

Introduction to geodynamic modelling

The marker-in-cell technique

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Goals of this lecture

Introduce the marker-in-cell technique as a partial solution to numerical diffusion

 Provide some examples of how various values are calculated for the markers



An example of the problem

 Let's check out an example of numerical diffusion at https://introgm.github.io/2020/day-4/numerical-diffusion.html



Numerical diffusion in advection problems

- In some cases we can handle a bit of diffusion, but in others diffusion is a problem
 - A bit of diffusion of heat may not be a major problem, but diffusion of rock types can be a serious issue
 - For example, it would not be acceptable for defined weak layer in a model, such as a evaporites in the stratigraphy of rocks in a fold-and-thrust belt, to diffuse into the surrounding, stronger bedrock
 - In this case, the weak layer that may be the principal detachment surface disappears and the fold-and-thrust belt may behave entirely differently



Reference frames: Eulerian versus Lagrangian

- We haven't said anything about it, but up to this point we have been performing our calculations on an Eulerian grid of points
 - Eulerian points are fixed in space
 - For example, we have calculated changes in temperature, but the locations of our grid points x_i have not changed
- Lagrangian points are free to move with time, following the velocity at which material is being advected, for example

- You can think of these as two different reference frames:
 - The Eulerian reference frame is <u>external</u>, the Lagrangian reference frame is <u>internal</u>



Material I

Material 2

A simple example

- Consider a ID heat transfer model simulating exhumation of rock as a result of erosion
 - The model has two materials and vertical advection of rock





Material

Material 2

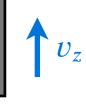
A simple example



 Previously, we were able to calculate temperatures including the thermal effects of advection

$$T_n^i = \left(\left(\alpha_n \frac{T_{n+1}^{i-1} - T_n^{i-1}}{\Delta x^2} - \alpha_{n-1} \frac{T_n^{i-1} - T_{n-1}^{i-1}}{\Delta x^2} + H_n \right) \middle/ \rho_n C_{p,n} - v_x \frac{T_{n+1}^{i-1} - T_{n-1}^{i-1}}{2\Delta x} \right) \Delta t + T_n^{i-1}$$

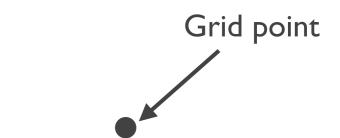
• This is OK, but as we know, erosion of rock and the resulting transport toward the surface should also advected the rocks (and their properties) toward the surface, not just temperature



A simple example Grid point (Eulerian) Material Moving interface (Lagrangian) One solution to this problem would be to track and store the location of the interface between the two rock types and use that to Material 2 determine the properties of the grid points This is a simple way to track material properties in ID, but tracking this kind of interface becomes more challenging in 2D or \mathcal{V}_z 3D, so alternative techniques are used Time 0 Time I Time 2



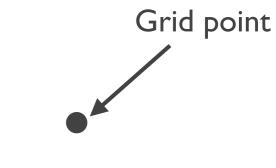
Marker-in-cell

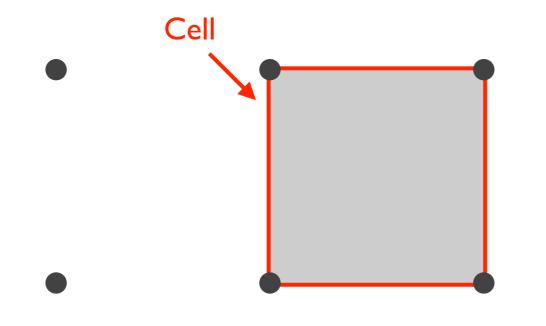


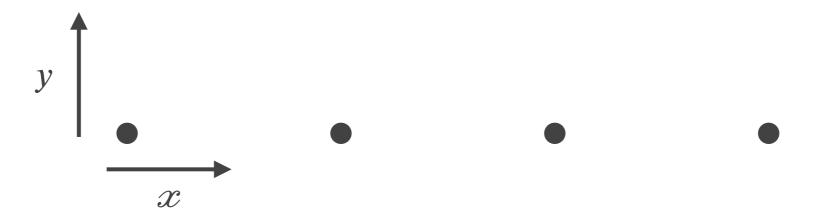




Marker-in-cell

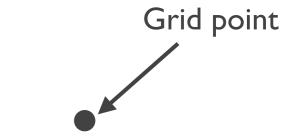


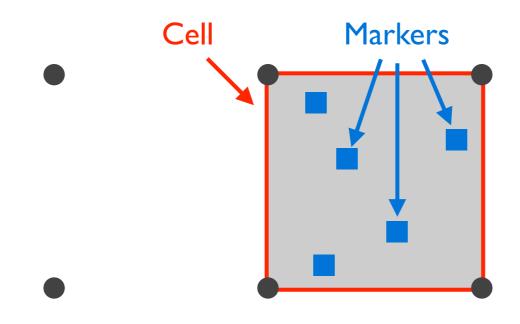


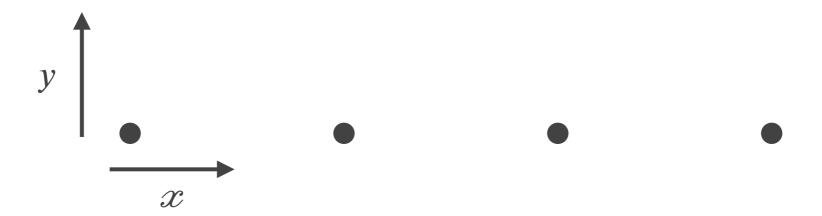




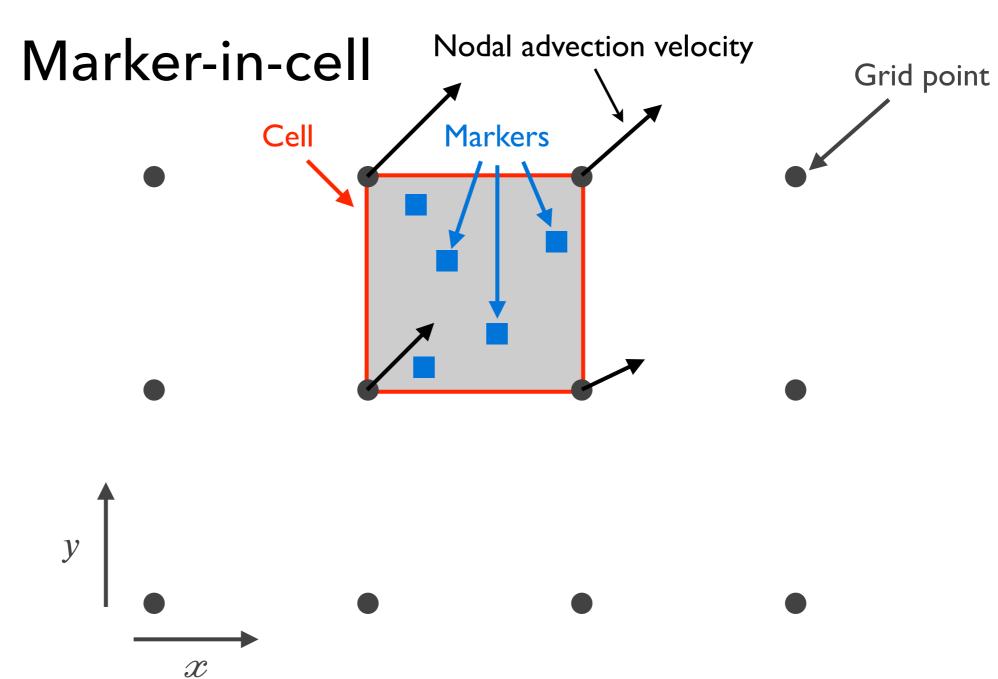
Marker-in-cell



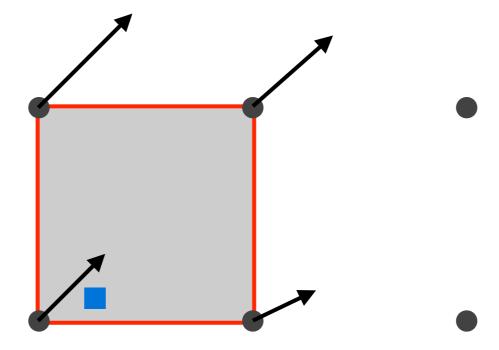








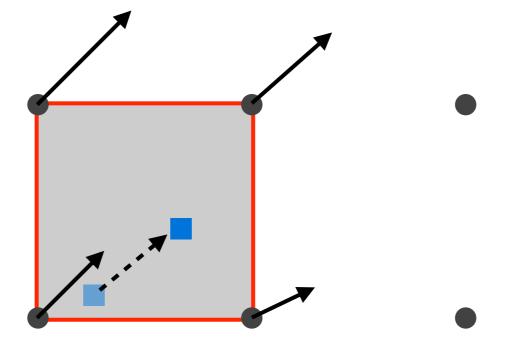


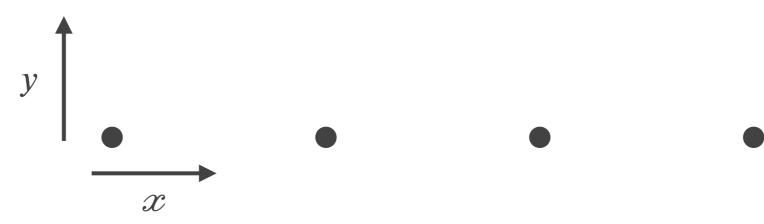




- Let's now consider a single marker in the model
 - Its position will change with time based on the velocities of material at the four surrounding nodes (the cell)

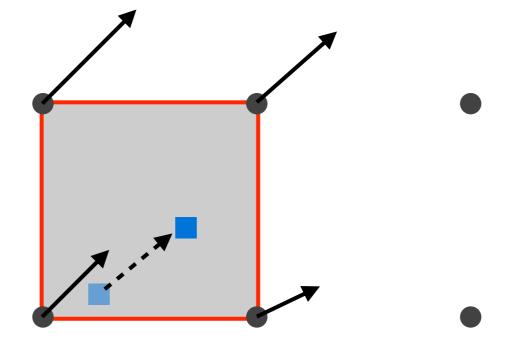


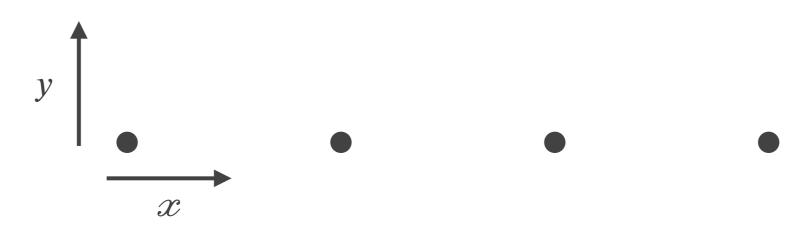




- The particle velocity is calculated using a bilinear interpolation of the nodal velocities
 - Obviously, this avoids numerical diffusion of the marker data

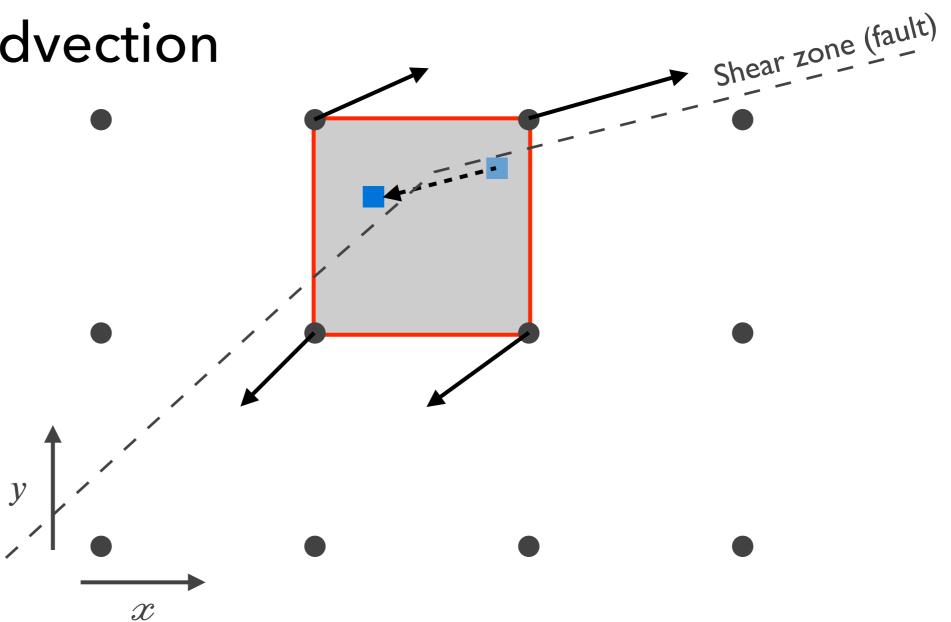






• It is important to note, though, that the velocity varies across a given cell, and error can be introduced by advecting a particle for an entire time step using the velocity at its origin





- Consider a particle in the footwall of a shear zone or fault
 - Here, because the initial velocity is used for the entire time step, material is transferred from the footwall to the hanging wall



Marker advection schemes

• The simple case from the previous slide is given below

$$x_A^i = x_A^{i-1} + v_{xA}\Delta t$$

$$y_A^i = y_A^{i-1} + v_{yA} \Delta t$$

where $x^i{}_A$ is the future x-coordinate of particle $x^{i\text{-l}}{}_A$ and v_x is the velocity in the x direction, and similarly for y

 This scheme is first-order accurate, and requires small time steps if there are large differences in velocity in the model

 An alternative is to consider a velocity that is not based solely on the velocity at the starting position of a marker, but one that considers more than a single point



Runge-Kutta advection

- The Runge-Kutta advection scheme uses 2-4 points to determine the velocity at which particles should be advected
- The second-order Runge-Kutta scheme is below

$$x_A^i = x_A^{i-1} + v_x^{\text{eff}} \Delta t$$

$$y_A^i = y_A^{i-1} + v_y^{\text{eff}} \Delta t$$

where

$$v_x^{\text{eff}} = v_{xB}$$
 $x_B = x_A^i + v_{xA} \frac{\Delta t}{2}$ $v_y^{\text{eff}} = v_{yB}$ $y_B = x_A^i + v_{yA} \frac{\Delta t}{2}$



Properties on the markers

- It is quite common to use the markers to carry a variety of values that should not diffuse
 - Material properties (density, conductivity, viscosity, friction angle, etc.)
 - Fields with a history, such as total strain

Intro to geodynamic modelling

 However, because the markers largely avoid typical numerical diffusion issues, they can be used to store nearly all of the information calculated in the numerical model (temperature, strain rates, etc.)

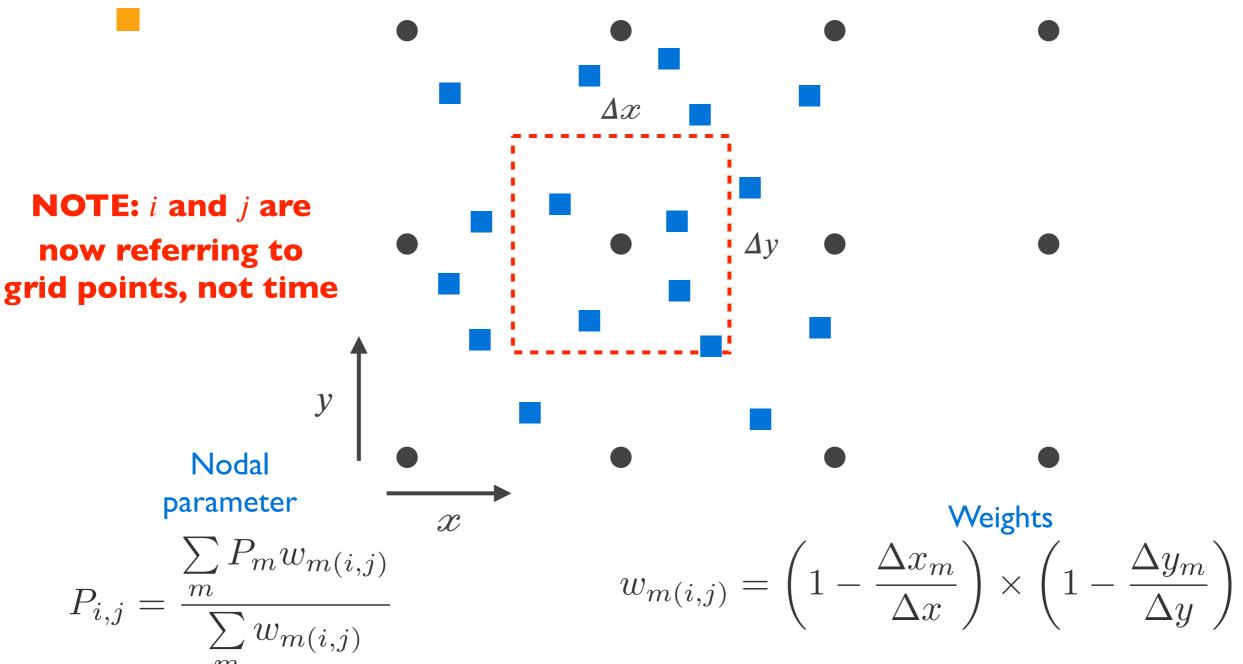


Retrieving information from the markers

- Now, we have established the benefits of the marker-in-cell method for avoiding numerical diffusion, but it is not as if it an method that is free of error
- In order to transfer information between the markers and the grid points, we need some kind of function
- A weighted difference averaging approach is one simple option



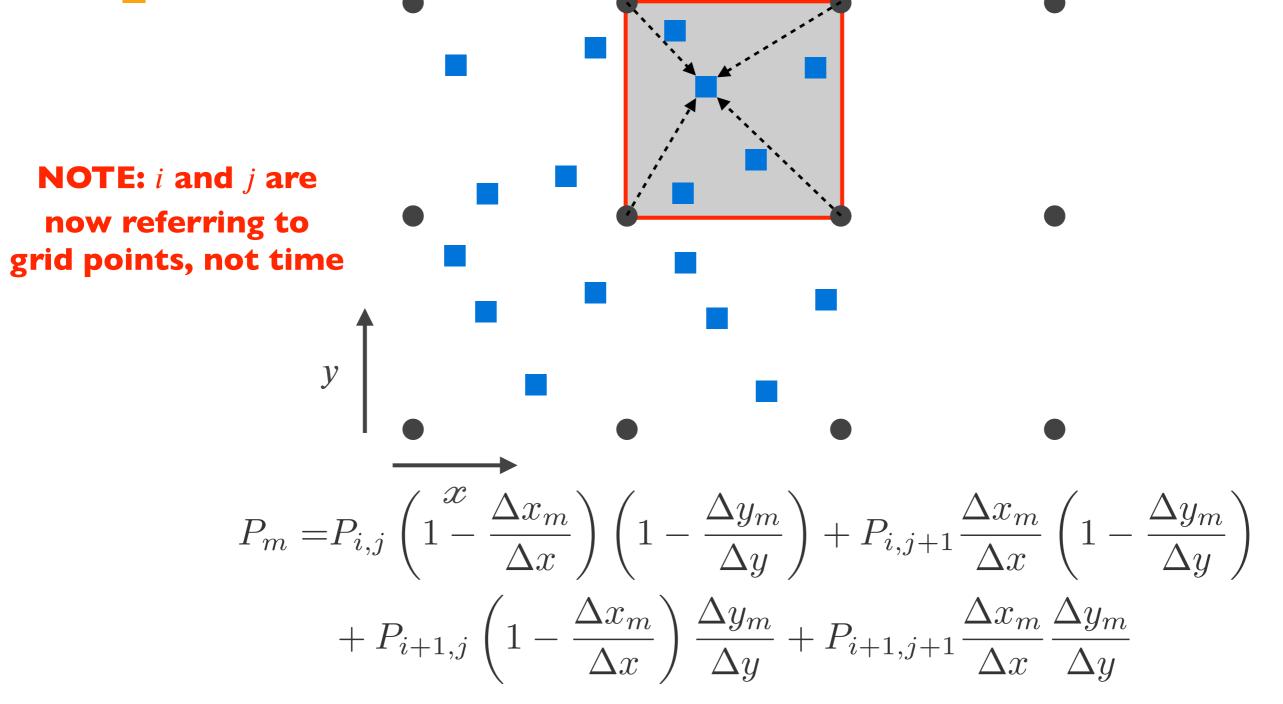
Retrieving information from the markers



• $w_{m(n)}$ is the weight of the m-th marker for grid point (i,j)

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Giving information to the markers



Interpolation of values of parameter P from grid points i,j to marker m 22



Some words of caution

- Even when using the marker-in-cell technique there are some important considerations that can introduce error
 - First, there is a threshold number of markers that must exist in a cell in order to provide a reasonable resolution of the field data
 - A minimum of perhaps 5 markers would be OK in 2D, perhaps 9 in 3D
 - A larger number is better, but calculations with a large "cloud" of markers can get heavy



Some words of caution

- When the number of markers falls below the minimum needed, many models (including the one you will use) allow markers to be injected into the cells
 - How this is best done, and the how values are transferred to the new markers is not completely obvious
 - This is a common issue for models where material is flowing into the model from the side and the markers are continually advected away from the side of the model
 - If you do not initialise the values on the markers carefully you can get unexpected results



Some words of caution

Other concerns?



Summary

We have seen that the marker-in-cell technique, which uses
 Lagrangian markers, is a useful way to avoid numerical diffusion

 The Lagrangian markers are a fairly intuitive concept, but there can be some challenges in using markers, including how material is advected, how markers are injected, and how marker values are initialised



References

Gerya, T. (2010). Introduction to numerical geodynamic modelling. Cambridge University Press.