 \approx 18^\circ$.

Finally, for reference, in Figure 4 we plot the **starry** model for our maximum likelihood fit (§4). The sharp rise at the beginning of each orbit in Sector 1 is due to the changing vantage point of *TESS* as the Earth transitions from a thin crescent to full phase.

3.4. Full Model and Likelihood Function

Our full model for each light curve is simply the product of the systematics model $\psi_n = \mathbf{B}\mathbf{w}_n$ (different for each target) and the **starry** model $\mathbf{s} = \mathbf{A}\mathbf{y}$ (shared by all targets). Since we know the exact distance \mathbf{r} between *TESS* and the Earth at all times (§3.1), we also include the inverse square of \mathbf{r} as a multiplicative term to scale the luminosity of the Earth to the flux observed by *TESS*. The model for the flux time series in the n^{th} postage stamp is therefore

$$\mathbf{m}_n = (\mathbf{r}^{-2}) \circ (\mathbf{B}\mathbf{w}_n) \circ (\mathbf{A}\mathbf{y}) \quad (9)$$

where \circ denotes the element-wise product of two vectors.

Equation (9) is linear in \mathbf{w}_n and in \mathbf{y} individually, but not in both simultaneously. To find the true optimum for both vectors, we must therefore perform a nonlinear optimization. However, we find that the dimensionality of the problem is so large that unless we initialize the optimizer with a good initial guess, it is unable to converge on a good solution. Fortunately, we are able to obtain an initial guess by exploiting the linearity of the problem and alternating between linear optimization of \mathbf{w}_n and linear optimization of \mathbf{y} . Then, in a second step, we run a nonlinear optimizer, applying additional nonlinear constraints to prevent overfitting and ensure an albedo in the range $[0, 1]$ everywhere on the surface.

In the first step, we initialize the **starry** model to unity and iteratively solve for \mathbf{w}_n and \mathbf{y} by solving the L2-regularized least squares problem (see, e.g., §2.1 in [Luger et al. 2018](#)). We assume zero-mean Gaussian priors for the weights:

$$\mathbf{w}_n \sim \mathcal{N}(0, \sigma_w^2) \quad \mathbf{y} \sim \mathcal{N}(0, \sigma_y(l)^2), \quad (10)$$

with $\sigma_w = 0.05$ and

$$\sigma_y(l) = \begin{cases} 1.75 \times 10^{-5} l^{\frac{3}{2}} & l < 2 \\ 1.40 \times 10^{-4} l^{-\frac{3}{2}} & l \geq 2 \end{cases}, \quad (11)$$

where l is the spherical harmonic degree. The latter prior was chosen by trial-and-error and was adjusted to keep the albedo positive everywhere on the surface. Note, importantly, that we do not fit for the coefficient of the constant $Y_{0,0}$ spherical harmonic, but instead fix it to unity. Since the solutions obtained in this step are merely used as a starting point for the second step (see below), the choice of prior at this stage does not significantly impact on our results.

In general, we find that the iterative scheme converges within about 50 iterations, taking about two minutes on a laptop computer. However, since it is strongly regularized, the model significantly underfits the data. In the second step, we remove our L2 prior on the spherical harmonic coefficients and instead regularize the actual surface albedo, which we compute on a uniform spherical grid with 50,000 points. We enforce a uniform prior in the range $[0, 1]$, with a small amount of Gaussian smoothing on either side. [Note that while the lower bound of this range corresponds to a true physical constraint—non-negative surface albedo—the upper bound of unity is somewhat *ad hoc*, since the true normalization of the flux is unknown. We find that varying this upper bound \(or removing it entirely\) does not significantly change our results, since we only report “normalized” albedos; we discuss this further below.](#)

We additionally impose a constraint on the power spectrum of the inferred map, requiring that it be drawn from a distribution whose mean is a decaying power law in the spherical harmonic degree l in order to suppress ringing and spurious high order features. [In practice, however, we find that this constraint is negligible compared to our prior on the albedo.](#) Our full negative log likelihood function in this step is therefore

$$\begin{aligned} -\log \mathcal{L} = & \frac{1}{2}(\mathbf{f} - \mathbf{m})^\top \Sigma^{-1}(\mathbf{f} - \mathbf{m}) + \\ & \frac{1}{2}\mathbf{w}^\top \Lambda_w^{-1}\mathbf{w} + \\ & \frac{1}{2}(\mathbf{a}_- + 1 - \mathbf{a}_+)^\top \Lambda_a^{-1}(\mathbf{a}_- + 1 - \mathbf{a}_+) + \\ & \frac{1}{2}(\boldsymbol{\rho} - \mathbf{l}^{-\gamma})^\top \Lambda_\rho^{-1}(\boldsymbol{\rho} - \mathbf{l}^{-\gamma}) + \\ & \frac{1}{2} \frac{(\gamma - \gamma_0)^2}{\sigma_\gamma^2}, \end{aligned} \quad (12)$$

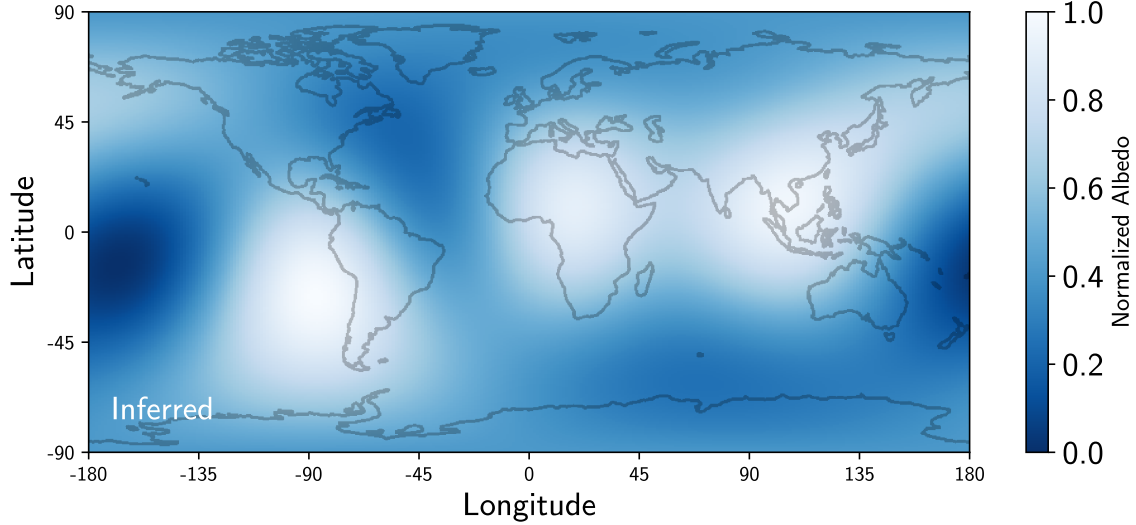


Figure 5. Maximum likelihood map of the Earth obtained during the first stage of our optimization procedure, in which an iterative analytic approach was used to solve the bilinear model, subject to very conservative regularization. Despite the low resolution of this map, bright features appear at the expected locations of South America, Africa, and Asia, and two dark features appear at locations consistent with the Pacific and Atlantic oceans. The colors correspond to a *normalized* albedo, scaled such that the dynamic range of the image spans $[0, 1]$. [🔗](#)

where \mathbf{f} is the measured flux across all targets, \mathbf{m} is the full model (Equation 9), Σ is the (diagonal) data covariance, \mathbf{w} are the systematics model weights for all targets, Λ_w is the prior covariance of the weights (Equation 10), \mathbf{a}_- and \mathbf{a}_+ are vectors containing the values of the albedo that are below zero and above one, respectively, Λ_a is the covariance of the Gaussian used to smooth the edges of the top-hat prior on \mathbf{a} , ρ is the power spectrum of the inferred spherical harmonic map (equal to the sum of the squares of the coefficients at each degree l), γ is the power law index of the power spectrum prior, Λ_ρ is the covariance of that prior, $\gamma_0 = 1.5$ is the mean of the prior on γ and $\sigma_\gamma^2 = 10^{-2}$ is the corresponding variance. For simplicity, all our prior covariances are diagonal and homoscedastic with variances equal to 2.5×10^{-3} for \mathbf{w} , 10^{-8} for \mathbf{a} , and 10^{-1} for ρ .

Unfortunately, the constraints outlined above break the linearity of the problem, and we must turn to a non-linear optimizer. We use the `AdamOptimizer` method in the `TensorFlow` package (Abadi et al. 2015) to maximize Equation (12). Starting with the values of the weights obtained in the first step, we run the `AdamOptimizer` for 1000 iterations, well past convergence. Since `TensorFlow` automatically computes and propagates gradient information, this step is fast, finishing in under one minute on a laptop.

4. RESULTS

Figure 5 shows the maximum likelihood map inferred during the first stage of our optimization procedure. Since the map coefficients were heavily regularized, the

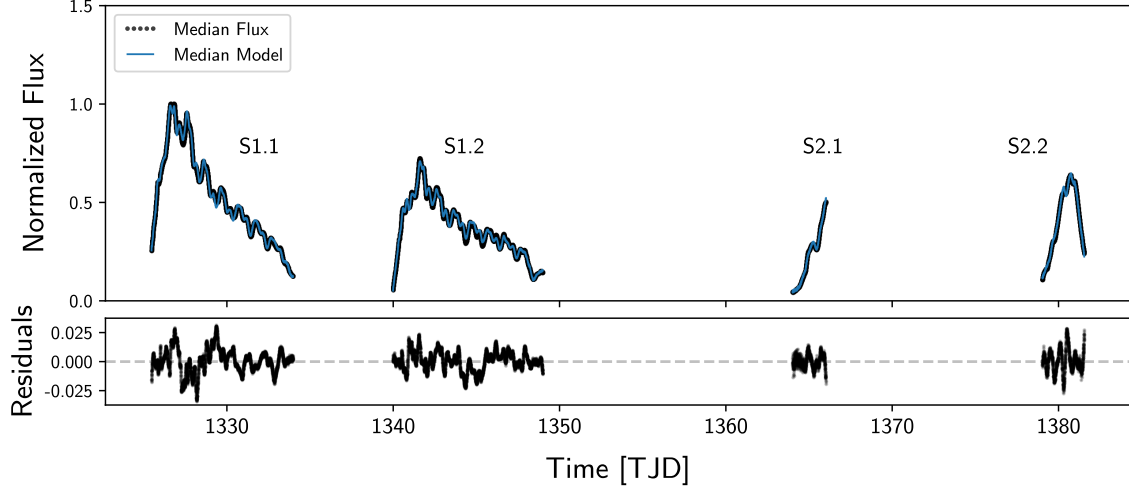


Figure 6. *Top:* Median flux across all postage stamps (black) and the median maximum likelihood model (blue) across Sectors 1 and 2. *Bottom:* Median residuals about the maximum likelihood fit. [🔗](#)

dynamic range of the albedo variations was less than 10%; we scaled it to span the range $[0, 1]$ for better visualization. Furthermore, although the maximum resolution of the spherical harmonic model is on the order of 18° , features typically span $\sim 45^\circ$; this is also an artefact of our regularization. However, broad features are already visible in the map: three bright regions and two dark regions (one of which wraps around the antimeridian). Overplotted on the inferred map are coastal outlines of the Earth’s continents. One can see that the three bright regions roughly track, from left to right, the continents of South America, Africa, and Asia. The dark regions appear at the locations of the Pacific ocean (centered at longitude $\pm 180^\circ$) and the Atlantic ocean (centered at longitude -45°).

As expected, the second, non-linear step in our optimization procedure yields a much better fit to the data. The top panel of Figure 6 shows the median of the maximum likelihood model in this second step across all 95 targets (blue) overplotted on the median flux (black). The median residuals are shown in the bottom panel; these have standard deviation of less than one percent. While there is significant correlated structure in the residuals, their power spectrum displays a broad peak extending between about 0.1 and 10 days, with a significant dip at 1.00 day. This suggests our model is fitting the rotational variability of the Earth quite well, but additional temporal variations—most likely due to cloud movement—cause aperiodic changes in the reflectivity of the surface. We revisit this point in §5.

In the top panel of Figure 7 we show the final, maximum likelihood albedo map inferred for Earth weighted by a visibility function. The visibility function is computed as the average of the product of the cosine of the illumination angle and the cosine of the angle between the observer vector and the vector normal to the surface of the Earth. This normalization has the effect of downweighting regions that are

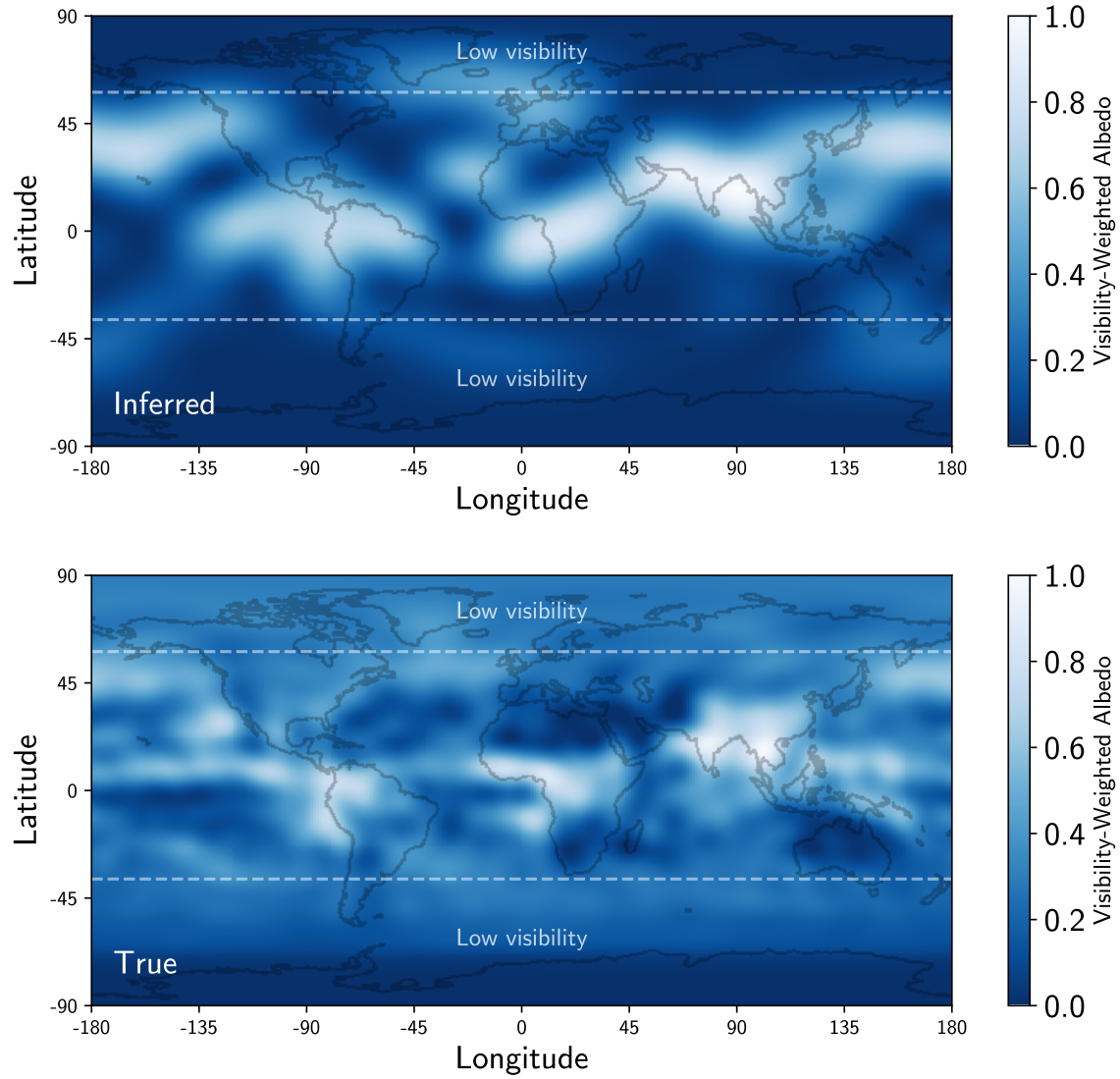


Figure 7. *Top:* The maximum likelihood model for the surface albedo of the Earth, weighted by the visibility across Sectors 1 and 2 and normalized. The dashed white lines indicate latitudes above/below which the visibility drops below 50%. As in Figure 5, coastlines are overplotted as black contours. *Bottom:* Approximate cloud coverage map based on corrected surface reflectance obtained by the *VIIRS* imager, averaged over the same date range and weighted by the same visibility function. [\[5\]](#)

either seldom illuminated (such as regions near the southern pole of the Earth, as the observations were taken during southern winter) or seldom in view (such as regions near the northern pole of the planet, which during Sectors 1 and 2 were mostly on the opposite side of the Earth from *TESS*). Because there is little data from these regions, their albedo is prior-dominated and tends to fluctuate substantially with small changes to the choices we make when fitting the data. On the other hand, we find that the large scale features close to the equator are mostly insensitive to our choices of prior. In the Figure, we use dashed lines to delineate the regions whose average visibility is larger than 0.5. Note that as before, the albedo is scaled to span the range $[0, 1]$. We emphasize that because of the unknown throughput of the optics that give rise to the Earthshine signal in *TESS*, the absolute scaling of the albedo cannot be inferred from the dataset alone.

The bottom panel of Figure 7 shows an approximate average cloud cover map of the Earth based on imagery taken by the Visible Infrared Imaging Radiometer Suite (*VIIRS*) instrument aboard the Suomi National Polar-orbiting Partnership (*S-NPP*) satellite⁶. The image was produced by analyzing the corrected true color reflectance in the I1 (red), M4 (blue), and M3 (green) bands and masking all pixels that were either dark or had significant variance among the three bands (i.e., inconsistent with a grey spectrum). We then averaged the images over the 28 days of data analyzed in this work, smoothed it with a Gaussian filter, and weighted it by the *TESS* visibility as in the top panel.

While the two panels display fundamentally different quantities—a scaled albedo and a cloudiness factor—we expect that the signal of the Earth in the *TESS* bandpass is dominated by cloud reflectivity, as water clouds have an albedo higher than soil, vegetation, or ocean in the optical/near infrared (e.g., Jedlovec 2009). We therefore expect large, coherent cloud structures to show up as the brightest regions in our inferred map, and in fact we find that this is generally the case. The dominant feature in our inferred map is a bright region spanning longitudes of $+50^\circ$ to $+100^\circ$ just northward of the equator. This is also the dominant feature in the *VIIRS* map, corresponding to the persistent weather system associated with summertime monsoons in the south of Asia.

Secondary features common to both maps are a permanent cloudy region over central Africa with a westward spur over the Atlantic Ocean, a veneer of clouds over the north Pacific, and a pile up of clouds on the western coast of northern South America. The dark regions in the inferred map rather loosely track the Pacific, Atlantic, and Indian Oceans, although the correspondence is not exact.

5. DISCUSSION

5.1. *An Ill-Posed Problem*

⁶ Data obtained via [this wget script](#).

The problem of inferring a two-dimensional map from a one-dimensional time series is famously ill-posed given its large null space and many complex degeneracies (e.g., Cowan et al. 2013). In fact, for a uniformly illuminated body, its rotational light curve encodes at most $2l_{\max}$ modes (one sine term and one cosine term per frequency), where l_{\max} is the highest spherical harmonic degree of features on the surface of the body. The number of modes *on the surface of the body*, however, increases quadratically as $(l_{\max} + 1)^2$. This means that for a body whose surface is perfectly described by a spherical harmonic map of degree $l_{\max} = 10$, which is described by 121 coefficients, only 20 linear combinations of those coefficients are actually projected onto the light curve. The problem is fundamentally non-invertible, since most of the information one wishes to extract never makes it into the light curve in the first place.

Fortunately, there are two aspects of the problem at hand that facilitate this (seemingly impossible) inversion. First, the surface of the Earth is not uniformly illuminated; instead, different parts of the surface are weighted by different amounts as the planet rotates and as the angle of incident starlight moves over the course of the planet’s year. This weighting is sufficient to break many of the degeneracies, which are rooted in the anti-symmetry of many of the spherical harmonic modes. Moreover, the discontinuous day/night terminator further breaks these degeneracies, as only modes that are anti-symmetric about the visible lune of the planet strictly remain in the null space. Second, the changing vantage point of *TESS* relative to Earth (and our exact knowledge of it) breaks degeneracies that are viewpoint-dependent. In particular, the motion of *TESS* changes the position and shape of the terminator on the projected disk of the Earth, enhancing the effect discussed above.

In fact, because of these points, there should actually be *no* null space in this particular problem. In practice, however, it is evident from Figure 7 that degeneracies still abound. While we infer a few of the dominant features of the cloud map of the Earth, several of them are skewed, shifted, or entirely out of place. This is the case, for instance, with the cloud bank west of South America, which extends too far westward, and the northern Atlantic Ocean, which appears to extend into the east coast of North America. We identify four broad reasons for this.

First, many of the issues arise because of the finite signal-to-noise of our measurements and various biases in our adopted model. While the measurement errors on our data points are small—since the signal of the Earth is both quite strong and independently measured at 95 different positions on the detector—our lack of understanding of the complex optics (and our attempt to naïvely model it with pixel-level decorrelation) likely introduces significant correlated noise when we de-trend. We discuss this and other sources of bias in more detail in §5.3 below.

Second, while we attempted to adopt agnostic priors on the map coefficients—i.e., we fit for a spectral slope rather than impose one—our inferred map is still heavily prior-dependent. Our choice of regularization strength for the systematics was based on trial-and-error and what “looked good,” as was the choice for the hyperprior on

the spectral slope and the maximum spherical harmonic degree of the map. Although the main features of the map (the monsoon over south Asia, the clouds over Africa and South America, and dark regions associated with the three major oceans) persist regardless of the choice of prior, their exact shapes and contrast vary significantly under different assumptions.

Third, inspection of Figure 2 shows that most of the time, the Earth is close to full phase as seen by *TESS*. With the terminator close to the limb of the planet, its ability to break degeneracies is considerably lessened. We are therefore somewhat limited in our ability to infer in particular the latitudinal position of features. The Earthshine present in upcoming sectors may contain longer segments in which the planet is seen in crescent, which should help mitigate these degeneracies.

Fourth, and most importantly, we have not performed Bayesian inference: that is, we simply found the maximum likelihood solution (subject to our constraints), with no estimate of the uncertainty anywhere on our map. A much better approach than simply maximizing the likelihood function is to marginalize over the nuisance parameters—such as the systematics model weights and the spectral slope. This is more costly, but certainly not intractable, and we will pursue this in future work.

5.2. *Confirmation Bias*

It is difficult to avoid being disingenuous in one’s solution when the answer (or the expected answer) to a problem is known ahead of time. Since we had access to “ground truth” from the beginning, many of the de-trending/processing/fitting choices we made were (however inadvertently) driven by the desire to come as close as possible to the correct answer. While many of these choices correspond to numbers we have explicitly quoted here, such as the prior variances on our model terms, several of them are implicit decisions on *how* we performed the analysis. For instance, in an early version of the analysis we used as regressors actual *TESS* housekeeping variables—such as the angle and elevation of the Earth and the Moon relative to each of the four camera boresights—instead of the PLD basis. However, the systematics model constructed from these quantities proved to be too inflexible, judging in part from the poor fidelity of the recovered maps. The results in this work must therefore be taken with a grain of salt. That said, we have attempted to offset the possibility of confirmation bias by making our model (Equation 9) as simple as possible in the end (a multiplicative baseline and a linear reflectance model) and thus limiting the number of knobs we could turn. We also opted to forego any use of real imaging data in our model; for instance, we fit for the spectral slope of the map as a free parameter rather than using the true power spectrum of the Earth. Moreover, the broad features in the strongly regularized map (Figure 5)—and in particular the monsoon clouds over southeast Asia—are insensitive to most of our model choices.

5.3. *Other Kinds of Bias*

One of the largest sources of bias is the fact that we assumed our systematics model is strictly multiplicative. Although we subtracted an estimate of the baseline non-Earth background flux from each target (measured from the mid-sector portions of the light curve where the planet was below the edge of the sunshade), we implicitly assumed that this constant baseline was the *only* source of additive noise. We know this to be strictly untrue, as reflected light from the Moon should also cause periodic modulations in the background flux of *TESS* pixels. Because of the Moon’s small size and very low (~ 0.1) albedo, we chose to ignore its effect, but it is likely to contaminate our signal at certain points in the orbit where *TESS* is closer to the Moon than to Earth. While this will likely bias the **starry** model at those times, the effect is particularly concerning because of its impact on the PLD basis. While PLD can be an extremely powerful de-trending technique, it is notorious for grossly overfitting signals in the presence of uncorrected background modulations (Luger et al. 2016). This happens because PLD strictly assumes the noise is multiplicative, which is not the case for signals with varying (additive) background levels. In principle, one could model the reflectance variability of the Moon in the same way as we did here, but since the optics are not well understood, this may be difficult in practice.

In our analysis, we also neglected the effects of limb darkening. Because of atmospheric attenuation, features near the limb of the planet should in reality contribute less to the flux than our model predicts. This could bias our results particularly in cases where the planet is seen at half phase (for instance, during days 1326 and 1341 in Figure 2). In the absence of limb darkening, the brightest regions on the planet are those close to the sub-stellar point, which is on the limb. Limb darkening should shift the regions of peak brightness closer to the center of the planet, adding an east-west bias in the location of the features we identify. Similarly, Rayleigh scattering and anisotropic scattering by clouds probably cause similar biases, as could specular reflection off the oceans (glint).

We also ignored the finite angular size of the Sun at Earth (about 0.5°), treating it instead as a point source. Regions immediately nightward of the terminator should therefore contribute a small amount of flux, which we neglect. Similarly, we used the timestamps reported in the *TESS* light curve FITS files, which are corrected to the barycenter of the Solar System. Since our signal originates from a source that is closer to *TESS* than the barycenter of the system, the timestamps may be wrong by up to 8 minutes, the light travel time from *TESS* to the Sun. Over the course of 8 minutes, Earth rotates by about 2° , so there should be a blurring of our inferred map at this scale. However, these two effects are well below the resolution of our map, so we expect them not to introduce significant bias.

Moreover, we implicitly assumed an isotropic prior on the spherical harmonic coefficients by imposing a power spectrum in only the degree l . This has the effect of lessening the extent to which features can be confined to certain latitudes. In particular, our prior implicitly disfavors banded cloud structure, which is quite common on

Earth. The bottom panel of Figure 7 shows such banded structure at the equator and at mid-latitudes due to the combined effects of north-south circulation and deflection by the Coriolis force.

Finally, and perhaps more importantly, our model assumed a *static* map for the surface of the Earth. We also know this to be incorrect, since clouds form and move on timescales of hours to days. We discuss this in detail in the following section.

5.4. Temporal Evolution

As we mentioned in §4, the residuals about our maximum likelihood fit show correlated structure with power at a large range of periods ranging from 0.1 up to 10 days. This is likely due to the fact that rotational variability is not the sole source of the signal of the Earth: additionally, there is temporal variation in the albedo of certain parts of the map due to cloud movement. In order to model this, we used **starry** to construct a time-variable spherical harmonic map, defined by a vector of spherical harmonic coefficients that varies slowly over time:

$$\mathbf{y}(t) = \sum_{n=0}^N \mathbf{y}_n c_n(t) \quad (13)$$

where the \mathbf{y}_n are the N spherical harmonic coefficient vectors that describe the temporal variability and the c_n are the components of the basis in time. Since this is a linear operation on the spherical harmonic coefficients, the **starry** model is still linear, and may be solved in the same way as before.

We choose the c_n vectors to be an orthogonal polynomial basis in time of degree 6 spanning all of Sector 1. Under this basis, one can interpret \mathbf{y}_0 as the static component of the map, \mathbf{y}_1 as the component that varies (quasi-)linearly in time, \mathbf{y}_2 as the component that varies (quasi-)quadratically, and so forth. In all, we fit for six of these map components.

We neglect the data taken in Sector 2, focusing instead on the portion of the data for which we have good temporal coverage. In order to mitigate overfitting due to the larger number of variables we are solving for, we add a term to our likelihood function that forces the coefficients of the $n > 0$ maps to remain small. We further evaluate the albedo of the map on a grid at six equally spaced points in time, enforcing our albedo and power spectrum priors separately at each of these points.

The mean of the temporal solution is shown in Figure 8, and a link to an animated version is provided in the caption. No significant differences are evident, although the residuals improved by $\sim 30\%$ and contain significantly less power on timescales longer than 1 day. We conclude that while this is evidence for cloud variability on these timescales, a more detailed investigation is warranted to say anything meaningful about the dynamics of clouds on Earth from *TESS* observations.

5.5. Implication for Future Observations

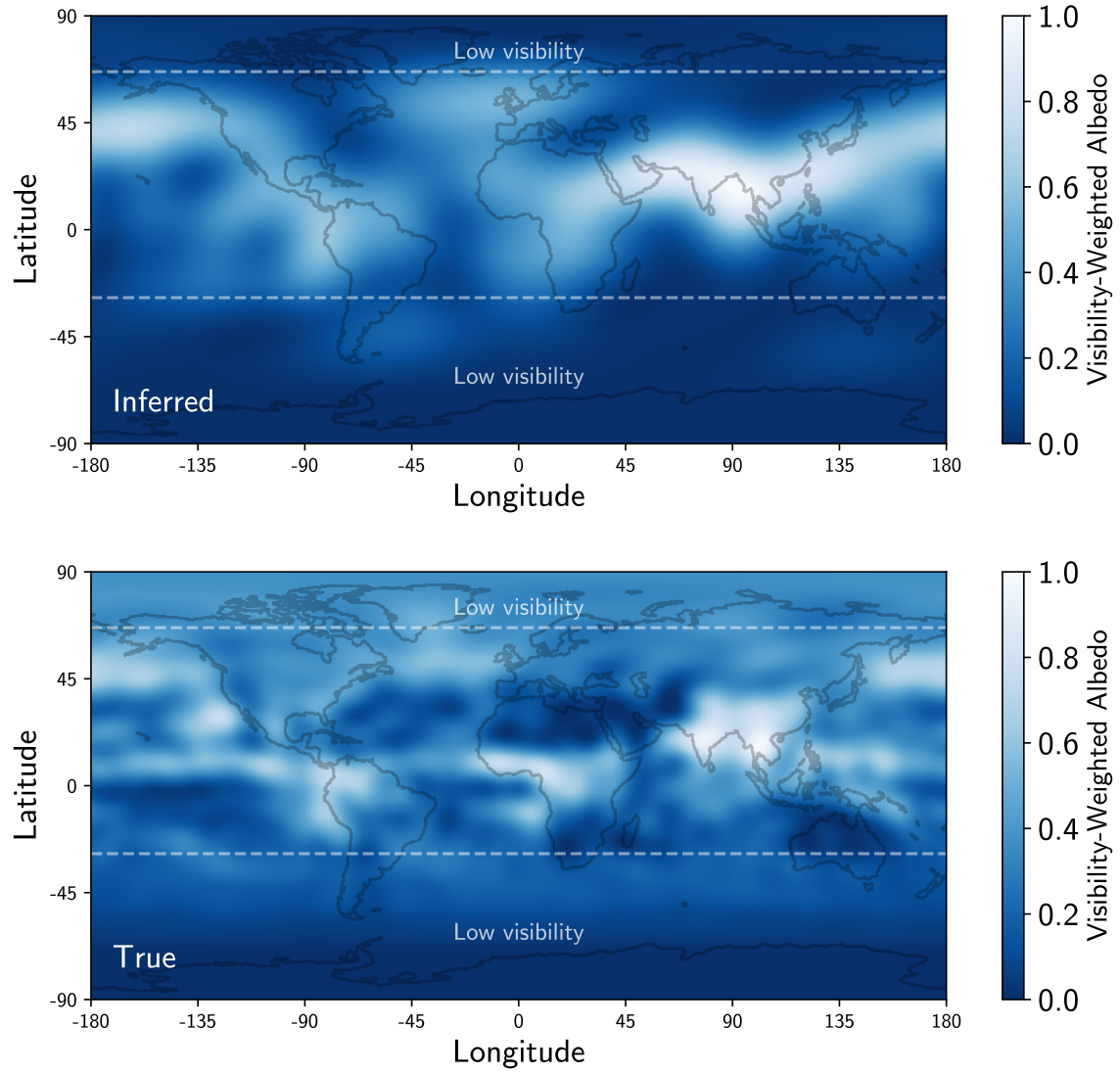



Figure 8. The inferred mean map when allowing for slight temporal variations in the spherical harmonic coefficients. Compare to Figure 7. [A 4 second animated version of the top panel is available in the HTML version of this article, showing the inferred cloud movement during sectors 1 and 2.](#) 

As we discussed above, light curve inversion is a hard problem, and we have shown that this is true even at high signal-to-noise and when the solution is known ahead of time. Presented with the top panel of Figure 7 (without the continent outlines), one may be hard-pressed to identify that as a map of our home planet. This is in part due to the biases in our optimization, which we discussed above, but also due to the simple fact that the reflectance of the Earth is dominated—by far—by time-variable clouds. While stationary cloud features on Earth are typically associated with continents (the most pronounced cloud features being over South America, Africa, and south Asia), persistent cloud cover can occur over open ocean, as seen in the north Pacific in Figure 7. Given a baseline of just a few weeks, it is not possible to tell the difference between highly reflective land features (such as ice or vegetation) and stationary clouds.

We should not expect other exoplanets in the habitable zone to be any different. Future missions such as LUVOIR or HabEx may enable us to collect data on the phase curves of such planets, but that data will be of significantly lower quality and poorer time resolution, and we will not know the planet’s rotation period or obliquity ahead of time. The work presented here can therefore be seen as an upper limit on what we may be able to infer about exoplanets in the habitable zone with next-generation facilities.

That said, the time baseline considered in this work was fairly limited—three weeks in Sector 1 and one week in Sector 2, spanning two months in total. Observations taken throughout an entire year on Earth—during which time cloud patterns should change with the seasons and the angle between the day/night terminator and the axis of rotation should precess—may help to disambiguate between clouds, oceans, and continents. Although the signal of the Earth mostly disappears from the *TESS* detector after Sector 2, it is expected to return in Sectors 14 and 15 when the planet will again be above the edge of the sunshade. We intend to perform this analysis in upcoming work once data from future *TESS* sectors becomes available.

Additionally, time-resolved observations in multiple bands could also help with this problem. This was discussed in Cowan et al. (2009), Kawahara & Fujii (2011), and Fujii & Kawahara (2012), all of which showed that multi-band photometry of the Earth enables one to disentangle the contributions from oceans, continents, and cloud cover. A robust detection of surface features on an exoplanet in the habitable zone will likely necessitate both approaches alongside a circumspect analysis of the uncertainties and degeneracies at play.

6. CONCLUSIONS

We have used *TESS* photometry to produce a global map of the Earth, treating it as an exoplanet with unknown surface features. We adapted the *starry* code to generate (and invert) light curves in reflected light, coupling it to a variant of pixel level decorrelation to model the systematics in the *TESS* detector.

We detect three large scale bright regions near the equator of the planet at longitudes of -60° , $+30^\circ$, and $+90^\circ$, which we interpret as persistent cloud structures near the Andes, sub-Saharan Africa, and the South Asian monsoon. Pronounced dark features are identified at longitudes of $\pm 180^\circ$, -30° , and $+60^\circ$ with significant latitudinal extent, likely corresponding to the Pacific, Atlantic, and Indian Oceans. Because of the vantage point of *TESS* during Sectors 1 and 2, we are not sensitive to high latitudes on the planet, although this may change with upcoming sectors.

We fit Earth’s light curve with both static and time-variable albedo maps. We find tentative evidence for variability in Earth’s cloud coverage over the course of Sector 1 observations. Upcoming observations with *TESS* will yield further information on the time variability of surface features on Earth on a variety of timescales.

The work presented here serves as a relevant case study for future efforts to characterize terrestrial exoplanets. While the circumstances of *TESS*’s observations of the Earth will not be replicated exactly for other planets, similar rotationally-modulated phase curves of reflected light may be feasible in the future. We show that a fast, analytic linear model can be successfully applied to such data. We also underscore the importance of high cadence *and* long duration for these observations, and we advocate for multi-wavelength coverage to help disentangle the photometric signatures of continents and persistent cloud features.

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Facility: TESS

Software: [Astropy](#) (Astropy Collaboration et al. 2013, 2018), [Matplotlib](#) (Hunter 2007), [Numpy](#) (van der Walt et al. 2011), [starry](#) (Luger et al. 2019), TensorFlow (Abadi et al. 2015), [spiceypy](#) (Annex et al. 2019)

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