

Title: Carbon cycling in mature and regrowth forests globally

Authors:

Kristina J. Anderson-Teixeira^{1,2*}

Valentine Herrmann¹

Becky Banbury Morgan^{1,3}

Ben Bond-Lamberty⁴

Susan C. Cook-Patton⁵

Abigail E. Ferson^{1,6}

Helene C. Muller-Landau²

Maria M. H. Wang^{1,7}

Author Affiliations:

1. Conservation Ecology Center; Smithsonian Conservation Biology Institute; National Zoological Park, Front Royal, VA 22630, USA

2. Center for Tropical Forest Science-Forest Global Earth Observatory; Smithsonian Tropical Research Institute; Panama, Republic of Panama

3. School of Geography, University of Leeds, Leeds, UK

4. Joint Global Change Research Institute, Pacific Northwest National Laboratory, College Park Maryland 20740, USA

5. The Nature Conservancy; Arlington VA 22203, USA

6. College of Natural Resources, University of Idaho; Moscow, Idaho 83843, USA

7. Grantham Centre for Sustainable Futures and Department of Animal and Plant Sciences, University of Sheffield, Western Bank, Sheffield, South Yorkshire S10 2TN, UK

*corresponding author: teixeirak@si.edu; +1 540 635 6546

Summary

Background. Forests are major components of the global carbon (C) cycle and thereby strongly influence atmospheric carbon dioxide (CO₂) and climate. However, efforts to incorporate forests into climate models and CO₂ accounting frameworks have been constrained by a lack of accessible, global-scale synthesis on how C cycling varies across forest types and stand ages.

Methods/Design. Here, we draw from the Global Forest Carbon Database, ForC, to provide a macroscopic overview of C cycling in the world's forests, giving special attention to stand age-related variation. Specifically, we use 11923 *ForC* records for 34 C cycle variables from 865 geographic locations to characterize ensemble C budgets for four broad forest types – tropical broadleaf evergreen, temperate broadleaf, temperate conifer, and boreal. We calculate statistics for both mature and regrowth (age <100 years) forests, and quantify trends with stand age in regrowth forests for all variables with sufficient data.

Review Results/ Synthesis. The rate of C cycling generally decreased from tropical to boreal regions in both mature and regrowth forests, whereas C stocks showed less directional variation. Net ecosystem production of mature forests was indistinguishable across biomes. The majority of flux variables, together with most live biomass pools, increased significantly with stand age when fit with logarithmic functions.

Discussion. As climate change accelerates, understanding and managing the carbon dynamics of forests is critical to forecasting, mitigation, and adaptation. This comprehensive and synthetic global overview of C stocks and fluxes across biomes and stand ages will help to advance these efforts.

Key words: forest ecosystems; carbon cycle; stand age; productivity; respiration; biomass; global

Background

Forest ecosystems are shaping the course of climate change through their influence on atmospheric carbon dioxide (CO_2 ; Bonan 2008, Friedlingstein *et al* 2019, IPCC 2018). Despite the centrality of forest C cycling in regulating atmospheric CO_2 , important uncertainties in climate models (Friedlingstein *et al* 2006, Krause *et al* 2018, Bonan *et al* 2019, Di Vittorio *et al* 2020) and CO_2 accounting frameworks (Pan *et al* 2011, IPCC 2019) can be traced to lack of understanding on how C cycling varies across forest types and in relation to stand history. This requires accessible, comprehensive, and large-scale databases with global coverage, which runs contrary to the traditional way forest C stocks and fluxes have been measured and published. Large-scale synthesis is critical to benchmarking model performance with global data (Clark *et al* 2017, Luo *et al* 2012), quantifying the the role of forests in the global C cycle (*e.g.*, Pan *et al* 2011), and using book-keeping methods to quantify actual or potential exchanges of CO_2 between forests and the atmosphere (Griscom *et al* 2017, Houghton 2020).

Forests in the global C cycle: current and future

A robust understanding of forest impacts on global C cycling is essential. Total annual photosynthesis in forests (gross primary productivity, *GPP*) is estimated at approximately 69 Gt C yr^{-1} (Badgley *et al* 2019), more than seven times the average annual fossil fuel emissions during 2009-2018 ($9.5 \pm 0.5 \text{ Gt C yr}^{-1}$; Friedlingstein *et al* 2019). Most of this enormous C uptake is counterbalanced by releases to the atmosphere through ecosystem respiration (R_{eco}) and fire, with forests globally dominant as sources of both soil respiration (Warner *et al* 2019) and fire (van der Werf *et al* 2017). In recent years, total forest C uptake has exceeded releases, such that globally forests have been a C sink (Harris *et al* 2021). Considering only areas remaining in forest, this C sink has averaged $3.2 \pm 0.6 \text{ Gt C yr}^{-1}$ for 2009-2018, offsetting 29% of anthropogenic fossil fuel emissions (Friedlingstein *et al* 2019). However, deforestation, estimated at $\sim 1 \text{ Gt C yr}^{-1}$ in recent decades (Pan *et al* 2011, Tubiello *et al* 2020), reduces the net forest sink to $\sim 1.1\text{-}2.2 \text{ Gt C yr}^{-1}$ (Friedlingstein *et al* 2019, Harris *et al* 2021).

The future of the current forest C sink is dependent both upon forest responses to climate change itself and human land use decisions, which will feedback and strongly influence the course of climate change (Friedlingstein *et al* 2006). Regrowing forests in particular will play an important role (Pugh *et al* 2019), as almost two-thirds of the world’s forests were secondary as of 2010 (FAO 2010). As anthropogenic and climate-driven disturbances impact an growing proportion of Earth’s forests (Andela *et al* 2017, McDowell *et al* 2020), understanding the carbon dynamics of regrowth forests is increasingly important (Anderson-Teixeira *et al* 2013). Although age trends in aboveground biomass have been well-studied and synthesized globally (Cook-Patton *et al* 2020), a relative dearth of data and synthesis on other C stocks and fluxes in secondary forests points to an under-filled need to characterize age-related trends in forest C cycling. Such understanding is particularly critical for reducing uncertainty regarding the potential for carbon uptake and climate change mitigation by regrowth forests (Krause *et al* 2018, Cook-Patton *et al* 2020). Understanding, modeling, and managing forest-atmosphere CO_2 exchange is thus central to efforts to mitigate climate change (Grassi *et al* 2017, Griscom *et al* 2017, Cavaleri *et al* 2015).

Evolution of forest C cycle research

For more than half a century, researchers have sought to understand how forest carbon cycling varies across stands, including those of different biomes (*e.g.*, Lieth 1973, Luyssaert *et al* 2007) and stand ages (*e.g.*,

Odum 1969, Luyssaert *et al* 2008). Over this time, an increasingly refined conceptual understanding of the elements of ecosystem C cycles has developed, as a growing number of variables have been defined (e.g., Chapin *et al* 2006), along with appropriate measurement methods (e.g., Clark *et al* 2001). New technology has also enabled researchers to directly measure an expanding set of variables, notably including the development of continuous measurements of soil CO₂ efflux (Kuzakov 2006) and ecosystem-atmosphere CO₂ exchange (Baldocchi *et al* 2001). Measurement techniques have been increasingly standardized; for example, of the biomass allometries that strongly influence estimates of most C cycle variables (e.g., Chave *et al* 2014). Further standardization has been made possible through research networks such as ForestGEO (Anderson-Teixeira *et al* 2015, Davies *et al* 2021), NEON (Schimel *et al* 2007), and FLUXNET (Baldocchi *et al* 2001, Novick *et al* 2018). Remote sensing technology has become increasingly useful for global- or regional-scale estimates of a few critical variables (e.g., aboveground biomass, B_{ag} : Saatchi *et al* 2011, Hu *et al* 2016, Spawn *et al* 2020, gross primary productivity, GPP : Li and Xiao 2019), yet measurement and validation of most forest C stocks and fluxes necessarily requires intensive on-the-ground data collection.

Alongside these conceptual and methodological developments, there has been a proliferation of measurements across the world’s forests. The result of decades of research on forest C cycling is tens of thousands of records distributed across thousands of scientific articles, varying in data formats, units, measurement methods, *etc.* To address global-scale questions, researchers began synthesizing data into increasingly large databases (e.g., Lieth 1973, Luyssaert *et al* 2007, Bond-Lamberty and Thomson 2010, Anderson-Teixeira *et al* 2016, 2018, Cook-Patton *et al* 2020). The current largest, most comprehensive database on forest C cycling is *ForC* (Anderson-Teixeira *et al* 2016, 2018), which contains published estimates of forest ecosystem C stocks and annual fluxes (>50 variables), with different variables capturing distinct ecosystem pools (e.g., woody, foliage, and root biomass; dead wood) and flux types (e.g, gross and net primary productivity; soil, root, and ecosystem respiration). These data represent ground-based measurements, and *ForC* contains associated data required for interpretation (e.g., stand history, measurement methods). Since its most recent publication (*ForC v2.0-Ecology*; Anderson-Teixeira *et al* 2018), *ForC* has grown 129%, primarily through the incorporation of two additional large databases that also synthesized published forest C data: the Global Soil Respiration Database (SRDB; Bond-Lamberty and Thomson 2010, Jian *et al* 2020) and the Global Reforestation Opportunity Assessment database (GROA; Cook-Patton *et al* 2020). Following these additions, *ForC* currently contains 39762 records from 10608 plots and 1532 distinct geographic areas representing all forested biogeographic and climate zones, making it ideal for assessing how forest C cycling varies across biomes and with respect to stand age.

Biome differences

Forest C cycling varies enormously across biomes, which categorize the world’s forests according to major differences in climate, vegetation, *etc.* Since the early 19th century, it has been recognized that climate plays a dominant role in shaping differences among forests on a global scale (Humboldt and Bonpland 1807, Holdridge 1947). Global scale data syntheses have shown that C fluxes including GPP , net primary productivity (NPP), and soil respiration (R_{soil}) decrease with latitude or, correspondingly, increase with mean annual temperature (Fig. 1; e.g., Lieth 1973, Luyssaert *et al* 2007, Hursh *et al* 2017, Banbury Morgan *et al* n.d.). C stocks of mature forests show less directional variation (Fig. 1). On average, aboveground biomass (B_{ag}) tends to decrease with latitude, but not as dramatically as fluxes, and with the highest B_{ag} forests in relatively cool, moist temperate regions (Keith *et al* 2009, Smithwick *et al* 2002, Hu *et al* 2016). In

contrast, standing and downed dead wood ($DW_{standing}$ and DW_{down} , respectively, summing to DW_{tot}) and the organic layer (OL) tend to accumulate more in colder climates where decomposition is slow relative to NPP (Harmon *et al* 1986, Allen *et al* 2002).

Correlative analyses relating C cycle variables to climate and other environmental variables have recently been taken to a new level through use of machine-learning algorithms that relate ground-based C cycle data to global maps of environmental covariates, making it possible to create fine-scale global maps of C cycling (e.g., Warner *et al* 2019, Cook-Patton *et al* 2020). This approach can be particularly effective when paired with satellite measurements that correlate to C cycle variables of interest; for example, solar-induced chlorophyll fluorescence is useful for fine-scale mapping of gross primary productivity (GPP ; Li and Xiao 2019), while LiDAR, radar, and optical imagery are being used to estimate B_{ag} at regional to global scales (e.g., Saatchi *et al* 2011, Hu *et al* 2016). Any such analysis is however constrained by the quality and coverage of ground-based estimates of forest C fluxes or stocks (e.g., Schepaschenko *et al* 2019). While estimates of some variables (e.g., B_{ag} , GPP , NPP , R_{soil}) are widely available, many remain poorly characterized (e.g., DW_{tot} ; OL ; autotrophic respiration, R_{auto}) –even at the coarse resolution of biomes. This is a critical limitation not only for understanding forest C cycling, but also for quantifying forest-based climate change mitigation across forest biomes or ecozones (e.g., IPCC 2019).

Age trends and their variation across biomes

Stand age is another important axis of variation in forest C cycling (Fig. 1). In 1969, E.P. Odum’s “The Strategy of Ecosystem Development” laid out predictions as to how forest energy flows and organic matter stocks vary with stand age (Odum 1969). Although the conceptualization of the C cycle in this paper was simplistic by current standards, the paper was foundational in framing the theory around which research on the subject still revolves (Corman *et al* 2019), and the basic framework still holds, albeit with modest modifications (Fig. 1; Anderson-Teixeira *et al* 2013). Following stand-clearing disturbance, GPP , NPP , and biomass of leaves ($B_{foliage}$) and fine roots ($B_{root-fine}$) increase rapidly and thereafter remain relatively stable ($B_{foliage}$, $B_{root-fine}$, sometimes GPP) or decline slightly (NPP , sometimes GPP ; e.g., Law *et al* 2003, Pregitzer and Euskirchen 2004, Amiro *et al* 2010, Gouliden *et al* 2011). The decline in NPP occurs because R_{auto} increases relative to GPP as forests age, corresponding to declining carbon use efficiency with stand age (DeLucia *et al* 2007, Collalti *et al* 2020). Heterotrophic respiration, most of which originates from the soil ($R_{het-soil}$) remains relatively constant with stand age (Law *et al* 2003, Pregitzer and Euskirchen 2004, Gouliden *et al* 2011), with the result that net ecosystem production ($NEP = GPP - R_{eco}$, where R_{eco} is total ecosystem respiration) is initially negative, increases to a maximum at intermediate ages, and declines—typically to a small positive value—thereafter (Law *et al* 2003, Pregitzer and Euskirchen 2004, Amiro *et al* 2010, Gouliden *et al* 2011, Luyssaert *et al* 2008). The result is that biomass accumulates rapidly in young forests, followed by a slow decline to near zero in old forests (e.g., Lichstein *et al* 2009, Yang *et al* 2011). While these trends have been subject of fairly recent qualitative review (Anderson-Teixeira *et al* 2013), there is need for a synthetic, quantitative review taking advantage of the greatly expanded data now available.

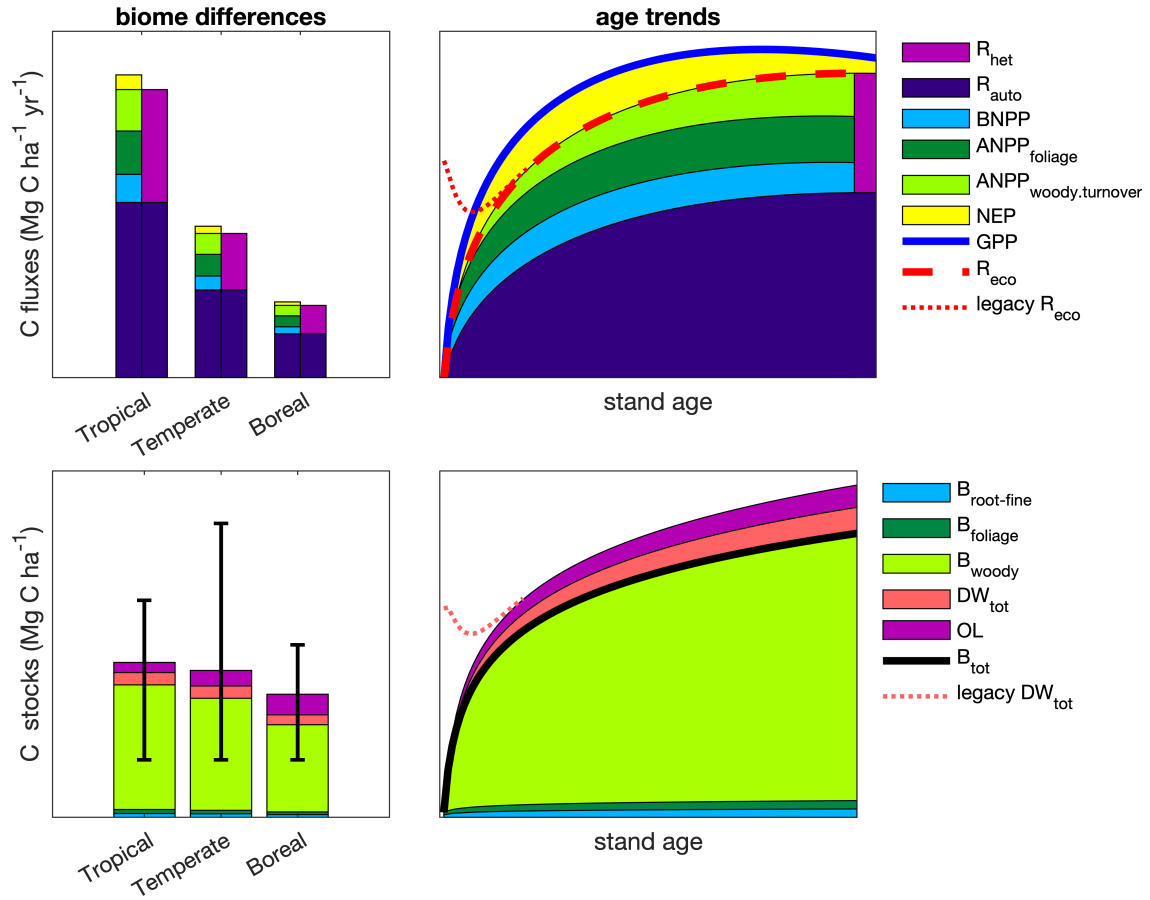


Figure 1 | Schematic diagram summarizing current understanding of biome differences and age trends in forest C cycling. Variables are defined in Table 1. Age trends, which represent idealized dynamics following a disturbance that removes all living vegetation, are an updated version of the classic figure from Odum (1969), with heavy lines corresponding to those in Odum's figure and NEP corresponding to Odum's 'net production'. Here, NEP consists primarily of woody aboveground net primary production ($ANPP_{woody}$), while $ANPP_{woody.turnover}$ is the sum of woody mortality and branch turnover. Dotted lines refer to decomposition of potential 'legacy' organic material produced prior to the disturbance and remaining at the site (e.g., standing and fallen dead wood, DW_{tot} ; soil organic matter). Error bars on C stocks plot represent within-biome variability, wherein mean biomass is highest in the tropics, but maximum biomass is highest in temperate regions.

In the past few decades, researchers have started asking how age trends—mostly in B_{ag} or total biomass (B_{tot}) accumulation—vary across biomes. Early research on this theme showed that biomass accumulation rates during secondary succession increase with temperature on a global scale (Johnson *et al* 2000, Anderson *et al* 2006) and with water availability in the neotropics (Poorter *et al* 2016). Most recently, Cook-Patton *et al* (2020) reinforced these earlier findings with a much larger dataset and created a high-resolution global map of estimated potential C accumulation rates. However, there has been little synthesis of cross-biome differences in variables other than biomass and its accumulation rate (but see Cook-Patton *et al* (2020) for DW , OL , and soil C accumulation in young stands). Given the important role of secondary forests in the current and future global C cycle, concrete understanding of age trends in C fluxes and stocks and how these vary across biomes is critical to better understanding of the global C cycle. Accurate estimates of C sequestration rates by regrowth forests are also critical for national greenhouse gas accounting under the IPCC framework (IPCC 2019, Requena Suarez *et al* 2019) and to quantifying the value of regrowth forests for climate change mitigation (Anderson-Teixeira and DeLucia 2011, Goldstein *et al* 2020).

Here, we conduct a data-based review of carbon cycling from a stand to global level, and by biome and stand age, using the largest global compilation of forest carbon data, which is available in our open source Global Carbon Forest database (*ForC*; Fig. 2). Our goal is to provide a comprehensive synthesis on broad trends in forest C cycling that can serve as a foundation for improved understanding of global forest C cycling and highlight where key sources of uncertainty still reside.

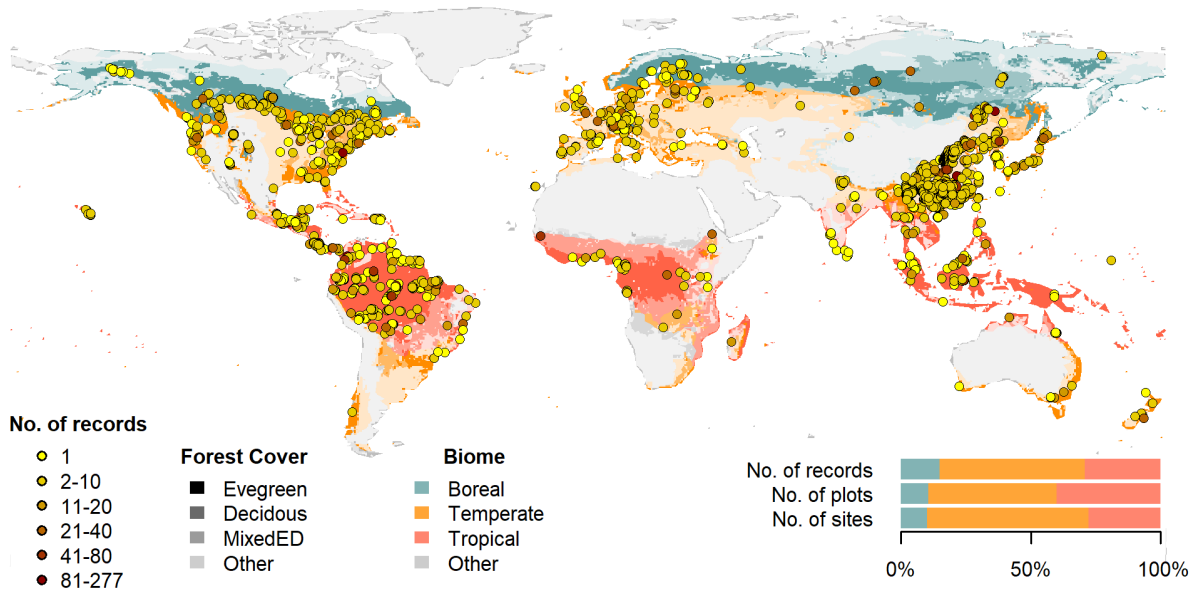


Figure 2 | Map of sites included in this analysis. Symbols are colored according to the number of records at each site. Underlying map shows coverage of evergreen, deciduous, and mixed forests (shading differences; data from Jung et al. 2006) and biomes (color differences). Distribution of sites, plots, and records among biomes is shown in the inset.

Methods/ Design

This review synthesizes data from the *ForC* database (Fig. 2; <https://github.com/forc-db/ForC>; Anderson-Teixeira *et al* 2016, 2018). *ForC* amalgamates numerous intermediary data sets (*e.g.*, Luyssaert *et al* 2007, Bond-Lamberty and Thomson 2010, Cook-Patton *et al* 2020) and original studies. Original publications were referenced to check values and obtain information not contained in intermediary data sets, although this process has not been completed for all records. The database was developed with goals of understanding how C cycling in forests varies across broad geographic scales and as a function of stand age. As such, there has been a focus on incorporating data from regrowth forests (*e.g.*, Anderson *et al* 2006, Martin *et al* 2013, Bonner *et al* 2013) and obtaining stand age data when possible (83% of records in v2.0; Anderson-Teixeira *et al* 2018). Particular attention was given to developing the database for tropical forests (Anderson-Teixeira *et al* 2016), which represented roughly one-third of records in *ForC* v2.0 (Anderson-Teixeira *et al* 2018). Since publication of *ForC* v2.0, we imported three large additional databases into *ForC* via a combination of R scripts and manual edits. First, we imported (via R script) the Global Database of Soil Respiration Database (*SRDB* v4, 9488 records; Bond-Lamberty and Thomson 2010), and corrections and improvements to *SRDB* arising from this process were incorporated in *SRDB* v5 (Jian *et al*

2020). Second, we imported (via R script) the Global Reforestation Opportunity Assessment database (*GROA* v1.0, 10116 records; Cook-Patton *et al* 2020, Anderson-Teixeira *et al* 2020), which itself had drawn on an earlier version of *ForC*. Because all records in *GROA* were checked against original publications, these records were given priority over duplicates in *ForC* (Appendix S1). Third, we incorporated records of annual *NEP*, *GPP*, and *R_{eco}* from the FLUXNET2015 dataset (Pastorello *et al* 2020), treating these records as authoritative when they duplicated earlier records (Appendix S1). We have also added data from individual publications, with a particular focus on productivity (e.g., Taylor *et al* 2017), dead wood, and ForestGEO sites (e.g., Lutz *et al* 2018, Johnson *et al* 2018). A record of data sets added to *ForC* over the course of its development is available at [https://github.com/forc-](https://github.com/forc-db/ForC/blob/master/database_management_records/ForC_data_additions_log.csv)
db/ForC/blob/master/database_management_records/ForC_data_additions_log.csv. The database version used for this analysis has been tagged as a new release on Github (v3.0) and assigned a DOI through Zenodo (DOI: TBD).

All measurements originally expressed in units of dry organic matter (*OM*) were converted to units of C using the IPCC default of $C = 0.47 * OM$ (IPCC 2018). Duplicate or otherwise conflicting records were purged as described in Appendix S1, resulting in a total of 22265 records (56% size of total database). Records were filtered to remove plots that had undergone significant anthropogenic management or major disturbance since the most recent stand initiation event. Specifically, we removed plots with any record of managements manipulating CO₂, temperature, hydrology, nutrients, or biota, as well as any plots whose site or plot name contained the terms “plantation”, “planted”, “managed”, “irrigated”, or “fertilized” (13.9% of duplicate-purged records). We also removed stands that had undergone any notable anthropogenic thinning or partial harvest (5.6% of duplicate-purged records). We retained sites that were grazed or had undergone low severity natural disturbances (<10% mortality) including droughts, major storms, fires, and floods. We removed all plots for which no stand history information had been retrieved (5.7% of duplicate-purged records). In total, this resulted in 17349 records (43.6% of the records in the database) being eligible for inclusion in the analysis.

We selected 23 annual flux and 11 C stock variables for inclusion in the analysis (Table 1). These different flux and stock variables represent different pools (e.g., aboveground biomass, root biomass, dead wood) and levels of combination (e.g., total net primary productivity, *NPP*, versus the individual elements of *NPP* such as foliage, roots, and branches). We did not analyze soil carbon, which is not a focus of the *ForC* database. Note that two flux variables, aboveground heterotrophic respiration (R_{het-ag}) and total respiration (R_{het}), were included for conceptual completeness but had no records in *ForC* (Table 1). Records for our focal variables represented 90.3% of the total records eligible for inclusion. For this analysis, we combined some of *ForC*’s specific variables into more broadly defined variables. Specifically, net ecosystem exchange (measured by eddy-covariance; Baldocchi *et al* 2001) and biometric estimates of *NEP* were combined into the single variable *NEP* (Table 1). Furthermore, for *NPP*, aboveground *NPP* (*ANPP*), and the litterfall component of *ANPP* (*ANPP_{litterfall}*), *ForC* variables specifying inclusion of different components were combined (e.g., measurements including or excluding fruit and flower production and herbivory). Throughout *ForC*, for all measurements drawing from tree census data (e.g., biomass, productivity), trees were censused down to a minimum diameter breast height (DBH) threshold of 10 cm or less. All records were measured directly or derived from field measurements.

We grouped forests into four broad biome types based on climate zones and dominant vegetation type (tropical broadleaf, temperate broadleaf, temperate needleleaf, and boreal needleleaf) and two age

Table 1. Carbon cycle variables included in this analysis, their sample sizes, and summary of biome differences and age trends.

Variable	Description	N records			biome differences*	age trend†
		records	plots	geographic areas		
Annual fluxes						
<i>NEP</i>	net ecosystem production or net ecosystem exchange (+ indicates C sink)	329	146	88	n.s.	+; xB
<i>GPP</i>	gross primary production ($NPP + R_{auto}$ or $NEP + R_{eco}$)	303	115	84	TrB > TeB ≥ TeN ≥ BoN	+; xB
<i>NPP</i>	net primary production ($ANPP + BNPP$)	214	112	74	TrB > TeB ≥ TeN > BoN	n.s.
<i>ANPP</i>	aboveground <i>NPP</i>	343	236	131	TrB > TeB ≥ TeN > BoN	+; xB
<i>ANPP_{woody}</i>	woody production ($ANPP_{stem} + ANPP_{branch}$)	64	53	37	n.s.	+
<i>ANPP_{stem}</i>	woody stem production	217	190	117	TrB > TeN ≥ TeB ≥ BoN	n.s.
<i>ANPP_{branch}</i>	branch turnover	69	59	42	TrB > TeB ≥ TeN	n.s.
<i>ANPP_{foliage}</i>	foliage production, typically estimated as annual leaf litterfall	162	132	88	TrB > TeB ≥ TeN > BoN	+
<i>ANPP_{litterfall}</i>	litterfall, including leaves, reproductive structures, twigs, and sometimes branches	82	70	55	n.s.	+
<i>ANPP_{repro}</i>	production of reproductive structures (flowers, fruits, seeds)	51	44	34	n.t.	n.t.
<i>ANPP_{folivory}</i>	foliar biomass consumed by folivores	20	12	11	n.t.	n.t.
<i>M_{woody}</i>	woody mortality—i.e., B_{ag} of trees that die	18	18	18	n.t.	n.t.
<i>BNPP</i>	belowground NPP ($BNPP_{coarse} + BNPP_{fine}$)	148	116	79	TrB > TeN ≥ TeB ≥ BoN	+
<i>BNPP_{coarse}</i>	coarse root production	77	56	36	TeN ≥ TrB	n.s.
<i>BNPP_{fine}</i>	fine root production	123	99	66	n.s.	+
<i>R_{eco}</i>	ecosystem respiration ($R_{auto} + R_{het}$)	213	98	70	TrB > TeB ≥ TeN	+
<i>R_{auto}</i>	autotrophic respiration ($R_{auto-ag} + R_{root}$)	24	23	15	n.t.	n.t.
<i>R_{auto-ag}</i>	aboveground autotrophic respiration (i.e., leaves and stems)	2	2	1	n.t.	n.t.
<i>R_{root}</i>	root respiration	181	139	95	TrB ≥ TeB	+
<i>R_{soil}</i>	soil respiration ($R_{het-soil} + R_{root}$)	627	411	229	TrB > TeB > TeN ≥ BoN	n.s.
<i>R_{het-soil}</i>	soil heterotrophic respiration	197	156	100	TrB > TeB ≥ TeN	n.s.
<i>R_{het-ag}</i>	aboveground heterotrophic respiration	0	0	0	-	-
<i>R_{het}</i>	heterotrophic respiration ($R_{het-ag} + R_{het-soil}$)	0	0	0	-	-
Stocks						
<i>B_{tot}</i>	total live biomass ($B_{ag} + B_{root}$)	188	157	87	TrB ≥ TeB > BoN	+; xB
<i>B_{ag}</i>	aboveground live biomass ($B_{ag-wood} + B_{foliage}$)	4466	4072	621	TrB ≥ TeN ≥ TeB > BoN	+; xB
<i>B_{ag-wood}</i>	woody component of aboveground biomass	115	102	64	TeN > TrB ≥ BoN	+; xB
<i>B_{foliage}</i>	foliage biomass	134	115	72	TeN > TrB ≥ BoN ≥ TeB	+; xB
<i>B_{root}</i>	total root biomass ($B_{root-coarse} + B_{root-fine}$)	2329	2298	360	n.s.	+; xB
<i>B_{root-coarse}</i>	coarse root biomass	134	120	73	TeN > TeB ≥ BoN	+; xB
<i>B_{root-fine}</i>	fine root biomass	226	180	109	n.s.	+; xB
<i>DW_{tot}</i>	deadwood ($DW_{standing} + DW_{down}$)	79	73	42	n.t.	+; xB
<i>DW_{standing}</i>	standing dead wood	36	35	22	n.t.	n.t.
<i>DW_{down}</i>	fallen dead wood, including coarse and sometimes fine woody debris	278	265	37	n.t.	+; xB
<i>OL</i>	organic layer / litter/ forest floor	474	413	115	n.s.	+; xB

* TrB: Tropical, TeB: Temperate Broadleaf, TeN: Temperate Needleleaf, BoN: Boreal, n.s.: no significant differences, n.t.: not tested

† + or -: significant positive or negative trend, xB: significant age x biome interaction, n.s.: no significant age trend, n.t.: not tested

235 classifications (young and mature). Climate zones (Fig. 2) were defined based on site geographic coordinates
 236 according to Köppen-Geiger zones (Rubel and Kottek 2010). We defined the tropical biome as including all

equatorial (A) zones, temperate biomes as including all warm temperate (C) zones and warmer snow climates (Dsa, Dsb, Dwa, Dw, Dfa, and Dfb), and the boreal biome as including the colder snow climates (Dsc, Dsd, Dwc, Dwd, Dfc, and Dfd). Any forests in dry (B) and polar (E) Köppen-Geiger zones were excluded from the analysis. We defined leaf type (broadleaf / needleleaf) based on descriptions in original publications (prioritized) or values extracted from a global map based on satellite observations (SYNMAP; Jung *et al* 2006). For young tropical forests imported from *GROA* but not yet classified by leaf type, we assumed that all were broadleaf, consistent with the rarity of naturally regenerating needleleaf forests in the tropics. We also classified forests as “young” (< 100 years) or “mature” (\geq 100 years or classified as “mature”, “old growth”, “intact”, or “undisturbed” in original publication). Assigning stands to these groupings required the exclusion of records for which *ForC* lacked geographic coordinates (0.4% of sites in full database) or records of stand age (5.7% of records in full database). We also excluded records of stand age = 0 year (0.8% of records in full database). In total, our analysis retained 11923 records. Numbers of records by biome and age class are given in Table S1.

Data were summarized to produce schematics of C cycling for mature forests of each biome. To obtain the values reported in the C cycle schematics, we first averaged any repeated measurements within a plot. Values were then averaged across geographically distinct areas, defined as plots clustered within 25 km of one another (*sensu* Anderson-Teixeira *et al* 2018), weighting by area sampled if available for all records. This step was taken to avoid pseudo-replication.

We tested whether the C budgets described above “closed”—*i.e.*, whether they were internally consistent. Specifically, we first defined relationships among variables: for example, $NEP = GPP - R_{eco}$, $BNPP = BNPP_{coarse} + BNPP_{fine}$, $DW_{tot} = DW_{standing} + DW_{down}$). Henceforth, we refer to the variables on the left side of the equation as “aggregate” fluxes or stocks, and those that are summed as “component” fluxes or stocks, noting that the same variable can take both aggregate and component positions in different relationships. We considered the C budget for a given relationship “closed” when component variables summed to within one standard deviation of the aggregate variable.

To test for differences across mature forest biomes, we also examined how stand age impacted fluxes and stocks, employing a mixed effects model (“lmer” function in “lme4” R package; Bates *et al* 2015) with biome as fixed effect and plot nested within geographic.area as random effects on the intercept. When Biome had a significant effect, we looked at a Tukey’s pairwise comparison to see which biomes were significantly different from one another. This analysis was run for variables with records for at least seven distinct geographic areas in more than one biome, excluding any biomes that failed this criteria (Table 1).

To test for age trends in young (<100yrs) forests, we employed a mixed effects model with biome and $\log_{10}[\text{stand age}]$ as fixed effects and plot nested within geographic area as a random effect on the intercept. This analysis was run for variables with records for at least three distinct geographic areas in more than one biome, excluding any biomes that failed this criteria (Table 1). When the effect of stand age was significant at $p \leq 0.05$ and when each biome had records for stands of at least 10 different ages, a biome \times stand.age interaction was included in the model. We note that the logarithmic function fit in this analysis does not always correspond to theoretical expectations (Fig. 1); however, data limitations did not support fitting of functions with more parameters or reliable comparison of different functional forms. Within the data constraints, we deemed a logarithmic function to be the most appropriate functional form for the majority of variables.

To facilitate the accessibility of our results and data, and to allow for rapid updates as additional data become available, we automated all database manipulation, analyses, and figure production in R (Team 2020).

Review Results/ Synthesis

Data Coverage

Of the 39762 records in *ForC* v3.0, 11923 met our strict criteria for inclusion in this study (Fig. 2). These records were distributed across 5062 plots in 865 distinct geographic areas. Of the 23 flux and 11 stock variables mapped in our C cycle diagrams (Figs. 3-6, S1-S4), *ForC* contained sufficient mature forest data for inclusion in our statistical analyses (*i.e.*, records from ≥ 7 distinct geographic areas) for 20 fluxes and 9 stocks in tropical broadleaf forests, 15 fluxes and 8 stocks in temperate broadleaf forests, 14 fluxes and 7 stocks in temperate conifer forests, and 8 fluxes and 7 stocks in boreal forests. For regrowth forests (<100 yrs), *ForC* contained sufficient data for inclusion in our statistical analyses (*i.e.*, records from ≥ 3 distinct geographic areas) for 11 fluxes and 10 stocks in tropical broadleaf forests, 16 fluxes and 10 stocks in temperate broadleaf forests, 16 fluxes and 10 stocks in temperate conifer forests, and 14 fluxes and 9 stocks in boreal forests.

C cycling in mature forests

Average C cycles for mature tropical broadleaf, temperate broadleaf, temperate conifer, and boreal forests ≥ 100 years old and with no known major natural or anthropogenic disturbance are presented in Figures 3-6 (and available in tabular format in the *ForC* release accompanying this publication: `ForC/numbers_and_facts/ForC_variable_averages_per_Biome.csv`).

For variables with records from ≥ 7 distinct geographic areas, these ensemble C budgets met our criteria for budget “closure”. That is, component variables summed to within one standard deviation of their respective aggregate variables in all but one instance. In the temperate conifer biome, the average composite measure of root biomass (B_{root}) was less than the combined average value of coarse and fine root biomass ($B_{root-coarse}$ and $B_{root-fine}$, respectively). This lack of closure was driven by very high estimates of $B_{root-coarse}$ from high-biomass forests of the US Pacific Northwest (Fig. S25).

Figure 3 | C cycle diagram for mature tropical broadleaf forests. Arrows indicate fluxes ($\text{Mg C ha}^{-1} \text{ yr}^{-1}$); boxes indicate stocks (Mg C ha^{-1}), with variables as defined in Table 1. Presented are mean \pm std, where geographically distinct areas are treated as the unit of replication. Dashed shape outlines indicate variables with records from <7 distinct geographic areas, and dashed arrows indicate fluxes with no data. To illustrate the magnitude of different fluxes, arrow width is proportional to the square root of the corresponding flux. Mean component fluxes do not necessarily add up to the mean total fluxes because different sets of sites are included depending on availability of data (Figs. S5-S30).

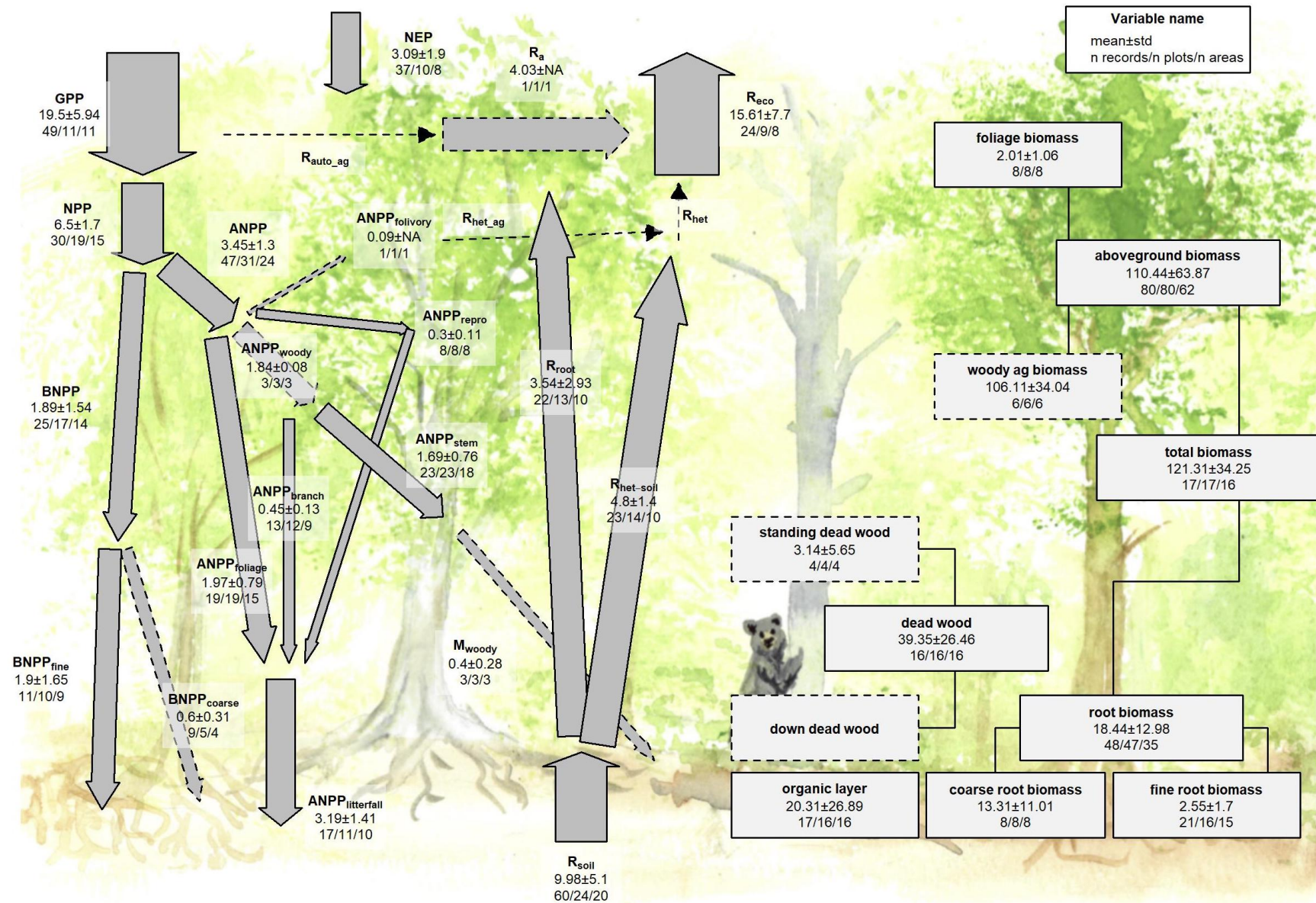


Figure 4 | C cycle diagram for mature temperate broadleaf forests. Arrows indicate fluxes (Mg C ha⁻¹ yr⁻¹); boxes indicate stocks (Mg C ha⁻¹), with variables as defined in Table 1. Presented are mean \pm std, where geographically distinct areas are treated as the unit of replication. Dashed shape outlines indicate variables with records from <7 distinct geographic areas, and dashed arrows indicate fluxes with no data. To illustrate the magnitude of different fluxes, arrow width is proportional to the square root of the corresponding flux. Mean component fluxes do not necessarily add up to the mean total fluxes because different sets of sites are included depending on availability of data (Figs. S5-S30).



Figure 5 | C cycle diagram for mature temperate conifer forests. Arrows indicate fluxes (Mg C ha⁻¹ yr⁻¹); boxes indicate stocks (Mg C ha⁻¹), with variables as defined in Table 1. Presented are mean ± std, where geographically distinct areas are treated as the unit of replication. Dashed shape outlines indicate variables with records from <7 distinct geographic areas, and dashed arrows indicate fluxes with no data. To illustrate the magnitude of different fluxes, arrow width is proportional to the square root of the corresponding flux. Mean component fluxes do not necessarily add up to the mean total fluxes because different sets of sites are included depending on availability of data (Figs. S5-S30). The temperate conifer biome in particular is subject to high variability, with highest fluxes and stocks in the high-biomass forests of the US Pacific Northwest. An asterisk after a variable name indicates lack of C cycle closure.

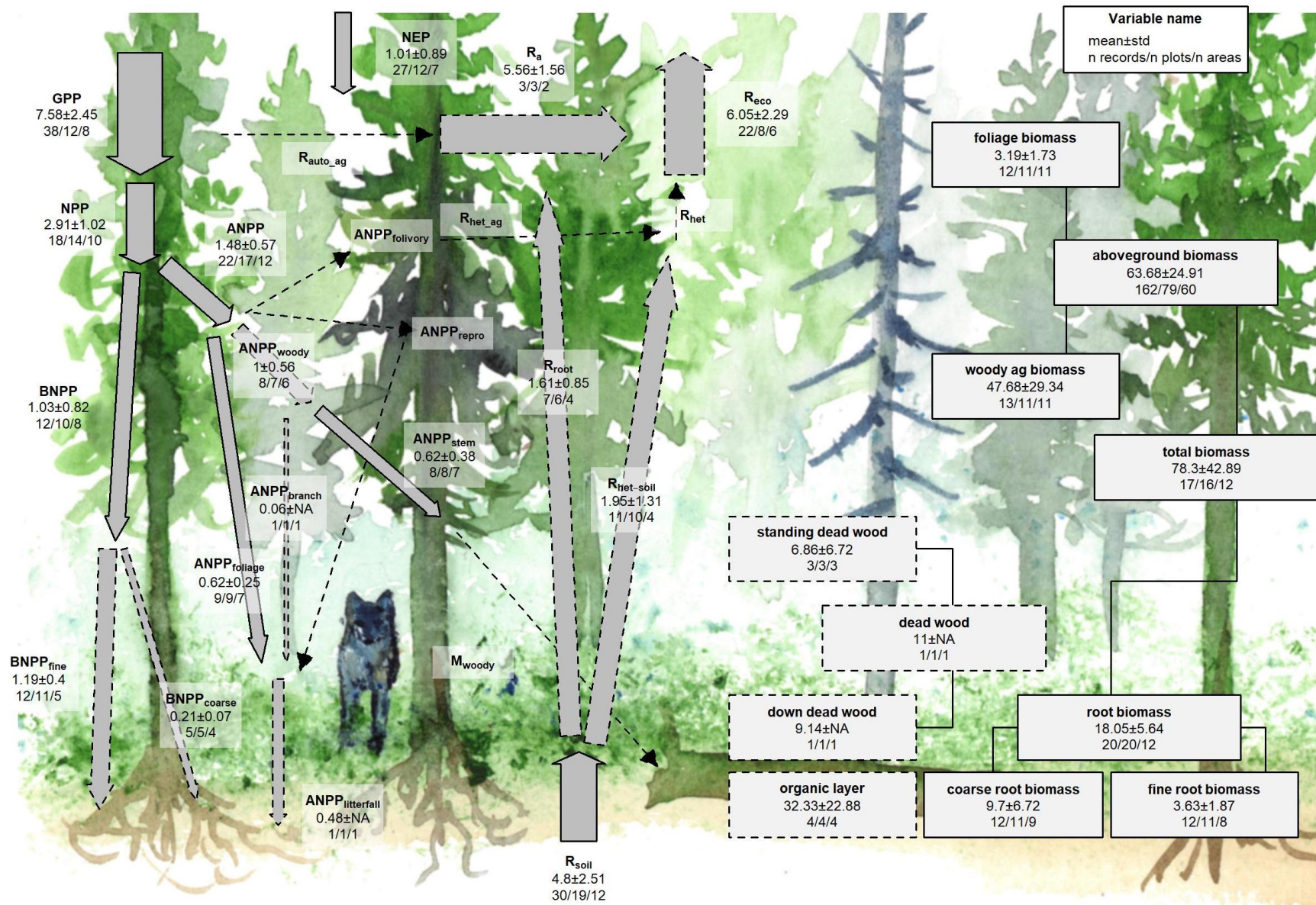


Figure 6 | C cycle diagram for mature boreal conifer forests. Arrows indicate fluxes (Mg C ha⁻¹ yr⁻¹); boxes indicate stocks (Mg C ha⁻¹), with variables as defined in Table 1. Presented are mean \pm std, where geographically distinct areas are treated as the unit of replication. Dashed shape outlines indicate variables with records from <7 distinct geographic areas, and dashed arrows indicate fluxes with no data. To illustrate the magnitude of different fluxes, arrow width is proportional to the square root of the corresponding flux. Mean component fluxes do not necessarily add up to the mean total fluxes because different sets of sites are included depending on availability of data (Figs. S5-S30).

There were sufficient data to assess differences among biomes in mature forest values for 15 flux variables, and significant differences among biomes were detected for 12 variables (Table 1). In all of these cases—including C fluxes into, within, and out of the ecosystem—C fluxes were highest in tropical forests, intermediate in temperate (broadleaf or conifer) forests, and lowest in boreal forests (Table 1, Figs. 7, S5-S19). Differences between tropical and boreal forests were always significant, with temperate forests intermediate and significantly different from one or both. Fluxes tended to be numerically greater in temperate broadleaf than temperate conifer forests, but the difference was never statistically significant. This pattern held for the following variables: GPP , NPP , $ANPP$, $ANPP_{stem}$, $ANPP_{branch}$, $ANPP_{foliage}$, $BNPP$, R_{eco} , R_{root} , R_{soil} , and $R_{het-soil}$. For two of the variables without significant differences among biomes ($ANPP_{litterfall}$ and $BNPP_{fine}$; Figs. S12 and S15, respectively), the same general trends applied but were not statistically significant. Another exception was for $BNPP_{root-coarse}$, where all records came from high-biomass forests in the US Pacific Northwest, resulting in marginally higher values for the temperate conifer biome (Fig. S14; differences significant in mixed effects model but not in post-hoc pairwise comparison).

The most notable exception to the pattern of decreasing flux per unit area from tropical to boreal biomes was NEP , with no significant differences across biomes but with the largest average in temperate broadleaf forests, followed by tropical, boreal, and temperate conifer forests (Figs. 7, S5). For all biomes, NEP was positive, with 95% confidence intervals excluding zero.

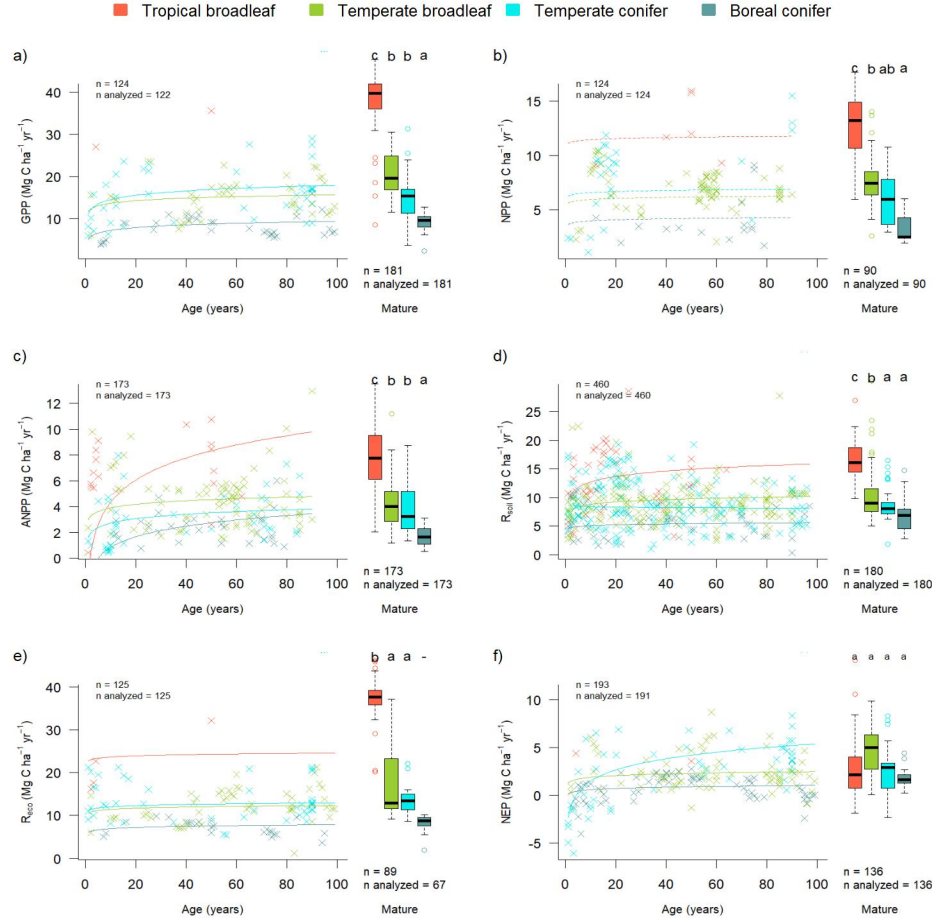


Figure 7 | Age trends and biome differences in some of the major C fluxes: (a) *GPP*, (b) *NPP*, (c) *ANPP*, (d) *R_{soil}*, (e) *R_{eco}*, and (f) *NEP*. In each panel, the left scatterplot shows age trends in forests up to 100 years old, as characterized by a linear mixed effects model with fixed effects of $\log_{10}(\text{age})$ and biome. The fitted line indicates the effect of age (solid lines: significant at $p < 0.05$, dashed lines: non-significant), and non-parallel lines indicate a significant $\log_{10}(\text{age}) \times \text{biome}$ interaction. The boxplot illustrates distribution across mature forests, with different letters indicating significant differences between biomes. Data from biomes that did not meet the sample size criteria (see Methods) are plotted, but lack regression lines (young forests) or test of differences across biomes (mature forests). Individual figures for each flux with sufficient data, along with maps showing geographic distribution of the data, are given in the Supplement (Figs. S5-S19).

Biome differences were less consistent across C stocks than fluxes (Figs. 8, S20-S30). There were sufficient data to assess mature forest biome differences for nine stock variables, and significant differences among biomes were detected for five variables (B_{tot} , B_{ag} , $B_{\text{ag-wood}}$, B_{foliage} , $B_{\text{root-coarse}}$; Table 1). For B_{tot} and B_{ag} , tropical broadleaf forests had the highest mean biomass and boreal forests the lowest, with intermediate means for temperate broadleaf and needleleaf forests (temperate needleleaf excluded from B_{tot} analysis because of insufficient data; Figs. S20, S21). However, maximum values for these variables – along with all other stocks including live or standing woody biomass ($B_{\text{ag-wood}}$, B_{root} , $B_{\text{root-coarse}}$, DW_{tot} , DW_{standing}) – consistently occurred in temperate biomes (Figs. 1, 8, S20-S30). For variables that were disproportionately sampled in such high-biomass forests ($B_{\text{ag-wood}}$, B_{foliage} , and $B_{\text{root-coarse}}$; disproportionately sampled in the US Pacific Northwest), temperate conifer forests had significantly higher stocks than the other biomes.

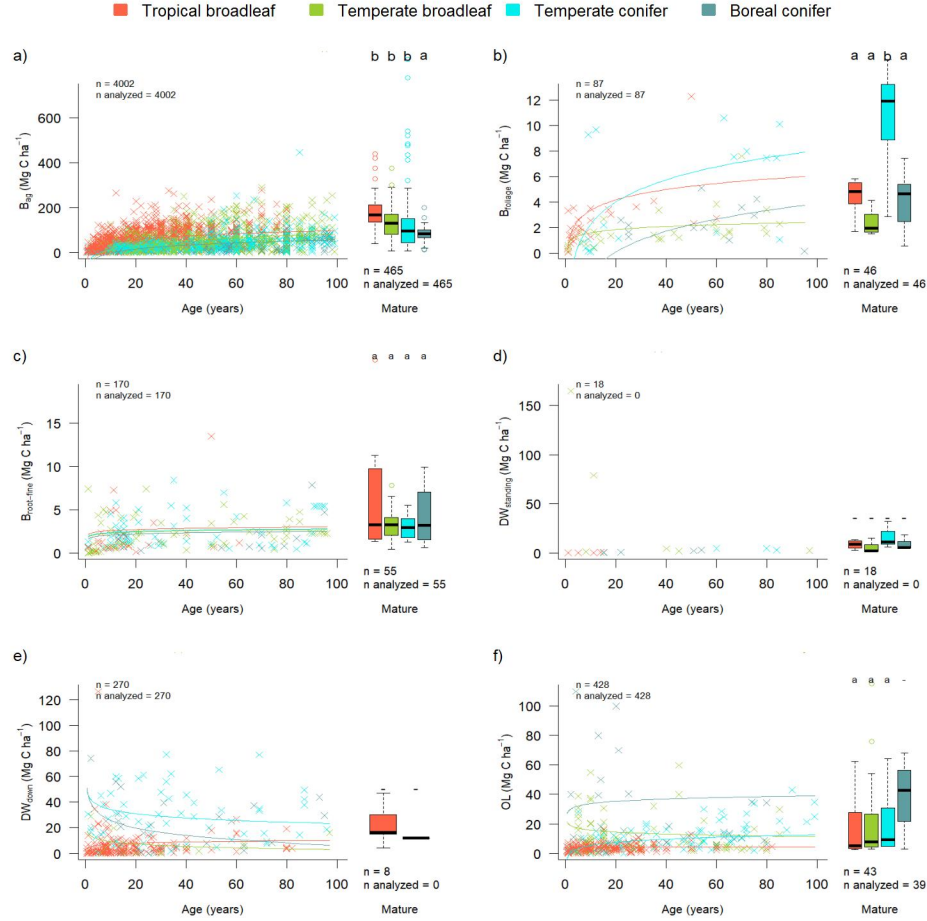


Figure 8 | Age trends and biome differences in some of the major forest C stocks: (a) aboveground biomass, (b) foliage, (c) fine roots, (d) dead wood. In each panel, the left scatterplot shows age trends in forests up to 100 years old, as characterized by a linear mixed effects model with fixed effects of $\log_{10}(\text{age})$ and biome. The fitted line indicates the effect of age (solid lines: significant at $p < 0.05$, dashed lines: non-significant), and non-parallel lines indicate a significant $\log_{10}(\text{age}) \times \text{biome}$ interaction. The boxplot illustrates distribution across mature forests, with different letters indicating significant differences between biomes. Data from biomes that did not meet the sample size criteria (see Methods) are plotted, but lack regression lines (young forests) or test of differences across biomes (mature forests). Individual figures for each stock with sufficient data, along with maps showing geographic distribution of the data, are given in the Supplement (Figs. S20-S30).

C cycling in young forests

C fluxes commonly increased significantly with stand age (Tables 1, S2, Figs. 7, 9, S5-S30). *ForC* contained 16 C flux variables with sufficient data for analyses of age trends in young forests (see Methods). Of these, ten increased significantly with $\log_{10}[\text{age}]$: NEP , GPP , $ANPP$, $ANPP_{\text{woody}}$, $ANPP_{\text{foliage}}$, $ANPP_{\text{litterfall}}$, $BNPP$, $BNPP_{\text{fine}}$, R_{eco} , and R_{root} . The remaining six – NPP , $ANPP_{\text{stem}}$, $ANPP_{\text{branch}}$, $BNPP_{\text{coarse}}$, R_{soil} , and $R_{\text{het-soil}}$ – displayed no significant relationship to stand age.

Differences in C fluxes across biomes typically paralleled those observed for mature forests, with C cycling generally most rapid in the tropics and slowest in boreal forests (Table 1, Figs. 7, S5-S30). The single exception was $ANPP_{\text{stem}}$, for which temperate broadleaf and conifer forests had flux rates similar to tropical forests. Notably, and in contrast to the lack of biome differences in NEP for mature forests (Fig. 7), the tendency for temperate forests to have greater fluxes than boreal forests held for NEP in regrowth forests

341 (tropical forests excluded because of insufficient data).

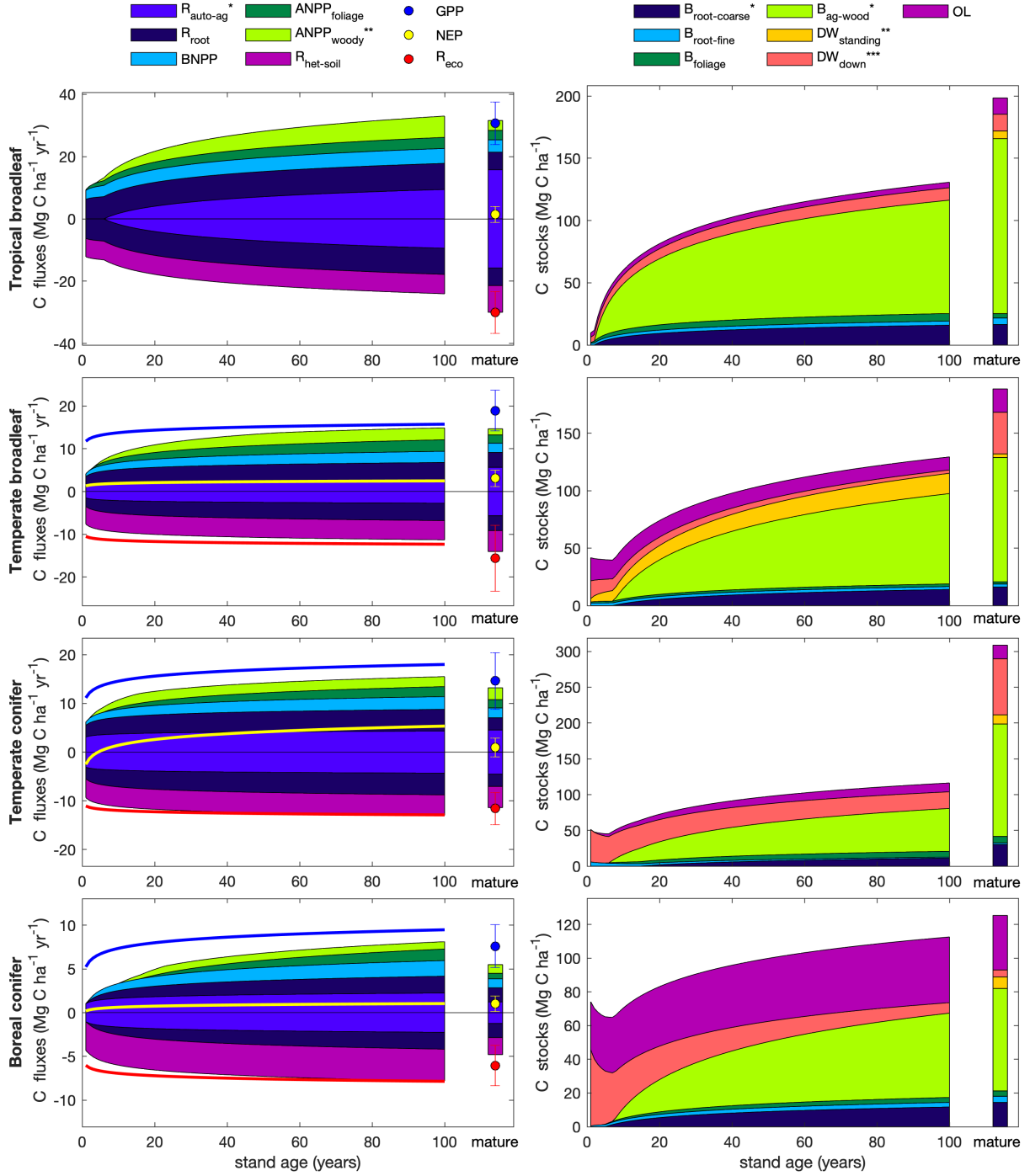


Figure 9 | Age trends in C cycling. Selection of variables for plotting seeks to maximize sample size and broad geographic representation while representing all elements of C cycle. Error bars on mature forest flux estimates indicate ± 1 standard deviation. Asterisks indicate variables whose age trends were calculated based on other variables (* young and mature forests; ** young forests only; *** mature forests only), as follows. For all forests: $B_{ag-wood} = \max(0, B_{ag} - B_{foliage})$, $B_{root-coarse} = \max(0, B_{root} - B_{root-fine})$, $DW_{standing} = \max(0, DW_{tot} - DW_{down})$. For tropical forests: $ANPP_{woody} = \max(0, ANPP - ANPP_{foliage})$, $R_{auto-ag} = R_{auto} - R_{root}$, where $R_{auto} = NPP(1/CUE - 1)$ and $CUE=0.46$ (Collati et al. 2020). For non-tropical forests: $ANPP_{woody} = \min(ANPP_{stem}, ANPP_{woody})$, $R_{auto-ag} = R_{eco} - R_{soil}$. Note that there remain substantial uncertainties as to the functional form of age trends and discrepancies in closure among related variables.

“Closure” and internal consistency of the C flux budget was less successful for young than mature forests (Figs. 9). Summed regression equations for $R_{soil-het}$ and R_{root} were generally very close to R_{soil} . In assessing the C budget of young forests, we calculated $R_{auto-ag}$ as the difference between R_{eco} and R_{soil} (except for tropical forests, which had insufficient R_{eco} data), effectively guaranteeing near-closure of the CO₂ efflux (respiration) portion of the budget (negative values in Figs. 9). In contrast, the CO₂ influx portion of the budget generally did not “close”: the sum of R_{auto} ($R_{root} + R_{auto-ag}$, as described above) and components of NPP consistently fell short of GPP , particularly in young stands (range across forest types and ages: 0.9-7.6 Mg C ha⁻¹ yr⁻¹). Moreover, there was not consistent budget closure among the components of NPP , and substantially different age trends resulting from the sum of components versus total NPP (Figs. 9). Although age trends of young forests often converged towards mature forest averages, there were also some discrepancies between young forest trends and mature forest averages (Figs. 7, 9, S5-S30), most notably including a tendency for higher fluxes in regrowth boreal forests than in their mature counterparts (Fig. 9).

In terms of C stocks, ten variables (all but standing deadwood, $DW_{standing}$) had sufficient data to test for age trends (Table 1, Figs. 8, S20-S30). All of these displayed a significant overall increase with with $\log_{10}[stand.age]$. Age \times biome interactions were also significant for all ten of these C stock variables (Table S2), with living C stocks tending to accumulate more rapidly during the early stages of forest regrowth in tropical forests (Figs. 8, 9, S20-S30). In the case of two non-living C stocks (DW_{down} and OL), age \times biome interactions were such that age trends were positive in some biomes and negative in others. Specifically, DW_{down} declined with age in temperate and boreal forests, compared to an increase with age in tropical forests (Figs. 8,9, S29). Similarly, OL declined slightly with age in temperate broadleaf forests, contrasting an increase in the other three biomes (Figs. 8, 9,S30). Again, there were some discrepancies between young forest trends and mature forests, most notably including generally higher C stocks in mature forests relative to their 100-year counterparts, particularly for temperate conifer forests (with discrepancies again driven by differences in geographic representation) and, to a lesser extent, tropical broadleaf forests (Fig. 9).

Discussion

ForC v3.0 provided unprecedented coverage of most major variables, yielding a broad picture of C cycling in the world's major forest biomes. Carbon cycling rates generally increased from boreal to tropical regions and with stand age (Figs. 1, 9). Specifically, most C fluxes were highest in tropical forests, intermediate in temperate (broadleaf or conifer) forests, and lowest in boreal forests – a pattern that generally held for regrowth as well as mature forests (Figs. 1, 7- 8, 9). The notable exception was mature forest *NEP*, which, as the difference between *GPP* and *R_{eco}*, was statistically indistinguishable across biomes (Fig. 7f). There was also little directional variation in mean mature forest C stocks (biomass, dead wood, and organic layer) across biomes, although maximum values for the majority of stocks (all including live or standing woody biomass) occurred in temperate biomes (Figs. 1, 3-6, 8). Consistent with theory and previous studies (Fig. 1), the majority of flux variables, together with most live biomass pools, increased significantly with stand age (Table 1; Figs. 7- 9, S5-S30). Together, these results indicate that, moving from cold to tropical climates and from young to old stands, there is a general acceleration of C cycling, whereas C stocks and *NEP* of mature forests, which are defined by the differences between in- and out- fluxes, do not vary systematically across biomes. Together, these results refine and expand out understanding of C cycling in mature forests, while providing the first global-scale analysis of age trends in multiple forest C stocks and fluxes (Figs. 9).

C cycling across biomes

Our analysis reveals that carbon cycling is most rapid in the tropics and slowest in boreal regions, including C fluxes into (*GPP*), within (e.g., *NPP* and its components), and out of (e.g., *R_{soil}*, *R_{eco}*) the ecosystem. For mature forests, this is consistent with a large body of previous work demonstrating that C fluxes generally decline with latitude – or increase with temperature – on a global scale (e.g., Luyssaert *et al* 2007, Gillman *et al* 2015, Li and Xiao 2019, Banbury Morgan *et al* n.d.). This consistency is not surprising, particularly given commonality in the data analyzed or used for calibration. The finding that these patterns hold consistently across numerous fluxes, while consistent with theoretical expectations (Fig. 1), is novel to this analysis (but see Banbury Morgan *et al* n.d. for nine autotrophic fluxes).

The notable exception to the pattern of fluxes decreasing from tropical to boreal regions is *NEP*, which showed no significant differences across biomes, albeit with the highest mean in temperate broadleaf forests (Fig. 7f). Unlike the other C flux variables, *NEP* does not characterize the rate at which C cycles through the ecosystem, but, as the balance between *GPP* and *R_{eco}*, represents net CO₂ sequestration (or release) by the ecosystem (Fig. 1). *NEP* tends to be relatively small in mature forest stands, which accumulate carbon slowly relative to younger stands, if at all (Fig. 1; Luyssaert *et al* 2008, Amiro *et al* 2010, Besnard *et al* 2018). It is therefore consistent with theory – and with previous research (Luyssaert *et al* 2007) – that there are no pronounced differences across biomes. Rather, variation in *NEP* of mature forests appears to be controlled less by climate and more by other factors including moderate disturbances (Curtis and Gough 2018) or disequilibrium of *R_{soil}* relative to C inputs (e.g., in peatlands where anoxic conditions inhibit decomposition; Wilson *et al* 2016). The fact that mature temperate broadleaf forests have a higher mean than the other biomes may reflect the fact that most of these forests are older secondary forests that, while classified here as mature, are still accumulating carbon (Curtis and Gough 2018).

In contrast to the patterns observed for *NEP* in mature stands, *NEP* of stands between 20 and 100 years of age varied across biomes, being lowest in boreal forests, intermediate in temperate broadleaf forests, and highest in temperate conifer forests (with insufficient data to assess tropical forests; Figs. 7, S5). This is

consistent with findings that live biomass accumulation rates (ΔB_{ag} or ΔB_{tot}) during early secondary succession decrease with latitude (Figs. 8a, S20-S30; Anderson *et al* 2006, Cook-Patton *et al* 2020). Note, though, that NEP includes not only ΔB_{tot} , but also changes in DW_{tot} , OL , and soil carbon (not analyzed here). Biome differences in the accumulation rates of DW , OL , and soil C have not been detected, in part because these variables do not consistently increase with stand age (Figs. 1, 8, S27-S30, and see discussion below; Cook-Patton *et al* 2020).

For regrowth forests, little was previously known about cross-biome differences in carbon fluxes, and we are not aware of any previous large-scale comparisons of C fluxes that have been limited to regrowth forests. Thus, this analysis was the first to examine flux trends in regrowth forests across biomes. The observed tendency for young forest fluxes to decrease from tropical to boreal regions paralleled patterns in mature forests (Figs. 7, 9, S5-S19), suggesting that regrowth forests follow latitudinal trends in carbon cycling similar to those of mature forests (e.g., Banbury Morgan *et al* n.d.).

In contrast to C fluxes and biomass accumulation rates in regrowth forests, stocks showed less systematic variation across biomes (c.f. Fig. 1). For aboveground biomass, which is the variable in $ForC$ with broadest geographical representation, the modest trend of declining biomass from tropical to boreal regions mirrors observations from spaceborne lidar that reveal a decline in aboveground biomass (for all forests, including secondary) with latitude across the N hemisphere (Hu *et al* 2016). The highest- biomass forests on Earth are, however, found in coastal temperate climates of both the southern and northern hemisphere (Figs. 1, 8a; Keith *et al* 2009, Smithwick *et al* 2002, Hu *et al* 2016). Disproportionate representation of forests in one such region—the US Pacific Northwest—inflated estimates of temperate conifer fluxes and stocks for some variables and was responsible for all of the anomalous results described here (e.g., lack of complete C budget closure, anomalous trend across biomes for $BNPP_{coarse}$). Thus, biome differences should always be interpreted relative to the geographic distribution of sampling, which only rarely covers the majority of forested area within a biome.

Whereas aboveground biomass can be remotely sensed (albeit with significant uncertainties; Ploton *et al* 2020) and receives significant research attention, far less is known about geographical variation in deadwood and organic layer (OL) across biomes, which has proved a limitation for C accounting efforts (Pan *et al* 2011). Although these stocks can be important—exceeding 100 Mg C ha⁻¹ in some stands (Figs. 8, S27-S29), this study is the first to synthesize deadwood data on a global scale (but see Cook-Patton *et al* 2020 for young forests). Unfortunately, data remain too sparse for statistical comparison across biomes (Figs. 8, S27-S29; but see below for age trends), pointing to a need for more widespread quantification of both standing and downed deadwood. $ForC$ coverage of OL stocks is more comprehensive, revealing no significant differences across temperate and tropical biomes, but a tendency towards higher OL in boreal forests, consistent with the idea that proportionally slower decomposition in colder climates results in more buildup of organic matter (Fig. 1; Allen *et al* 2002, Anderson-Teixeira *et al* 2011). Further research on non-living C stocks in the world’s forests will be essential to completing the picture.

Age trends in C cycling

Our study reveals that most C fluxes quickly increase and then decelerate as stands age (Figs. 7, 9), consistent with current understanding of age trends in forest C cycling (Fig. 1; e.g., Anderson-Teixeira *et al* 2013, Amiro *et al* 2010, Magnani *et al* 2007). While limited records in very young (*i.e.*, <5 year old) stands resulted in poor resolution of the earliest phases of this increase for many variables (sometimes detecting no

age trend; Table 1), any autotrophic C flux (e.g., GPP , NPP and its components, R_{auto}) would be minimal immediately following a stand-clearing disturbance (Fig. 1). These would be expected to increase rapidly, along with the most metabolically active components of biomass, foliage and fine roots, which also increase rapidly with stand age (Figs. 1, 7-9). In contrast, soil heterotrophic respiration ($R_{het-soil}$) and total soil respiration (R_{soil})—and therefore R_{eco} are expected to be non-zero following stand-clearing disturbance (Fig. 1), although these may decrease with a reduction of root respiration (R_{soil} only) and C exudates or increase in response to an influx of dead roots and litter (Ribeiro-Kumara *et al* 2020, Maurer *et al* 2016, Bond-Lamberty *et al* 2004). In this study, we detect no significant age trends in either $R_{het-soil}$ or R_{soil} , consistent with previous findings (Law *et al* 2003, Pregitzer and Euskirchen 2004, Goulden *et al* 2011).

Notably, net carbon sequestration (NEP) exhibits an overall increase with age across the first 100 years of stand development, with more pronounced patterns in temperate than boreal forests (Fig. 7). This finding is consistent with previous studies showing an increase in NEP across relatively young stand ages (Pregitzer and Euskirchen 2004, Baldocchi *et al* 2001, Luyssaert *et al* 2008). However, NEP is theoretically expected to peak in intermediate-aged stands and thereafter decline, consistent with decelerating C accumulation as stands age (Fig. 9; Odum 1969), and such declines have been observed (Law *et al* 2003, Luyssaert *et al* 2008). The fact that NEP values estimated by our models for 100-year-old stands were not systematically different from those of mature stands (lower for temperate broadleaf, higher for temperate conifer, and equal for boreal; Fig. 9) may be driven by differences in geographical representation across age classes or by the fitting of an inappropriate functional form. Moreover, both biomass and non-living C stocks often continue to increase well beyond the 100-yr threshold used here to delimit young and mature stands (Luyssaert *et al* 2008, McGarvey *et al* 2014, Lichstein *et al* 2009). Additional data, including on age trends of deadwood, the organic layer, and soil C will be important to parsing the timing and extent of an age-related NEP decrease across forest biomes.

In terms of stocks, our study reveals consistent increases in live biomass stocks with stand age, a pattern that is well-known and expected (e.g., Lichstein *et al* 2009, Yang *et al* 2011), contrasting with more variable age trends in deadwood and the organic layer (Fig. 9). The latter are particularly sensitive to the type of disturbance, where disturbances that remove most organic material (e.g., logging, agriculture) result in negligible deadwood in young stands, followed by a buildup over time (e.g., tropical stands in Figs. 8, 9; e.g., Vargas *et al* 2008). In contrast, natural disturbances (e.g., fire, drought) can produce large amounts of deadwood (mostly $DW_{standing}$) that slowly decomposes as the stand recovers, resulting in declines across young stand ages (e.g., temperate and boreal stands in Figs. 8, 9; Carmona *et al* 2002). Again, further study and synthesis of non-living C stocks across biomes and stand ages will be valuable to giving a more comprehensive picture.

C variable coverage and budget closure

The large number of C cycle variables covered by ForC, and the relatively high consistency among them (Figs. 3-6, 9), provide confidence that our reported mature forest means provide useful baselines for analysis – with the caveats that they are unlikely to be accurate representations of C cycling for any particular forest, and that these sample means almost certainly do not represent true biome means (particularly for temperate conifer forests where high-biomass stands are over-represented in *ForC*).

In this analysis, the C cycle budgets for mature forests usually come close to closure—that is, the sums of component variables do not differ from the larger fluxes by more than one standard deviation (Figs. 3-6, 9).

On the one hand, this reflects the general fact that ecosystem-scale measurements tend to close the C budget more easily and consistently than, for example, for energy balance (Stoy *et al* 2013). On the other, however, *ForC* derives data from multiple heterogeneous sources, and standard deviations within each biome are high; as a result, the standard for C closure is relatively loose (*c.f.* Houghton 2020). The one instance where the C budgets does not close according to our criteria is likely due to differences in the representation of forest types (*i.e.*, disproportionate representation of US Pacific NW for $B_{root-coarse}$ relative to B_{root} ; Fig. 5) rather than issues of methodological accuracy. The overall high degree of closure implies that *ForC* gives an at least roughly consistent picture of C cycling within biomes for mature forests. This is an important and useful test, because it allows for consistency checks within the C cycle, for example leveraging separate and independently-measured fluxes to constrain errors in another (Phillips *et al* 2017, Williams *et al* 2014, Harmon *et al* 2011), or producing internally consistent global data products (Wang *et al* 2018).

In contrast, age trends for young forests generally remain less clearly defined, in large part because their data records remain relatively sparse (*i.e.*, have low representation of different geographical regions for any given age) for most variables, particularly in the tropics (Anderson-Teixeira *et al* 2016). While this review provides a first analysis of age trends in forest C cycling for multiple variables at a global scale, improved resolution of these trends will require additional data.

There are of course notable holes in the *ForC* variable coverage that limit the scope of our inferences here. Notably, *ForC* currently has sparse—if any—coverage of fluxes to herbivores and higher consumers, along with the woody mortality (M_{woody}) and DW (Table 1, Figs. S27-S29). *ForC* does not include soil carbon, which is covered by other efforts (e.g., Köchy *et al* 2015). *ForC* is not intended to replace databases that are specialized for particular parts of the C cycle analyses, e.g., aboveground biomass (Spawn *et al* 2020), land-atmosphere fluxes (Baldocchi *et al* 2001), soil respiration (Jian *et al* 2020), or the human footprint in global forests (Magnani *et al* 2007).

Importantly, *ForC* and the analyses presented here cover the forests that have received research attention, which are not a representative sample of the world’s existing forests, either geographically or in terms of human impacts (Martin *et al* 2012). Geographically, all variables are poorly covered in Africa and Siberia (Fig. 2), a common problem in the carbon-cycle community (Xu and Shang 2016, Schimel *et al* 2015). In terms of human impacts, research efforts tend to focus on interior forest ecosystems (Martin *et al* 2012), often in permanently protected areas (e.g., Davies *et al* 2021). Studies of regrowth forests tend to focus on sites where recurring anthropogenic disturbance is not a confounding factor. Yet, fragmentation and degradation impact a large and growing proportion of Earth’s forests (FAO and UNEP 2020). Fragmentation and the creation of edges strongly impacts forest C cycling (e.g., Chaplin-Kramer *et al* 2015, Remy *et al* 2016, Reinmann and Hutrya 2017, Smith *et al* 2019, Reinmann *et al* 2020, Ordway and Asner 2020). Partial logging and other forms of non-stand clearing anthropogenic disturbance also alter forest C cycling (e.g., Huang and Asner 2010, Piponirot *et al* 2016), but are under-studied (Sist *et al* 2015) and excluded from this analysis. Fragmented and degraded forests do not fit the idealized conceptual framework around which this review is structured (Fig. 1), yet their representation in models, sustainability assessments, and C accounting systems is critical to accurate accounting of C cycling in Earth’s forests (e.g., Huang and Asner 2010, Reinmann and Hutrya 2017, Smith *et al* 2019, Piponirot *et al* 2019). Finally, plantation forests account for approximately 3% of Earth’s forests (FAO and UNEP 2020) but are not included in this analysis. While it is known that these tend to accumulate biomass faster than naturally regenerating forests (Anderson *et al* 2006, Bonner *et al* 2013), their global scale C cycling patterns remain less clearly understood (*c.f.* Cook-Patton *et*

al 2020). Additional research and synthesis are needed to fill these critical gaps in our understanding of forest C cycling.

Relevance for climate change prediction and mitigation

The future of forest C cycling (Song *et al* 2019) will shape trends in atmospheric CO₂ and the course of climate change (Schimel *et al* 2015). Our findings, and more generally the data contained in *ForC* and summarized here, can help to meet two major challenges.

First, improved representation of forest C cycling in models is essential to improving predictions of the future course of climate change, for the simple reason that by definition future projections extend our existing observations and understanding to conditions that do not currently exist on Earth (McDowell *et al* 2018, Bonan and Doney 2018, Gustafson *et al* 2018). To ensure that models are giving the right answers for the right reasons (Sulman *et al* 2018), it is important to benchmark against multiple components of the C cycle that are internally consistent with each other (Collier *et al* 2018, Wang *et al* 2018). *ForC*'s tens of thousands of records are readily available in a standardized format, along with all code used in the analyses presented here, and we recommend that researchers use these resources to identify and summarize data specific to the analysis at hand. Integration of *ForC* with models will be valuable to improving the accuracy and reliability of models (Fer *et al* 2021).

Second, *ForC* can serve as a pipeline through which information can flow efficiently from forest researchers to decision-makers working to implement forest conservation strategies at global, national, or landscape scales. This is already happening: *ForC* has contributed to updating the IPCC guidelines for carbon accounting in forests (IPCC 2019, Requena Suarez *et al* 2019), mapping C accumulation potential from natural forest regrowth globally (Cook-Patton *et al* 2020), and informing ecosystem conservation priorities (Goldstein *et al* 2020).

It is also interesting to consider the complementary utility of global-scale but spatially discontinuous databases such as *ForC* and wall-to-wall remote sensing products. The latter provide insight into aboveground carbon stocks, but less constraint on belowground stocks or carbon fluxes in general (Bond-Lamberty *et al* 2016, Anav *et al* 2015). Combining observational data and remote observations may provide a much more comprehensive and accurate picture of global forest C cycling, particularly when used in formal data assimilation systems (Konings *et al* 2019, Liu *et al* 2018). Biomass is the largest C stock in most forests, and most of the emphasis has traditionally been on this variable. Remote-sensing driven aboveground biomass estimates (e.g., Saatchi *et al* 2011), calibrated based on high-quality ground-based data (Schepaschenko *et al* 2019, Chave *et al* 2019), provide the most promising approach, but significant uncertainties remain (Ploton *et al* 2020). Note, however, that factors such as stand age and disturbance history are difficult, if not impossible, to detect remotely, and can only be characterized for very recent decades (Hansen *et al* 2013, Song *et al* 2018, Curtis *et al* 2018). Ground-based data such as *ForC* are therefore valuable in defining age-based trajectories in biomass, as in Cook-Patton *et al* (2020), and thus constraining variables such as carbon sink potential (Luyssaert *et al* 2008).

In contrast, carbon allocation within forest ecosystems and respiration fluxes cannot be remotely sensed. Efforts such as the Global Carbon Project (Friedlingstein *et al* 2019) and NASA's Carbon Monitoring System (Liu *et al* 2018) typically compute respiration as residuals of all other terms (Bond-Lamberty *et al* 2016, Harmon *et al* 2011). This means that the errors on respiration outputs are likely to be large and

certainly poorly constrained, offering a unique opportunity for databases such as ForC and SRDB (Jian *et al* 2020) to provide observational benchmarks. For example, Konings *et al* (2019) produced a unique top-down estimate of global heterotrophic respiration that can both be compared with extant bottom-up estimates (Bond-Lamberty 2018) and used as an internal consistency check on other parts of the carbon cycle (Phillips *et al* 2017).

Conclusions

As climate change accelerates, understanding and managing the carbon dynamics of forests— including dynamics and fluxes that cannot be observed by satellites—is critical to forecasting, mitigation, and adaptation. The C data in *ForC*, as summarized here, will be valuable to these efforts. Notably, the fact that tropical forests tend to have both the highest rates of C sequestration in young stands (Fig. 8; Cook-Patton *et al* 2020), fueled by their generally high C flux rates (Table 1; Fig. 7), and the highest mean biomass (Fig. 8; Table 1; Hu *et al* 2016, Jian *et al* 2020) reinforces the concept that conservation and restoration of these forests is a priority for climate change mitigation, along with high-biomass old-growth temperate stands (Grassi *et al* 2017, Goldstein *et al* 2020). It is also important to note the trade-off in climate mitigation potential of restoration of young forests, with high rates of CO₂ sequestration (*NEP*; Cook-Patton *et al* 2020), versus conservation and management of mature forests, with low *NEP* but high C stocks that could not be recovered on a time scale relevant to climate change mitigation (Goldstein *et al* 2020). Generally speaking, the conservation of mature forests will yield greater climate benefits (Anderson-Teixeira and DeLucia 2011), but both approaches are critical to avoiding catastrophic climate change (IPCC 2018).

Acknowledgements

Thanks to all researchers whose data are included in *ForC* and this analysis, to Jennifer McGarvey and Ian McGregor for help with the database, and to Norbert Kunert for helpful discussion. Three anonymous reviewers provided helpful feedback. Funding sources included a Smithsonian Scholarly Studies grant to KAT and HML and a Smithsonian Working Land and Seascapes grant to KAT.

Data availability statement

The data that support the findings of this study are openly available. Materials required to fully reproduce these analyses, including data, R scripts, and image files, are archived in Zenodo (DOI: TBD). Data, scripts, and results presented here are also available through the open-access *ForC* GitHub repository (<https://github.com/forc-db/ForC>), where many will be updated as the database develops.

References

- Allen A, Brown J and Gillooly J 2002 Global biodiversity, biochemical kinetics, and the energetic-equivalence rule *SCIENCE* **297** 1545–8
- Amiro B, Barr A, Barr J, Black T, Bracho R, Brown M, Chen J, Clark K, Davis K, Desai A, Dore S, Engel V, Fuentes J, Goldstein A, Goulden M, Kolb T, Lavigne M, Law B, Margolis H, Martin T, McCaughey J, Misson L, Montes-Helu M, Noormets A, Randerson J, Starr G and Xiao J 2010 Ecosystem carbon dioxide fluxes after disturbance in forests of North America *J. Geophys. Res.* **115** G00K02

607 Anav A, Friedlingstein P, Beer C, Ciais P, Harper A, Jones C, Murray-Tortarolo G, Papale D, Parazoo N C,
 608 Peylin P, Piao S, Sitch S, Viovy N, Wiltshire A and Zhao M 2015 Spatiotemporal patterns of terrestrial gross
 609 primary production: A review *Reviews of Geophysics* **53** 785–818

610 Andela N, Morton D C, Giglio L, Chen Y, van der Werf G R, Kasibhatla P S, DeFries R S, Collatz G J,
 611 Hantson S, Kloster S, Bachelet D, Forrest M, Lasslop G, Li F, Mangeon S, Melton J R, Yue C and
 612 Randerson J T 2017 A human-driven decline in global burned area *Science* **356** 1356–62

613 Anderson K, Allen A, Gillooly J and Brown J 2006 Temperature-dependence of biomass accumulation rates
 614 during secondary succession *Ecology Letters* **9** 673–82

615 Anderson-Teixeira K, Herrmann V, CookPatton, Ferson A and Lister K 2020 Forc-db/GROA: Release with
 616 Cook-Patton et al. 2020, Nature.

617 Anderson-Teixeira K J, Davies S J, Bennett A C, Gonzalez-Akre E B, Muller-Landau H C, Joseph Wright S,
 618 Abu Salim K, Almeyda Zambrano A M, Alonso A, Baltzer J L, Basset Y, Bourg N A, Broadbent E N,
 619 Brockelman W Y, Bunyavejchewin S, Burslem D F R P, Butt N, Cao M, Cardenas D, Chuyong G B, Clay K,
 620 Cordell S, Dattaraja H S, Deng X, Detto M, Du X, Duque A, Erikson D L, Ewango C E, Fischer G A,
 621 Fletcher C, Foster R B, Giardina C P, Gilbert G S, Gunatilleke N, Gunatilleke S, Hao Z, Hargrove W W,
 622 Hart T B, Hau B C, He F, Hoffman F M, Howe R W, Hubbell S P, Inman-Narahari F M, Jansen P A, Jiang
 623 M, Johnson D J, Kanzaki M, Kassim A R, Kenfack D, Kibet S, Kinnaird M F, Korte L, Kral K, Kumar J,
 624 Larson A J, Li Y, Li X, Liu S, Lum S K, Lutz J A, Ma K, Maddalena D M, Makana J-R, Malhi Y, Marthews
 625 T, Mat Serudin R, McMahon S M, McShea W J, Memiaghe H R, Mi X, Mizuno T, Morecroft M, Myers J A,
 626 Novotny V, de Oliveira A A, Ong P S, Orwig D A, Ostertag R, den Ouden J, Parker G G, Phillips R P, Sack
 627 L, Sainge M N, Sang W, Sri-ngernyuang K, Sukumar R, Sun I-F, Sungpalee W, Suresh H S, Tan S, Thomas
 628 S C, Thomas D W, Thompson J, Turner B L, Uriarte M, Valencia R, et al 2015 CTFs-ForestGEO : A
 629 worldwide network monitoring forests in an era of global change *Global Change Biology* **21** 528–49

630 Anderson-Teixeira K J, Delong J P, Fox A M, Brese D A and Litvak M E 2011 Differential responses of
 631 production and respiration to temperature and moisture drive the carbon balance across a climatic gradient
 632 in New Mexico *Global Change Biology* **17** 410–24

633 Anderson-Teixeira K J and DeLucia E H 2011 The greenhouse gas value of ecosystems *Global Change Biology*
 634 **17** 425–38

635 Anderson-Teixeira K J, Miller A D, Mohan J E, Hudiburg T W, Duval B D and DeLucia E H 2013 Altered
 636 dynamics of forest recovery under a changing climate *Global Change Biology* **19** 2001–21

637 Anderson-Teixeira K J, Wang M M H, McGarvey J C, Herrmann V, Tepley A J, Bond-Lamberty B and
 638 LeBauer D S 2018 ForC : A global database of forest carbon stocks and fluxes *Ecology* **99** 1507–7

639 Anderson-Teixeira K J, Wang M M H, McGarvey J C and LeBauer D S 2016 Carbon dynamics of mature
 640 and regrowth tropical forests derived from a pantropical database (TropForC-db) *Global Change Biology* **22**
 641 1690–709

642 Badgley G, Anderegg L D L, Berry J A and Field C B 2019 Terrestrial gross primary production: Using
 643 NIRV to scale from site to globe *Global Change Biology* **25** 3731–40

644 Baldocchi D, Falge E, Gu L, Olson R, Hollinger D, Running S, Anthoni P, Bernhofer C, Davis K, Evans R,
 645 Fuentes J, Goldstein A, Katul G, Law B, Lee X, Malhi Y, Meyers T, Munger W, Oechel W, Paw K T,

646 Pilegaard K, Schmid H P, Valentini R, Verma S, Vesala T, Wilson K and Wofsy S 2001 FLUXNET : A New
647 Tool to Study the Temporal and Spatial Variability of EcosystemScale Carbon Dioxide, Water Vapor, and
648 Energy Flux Densities *Bulletin of the American Meteorological Society* **82** 2415–34

649 Banbury Morgan B, Herrmann V, Kunert N, Bond-Lamberty B, Muller-Landau H C and Anderson-Teixeira
650 K Global patterns of forest autotrophic carbon fluxes *Global Change Biology*

651 Bates D, Mächler M, Bolker B and Walker S 2015 Fitting Linear Mixed-Effects Models Using **Lme4** *Journal*
652 *of Statistical Software* **67**

653 Besnard S, Carvalhais N, Arain M A, Black A, de Bruin S, Buchmann N, Cescatti A, Chen J, Clevers J G P
654 W, Desai A R, Gough C M, Havrankova K, Herold M, Hörtnagl L, Jung M, Knohl A, Kruijt B, Krupkova L,
655 Law B E, Lindroth A, Noormets A, Rouspard O, Steinbrecher R, Varlagin A, Vincke C and Reichstein M
656 2018 Quantifying the effect of forest age in annual net forest carbon balance *Environmental Research Letters*
657 **13** 124018

658 Bonan G 2008 Forests and Climate Change: Forcings, Feedbacks, and the Climate Benefits of Forests *Science*
659 **320** 1444–9

660 Bonan G B and Doney S C 2018 Climate, ecosystems, and planetary futures: The challenge to predict life in
661 Earth system models *Science* **359**

662 Bonan G B, Lombardozzi D L, Wieder W R, Oleson K W, Lawrence D M, Hoffman F M and Collier N 2019
663 Model Structure and Climate Data Uncertainty in Historical Simulations of the Terrestrial Carbon Cycle
664 (18502014) *Global Biogeochemical Cycles* **33** 1310–26

665 Bond-Lamberty B 2018 New Techniques and Data for Understanding the Global Soil Respiration Flux
666 *Earth's Future* **6** 1176–80

667 Bond-Lamberty B, Epron D, Harden J, Harmon M E, Hoffman F, Kumar J, David McGuire A and Vargas R
668 2016 Estimating heterotrophic respiration at large scales: Challenges, approaches, and next steps *Ecosphere* **7**

669 Bond-Lamberty B and Thomson A 2010 A global database of soil respiration data *Biogeosciences* **7** 1915–26

670 Bond-Lamberty B, Wang C and Gower S 2004 Contribution of root respiration to soil surface CO₂ flux in a
671 boreal black spruce chronosequence *Tree Physiology* **24** 1387–95

672 Bonner M T L, Schmidt S and Shoo L P 2013 A meta-analytical global comparison of aboveground biomass
673 accumulation between tropical secondary forests and monoculture plantations *Forest Ecology and*
674 *Management* **291** 73–86

675 Carmona M R, Armesto J J, Aravena J C and Pérez C A 2002 Coarse woody debris biomass in successional
676 and primary temperate forests in Chiloé Island, Chile *Forest Ecology and Management* **164** 265–75

677 Cavaleri M A, Reed S C, Smith W K and Wood T E 2015 Urgent need for warming experiments in tropical
678 forests *Global Change Biology* **21** 2111–21

679 Chapin F, Woodwell G, Randerson J, Rastetter E, Lovett G, Baldocchi D, Clark D, Harmon M, Schimel D,
680 Valentini R, Wirth C, Aber J, Cole J, Goulden M, Harden J, Heimann M, Howarth R, Matson P, McGuire A,
681 Melillo J, Mooney H, Neff J, Houghton R, Pace M, Ryan M, Running S, Sala O, Schlesinger W and Schulze
682 E 2006 Reconciling Carbon-cycle Concepts, Terminology, and Methods *Ecosystems* **9** 1041–50

683 Chaplin-Kramer R, Ramler I, Sharp R, Haddad N, Gerber J, West P, Mandle L, Engstrom P, Baccini A, Sim
684 S, Mueller C and King H 2015 Degradation in carbon stocks near tropical forest edges *Nature*
685 *Communications* **6**

686 Chave J, Davies S J, Phillips O L, Lewis S L, Sist P, Schepaschenko D, Armston J, Baker T R, Coomes D,
687 Disney M, Duncanson L, Hérault B, Labrière N, Meyer V, Réjou-Méchain M, Scipal K and Saatchi S 2019
688 Ground Data are Essential for Biomass Remote Sensing Missions *Surveys in Geophysics*

689 Chave J, Réjou-Méchain M, Búrquez A, Chidumayo E, Colgan M S, Delitti W B, Duque A, Eid T, Fearnside
690 P M, Goodman R C, Henry M, Martínez-Yrizar A, Mugasha W A, Muller-Landau H C, Mencuccini M,
691 Nelson B W, Ngomanda A, Nogueira E M, Ortiz-Malavassi E, Péliissier R, Ploton P, Ryan C M, Saldarriaga
692 J G and Vieilledent G 2014 Improved allometric models to estimate the aboveground biomass of tropical
693 trees *Global Change Biology* n/a-a

694 Clark D A, Asao S, Fisher R, Reed S, Reich P B, Ryan M G, Wood T E and Yang X 2017 Field data to
695 benchmark the carbon-cycle models for tropical forests *Biogeosciences Discussions* 1–44

696 Clark D, Brown S, Kicklighter D, Chambers J, Thomlinson J and Ni J 2001 Measuring net primary
697 production in forests: Concepts and field methods *Ecological Applications* **11** 356–70

698 Collalti A, Ibrom A, Stockmarr A, Cescatti A, Alkama R, Fernández-Martínez M, Matteucci G, Sitch S,
699 Friedlingstein P, Ciais P, Goll D S, Nabel J E M S, Pongratz J, Arneth A, Haverd V and Prentice I C 2020
700 Forest production efficiency increases with growth temperature *Nature Communications* **11** 5322

701 Collier N, Hoffman F M, Lawrence D M, Keppel-Aleks G, Koven C D, Riley W J, Mu M and Randerson J T
702 2018 The International Land Model Benchmarking (ILAMB) System: Design, Theory, and Implementation
703 *Journal of Advances in Modeling Earth Systems* **10** 2731–54

704 Cook-Patton S C, Leavitt S M, Gibbs D, Harris N L, Lister K, Anderson-Teixeira K J, Briggs R D, Chazdon
705 R L, Crowther T W, Ellis P W, Griscom H P, Herrmann V, Holl K D, Houghton R A, Larrosa C, Lomax G,
706 Lucas R, Madsen P, Malhi Y, Paquette A, Parker J D, Paul K, Routh D, Roxburgh S, Saatchi S, van den
707 Hoogen J, Walker W S, Wheeler C E, Wood S A, Xu L and Griscom B W 2020 Mapping carbon
708 accumulation potential from global natural forest regrowth *Nature* **585** 545–50

709 Corman J R, Collins S L, Cook E M, Dong X, Gherardi L A, Grimm N B, Hale R L, Lin T, Ramos J,
710 Reichmann L G and Sala O E 2019 Foundations and Frontiers of Ecosystem Science: Legacy of a Classic
711 Paper (Odum 1969) *Ecosystems* **22** 1160–72

712 Curtis P G, Slay C M, Harris N L, Tyukavina A and Hansen M C 2018 Classifying drivers of global forest
713 loss *Science* **361** 1108–11

714 Curtis P S and Gough C M 2018 Forest aging, disturbance and the carbon cycle *New Phytologist*

715 Davies S J, Abiem I, Abu Salim K, Aguilar S, Allen D, Alonso A, Anderson-Teixeira K, Andrade A, Arellano
716 G, Ashton P S, Baker P J, Baker M E, Baltzer J L, Basset Y, Bissengou P, Bohlman S, Bourg N A,
717 Brockelman W Y, Bunyavechewin S, Burslem D F, Cao M, Cárdenas D, Chang L-W, Chang-Yang C-H,
718 Chao K-J, Chao W-C, Chapman H, Chen Y-Y, Chisholm R A, Chu C, Chuyong G, Clay K, Comita L S,
719 Condit R, Cordell S, Dattaraja H S, de Oliveira A A, den Ouden J, Detto M, Dick C, Du X, Duque Á,
720 Ediriweera S, Ellis E C, Obiang N L E, Esufali S, Ewango C E, Fernando E S, Filip J, Fischer G A, Foster R,
721 Giambelluca T, Giardina C, Gilbert G S, Gonzalez-Akre E, Gunatilleke I, Gunatilleke C, Hao Z, Hau B C,

722 He F, Ni H, Howe R W, Hubbell S P, Huth A, Inman-Narahari F, Itoh A, Janík D, Jansen P A, Jiang M,
723 Johnson D J, Jones F A, Kanzaki M, Kenfack D, Kiratiprayoon S, Král K, Krizel L, Lao S, Larson A J, Li Y,
724 Li X, Litton C M, Liu Y, Liu S, Lum S K, Luskin M S, Lutz J A, Luu H T, Ma K, Makana J-R, Malhi Y,
725 Martin A, McCarthy C, McMahon S M, McShea W J, Memiaghe H, Mi X, Mitre D, Mohamad M, et al 2021
726 ForestGEO: Understanding forest diversity and dynamics through a global observatory network *Biological*
727 *Conservation* **253** 108907

728 DeLucia E, Drake J, Thomas R and Gonzalez-Meler M 2007 Forest carbon use efficiency: Is respiration a
729 constant fraction of gross primary production? *Global Change Biology* **13** 1157–67

730 Di Vittorio A, Shi X, Bond-Lamberty B, Calvin K and Jones A 2020 Initial Land Use/Cover Distribution
731 Substantially Affects Global Carbon and Local Temperature Projections in the Integrated Earth System
732 Model *Global Biogeochemical Cycles* **34**

733 FAO 2010 *Global Forest Resources Assessment 2010* (Rome, Italy: Food and Agriculture Organization of the
734 United Nations)

735 FAO and UNEP 2020 *The State of the World's Forests 2020: Forests, biodiversity and people* (Rome, Italy:
736 FAO and UNEP)

737 Fer I, Gardella A K, Shiklomanov A N, Campbell E E, Cowdery E M, Kauwe M G D, Desai A, Duveneck M
738 J, Fisher J B, Haynes K D, Hoffman F M, Johnston M R, Kooper R, LeBauer D S, Mantooth J, Parton W J,
739 Poulter B, Quaife T, Raiho A, Schaefer K, Serbin S P, Simkins J, Wilcox K R, Viskari T and Dietze M C
740 2021 Beyond ecosystem modeling: A roadmap to community cyberinfrastructure for ecological data-model
741 integration *Global Change Biology* **27** 13–26

742 Friedlingstein P, Cox P, Betts R, Bopp L, von Bloh W, Brovkin V, Cadule P, Doney S, Eby M, Fung I, Bala
743 G, John J, Jones C, Joos F, Kato T, Kawamiya M, Knorr W, Lindsay K, Matthews H D, Raddatz T, Rayner
744 P, Reick C, Roeckner E, Schnitzler K-G, Schnur R, Strassmann K, Weaver A J, Yoshikawa C and Zeng N
745 2006 ClimateCarbon Cycle Feedback Analysis: Results from the C4MIP Model Intercomparison *Journal of*
746 *Climate* **19** 3337–53

747 Friedlingstein P, Jones M W, O'Sullivan M, Andrew R M, Hauck J, Peters G P, Peters W, Pongratz J, Sitch
748 S, Quéré C L, Bakker D C E, Canadell J G, Ciais P, Jackson R B, Anthoni P, Barbero L, Bastos A,
749 Bastrikov V, Becker M, Bopp L, Buitenhuis E, Chandra N, Chevallier F, Chini L P, Currie K I, Feely R A,
750 Gehlen M, Gilfillan D, Gkritzalis T, Goll D S, Gruber N, Gutekunst S, Harris I, Haverd V, Houghton R A,
751 Hurtt G, Ilyina T, Jain A K, Joetzjer E, Kaplan J O, Kato E, Klein Goldewijk K, Korsbakken J I,
752 Landschützer P, Lauvset S K, Lefèvre N, Lenton A, Lienert S, Lombardozzi D, Marland G, McGuire P C,
753 Melton J R, Metzl N, Munro D R, Nabel J E M S, Nakaoka S-I, Neill C, Omar A M, Ono T, Peregon A,
754 Pierrot D, Poulter B, Rehder G, Resplandy L, Robertson E, Rödenbeck C, Séférian R, Schwinger J, Smith N,
755 Tans P P, Tian H, Tilbrook B, Tubiello F N, Werf G R van der, Wiltshire A J and Zaehle S 2019 Global
756 Carbon Budget 2019 *Earth System Science Data* **11** 1783–838

757 Gillman L N, Wright S D, Cusens J, McBride P D, Malhi Y and Whittaker R J 2015 Latitude, productivity
758 and species richness *Global Ecology and Biogeography* **24** 107–17

759 Goldstein A, Turner W R, Spawn S A, Anderson-Teixeira K J, Cook-Patton S, Fargione J, Gibbs H K,
760 Griscom B, Hewson J H, Howard J F, Ledezma J C, Page S, Koh L P, Rockström J, Sanderman J and Hole
761 D G 2020 Protecting irrecoverable carbon in Earth's ecosystems *Nature Climate Change* **10** 287–95

Goulden M, McMillan A, Winston G, Rocha A, Manies K, Harden J and Bond-Lamberty B 2011 Patterns of NPP, GPP, respiration, and NEP during boreal forest succession *Global Change Biology* **17** 855–71

Grassi G, House J, Dentener F, Federici S, den Elzen M and Penman J 2017 The key role of forests in meeting climate targets requires science for credible mitigation *Nature Climate Change* **7** 220–6

Griscom B W, Adams J, Ellis P W, Houghton R A, Lomax G, Miteva D A, Schlesinger W H, Shoch D, Siikamäki J V, Smith P, Woodbury P, Zganjar C, Blackman A, Campari J, Conant R T, Delgado C, Elias P, Gopalakrishna T, Hamsik M R, Herrero M, Kiesecker J, Landis E, Laestadius L, Leavitt S M, Minnemeyer S, Polasky S, Potapov P, Putz F E, Sanderman J, Silvius M, Wollenberg E and Fargione J 2017 Natural climate solutions *Proceedings of the National Academy of Sciences* **114** 11645–50

Gustafson E J, Kubiske M E, Miranda B R, Hoshika Y and Paoletti E 2018 Extrapolating plot-scale CO₂ and ozone enrichment experimental results to novel conditions and scales using mechanistic modeling *Ecological Processes* **7** 31

Hansen M C, Potapov P V, Moore R, Hancher M, Turubanova S A, Tyukavina A, Thau D, Stehman S V, Goetz S J, Loveland T R, Kommareddy A, Egorov A, Chini L, Justice C O and Townshend J R G 2013 High-Resolution Global Maps of 21st-Century Forest Cover Change *Science* **342** 850–3

Harmon M E, Bond-Lamberty B, Tang J and Vargas R 2011 Heterotrophic respiration in disturbed forests: A review with examples from North America *Journal of Geophysical Research* **116**

Harmon M E, Franklin J F, Swanson F J, Sollins P, Gregory S V, Lattin J D, Anderson N H, Cline S P, Aumen N G, Sedell J R, Lienkaemper G W, Cromack K and Cummins K W 1986 Ecology of Coarse Woody Debris in Temperate Ecosystems *Advances in Ecological Research* vol 15, ed A MacFadyen and E D Ford (Academic Press) pp 133–302

Harris N L, Gibbs D A, Baccini A, Birdsey R A, Bruin S de, Farina M, Fatoyinbo L, Hansen M C, Herold M, Houghton R A, Potapov P V, Suarez D R, Roman-Cuesta R M, Saatchi S S, Slay C M, Turubanova S A and Tyukavina A 2021 Global maps of twenty-first century forest carbon fluxes *Nature Climate Change* 1–7

Holdridge L R 1947 Determination of World Plant Formations From Simple Climatic Data *Science* **105** 367–8

Houghton R A 2020 Terrestrial fluxes of carbon in GCP carbon budgets *Global Change Biology* **26** 3006–14

Hu T, Su Y, Xue B, Liu J, Zhao X, Fang J and Guo Q 2016 Mapping Global Forest Aboveground Biomass with Spaceborne LiDAR, Optical Imagery, and Forest Inventory Data *Remote Sensing* **8** 565

Huang M and Asner G P 2010 Long-term carbon loss and recovery following selective logging in Amazon forests *Global Biogeochemical Cycles* **24**

Humboldt A von and Bonpland A 1807 *Essay on the Geography of Plants*

Hursh A, Ballantyne A, Cooper L, Maneta M, Kimball J and Watts J 2017 The sensitivity of soil respiration to soil temperature, moisture, and carbon supply at the global scale *Global Change Biology* **23** 2090–103

IPCC 2019 *2019 Refinement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories*

IPCC 2018 *Global Warming of 1.5°C. An IPCC Special Report on the impacts of global warming of 1.5°C above pre-industrial levels and related global greenhouse gas emission pathways, in the context of strengthening the global response to the threat of climate change, sustainable development, and efforts to*

- eradicate poverty [Masson-Delmotte, V., P. Zhai, H.-O. Pörtner, D. Roberts, J. Skea, P.R. Shukla, A. Pirani, W. Moufouma-Okia, C. Péan, R. Pidcock, S. Connors, J.B.R. Matthews, Y. Chen, X. Zhou, M.I. Gomis, E. Lonnoy, T. Maycock, M. Tignor, and T. Waterfield (eds.)].
- Jian J, Vargas R, Anderson-Teixeira K, Stell E, Herrmann V, Horn M, Kholod N, Manzon J, Marchesi R, Paredes D and Bond-Lamberty B 2020 *A restructured and updated global soil respiration database (SRDB-V5)* (Data, Algorithms, and Models)
- Johnson C, Zarin D and Johnson A 2000 Post-disturbance aboveground biomass accumulation in global secondary forests *Ecology* **81** 1395–401
- Johnson D J, Needham J, Xu C, Massoud E C, Davies S J, Anderson-Teixeira K J, Bunyavejchewin S, Chambers J Q, Chang-Yang C-H, Chiang J-M, Chuyong G B, Condit R, Cordell S, Fletcher C, Giardina C P, Giambelluca T W, Gunatilleke N, Gunatilleke S, Hsieh C-F, Hubbell S, Inman-Narahari F, Kassim A R, Katabuchi M, Kenfack D, Litton C M, Lum S, Mohamad M, Nasardin M, Ong P S, Ostertag R, Sack L, Swenson N G, Sun I F, Tan S, Thomas D W, Thompson J, Umaña M N, Uriarte M, Valencia R, Yap S, Zimmerman J, McDowell N G and McMahon S M 2018 Climate sensitive size-dependent survival in tropical trees *Nature Ecology & Evolution* **1**
- Jung M, Henkel K, Herold M and Churkina G 2006 Exploiting synergies of global land cover products for carbon cycle modeling *Remote Sensing of Environment* **101** 534–53
- Keith H, Mackey B and Lindenmayer D 2009 Re-evaluation of forest biomass carbon stocks and lessons from the world’s most carbon-dense forests *Proceedings of the National Academy of Sciences* **106** 11635–40
- Konings A G, Bloom A A, Liu J, Parazoo N C, Schimel D S and Bowman K W 2019 Global satellite-driven estimates of heterotrophic respiration *Biogeosciences* **16** 2269–84
- Köchy M, Hiederer R and Freibauer A 2015 Global distribution of soil organic carbon 1: Masses and frequency distributions of SOC stocks for the tropics, permafrost regions, wetlands, and the world *SOIL* **1** 351–65
- Krause A, Pugh T A M, Bayer A D, Li W, Leung F, Bondeau A, Doelman J C, Humpenöder F, Anthoni P, BDIRSKY B L, Ciais P, Müller C, Murray-Tortarolo G, Olin S, Popp A, Sitch S, Stehfest E and Arneth A 2018 Large uncertainty in carbon uptake potential of land-based climate-change mitigation efforts *Global Change Biology* **24** 3025–38
- Kuzyakov Y 2006 Sources of CO₂ efflux from soil and review of partitioning methods *Soil Biology and Biochemistry* **38** 425–48
- Law B, Sun O, Campbell J, Tuyl S and Thornton P 2003 Changes in carbon storage and fluxes in a chronosequence of ponderosa pine *Global Change Biology* **9** 510–24
- Li X and Xiao J 2019 Mapping Photosynthesis Solely from Solar-Induced Chlorophyll Fluorescence: A Global, Fine-Resolution Dataset of Gross Primary Production Derived from OCO-2 *Remote Sensing* **11** 2563
- Lichstein J W, Wirth C, Horn H S and Pacala S W 2009 Biomass Chronosequences of United States Forests: Implications for Carbon Storage and Forest Management *Old-Growth Forests* Ecological Studies ed C Wirth, G Gleixner and M Heimann (Springer Berlin Heidelberg) pp 301–41
- Lieth H 1973 Primary production: Terrestrial ecosystems *Human Ecology* **1** 303–32

- 837 Liu J, Bowman K, Parazoo N C, Bloom A A, Wunch D, Jiang Z, Gurney K R and Schimel D 2018 Detecting
838 drought impact on terrestrial biosphere carbon fluxes over contiguous US with satellite observations
839 *Environmental Research Letters* **13** 095003
- 840 Luo Y Q, Randerson J T, Abramowitz G, Bacour C, Blyth E, Carvalhais N, Ciais P, Dalmonech D, Fisher J
841 B, Fisher R, Friedlingstein P, Hibbard K, Hoffman F, Huntzinger D, Jones C D, Koven C, Lawrence D, Li D
842 J, Mahecha M, Niu S L, Norby R, Piao S L, Qi X, Peylin P, Prentice I C, Riley W, Reichstein M, Schwalm
843 C, Wang Y P, Xia J Y, Zaehle S and Zhou X H 2012 A framework for benchmarking land models
844 *Biogeosciences* **9** 3857–74
- 845 Lutz J A, Furniss T J, Johnson D J, Davies S J, Allen D, Alonso A, Anderson-Teixeira K J, Andrade A,
846 Baltzer J, Becker K M L, Blomdahl E M, Bourg N A, Bunyavejchewin S, Burslem D F R P, Cansler C A,
847 Cao K, Cao M, Cárdenas D, Chang L-W, Chao K-J, Chao W-C, Chiang J-M, Chu C, Chuyong G B, Clay K,
848 Condit R, Cordell S, Dattaraja H S, Duque A, Ewango C E N, Fischer G A, Fletcher C, Freund J A,
849 Giardina C, Germain S J, Gilbert G S, Hao Z, Hart T, Hau B C H, He F, Hector A, Howe R W, Hsieh C-F,
850 Hu Y-H, Hubbell S P, Inman-Narahari F M, Itoh A, Janík D, Kassim A R, Kenfack D, Korte L, Král K,
851 Larson A J, Li Y, Lin Y, Liu S, Lum S, Ma K, Makana J-R, Malhi Y, McMahon S M, McShea W J,
852 Memiaghe H R, Mi X, Morecroft M, Musili P M, Myers J A, Novotny V, Oliveira A de, Ong P, Orwig D A,
853 Ostertag R, Parker G G, Patankar R, Phillips R P, Reynolds G, Sack L, Song G-Z M, Su S-H, Sukumar R,
854 Sun I-F, Suresh H S, Swanson M E, Tan S, Thomas D W, Thompson J, Uriarte M, Valencia R, Vicentini A,
855 Vrška T, Wang X, Weiblen G D, Wolf A, Wu S-H, Xu H, Yamakura T, Yap S and Zimmerman J K 2018
856 Global importance of large-diameter trees *Global Ecology and Biogeography* **27** 849–64
- 857 Luyssaert S, Inglis I, Jung M, Richardson A D, Reichstein M, Papale D, Piao S L, Schulze E-D, Wingate L,
858 Matteucci G, Aragao L, Aubinet M, Beer C, Bernhofer C, Black K G, Bonal D, Bonnefond J-M, Chambers J,
859 Ciais P, Cook B, Davis K J, Dolman A J, Gielen B, Goulden M, Grace J, Granier A, Grelle A, Griffis T,
860 Grünwald T, Guidolotti G, Hanson P J, Harding R, Hollinger D Y, Huttyra L R, Kolari P, Kruijt B, Kutsch
861 W, Lagergren F, Laurila T, Law B E, Maire G L, Lindroth A, Loustau D, Malhi Y, Mateus J, Migliavacca M,
862 Misson L, Montagnani L, Moncrieff J, Moors E, Munger J W, Nikinmaa E, Ollinger S V, Pita G, Rebmann
863 C, Roupsard O, Saigusa N, Sanz M J, Seufert G, Sierra C, Smith M-L, Tang J, Valentini R, Vesala T and
864 Janssens I A 2007 CO₂ balance of boreal, temperate, and tropical forests derived from a global database
865 *Global Change Biology* **13** 2509–37
- 866 Luyssaert S, Schulze E, Borner A, Knohl A, Hessenmoller D, Law B, Ciais P and Grace J 2008 Old-growth
867 forests as global carbon sinks *Nature* **455** 213
- 868 Magnani F, Mencuccini M, Borghetti M, Berbigier P, Berninger F, Delzon S, Grelle A, Hari P, Jarvis P,
869 Kolari P, Kowalski A, Lankreijer H, Law B, Lindroth A, Loustau D, Manca G, Moncrieff J, Rayment M,
870 Tedeschi V, Valentini R and Grace J 2007 The human footprint in the carbon cycle of temperate and boreal
871 forests *Nature* **447** 849–51
- 872 Martin L J, Blossey B and Ellis E 2012 Mapping where ecologists work: Biases in the global distribution of
873 terrestrial ecological observations *Frontiers in Ecology and the Environment* **10** 195–201
- 874 Martin P A, Newton A C and Bullock J M 2013 Carbon pools recover more quickly than plant biodiversity
875 in tropical secondary forests *Proceedings of the Royal Society B: Biological Sciences* **280** 20132236–6
- 876 Maurer G E, Chan A M, Trahan N A, Moore D J P and Bowling D R 2016 Carbon isotopic composition of

877 forest soil respiration in the decade following bark beetle and stem girdling disturbances in the Rocky
878 Mountains *Plant, Cell & Environment* **39** 1513–23

879 McDowell N G, Allen C D, Anderson-Teixeira K, Aukema B H, Bond-Lamberty B, Chini L, Clark J S,
880 Dietze M, Grossiord C, Hanbury-Brown A, Hurr G C, Jackson R B, Johnson D J, Kueppers L, Lichstein J
881 W, Ogle K, Poulter B, Pugh T A M, Seidl R, Turner M G, Uriarte M, Walker A P and Xu C 2020 Pervasive
882 shifts in forest dynamics in a changing world *Science* **368**

883 McDowell N G, Michaletz S T, Bennett K E, Solander K C, Xu C, Maxwell R M and Middleton R S 2018
884 Predicting Chronic Climate-Driven Disturbances and Their Mitigation *Trends in Ecology & Evolution* **33**
885 15–27

886 McGarvey J C, Thompson J R, Epstein H E and Shugart H H 2014 Carbon storage in old-growth forests of
887 the Mid-Atlantic: Toward better understanding the eastern forest carbon sink *Ecology* **96** 311–7

888 Novick K A, Biederman J A, Desai A R, Litvak M E, Moore D J P, Scott R L and Torn M S 2018 The
889 AmeriFlux network: A coalition of the willing *Agricultural and Forest Meteorology* **249** 444–56

890 Odum E 1969 The strategy of ecosystem development *Science* **164** 262–70

891 Ordway E M and Asner G P 2020 Carbon declines along tropical forest edges correspond to heterogeneous
892 effects on canopy structure and function *Proceedings of the National Academy of Sciences* **117** 7863–70

893 Pan Y, Birdsey R, Fang J, Houghton R, Kauppi P, Kurz W, Phillips O, Shvidenko A, Lewis S, Canadell J,
894 Ciais P, Jackson R, Pacala S, McGuire A, Piao S, Rautiainen A, Sitch S and Hayes D 2011 A Large and
895 Persistent Carbon Sink in the World's Forests *Science* **333** 988–93

896 Pastorello G, Trotta C, Canfora E, Chu H, Christianson D, Cheah Y-W, Poindexter C, Chen J, Elbashandy
897 A, Humphrey M, Isaac P, Polidori D, Ribeca A, van Ingen C, Zhang L, Amiro B, Ammann C, Arain M A,
898 Ardö J, Arkebauer T, Arndt S K, Arriga N, Aubinet M, Aurela M, Baldocchi D, Barr A, Beamesderfer E,
899 Marchesini L B, Bergeron O, Beringer J, Bernhofer C, Berveiller D, Billesbach D, Black T A, Blanken P D,
900 Bohrer G, Boike J, Bolstad P V, Bonal D, Bonnefond J-M, Bowling D R, Bracho R, Brodeur J, Brümmer C,
901 Buchmann N, Burban B, Burns S P, Buysse P, Cale P, Cavagna M, Cellier P, Chen S, Chini I, Christensen T
902 R, Cleverly J, Collalti A, Consalvo C, Cook B D, Cook D, Coursolle C, Cremonese E, Curtis P S, D'Andrea
903 E, da Rocha H, Dai X, Davis K J, De Cinti B, de Grandcourt A, De Ligne A, De Oliveira R C, Delpierre N,
904 Desai A R, Di Bella C M, di Tommasi P, Dolman H, Domingo F, Dong G, Dore S, Duce P, Dufrêne E, Dunn
905 A, Dušek J, Eamus D, Eichelmann U, ElKhidir H A M, Eugster W, Ewenz C M, Ewers B, Famulari D, Fares
906 S, Feigenwinter I, Feitz A, Fensholt R, Filippa G, Fischer M, Frank J, Galvagno M, Gharun M, et al 2020 The
907 FLUXNET2015 dataset and the ONEFlux processing pipeline for eddy covariance data *Scientific Data* **7** 225

908 Phillips C L, Bond-Lamberty B, Desai A R, Lavoie M, Risk D, Tang J, Todd-Brown K and Vargas R 2017
909 The value of soil respiration measurements for interpreting and modeling terrestrial carbon cycling *Plant and*
910 *Soil* **413** 1–25

911 Piconiot C, Rödig E, Putz F E, Rutishauser E, Sist P, Ascarrunz N, Blanc L, Derroire G, Descroix L,
912 Guedes M C, Coronado E H, Huth A, Kanashiro M, Licona J C, Mazzei L, d'Oliveira M V N, Peña-Claros
913 M, Rodney K, Shenkin A, Souza C R de, Vidal E, West T A P, Wortel V and Hérault B 2019 Can timber
914 provision from Amazonian production forests be sustainable? *Environmental Research Letters* **14** 064014

915 Piconiot C, Sist P, Mazzei L, Peña-Claros M, Putz F E, Rutishauser E, Shenkin A, Ascarrunz N, de Azevedo

916 C P, Baraloto C, França M, Guedes M, Honorio Coronado E N, d'Oliveira M V, Ruschel A R, da Silva K E,
917 Doff Sotta E, de Souza C R, Vidal E, West T A and Hérault B 2016 Carbon recovery dynamics following
918 disturbance by selective logging in Amazonian forests *eLife* **5** e21394

919 Ploton P, Mortier F, Réjou-Méchain M, Barbier N, Picard N, Rossi V, Dormann C, Cornu G, Viennois G,
920 Bayol N, Lyapustin A, Gourlet-Fleury S and Péliissier R 2020 Spatial validation reveals poor predictive
921 performance of large-scale ecological mapping models *Nature Communications* **11** 4540

922 Poorter L, Bongers F, Aide T M, Zambrano A M A, Balvanera P, Becknell J M, Boukili V, Brancalion P H
923 S, Broadbent E N, Chazdon R L, Craven D, Almeida-Cortez J S de, Cabral G A L, Jong B H J de, Denslow
924 J S, Dent D H, DeWalt S J, Dupuy J M, Durán S M, Espírito-Santo M M, Fandino M C, César R G, Hall J
925 S, Hernandez-Stefanoni J L, Jakovac C C, Junqueira A B, Kennard D, Letcher S G, Licona J-C, Lohbeck M,
926 Marín-Spiotta E, Martínez-Ramos M, Massoca P, Meave J A, Mesquita R, Mora F, Muñoz R, Muscarella R,
927 Nunes Y R F, Ochoa-Gaona S, Oliveira A A de, Orihuela-Belmonte E, Peña-Claros M, Pérez-García E A,
928 Piotto D, Powers J S, Rodríguez-Velázquez J, Romero-Pérez I E, Ruíz J, Saldarriaga J G, Sanchez-Azofeifa
929 A, Schwartz N B, Steininger M K, Swenson N G, Toledo M, Uriarte M, Breugel M van, Wal H van der,
930 Veloso M D M, Vester H F M, Vicentini A, Vieira I C G, Bents T V, Williamson G B and Rozendaal D M
931 A 2016 Biomass resilience of Neotropical secondary forests *Nature* **530** 211–4

932 Pregitzer K and Euskirchen E 2004 Carbon cycling and storage in world forests: Biome patterns related to
933 forest age *Global Change Biology* **10** 2052–77

934 Pugh T A M, Lindeskog M, Smith B, Poulter B, Arneth A, Haverd V and Calle L 2019 Role of forest
935 regrowth in global carbon sink dynamics *Proceedings of the National Academy of Sciences* **116** 4382–7

936 Reinmann A B and Hutrya L R 2017 Edge effects enhance carbon uptake and its vulnerability to climate
937 change in temperate broadleaf forests *Proceedings of the National Academy of Sciences* **114** 107–12

938 Reinmann A B, Smith I A, Thompson J R and Hutrya L R 2020 Urbanization and fragmentation mediate
939 temperate forest carbon cycle response to climate *Environmental Research Letters* **15** 114036

940 Remy E, Wuyts K, Boeckx P, Ginzburg S, Gundersen P, Demey A, Van Den Bulcke J, Van Acker J and
941 Verheyen K 2016 Strong gradients in nitrogen and carbon stocks at temperate forest edges *Forest Ecology
942 and Management* **376** 45–58

943 Requena Suarez D, Rozendaal D M A, Sy V D, Phillips O L, Alvarez-Dávila E, Anderson-Teixeira K,
944 Araujo-Murakami A, Arroyo L, Baker T R, Bongers F, Brienen R J W, Carter S, Cook-Patton S C,
945 Feldpausch T R, Griscom B W, Harris N, Hérault B, Coronado E N H, Leavitt S M, Lewis S L, Marimon B
946 S, Mendoza A M, N'dja J K, N'Guessan A E, Poorter L, Qie L, Rutishauser E, Sist P, Sonké B, Sullivan M J
947 P, Vilanova E, Wang M M H, Martius C and Herold M 2019 Estimating aboveground net biomass change for
948 tropical and subtropical forests: Refinement of IPCC default rates using forest plot data *Global Change
949 Biology* **25** 3609–24

950 Ribeiro-Kumara C, Köster E, Aaltonen H and Köster K 2020 How do forest fires affect soil greenhouse gas
951 emissions in upland boreal forests? A review *Environmental Research* **184** 109328

952 Saatchi S S, Harris N L, Brown S, Lefsky M, Mitchard E T A, Salas W, Zutta B R, Buermann W, Lewis S L,
953 Hagen S, Petrova S, White L, Silman M and Morel A 2011 Benchmark map of forest carbon stocks in
954 tropical regions across three continents *Proceedings of the National Academy of Sciences* **108** 9899–904

955 Schepaschenko D, Chave J, Phillips O L, Lewis S L, Davies S J, Réjou-Méchain M, Sist P, Scipal K, Perger
 956 C, Herault B, Labrière N, Hofhansl F, Affum-Baffoe K, Aleinikov A, Alonso A, Amani C, Araujo-Murakami
 957 A, Armston J, Arroyo L, Ascarrunz N, Azevedo C, Baker T, Balazy R, Bedeau C, Berry N, Bilous A M,
 958 Bilous S Y, Bissengou P, Blanc L, Bobkova K S, Braslavskaya T, Brien R, Burslem D F R P, Condit R,
 959 Cuni-Sanchez A, Danilina D, Torres D del C, Derroire G, Descroix L, Sotta E D, d'Oliveira M V N, Dresel C,
 960 Erwin T, Evdokimenko M D, Falck J, Feldpausch T R, Foli E G, Foster R, Fritz S, Garcia-Abril A D,
 961 Gornov A, Gornova M, Gothard-Bassébé E, Gourlet-Fleury S, Guedes M, Hamer K C, Susanty F H, Higuchi
 962 N, Coronado E N H, Hubau W, Hubbell S, Ilstedt U, Ivanov V V, Kanashiro M, Karlsson A, Karminov V N,
 963 Killeen T, Koffi J-C K, Konovalova M, Kraxner F, Krejza J, Krisnawati H, Krivobokov L V, Kuznetsov M A,
 964 Lakyda I, Lakyda P I, Licona J C, Lucas R M, Lukina N, Lussetti D, Malhi Y, Manzanera J A, Marimon B,
 965 Junior B H M, Martinez R V, Martynenko O V, Matsala M, Matyashuk R K, Mazzei L, Memiaghe H,
 966 Mendoza C, Mendoza A M, Moroziuk O V, Mukhortova L, Musa S, Nazimova D I, Okuda T, Oliveira L C,
 967 et al 2019 The Forest Observation System, building a global reference dataset for remote sensing of forest
 968 biomass *Scientific Data* **6** 1–11

969 Schimel D, Hargrove W, Hoffman F and MacMahon J 2007 NEON: A hierarchically designed national
 970 ecological network *Frontiers in Ecology and the Environment* **5** 59–9

971 Schimel D, Stephens B B and Fisher J B 2015 Effect of increasing CO₂ on the terrestrial carbon cycle
 972 *Proceedings of the National Academy of Sciences* **112** 436–41

973 Sist P, Rutishauser E, Peña-Claros M, Shenkin A, Hérault B, Blanc L, Baraloto C, Baya F, Benedet F, Silva
 974 K E da, Descroix L, Ferreira J N, Gourlet-Fleury S, Guedes M C, Harun I B, Jalonen R, Kanashiro M,
 975 Krisnawati H, Kshatriya M, Lincoln P, Mazzei L, Medjibé V, Nasi R, d'Oliveira M V N, Oliveira L C de,
 976 Picard N, Pietsch S, Pinard M, Priyadi H, Putz F E, Rodney K, Rossi V, Roopsind A, Ruschel A R, Shari N
 977 H Z, Souza C R de, Susanty F H, Sotta E D, Toledo M, Vidal E, West T A P, Wortel V and Yamada T 2015
 978 The Tropical managed Forests Observatory: A research network addressing the future of tropical logged
 979 forests *Applied Vegetation Science* **18** 171–4

980 Smith I A, Hutryra L R, Reinmann A B, Thompson J R and Allen D W 2019 Evidence for Edge
 981 Enhancements of Soil Respiration in Temperate Forests *Geophysical Research Letters* **46** 4278–87

982 Smithwick E A H, Harmon M E, Remillard S M, Acker S A and Franklin J F 2002 Potential upper bounds of
 983 carbon stores in forests of the Pacific Northwest *Ecological Applications* **12** 1303–17

984 Song J, Wan S, Piao S, Knapp A K, Classen A T, Vicca S, Ciais P, Hovenden M J, Leuzinger S, Beier C,
 985 Kardol P, Xia J, Liu Q, Ru J, Zhou Z, Luo Y, Guo D, Adam Langley J, Zscheischler J, Dukes J S, Tang J,
 986 Chen J, Hofmockel K S, Kueppers L M, Rustad L, Liu L, Smith M D, Templer P H, Quinn Thomas R,
 987 Norby R J, Phillips R P, Niu S, Fatichi S, Wang Y, Shao P, Han H, Wang D, Lei L, Wang J, Li X, Zhang Q,
 988 Li X, Su F, Liu B, Yang F, Ma G, Li G, Liu Y, Liu Y, Yang Z, Zhang K, Miao Y, Hu M, Yan C, Zhang A,
 989 Zhong M, Hui Y, Li Y and Zheng M 2019 A meta-analysis of 1,119 manipulative experiments on terrestrial
 990 carbon-cycling responses to global change *Nature Ecology & Evolution* **3** 1309–20

991 Song X-P, Hansen M C, Stehman S V, Potapov P V, Tyukavina A, Vermote E F and Townshend J R 2018
 992 Global land change from 1982 to 2016 *Nature* **560** 639–43

993 Spawn S A, Sullivan C C, Lark T J and Gibbs H K 2020 Harmonized global maps of above and belowground
 994 biomass carbon density in the year 2010 *Scientific Data* **7** 112

- Stoy P C, Mauder M, Foken T, Marcolla B, Boegh E, Ibrom A, Arain M A, Arneth A, Aurela M, Bernhofer C, Cescatti A, Dellwik E, Duce P, Gianelle D, van Gorsel E, Kiely G, Knohl A, Margolis H, McCaughey H, Merbold L, Montagnani L, Papale D, Reichstein M, Saunders M, Serrano-Ortiz P, Sottocornola M, Spano D, Vaccari F and Varlagin A 2013 A data-driven analysis of energy balance closure across FLUXNET research sites: The role of landscape scale heterogeneity *Agricultural and Forest Meteorology* **171-172** 137–52
- Sulman B N, Moore J A M, Abramoff R, Averill C, Kivlin S, Georgiou K, Sridhar B, Hartman M D, Wang G, Wieder W R, Bradford M A, Luo Y, Mayes M A, Morrison E, Riley W J, Salazar A, Schimel J P, Tang J and Classen A T 2018 Multiple models and experiments underscore large uncertainty in soil carbon dynamics *Biogeochemistry* **141** 109–23
- Taylor P G, Cleveland C C, Wieder W R, Sullivan B W, Doughty C E, Dobrowski S Z and Townsend A R 2017 Temperature and rainfall interact to control carbon cycling in tropical forests ed L Liu *Ecology Letters* **20** 779–88
- Team R C 2020 R : A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria. URL <http://www.R-project.org/>.
- Tubiello F N, Pekkarinen A, Marklund L, Wanner N, Conchedda G, Federici S, Rossi S and Grassi G 2020 Carbon Emissions and Removals by Forests: New Estimates 1990–2020 *Earth System Science Data Discussions* 1–21
- van der Werf G R, Randerson J T, Giglio L, van Leeuwen T T, Chen Y, Rogers B M, Mu M, van Marle M J E, Morton D C, Collatz G J, Yokelson R J and Kasibhatla P S 2017 Global fire emissions estimates during 1997–2016 *Earth System Science Data* **9** 697–720
- Vargas R, Allen M F and Allen E B 2008 Biomass and carbon accumulation in a fire chronosequence of a seasonally dry tropical forest *Global Change Biology* **14** 109–24
- Wang Y, Ciais P, Goll D, Huang Y, Luo Y, Wang Y-P, Bloom A A, Broquet G, Hartmann J, Peng S, Penuelas J, Piao S, Sardans J, Stocker B D, Wang R, Zaehle S and Zechmeister-Boltenstern S 2018 GOLUM-CNP v1.0: A data-driven modeling of carbon, nitrogen and phosphorus cycles in major terrestrial biomes *Geoscientific Model Development* **11** 3903–28
- Warner D L, Bond-Lamberty B, Jian J, Stell E and Vargas R 2019 Spatial Predictions and Associated Uncertainty of Annual Soil Respiration at the Global Scale *Global Biogeochemical Cycles* **33** 1733–45
- Williams C A, Collatz G J, Masek J, Huang C and Goward S N 2014 Impacts of disturbance history on forest carbon stocks and fluxes: Merging satellite disturbance mapping with forest inventory data in a carbon cycle model framework *Remote Sensing of Environment* **151** 57–71
- Wilson R M, Hopple A M, Tfaily M M, Sebestyen S D, Schadt C W, Pfeifer-Meister L, Medvedeff C, McFarlane K J, Kostka J E, Kolton M, Kolka R K, Kluber L A, Keller J K, Guilderson T P, Griffiths N A, Chanton J P, Bridgman S D and Hanson P J 2016 Stability of peatland carbon to rising temperatures *Nature Communications* **7** 13723
- Xu M and Shang H 2016 Contribution of soil respiration to the global carbon equation *Journal of Plant Physiology* **203** 16–28
- Yang Y, Luo Y and Finzi A 2011 Carbon and nitrogen dynamics during forest stand development: A global

