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Supplementary Materials

Supplementary Table 1. Starting basalt composition as compared to hydrous starting materials in previous basalt-limestone interaction studies

	This Study	F08	IM07	M10	J13	IM08	IM08	D10
Starting Material	Primary arc basalt*	AH-7a plagioclase-free phonotephrite	AH7a potassic basalt	Primitive K-basalt	Primitive shoshonite	PST9 high-K calc-alkaline basalt	ST18 primitive shoshonite	M-94-a-107 basaltic-andesite
Source		Mt Mellone, Alban Hills, Italy	Mt Mellone, Alban Hills, Italy	Roman Province synthetic, Italy	Mt. Vesuvius, Italy	Stromboli, Italy	Stromboli, Italy	Mt Merapi, Indonesia
SiO₂	47.88	47.96	48.05	48.3	49.85(0.21)	50.38	52.07	51.83(0.43)
TiO₂	0.75	0.94	0.96	0.7	1.03(0.08)	0.86	0.93	0.89(0.05)
Al₂O₃	17.95	14.72	14.60	11.4	15.95(0.67)	15.36	16.71	18.08(0.24)
FeO_T	9.73	7.77	7.85	7.5	7.98(0.25)	7.91	7.61	8.17(0.16)
MnO	0.20	0.14	-	-	0.14(0.2)	-	-	-
MgO	5.99	6.38	6.45	13.7	6.03(0.33)	7.82	5.74	2.97(0.08)
CaO	10.97	11.32	11.40	14.0	9.98(0.29)	12.83	9.77	9.19(0.16)
Na₂O	2.00	1.67	1.63	1.2	2.35(0.07)	2.26	2.72	3.48(0.09)
K₂O	0.50	6.97	7.30	3.3	4.01(0.19)	1.91	3.65	2.05(0.03)
P₂O₅	-	0.50	0.51	-	0.7(0.09)	0.66	0.79	0.34(0.04)
Sum	95.97	98.97	98.75	100.1	98.02	97.48	97.83	97.2
NaO+K₂O	2.49	8.64	8.93	4.5	6.36	4.17	6.37	5.53
CaO/Al₂O₃	0.6	0.8	0.8	1.2	0.6	0.8	0.6	0.6
H₂O	4.03	0.7 to 1.7	1.25	1 and 5	0 to 2	1 to 6	1 to 6	2.23
Proportion of Carbonate	50	1 to 7	0 to 18	5 to 10	16.7 to 21.9	0 to 19	0 to 19	17 to 20

All values are in weight percent; for previous studies, number in (parentheses) are standard deviation where reported

FeO_T is total iron content (FeO+Fe₂O₃)

Primary arc basalt* composition determined based on weighed proportion of reagents (see Methods)

F08=Freda et al. (2008); IM07=Iacono-Marziano et al. (2007); M10=Mollo et al. (2010); J13=Jolis et al. (2013); IM08=Iacono Marziano et al. (2008); D10=Deegan et al. (2010)

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Supplementary Table 2. Major element composition of experimental clinopyroxene (in wt.%)

Run no.	B252*	B291*	B294*	B292	B306	B293	B263	B274	B297	B264	B283	B266	B267	B295	B298	B280
P (GPa)	1.0	1.0	1.0	1.0	1.0	1.0	1.0	0.8	0.8	0.8	0.8	0.8	0.5	0.5	0.5	0.5
T (°C)	1100	1150	1175	1100	1150	1175	1200	1100	1125	1150	1175	1200	1100	1125	1150	1175
<i>n</i>	7	7	9	13	7	9	7	8	8	6	5	7	2	4	9	5
SiO ₂	47.5(8)	49.2(8)	48.1(8)	46(1)	49(2)	41.3(7)	41.6(5)	45.4(1)	46(1)	40.8(1)	41(2)	41(2)	37.7(3)	41(1)	39.5(1)	38.3(2)
TiO ₂	0.61(9)	0.42(8)	0.40(4)	0.7(2)	0.8(2)	0.61(1)	0.7(2)	0.9(2)	0.74(8)	0.64(3)	0.7(2)	0.7(1)	0.63(9)	0.8(1)	0.71(6)	0.7(4)
Al ₂ O ₃	7.9(9)	6.7(9)	6.9(8)	8(1)	7(2)	15.1(8)	18.0(9)	8.3(9)	10(1)	15(1)	15(2)	18(2)	15.9(2)	15(1)	15(1)	20(1)
FeO _T	12.0(7)	10.6(3)	8.7(3)	11.4(4)	11.5(6)	8.1(2)	8.0(2)	12.4(7)	10.3(3)	8.5(6)	9.5(9)	7.4(6)	11.0(9)	9.0(4)	9.6(1)	4.3(1)
MnO	0.40(5)	0.34(3)	0.27(5)	0.35(6)	0.37(7)	0.10(1)	0.14(5)	0.22(4)	0.22(6)	0.12(4)	0.10(4)	0.08(4)	0.10(2)	0.14(5)	0.16(4)	0.04(3)
MgO	13.2(5)	14.1(4)	14.4(5)	11.0(5)	12(2)	8.1(3)	7.7(4)	8.3(6)	10.7(1)	8.5(5)	8(1)	8.0(9)	7.9(2)	7.6(6)	7.3(4)	8.3(3)
CaO	17.1(8)	17.1(4)	18.8(3)	19(1)	17(2)	23.9(2)	24.2(2)	23.6(4)	21(1)	23.9(3)	23.3(6)	24.5(2)	25.2(2)	25.1(4)	26(1)	24.8(3)
Na ₂ O	0.64(4)	0.54(4)	0.59(5)	0.72(6)	0.7(2)	0.44(8)	0.32(4)	0.25(6)	0.6(2)	0.30(7)	0.19(6)	0.22(5)	0.06(3)	0.22(9)	0.2(1)	0.07(5)
K ₂ O	0.012(9)	0.002(4)	0.02(1)	0.01(1)	0.1(2)	0.012(9)	0.004(4)	0.006(9)	0.04(7)	0.005(5)	0.011(8)	0.003(4)	#NUM!	0.02(2)	0.03(2)	0.006(8)
Sum	99.2(5)	99.0(6)	98.3(6)	98(2)	99(1)	97.7(7)	100.6(8)	99.4(5)	99.2(7)	97.8(5)	97.2(8)	99.4(3)	98.5(7)	99.0(6)	97.8(9)	96.4(7)
Fs+En	65	66	63	58	63	38	35	48	53	39	41	33	45	39	42	24
Wo	24	25	24	30	29	36	38	39	31	34	35	38	25	35	30	36
CaTs	11	9	13	12	8	26	26	13	16	27	23	28	30	26	28	40
Total	100	100	100	100	100	100	99	100	100	100	100	99	100	100	100	100

*Experiments with starting hydrous basalt composition only; all others performed with 50 wt.% basalt and 50 wt.% calcite

Number in (parentheses) is one sigma standard deviation for *n* number of EMPA spot analyses averaged for each phase reported in least digits cited, i.e., 47.5(8) should be read as 47.5 ± 0.8 wt.%

FeO_T is total iron content (FeO+Fe₂O₃)

End-members (En=enstatite, Fs=ferrosilite, Wo=wollastonite, CaTs=Ca-Tschermak) in % by atomic proportion

Supplementary Table 3. Major element composition of experimental plagioclase (in wt.%)

Run no.	B252*	B291*	B294*	B292	B306	B274	B297	B283	B267	B295	B298
P (GPa)	1.0	1.0	1.0	1.0	1.0	0.8	0.8	0.8	0.5	0.5	0.5
T (°C)	1100	1150	1175	1100	1150	1100	1125	1175	1100	1125	1150
<i>n</i>	4	4	8	3	4	6	10	8	8	5	5
SiO ₂	51.8(1)	50(1)	50.0(9)	49.9(5)	53(1)	51(1)	51(1)	50.9(5)	48(2)	49(1)	52.7(9)
TiO ₂	-	0.2(1)	0.05(5)	0.3(1)	0.2(1)	0.15(9)	0.09(6)	0.06(4)	0.5(2)	0.23(5)	0.09(2)
Al ₂ O ₃	30(1)	28(1)	30.0(5)	27(2)	29.9(9)	28.6(9)	30.5(9)	31.1(3)	25(3)	30.5(9)	28.3(5)
FeO _T	2.1(4)	3(2)	0.9(2)	2.4(7)	1.2(6)	1.3(4)	1.3(4)	0.9(1)	3(1)	1.5(2)	0.9(1)
MnO	-	0.07(5)	0.02(2)	0.10(5)	0.02(2)	0.06(6)	0.04(3)	0.02(2)	0.14(7)	0.09(3)	0.02(1)
MgO	0.4(2)	1.0(7)	0.2(2)	0.7(2)	0.2(2)	0.3(1)	0.17(8)	0.10(5)	0.9(4)	0.32(5)	0.2(2)
CaO	12.6(2)	13.2(5)	14.2(4)	12.3(4)	13(1)	12.9(9)	14.1(9)	13.9(5)	18.0(8)	16.3(7)	12(1)
Na ₂ O	3.5(3)	3.1(5)	3.4(2)	3.4(2)	3.3(3)	3.6(3)	3.4(6)	3.2(2)	3.1(5)	2.8(4)	4.6(4)
K ₂ O	0.26(5)	0.21(3)	0.13(2)	0.4(1)	0.3(2)	0.6(2)	0.25(8)	0.22(3)	0.9(3)	0.32(8)	0.42(4)
Sum	100(2)	99(1)	98.8(7)	96(1)	101.1(7)	100(1)	99.4(3)	100(2)	100(8)	101(4)	99(3)
An	49	53	53	48	51	47	53	53	57	60	41
Ab	49	45	46	48	46	48	45	45	36	37	56
Or	2	2	1	4	3	5	2	2	7	3	3
Total	100	100	100	100	100	100	100	100	100	100	100

*Experiments with starting hydrous basalt composition only; all others performed with 50 wt.% basalt and 50 wt.% calcite
Number in (parentheses) is one sigma standard deviation for *n* number of EMPA spot analyses averaged for each phase
reported in least digits cited, i.e., 51.8(1) should be read as 51.8 ± 0.1 wt.%

FeO_T is total iron content (FeO+Fe₂O₃)

End-members (Ab=albite, An=anorthite, Or=orthoclase) in % by atomic proportion of Na²⁺, Ca²⁺, and K⁺, respectively.

Supplementary Table 4. Major element composition of experimental scapolite (in wt.%)

Run no.	B306	B293	B263	B297	B264	B283	B266	B295	B298
P (GPa)	1.0	1.0	1.0	0.8	0.8	0.8	0.8	0.5	0.5
T (°C)	1150	1175	1200	1125	1150	1175	1200	1125	1150
<i>n</i>	4	7	10	5	7	4	4	7	5
SiO ₂	43(3)	40.6(5)	37(1)	37.4(9)	37.4(9)	41(1)	39.8(3)	40.4(5)	39(1)
TiO ₂	0.05(2)	0.02(2)	0.03(2)	0.03(2)	0.03(2)	0.04(3)	0.03(3)	0.02(1)	0.03(3)
Al ₂ O ₃	30(3)	29.3(3)	29.4(9)	26.1(8)	26.1(8)	29.0(8)	30.7(2)	30.8(5)	31.3(8)
FeO _T	0.8(1)	0.6(1)	0.51(5)	1.0(2)	1.0(2)	0.59(9)	0.62(5)	0.49(3)	0.7(1)
MnO	0.03(1)	0.03(3)	0.02(3)	0.02(2)	0.02(2)	0.01(2)	0.02(2)	0.02(2)	0.01(1)
MgO	0.2(1)	0.14(1)	0.17(2)	0.18(8)	0.18(8)	0.15(2)	0.16(3)	0.17(2)	0.13(5)
CaO	21(1)	21.9(4)	21.8(4)	22.8(4)	22.8(4)	21.1(5)	22.4(1)	22.2(3)	22.8(9)
Na ₂ O	1.1(7)	1.0(2)	0.86(5)	0.8(2)	0.8(2)	1.5(3)	0.47(4)	0.73(9)	0.4(1)
K ₂ O	0.12(5)	0.07(1)	0.06(2)	0.06(4)	0.06(4)	0.11(1)	0.04(2)	0.05(2)	0.02(2)
Est. CO ₂ ^a	4(1)	6.4(1)	10(2)	12(2)	12(2)	6(2)	5.8(6)	5(1)	5.6(9)
Sum	98.4(5)	100.3(6)	100.4(4)	100.8(8)	100.8(8)	100(5)	100(1)	100(2)	100(4)
Eq An	95	93	93	94	94	89	96	94	97
% Me (S60)	95	92	93	94	94	89	96	94	97
% Me (E69)	83	84	92	80	80	81	91	89	95

Number in (parentheses) is one sigma standard deviation for *n* number of EMPA spot analyses averaged for each phase reported in least digits cited, i.e., 43(3) should be read as 43 ± 3 wt.%

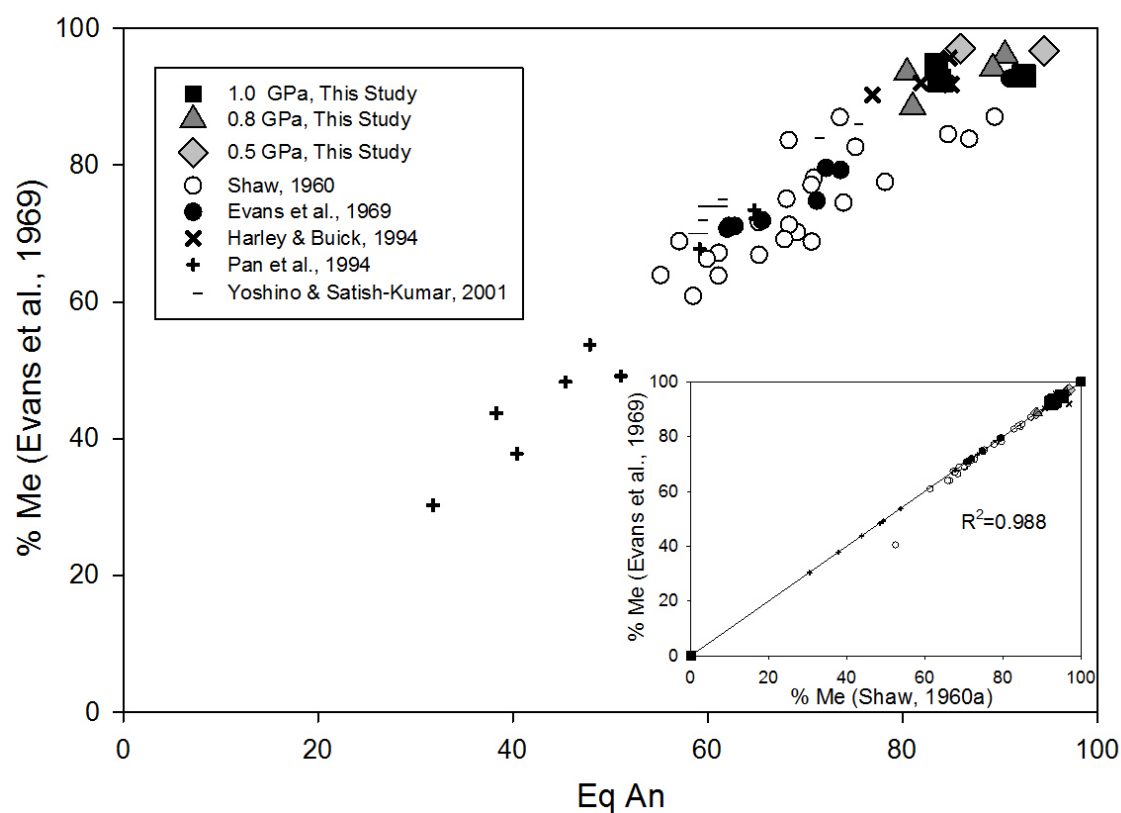
FeO_T is total iron content (FeO+Fe₂O₃)

Est. CO₂^a is estimated stoichiometrically, normalized to (Al+Si)=12 cations

Eq An=Equivalent Anorthite content calculated after the method of Moecher and Essene (1990)

% Me (E69)=percent Meionite following the method of Evans et al. (1969); % Me (S60)=percent Meionite following the method of Shaw (1960) (see Supplementary Fig. 1).

Supplementary Fig. 1. Experimental scapolites analyzed in this study (following the symbols used in Fig. 5) plotted by Equivalent Anorthite content (Eq An; $[Al-3]/3 \times 100$) after Moecher and Essene (1990) and by Meiorite content (% Me) following the equation of Evans *et al.* (1969; $Ca/[Ca+Na+K] \times 100$). The insert shows the close 1:1 correspondence between this method and that of Shaw (1960; $[Ca+Mg+Fe+Mn+Ti]/[Na+K+Ca+Mg+Fe+Mn+Ti] \times 100$). Natural data are from compilations of scapolite analyses (open circles – Shaw, 1960: Helsinki, USA (Massachusetts, Pennsylvania, New York, New Hampshire), Canada (Ontario and Quebec), the former USSR, Italy (Vesuvius), Finland, Sweden, Norway, Germany (Laacher See), Austria, Moravia; x – Harley and Buick, 1992: Antarctica; closed circles – Evans *et al.*, 1969: Tanzania, USA (Massachusetts, Pennsylvania, New York), and Italy (Somma-Vesuvius); crosses – Pan *et al.*, 1994: Canada; dashes – Yoshino and Satish-Kumar, 2001: the Himalayas).



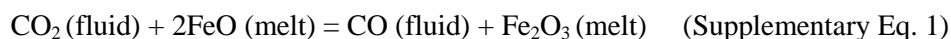
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Additional reference information for Fig. 7

Data for natural ultracalcic melt inclusions shown as the grey region were collected from: 1) Batan Island, Philippines (Schiano et al., 2000); 2) Iceland (Sigurdsson et al., 2000); 3) 43°N (Kamenetsky et al., 1998); 4) Bulgaria (Marchev et al., 2009); 5) Tonga arc (Falloo and Green, 1986); 6) Srednogorie arc, SW Europe (Georgiev et al., 2009). Previous carbonate assimilation experiments performed at 0.5 GPa (labeled ‘C09’ and ‘IM09’) from Conte et al. (2009) and Iacono Marziano et al. (2008) are plotted as open and closed circles, respectively. The ultracalcic melts of Mollo and Vona (2014) are similar but not plotted for comparison as they were performed at atmospheric pressures. Experimental lherzolite data (‘+’ symbols, labeled ‘G04’ and ‘S04’) are compiled from (Green et al., 2004; Schmidt et al., 2004) and wehrlite data (‘x’ symbols, labeled ‘M06’) from the experiments of (Medard, 2006).

Oxygen fugacity determination

It has been speculated that preferential incorporation of Al³⁺ for Si⁴⁺ in the tetrahedral site in cpx, seen with increasing assimilation in this study, is perhaps influenced by an enhanced incorporation of Fe³⁺ in the cpx, owing to oxidative capacity of CO₂, following the reaction such as:



(Conte *et al.*, 2009; Mollo and Vona, 2014b and references therein). Oxygen fugacity of our experiments were not buffered, but the estimated oxygen fugacity within the capsules, calculated using the equilibrium between iron content dissolved by the AuPd capsule and FeO content of coexisting silicate glass, was higher when calcite was present at 1.0 GPa, 1200 °C ($\Delta\text{FMQ} +1.8 \pm 0.2$ or 2.8 ± 0.1 with calcite; $\Delta\text{FMQ} +1.4 \pm 0.3$ or 2.4 ± 0.3 without calcite, following to the $f\text{O}_2$ parameterizations of Barr and Grove, 2010 and Balta *et al.*, 2011 respectively). Using the model of Barr and Grove (2010) who determined and included a temperature-dependence with their empirical equation, we calculate an increase in oxygen

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fugacity with increased calcite consumption at higher temperatures ($\Delta\text{FMQ} -0.30 \pm 0.1$ at 1100 °C; $\Delta\text{FMQ} + 0.41 \pm 0.1$ at 1200 °C, 0.5 GPa, Table 1). Additionally, Fe₂O₃/FeO_T in cpx grains was estimated using atomic site assignment calculations. The ratio of ferric to total iron rapidly increases with calcite presence (0.30 without calcite; 0.49 with calcite at 1.0 GPa, 1100 °C) and with increasing temperature and decreasing pressure in calcite-bearing experiments (as much as ~1.00 at 0.5 GPa, 1175 °C). Thus our experimental determination of $f\text{O}_2$ is broadly in agreement with the notion that increased CO₂ release by carbonate assimilation can also lead to oxidation of the silicate subsystem, thereby promoting formation of more CaTs-rich clinopyroxenes.

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