

Albedo iteration in 3D setups

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## 1 Introduction

In this document the albedo iterations are explained that make sure that also with a varying albedo between day and nightside of the planet, the total energy is conserved. The problem is here that the 3D scheme in principle uses a  $\beta$ -map that is not aware of if the local PT point is on the day side or on the night side. This creates a problem when the night side has a different reflectivity as the day side as the night side cannot reflect stellar radiation.

## 1.1 Equations

As explained in the WASP-43b paper (Chubb & Min), the  $\beta$ -map represents a way of redistributing heat from the day to the night side. We have first the static  $\beta$ -map:

$$\beta_{\star} = \cos \Lambda \cos \Phi,\tag{1}$$

on the dayside and

$$\beta_{\star} = 0, \tag{2}$$

on the nightside. Here  $\Lambda$  and  $\Phi$  are the longitude and latitude respectively. The  $\beta$  map is computed from a diffusion equation using  $\beta_{\star}$  as source term. Note that  $\beta$  and  $\beta_{\star}$  have the property that:

$$\int_{\text{planet}} \beta = \int_{\text{planet}} \beta_{\star} = \int_{\text{day}} \beta_{\star} = 1$$
 (3)

The total amount of reflected light from the planet is:

$$f_{\rm ref} = \int_{\rm day} \beta_{\star} \omega \tag{4}$$

If the albedo,  $\omega$ , is a constant over the day and the night side, the fraction of starlight reflected taken into the computation of the P-T structure is:

$$f'_{\text{ref}} = \int_{\text{planet}} \beta \omega = \omega \int_{\text{planet}} \beta = \int_{\text{dav}} \beta_{\star} \omega = \omega$$
 (5)

This only works if the albedo is constant. If the albedo on the day and the night side is different  $f'_{ref} \neq f_{ref}$ , which makes the energy balance incorrect.

To solve this issue we can add a scaling to the  $\beta$ -map. The total emission needs to be  $1 - f_{\text{ref}}$ . If we take the new  $\beta$ -map to be:

$$\beta' = \gamma \beta \tag{6}$$

with  $\gamma$  a constant. We now have that the emission we use for the computation of the PT structure is equal to

$$\gamma - \int_{\text{planet}} \gamma \beta(\omega) \tag{7}$$

To have the right energy balance we need to make sure that

$$\gamma - \int_{\text{planet}} \gamma \beta \omega = 1 - \int_{\text{day}} \beta_{\star} \omega \tag{8}$$

which gives for the scaling factor

$$\gamma = \frac{1 - \int_{\text{day}} \beta_{\star} \omega}{1 - \int_{\text{planet}} \beta \omega} \tag{9}$$

Note that for constant  $\omega$  we can use Eq. 3 to show that  $\gamma = 1$  (as expected).

To compute  $\gamma$  we need to know the albedo at each location of the planet. For this we start with  $\gamma=1$  and compute the albedo everywhere. We use this to compute a new estimate of  $\gamma$ . Usually this is already enough and the value of  $\gamma$  does not change after one iteration. In the case of self-consistent cloud formation the albedo might depend more heavily on the PT structure and more than 1 iteration is required for  $\gamma$  to converge.