

Manuscript Details

Manuscript number	CAGEO_2017_297
Title	PAF: A software tool to estimate free-geometry extended bodies of anomalous pressure from surface deformation data
Article type	Research Paper

Abstract

We present a software package, namely PAF, to carry out inversions of surface deformation data (any combination of InSAR, GPS, and terrestrial data) as produced by 3D free-geometry extended bodies with anomalous pressure changes. The anomalous structures are described as aggregation of elementary cells (point sources) in an elastic half space. The linear inverse problem (supply with some simple regularization conditions) is solved by means of an exploratory approach. This software represents the open implementation of a previously published methodology. It can be freely downloaded from the journal web page or asked to the corresponding author, and used both with large data sets (e.g. InSAR data sets) or with data coming from small control networks (e.g. GPS monitoring data), mainly in volcanic areas to estimate the expected causative pressure bodies. Here, PAF software is applied to some real test cases.

Keywords	software, surface deformation, pressure sources, volcano monitoring
Taxonomy	Volcanology, Modelling
Manuscript category	Surveying and monitoring
Corresponding Author	Jose Fernandez
Corresponding Author's Institution	CSIC
Order of Authors	Antonio G. Camacho, Jose Fernandez, Flavio Cannavo
Suggested reviewers	Kristy Tiampo, Pietro Tizzani, Maurizio Battaglia

Submission Files Included in this PDF

File Name [File Type]

Cover_letter_27-03-2017.docx [Cover Letter]

Camacho_et_al_PAF-Software.pdf [Manuscript File]

To view all the submission files, including those not included in the PDF, click on the manuscript title on your EVISE Homepage, then click 'Download zip file'.

Dear Editor:

Please find attached the manuscript "PAF: A software tool to estimate free-geometry extended bodies of anomalous pressure from surface deformation data." by A.G. Camacho, J. Fernández and F. Cannavó, we would like to publish in Computers & Geosciences journal.

It is a re-submission of the manuscript CAGEO_2017_158, after providing in the text where the source code of the software can be found, under an open source license. The source codes are also available for reviewers if they need. See last sentence of conclusions The editor who managed the previous manuscript was Prof. Edzer Pebesma.

The manuscript have been not submitted for consideration to any other journal.

Yours,

José Fernández

Correponding author

PAF: A software tool to estimate free-geometry extended bodies of anomalous pressure from surface deformation data.

A.G. Camacho ⁽¹⁾, J. Fernández ^(1,*), F. Cannavó ⁽²⁾

⁽¹⁾ Institute of Geosciences (CSIC-UCM), Madrid, Spain

⁽²⁾ Osservatorio Etneo, Istituto Nazionale di Geofisica e Vulcanologia, Catania, Italy.

*Corresponding author:

Institute of Geosciences (CSIC-UCM), Fac. C. Matemáticas, Plaza de Ciencias, 3. 28040-Madrid, Spain. e-mail: jft@mat.ucm.es; phone: +34-913944632

Abstract: We present a software package, namely PAF, to carry out inversions of surface deformation data (any combination of InSAR, GPS, and terrestrial data) as produced by 3D free-geometry extended bodies with anomalous pressure changes. The anomalous structures are described as aggregation of elementary cells (point sources) in an elastic half space. The linear inverse problem (supply with some simple regularization conditions) is solved by means of an exploratory approach. This software represents the open implementation of a previously published methodology. It can be freely downloaded from the journal web page or asked to the corresponding author, and used both with large data sets (e.g. InSAR data sets) or with data coming from small control networks (e.g. GPS monitoring data), mainly in volcanic areas to estimate the expected causative pressure bodies. Here, PAF software is applied to some real test cases.

Keywords: software, surface deformation, pressure sources, volcano monitoring.

23 **1. Introduction.**

24 In a volcanic context, surface deformation data are normally inverted to infer information about
25 the intrusive pressure sources (e.g., Rymer and Williams-Jones, 2000; Dzurisin, 2003). The surface
26 deformation is related to the dynamics of volcanic plumbing systems, such as the shape of magma
27 intrusions, magma pressure, and emplacement mechanisms. Normally, regular geometries (point
28 sources, disks, prolate or oblate spheroids, etc.) are assumed at the initial stages (Lisowski, 2007). The
29 inversion approach for the deformation data is usually based on the hypothesis of spherical, ellipsoidal
30 or general extended causative bodies. Other types of analytical models, as fault, dike, laminar or
31 cylindrical structures, are sometimes employed. Moreover, the mathematical model for the source
32 structure must consider some elastic properties to account for the deformation phenomena. Usual
33 analytical modeling assumes an elastic, homogeneous and isotropic crust, but it can take into account
34 effects from several source geometries, topography relief and gravity background (Williams and
35 Wadge, 1998; Charco et al., 2007; Battaglia and Hill, 2009).

36 Camacho et al. (2011) presented an original methodology for simultaneous inversion of vertical
37 (Up), east-west (EW), and north-south (NS) deformation components and/or LOS InSAR
38 displacement, by means of 3D extended bodies, with free geometry, for anomalous pressure. Assuming
39 homogenous elastic conditions, the approach determines a general geometrical configuration of
40 pressurized sources corresponding to prescribed values of anomalous pressure. The sources are
41 described as an aggregate of pressure point sources, and they fit the entire data within some regularity
42 conditions. The approach works in a step-by-step growth process that allows retrieving very general
43 geometrical configurations.

44 This approach provides interesting results for volcanic areas (Camacho et al., 2011; Samsonov et
45 al., 2014; and Cannavó et al., 2015b) when deformations come from pressure sources. The subsurface
46 volume is divided into a 3D partition of (thousands) elemental cells. The aggregation of elemental

47 sources (with superposition of their strain contribution) forms the geometry of the extended causative
48 bodies.

49 The observation equations are [see Camacho et al. (2011) for a detailed description]:

$$50 \quad \mathbf{ds} = \mathbf{ds}^c + \mathbf{v} \quad (1)$$

51 where \mathbf{ds} , \mathbf{ds}^c represent the vector of observed and calculated three component (3D) of the
52 displacement, and \mathbf{v} is the vector for residual values coming from uncertainties in the observation
53 process and the imperfect model fit.

54 The surface deformations, \mathbf{ds}^c , due to a buried over pressure structure are calculated as the
55 aggregated effect of different point sources, as due to the deformation effects from the incremental
56 pressure p_k and expansion radius within the elastic semi space, described by the Mogi model (Mogi,
57 1958; Rymer and Williams-Jones, 2000).

58 The inversion equations are solved by means of adding a regularization condition:

$$59 \quad \mathbf{v}^T \mathbf{Q}_D^{-1} \mathbf{v} + \lambda \mathbf{m}^T \mathbf{Q}_M^{-1} \mathbf{m} = \min. \quad (2)$$

60 where model vector \mathbf{m} is constituted by the pressure values p_k , $k=1, \dots, m$, for the m cells of the model,
61 \mathbf{Q}_D is the covariance matrix of data, \mathbf{Q}_M is a suitable covariance matrix corresponding to the physical
62 configuration and λ is a smoothing factor that balances the goodness of fit and smoothness of the
63 model. The inversion approach is a non-linear problem.

64 The final source is determined as a free aggregation of a large number of small over (under)
65 pressured sources. The inversion algorithm carries out a step-by-step process of growth of the 3D
66 models, using an exploratory technique to sequentially find the new cell/point source to be set as
67 (de)pressurized and aggregated to the model to improve the fit to the data (Camacho et al., 2011).

68 In fact, at the k -th step of the growth process, k cells are filled with the prescribed anomalous values
 69 for pressure, giving rise to the modelled values \mathbf{ds}^c . Successively, at the new $(k+1)$ -th step, the
 70 algorithm searches for a new cell to fill to improve the fit following the system:

$$71 \quad \mathbf{ds} = f \mathbf{ds}^c + \mathbf{v} \quad (3)$$

$$72 \quad misfit = \mathbf{v}^T \mathbf{Q}_D^{-1} \mathbf{v} + \lambda f^2 \mathbf{m}^T \mathbf{Q}_M^{-1} \mathbf{m}, \quad (4)$$

73 where $f > 1$ is a scale factor, estimated during the inversion, to allow for a fit between the anomaly of
 74 the provisional model and the observed anomaly.

75 This inversion methodology has been tested by means of several synthetic tests (Camacho et al,
 76 2011), and also by real application to volcanic environments: Campi Flegrei (Italy) (Camacho et al,
 77 2011; Samsonov et al., 2014) and Mount Etna (Italy) (Cannavo et al, 2015b).

78 In the next section, we describe the proposed software tool (PAF-package) that allows for a simple
 79 and nearly automatic application of this methodology. Then, we present two real applications that
 80 illustrate the selection of parameters and some features of the adjusted model for over pressure bodies.
 81 Most of the following figures are created by the PAF software.

82 **2. The PAF Software.**

83 The software package consists of two executable files: **ConfigPAF.exe** and **InverPAF.exe**. The
 84 first one determines a 3D partition of the subsurface volume into a grid of small parallelepiped cells
 85 through a graphical interface for the input of the inversion parameters. It creates an intermediate file
 86 (**CellsConfig.txt**) with the information of cell partitions. The second one, **InverPAF.exe**, reads the
 87 displacement data and the intermediate file (**CellsConfig.txt**), and runs the inversion to model the
 88 observed data with some 3D extended bodies of anomalous pressure described as aggregation of filled
 89 cells.

90 The PAF software can work in two different ways: (1) statically, considering large displacement
91 data files describing the deformation field (for instance, thousands of pixels with displacement values
92 from SAR interferometry, or combined with GPS data) for a concrete time period; and (2) dynamically,
93 sequential inversion of successive stages of a deformation process (for instance, successive GPS
94 displacement data obtained at different epochs from a permanent monitoring network).

95 **2.1.Input files and data.**

96 The 3D deformation data from the several sites are collected in the file **DeforData.txt**. This file
97 contains, for each data point, coordinates UTM East-North (m), altitude (m), and deformation values
98 dz (cm, positive upward), dx (cm, positive eastward) and dy (cm, positive northward). The value in the
99 adjacent column, associated to each measurement, indicates whether the corresponding deformation
100 value has to be included in the inversion process (one) or not (zero). This data file can optionally be
101 terminated with an end-of-file character or with a line of zeros.

102 A secondary data file is **map.blm**, which can be optionally used to draw the cartography of the
103 studied area. It includes numerical values of polygonal lines (described by the coordinates of the
104 vertexes) with topographical details (roads, etc.).

105 In the case of sequential application, the successive data are provided on successive files, named
106 **Dn.txt**, where n is a successive integer number, and with the same format as **DeforData.txt**.

107 **2.2.Software Usage.**

108 Once the data are available, the first step to perform an inversion with PAF is to generate a suitable
109 3D partition of the subsurface volume into a grid of small parallelepiped cells. This step is carried out
110 by using the **ConfigPAF.exe** tool. This program reads the data in **DeforData.txt** (Figure 1) and sets
111 several parameters by means of a friendly dialog window (Figure 2), offering default values.

112

425575	4527988	91.6	0.38	1	-0.23	1	0	0
425844	4527985	95.7	0.37	1	-1.53	1	0	0
426114	4527982	105.0	0.94	1	-0.21	1	0	0
426383	4527980	102.7	0.02	1	-0.48	1	0	0
425103	4527948	90.7	1.72	1	1.80	1	0	0
426653	4527932	111.6	0.95	1	-0.86	1	0	0
426922	4527930	130.7	1.56	1	0.52	1	0	0
424799	4527907	97.7	0.40	1	-1.91	1	0	0
427191	4527883	132.0	0.37	1	0.86	1	0	0
425338	4527857	85.9	0.55	1	1.08	1	0	0
427460	4527835	154.1	1.04	1	0.09	1	0	0
424326	4527822	80.8	0.94	1	0.32	1	0	0

Figure 1. Input data file DeforData.txt containing coordinates UTM (m), altitude (m), and deformation values dz (cm), dx (cm) and dy (cm) for the data points. Values of 0 and 1 are used to indicate whether the values are included or not, respectively.

Figure 2. Input dialog window for the parameters of the 3D grid of cells to prepare the inversion calculus.

In the upper part of the Graphical User Interface (GUI) (Figure 2), the user can revise the default values for the geometrical parameters of the partition: depth (m above sea level) of the top of the

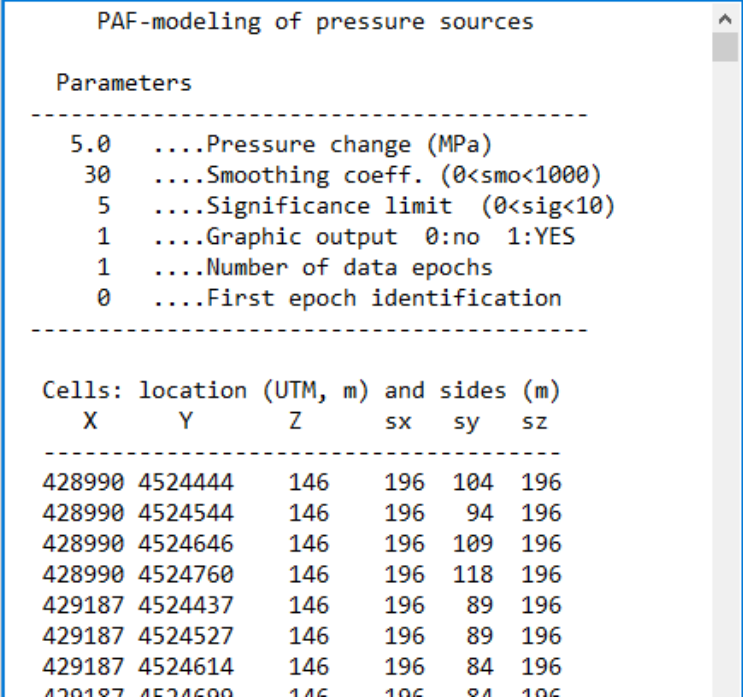
122 partition volume, depth (m, above sea level) of the bottom of the partition volume, and mean side size
123 (m) of the cells. Once these values are set (pressing Run/OK), the program calculates the resulting
124 number of cells.

125 Then, in the bottom part of the GUI (Figure 2), the program requires two more parameters: (a) the
126 dimensionless smoothing coefficient λ (ranging between 0 and 1000), and, (b) a value for the (positive
127 or negative) pressure contrast throughout the entire anomalous model, Δp (MPa). The λ coefficient
128 regulates the balance between data fit and model complexity in the inversion approach. For a low λ
129 value (close to 0) the resulting model becomes very simple, regular and compact, but the data fit can
130 be weak. Conversely, a high λ value (close to 1000) the data fit is very good (even by fitting some
131 noise component), but the resulting model can be very complex, sometimes even with artefacts. We
132 have imagined some objective criterion (see for instance, Camacho et al. 2007, in other context), but
133 for practical applications, we suggest to try different values in a trial and error manner, and to select
134 the most appropriate one considering the resulting models. Some criteria for choosing this value are:
135 be sure that the program finishes the inversion (for too high or too low values of the parameter, the fit
136 conditions are invalid and the program stops); auto-correlated components or significant signal in the
137 residuals should be avoided; the resulting model should be regular and simple (avoiding very small
138 and sparse shallow body for noise inversion). For more details, see the two application examples
139 below.

140 The pressure contrast value Δp can be selected trial and error also by doing some iterative running
141 of the software. For very high values, the model becomes very condensed and compact, and some
142 geometrical details can be lost. Conversely, for very low values the sources model becomes larger,
143 with rounding inflated shape. In general, big displacements require strong pressure contrasts and small
144 ones require low-pressure contrasts. We suggest, again, some iteration trying different values and
145 observing the resulting anomalous geometry. Nevertheless, this is not a critical parameter. It concerns
146 mostly to rather aesthetic aspects of the model. See also application examples below.

147 Once the **ConfigPAF.exe** is completed (by pressing Run/OK), it creates a new file, namely
 148 **CellsConfig.txt**. This is an intermediate file containing (see Figure 3): (a) the assumed values for the
 149 inversion parameters (smoothing coefficient, pressure contrast, etc.), and (b) the geometrical
 150 parameters (location and sides) of the parallelepiped cells.

151



The screenshot shows a text file titled "PAF-modeling of pressure sources". It contains two main sections: "Parameters" and "Cells: location (UTM, m) and sides (m)".

Parameters section:

```

5.0 ....Pressure change (MPa)
30 ....Smoothing coeff. (0<smo<1000)
5 ....Significance limit (0<sig<10)
1 ....Graphic output 0:no 1:YES
1 ....Number of data epochs
0 ....First epoch identification
  
```

Cells section:

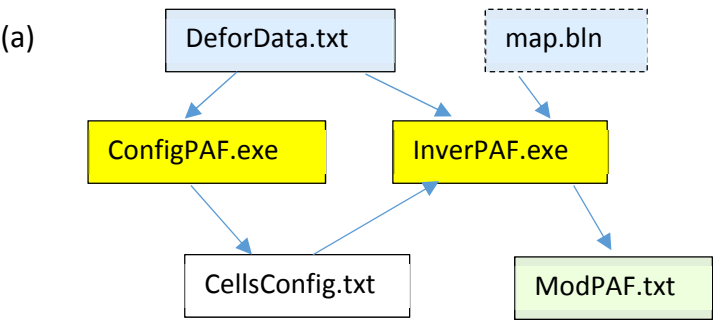
X	Y	Z	sx	sy	sz
428990	4524444	146	196	104	196
428990	4524544	146	196	94	196
428990	4524646	146	196	109	196
428990	4524760	146	196	118	196
429187	4524437	146	196	89	196
429187	4524527	146	196	89	196
429187	4524614	146	196	84	196
429187	4524600	146	196	84	196

152 **Figure 3.** Intermediate file *CellsConfig.txt* containing the assumed values for the inversion
 153 parameters, and the geometrical parameters (location and sides) of the parallelepiped cells.

154 The values of the parameters contained in this file can be manually modified to adapt to the
 155 modeller necessities. For instance, the third parameter, “*Significance limit (0<sig<10)*”, gives the
 156 desired threshold between the significant cells (for those close and covered by the data points) from
 157 the non-significant cells (for those far from the data points). Its default value is 5, but the user can set
 158 optionally a different value (for instance, for a low significance threshold, nearly all cells will be
 159 considered as significant). The fourth parameter (“*Graphic output 0:no 1:YES*”) is used to switch the

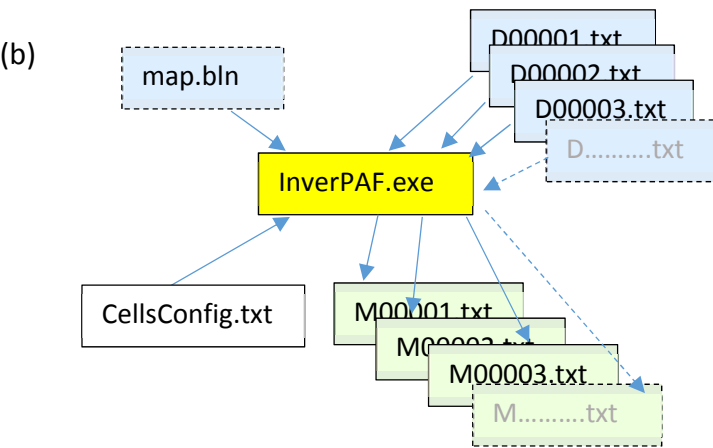
graphical output during the inversion approach (see below the content of this graphical output). The fifth parameter indicates that the data correspond to only one epoch, or conversely that the inversion has to be carried out in a sequential process of a number of epochs, starting at the epoch and with the data file identified by the sixth parameter.

164



165

166



167

Figure 4. Scheme of the application of the PAF software. (a) Case of deformation data for one epoch (for instance, thousands of pixel for a SAR interferogram), (b) Sequential application of the inverse approach for successive data files (in blue) producing successive model files (in green) and graphic pictures (case of a monitoring network)

172 Once the file **CellsConfig.txt** is created (and optionally revised), the inversion approach by
173 **InverPAF.exe** can start (Figure 4). It runs automatically, without any other requirement, and produce
174 an optional visualization (fourth parameter in CellCnfig.txt) of the inversion process and of the results
175 on the screen. In the case of sequential application (Figure 4b), the **InverPAF.exe** process works
176 continuously, epoch by epoch, until it inverts the data of the last epoch. As final product, the inversion
177 approach creates an output file named **ModPAF.txt** containing the geometrical description of the
178 resulting model, several others parameters of this model, and a detailed information about modelled
179 and residual values. They will be described further on.

180 **2.3. Output files and pictures.**

181 Figure 5 shows an example of the graphical output that can be optionally displayed throughout the
182 inversion approach (see Figure 4 and corresponding description text). This visual information contains
183 some interesting graphics about the model growth process and about the final results. Statistics about
184 the model growth process are: *misfit* value (see equation 4) evolution, pressure value evolution, planar
185 and vertical cuts views (EW and NS) of the aggregation process (significant cells in red and dark blue,
186 no significant cells in yellow and light blue, cells of a previous epoch in green), observed and modelled
187 deformations (along EW and along NS), inversion parameters (step number, pressure value, root mean
188 square (rms) residuals, cell depth). Statistics about the final results are: *misfit* value (equation 4),
189 number of rejected data, number of filled cells, and number of significant cells.

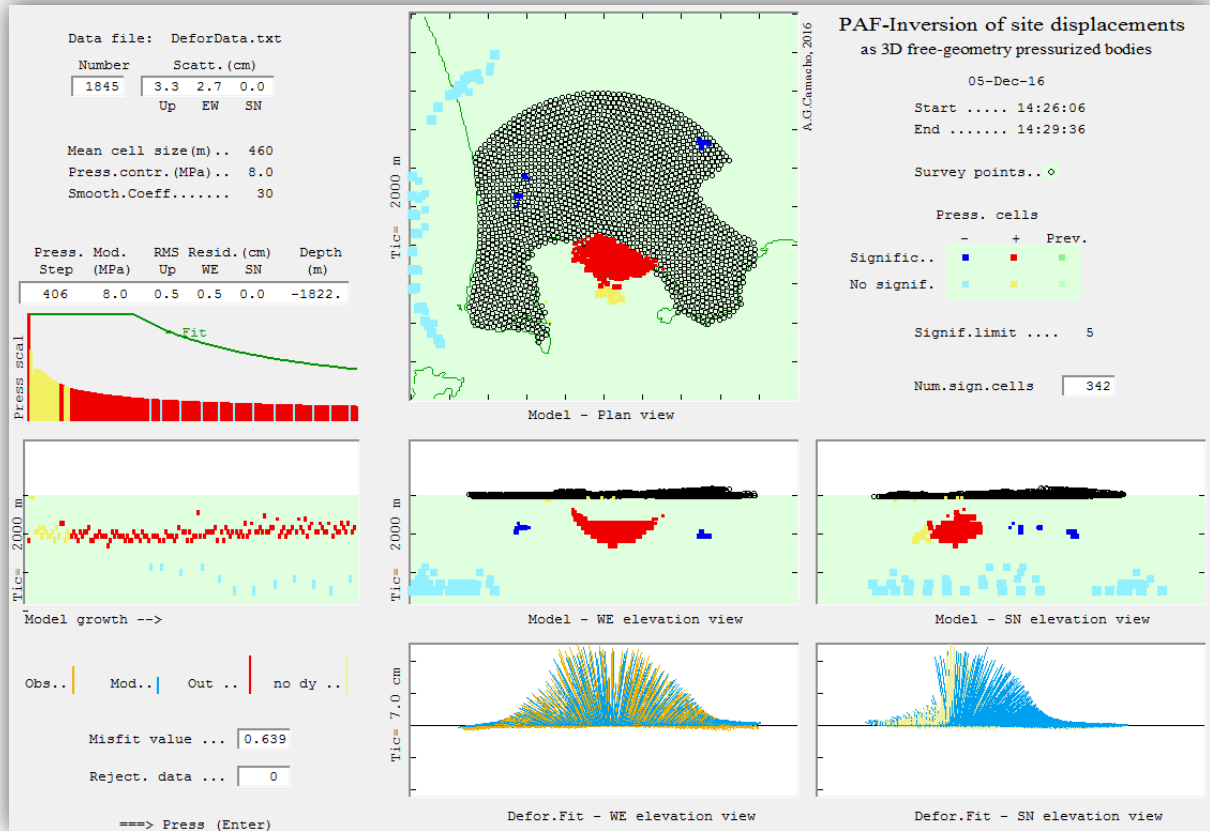
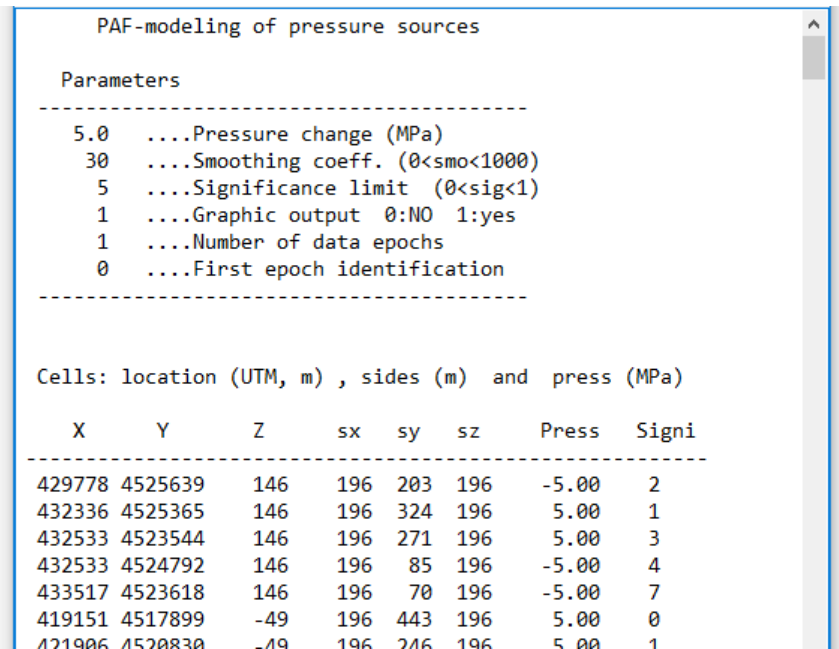


Figure 5. Example of the (optional) screen drawing along the inversion approach. It informs about the evolution of the model growth process and about some results.

The main product of the PAF inversion is the file **ModPAF.txt**. This file contains the following information:

- (1) A copy of the assumed inversion parameters (Figure 6) similar to that in file **CellConfig.txt**.
- (2) A list of the filled cells (those with no-null pressure values) containing (see Figure 6); for each cells, the UTM coordinates X and Y (m), and the depth Z (m, above sea level) of its geometric center, their sides sizes s_x , s_y , s_z (m), the corresponding positive or negative pressure contrast value, and a value for the significance (given by the inverse of the root mean square distance of the cell with respect

200 to the data points, normalized to range from 0 to 20). Higher values of significance correspond to more
 201 sensitive cells. This part can be used for further, more sophisticated, drawings of the resulting model.



PAF-modeling of pressure sources

Parameters

5.0Pressure change (MPa)
 30Smoothing coeff. (0<smo<1000)
 5Significance limit (0<sig<1)
 1Graphic output 0:NO 1:yes
 1Number of data epochs
 0First epoch identification

Cells: location (UTM, m) , sides (m) and press (MPa)

X	Y	Z	sx	sy	sz	Press	Signi
429778	4525639	146	196	203	196	-5.00	2
432336	4525365	146	196	324	196	5.00	1
432533	4523544	146	196	271	196	5.00	3
432533	4524792	146	196	85	196	-5.00	4
433517	4523618	146	196	70	196	-5.00	7
419151	4517899	-49	196	443	196	5.00	0
421906	4520830	-19	196	216	196	5.00	1

203 **Figure 6.** Upper content of the output file **ModPAF.txt**. First, the modelling parameters are shown,
 204 Then, the list of the filled cells (coordinates, sides, significance) are included.

205 (3) After that, some additional parameters and results about the inversion process are given,
 206 followed by a list of the observed, modelled and residual values for the data points (see Figure 7). For
 207 each point, the file contains UTM coordinates (m), altitude (m), observed, modelled and residual values
 208 (cm) for each component (dz , dx , dy), and an additional value for relative quality weighting (mean
 209 value 1) according to the resulting residual values for the three components. This part of the output file
 210 can be used for further drawings and statistical analysis of the inversion residuals (see examples
 211 below).

212 In the case of sequential application of the PAF inversion approach, for each successive data file
 213 **D.....txt** (e.g., **D00016.txt**) corresponding to successive monitoring epochs (respective sixteenth

epoch), the program generates a model file **M.....txt** (e.g., **M00016.txt**) with the same content as **ModPAF.txt**.

```

421443 4532415 -6109 439 1277 439 -5.00 1
422763 4532598 -6109 439 1211 439 -5.00 1
437719 4527158 -6109 439 1277 439 5.00 1
----- Date:08-Nov-16-----

Num. data points = 1061
Num. total cells= 50599
Num.filled cells= 469 Neg, Pos = 160 309
Medium param.: Poisson=0.25 Share Mod.= 10.GPa
Random explor.coeff.= 10
Pressure model: Press. contrast (-+)= 5.00 MPa
                  Press*vol: 150.1 MPA*Km3 Mean mod.depth = -3875.
RMS residuals (cm): Up= 0.46 WE= 0.58 SN= 0.33 Misfit= 0.3390
Initial and final exec.times: 09:55:38 09:56:51

Observed, modeled, and residual values

Data point loc(UTM, m)      dz (cm)      dx(cm)      dy(cm)      Weight
  X      Y      Z      obs mod res      obs mod res      obs mod res
-----
426261 4528962 96 -0.11 -0.74 0.63 -1.70 -1.50 -0.20 2.00 1.68 0.32 1.0
426611 4528958 102 0.27 -0.67 0.94 -1.60 -1.35 -0.25 2.00 1.62 0.38 0.2
425351 4528878 90 0.42 -1.01 1.43 -1.50 -2.09 0.59 1.60 1.69 -0.09 0.0
426960 4528862 111 0.50 -0.59 1.09 -1.00 -1.18 0.18 2.10 1.64 0.46 0.0
427309 4528859 110 0.01 -0.55 0.56 -2.10 -0.98 -1.12 2.10 1.64 0.46 1.0
425700 4528783 99 0.48 -0.79 1.27 -1.40 -1.93 0.53 2.10 1.95 0.15 0.0
426049 4528687 102 0.86 -0.63 1.49 -0.20 -1.42 1.22 2.50 2.00 0.50 0.0
426398 4528684 108 0.11 -0.65 0.76 -0.90 -1.30 0.40 2.20 1.70 0.50 0.8

```

Figure 7. Last part content of the output file *ModPAF.txt*. Top: End of the list of the filled cells (coordinates, sides, significance). Mean: Some parameters and results about the inversion process. Bot: List of the observed, modelled and residual values (cm) for each data point.

The code is in Fortran language and the executable compiled files are obtained with Microsoft Visual Studio Community 2015 for Windows 10 64 bits operative system.

3. Application examples

3.1. Campi Flegrei inflation case.

Campi Flegrei, one of the most hazardous volcanic areas in the world because of its close proximity to the city of Naples, is the subject of many studies and surveys. Samsonov et al. (2014) apply the

226 multidimensional small baseline subset (MSBAS) differential interferometric synthetic aperture radar
227 (DInSAR) technique to obtain vertical (Up) and horizontal components (EW) of ground deformation
228 for Campi Flegrei at high spatial and temporal resolutions spanning along 20 years. Their results show
229 that the area underwent continuous subsidence from 1993 through 1999, followed by several alternate
230 periods, and a moderate uplift which began in 2007 and increased through 2012, reaching to about 17
231 cm by 2013. We study the uplift period from 2007/10/03 to 2013/06/15, taking as input data the
232 accumulated 2D deformation (Up and EW components).

233 From the thousands of pixels obtained by Samsonov et al. (2014) we have selected 1845 pixels
234 with mutual distance larger than 250 m and with a distance to the centre of deformation smaller than
235 8 km. Figure 8 shows the distribution of the selected pixels and the feature of the vertical component
236 (with more than 12 cm accumulated in the central area).

237 Once the 3D deformation data for the 1845 pixels is arranged in the file **DeforData.txt**, we can
238 apply the described process and software. The program **ConfigPAF.exe** suggests some default values:
239 437 m for altitude of the model top, -5915 m for depth of the model bottom, and 231 m for mean cell
240 side. We accept the first two values, but we try 180 m for mean cell side, looking for higher resolution.
241 For this new value, the resulting number of cells is 72017, as provided by the program.

242 Then, the program suggests the usual default values 200 for smoothing coefficient and 0.5 MPa for
243 pressure contrast. They are very general values. For this particular case of Campi Flegrei we have a
244 very strong accumulated deformation field, and it requires a higher pressure contrast. After some trials
245 we select 10.0 MPa (accumulated pressure for the anomalous volume elements) as tentative model
246 contrast (for smaller values the anomalous structure becomes very inflated).

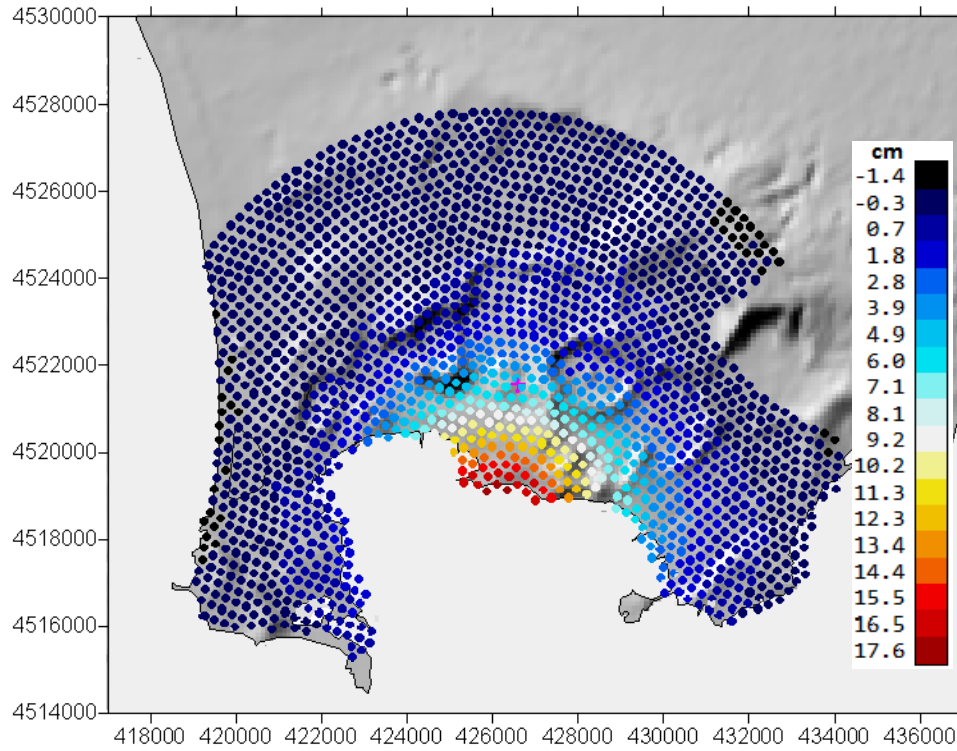
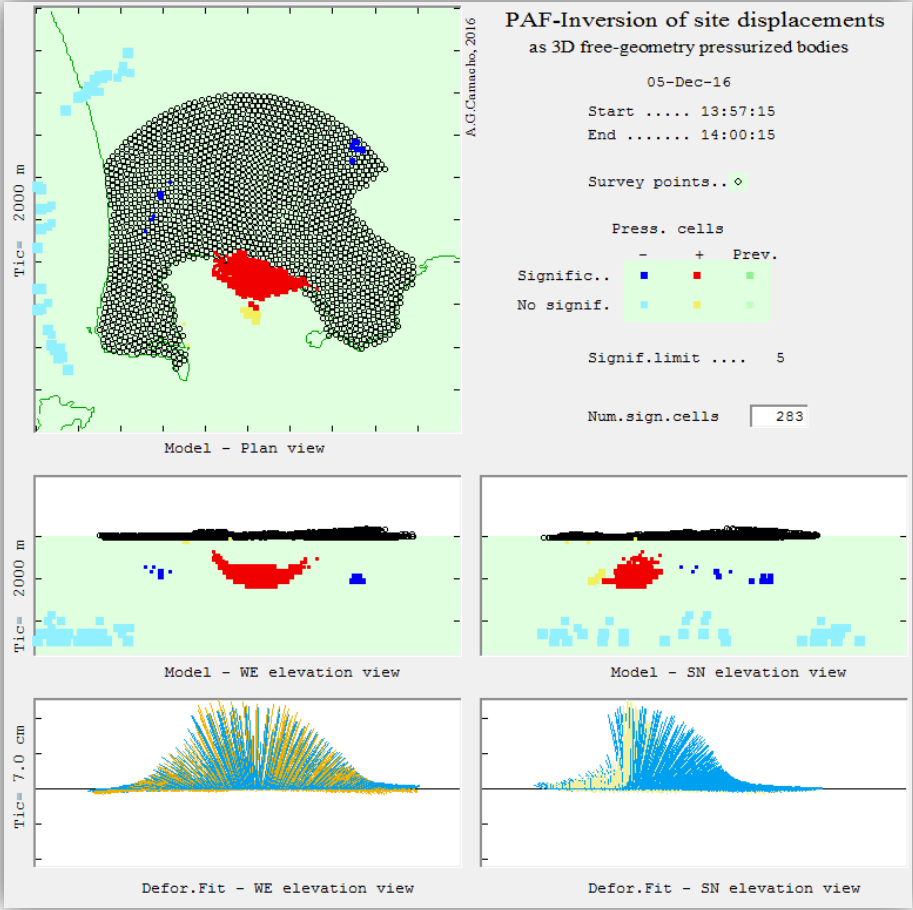


Figure 8. Selected pixels and vertical deformation (cm) (annual linear rate) for an inflation episode in Campi Flegrei (Italy). UTM coordinates. Uplift period 10/2007-06/2013.

About the smoothing coefficient λ , the general value 200 can be acceptable for usual cases. In this case of strong deformation, the signal/noise ratio is very high (see below, final residual level with respect to data scattering level), and the model does not require so high smoothing value. After some trials, we select 10 as suitable smoothing value. Nevertheless, the difference between models obtained with different inversion values is not critical (as a way of example, Figure 5 shows the model obtained with a value of 30 for smoothing and 8.0 MPa for accumulated pressure).

Once these values are selected, the rest of the inversion approach is nearly automatic. The number of epochs is one. We accept the general value 5 for the significance limit. Figure 9 shows the resulting model by means of planar and vertical cuts views. The model consists mostly of an anomalous body, at a mean depth about 1.8 km, composed of overpressure cells (red colour for significant cells and yellow for non-significant cells). Morphology of the anomalous structure as a “partially filled parabolic

261 glass” suggest a shallow (depth 1-2 km) hydrothermal system confined to the caldera fill materials.
 262 This result is very similar to that obtained in Camacho et al. (2011) for the uplift period 1992-2000.



263
 264 **Figure 9.** Resulting inversion model for the Campi Flegrei inflation data as subsurface anomalous
 265 structure (about 1.8 km mean depth) of aggregated overpressure cells.

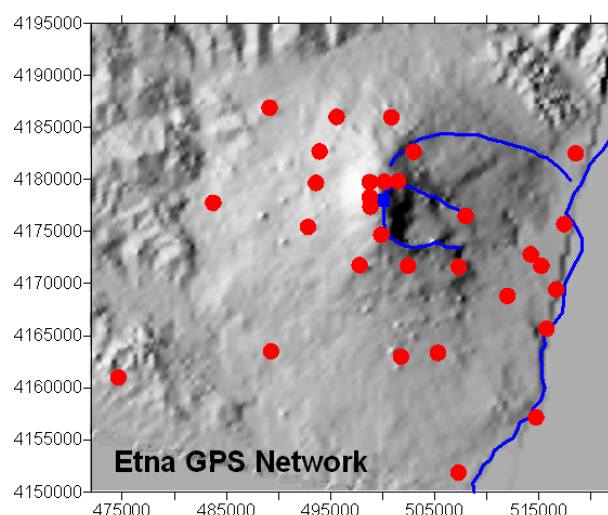
266 The final residual values after the inversion approach (root mean square values of 0.5 cm for Up,
 267 and EW components) are quite small with respect to the total scattering of the deformation data (3.3
 268 cm and 2.7 cm accumulated for Up and EW, respectively), confirming the suspected good ratio
 269 signal/noise in the present deformation case.

270 Figure 9 also shows two vertical views of the fit between observed (orange lines) and modelled
 271 (blue lines) deformations in the bottom. (Light yellow lines correspond to dz data without dy values,

272 the approach determines modelled values also for dy). In this figure, we can also observe the presence
 273 of several areas with non-significant cells (light blue and yellow ones), mostly in peripheral areas.
 274 These fictitious structures are introduced by the inverse approach to absorb distortion effects contained
 275 in the interferometric data. These fictitious bodies allow for getting the significant structures nearly
 276 free of distortion components in the data.

277 **3.2. Etna monitoring case.**

278 A second example regards the sequential application of the PAF software to a network of 31
 279 permanent GPS stations close to Etna mountain (Sicily, Italy) (Cannavò et al., 2015b and references
 280 on it), see Figure 10. Mt. Etna, situated on the eastern coast of Sicily, is a large basaltic volcano built
 281 up in a geodynamic setting generated during the Neogene convergence between the African and
 282 European plates (Allard et al., 2006; Branca et al., 2011) and is one of the most active volcanoes in the
 283 world. Its activity comprises strombolian activity, which may evolve into lava fountains and effusive
 284 events, and lateral eruptions occurring along fractures (Aloisi et al., 2006; Cannavò et al., 2015a).



285
 286 **Figure 10.** Location of a network of 31 permanent GPS stations close to Mount Etna volcano
 287 (Sicily, Italy). UTM coordinates are indicated.

288 We analyse the three component (Up, EW, NS) deformation data corresponding to 8 successive
289 fortnightly periods, from May/2007 to September/2007 (Cannavò et al., 2015b). The data are arranged
290 into sequential files, one for each period or epoch, containing the deformation at the network stations
291 for each period. Operatively, we could imagine receiving the data files just after finishing to collect
292 and process them in a quasi-real time manner.

293 The PAF software is applied taking into account the peculiarities of the case. The **CellsConfig.exe**
294 program suggests the following values: depth of the model top = 3131 m (above sea level), depth of
295 the model bottom = -15859 m (positive above sea level), and mean side of cells = 658 m. We accept
296 these default suggestions. Then, the resulting number of cells is 33250. Next, the program suggests the
297 usual general values for the smoothing coefficient (200) and for the pressure contrast (0.5MPa).

298 The suggested pressure value 0.5 MPa is small, and we accept it as suitable for this case
299 (deformations are not very large, and therefore causative anomalous pressure should be neither very
300 large).

301 However, in this case, for the considered monitoring network the general smoothing value 200
302 looks too small. In fact, in this case the number of data points is very reduced (31 stations), the
303 measured displacements are not large (except just in the case of eruption), and GPS data involve a
304 significant inaccuracy level (estimated about 0.1 - 0.3 cm in the fortnightly period) with respect to the
305 deformation level (displacement data scattering is about 0.3 - 0.6 cm) for the considered time periods.
306 Therefore, we need to introduce a bigger smoothing coefficient to control the high noise level and
307 avoid fictitious structures. After some trial and error running of the code we have select the value 600
308 as a suitable smoothing.

309 Then, after these selections, the inversion approach is nearly automatic. We use the usual
310 significance limit 5, and in this case we analyse 8 epochs, starting in the epoch 10 (so our data files
311 will be **D00010.txt**, **D00011.txt**,, **D00017.txt**). The program **InverPAF.exe** takes only some

seconds to carry out the inversion for each period, and to produce the drawings and the respective output files.

Usually, the displacement results from the monitoring network, for most of the deformation periods without an eruptive episode, do not show a significant deformation. However, we have selected a particular sequence of fortnightly periods (8 successive fortnightly periods, from May/2007 to September/2007) that shown an apparent intrusive episode. See Cannavò et al. (2015) and references therein. Figure 11 shows some pictures (coming from the program display) of the resulting model for the sequential study. We can observe some features. For initial epochs (e.g. epochs (a) and (b)), the model does not show significant pressure cells (red, for pressure increase, or dark blue, for pressure decrease). There are many filled cells, but in non-significant locations (yellow or light blue) (low-sensitive cells), which look artefact sources, and are introduced by the inversion program to model some global noisy pattern.

For the next epochs, (c) to (f), we appreciate the appearance of a significant body of high pressure. It ascends in the central zone with a certain inclination towards the east, reaching a minimum depth about 3 km below sea level. This structure is similar to that modelled in Cannavò et al. (2015). In addition, we appreciate a previous alignment of deep positive (but non-significant) cells from the west to the base of the significant body. Green cells in the figure indicates cells from the previous epoch, allowing to follow the evolution of the anomalous structures from one to another epoch.

Epochs (g) and (h) return to the non-significant values of pressure changes. For epoch (h), the program detects and reject an outlier observation data (denoted with red colour in the corresponding data panel in Figure 11).

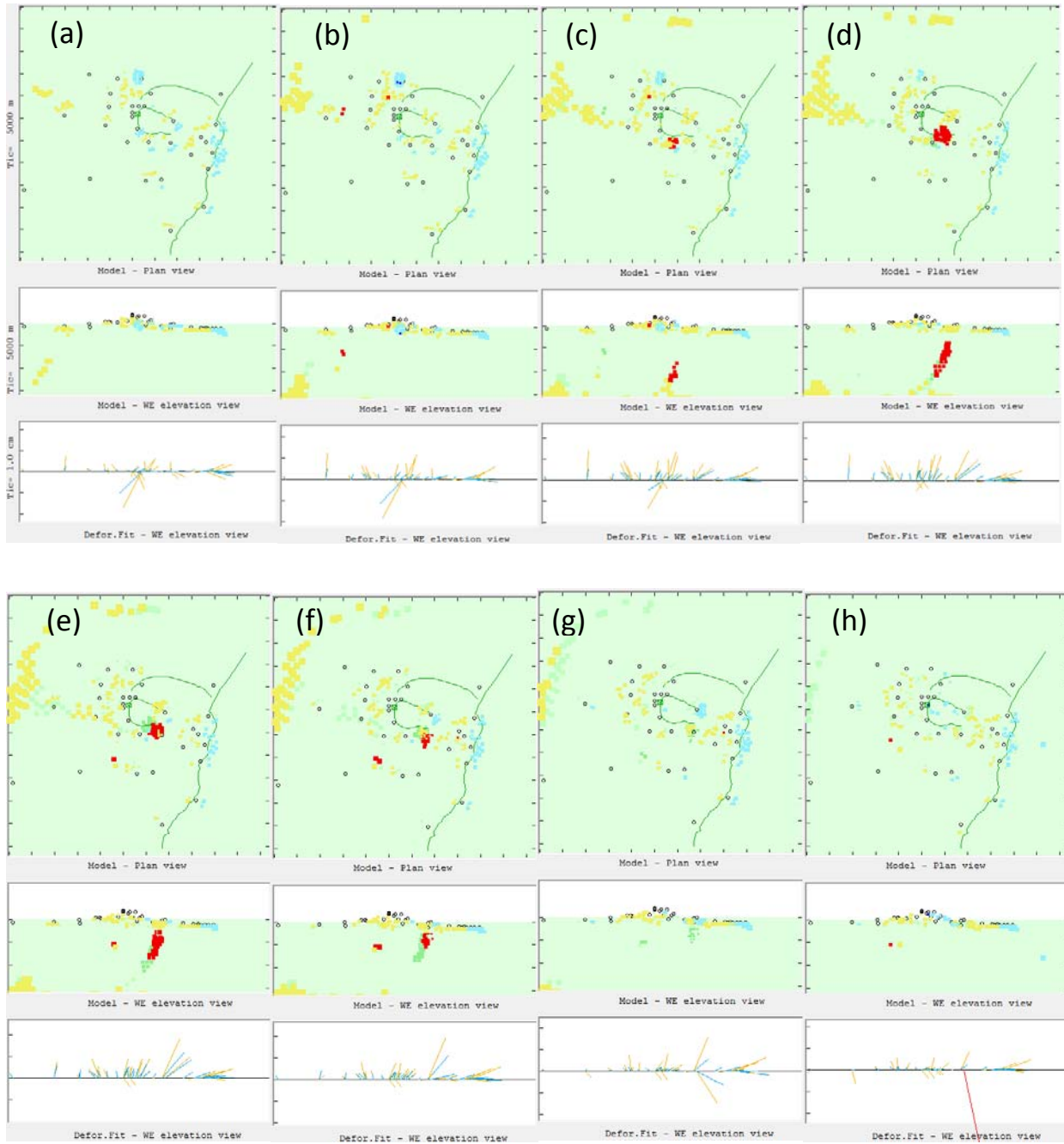


Figure 11. Planar and WE elevation views of the successive inversion models obtained in the sequential monitoring for the Etna network (biweekly data, from May to September 2007). Yellow and light blue cells indicate non-significant cells (fictitious structures to absorb data distortions). Epochs (c) to (f), with a significant high pressure body (red colour), suggest an intrusive episode. Tic step in axes correspond to 5 km.

4. Conclusions

PAF software allows carrying out an inversion, in a nearly automatic mode, of displacement data (1D to 3D) as extended 3D sources of overpressure bodies with free geometry. An elastic half-space response is assumed for the deformation calculus for each elementary source. The source bodies are described as aggregation of small cells filled with the prescribed pressure contrast. It works properly for volcanic areas, where deformations can be supposed as (mainly) due to subsurface pressure sources.

The software requires to be tuned by choosing suitable values for two main parameters in the inversion approach: the smoothing coefficient λ (which controls the balance between data fit and model regularity in the regularity conditions) and the basic pressure contrast (in MPa). Suitable values for both parameters are normally selected in a trial and error way, looking for good model features (regularity, size, etc.) and good residual distribution (null autocorrelation of residuals). Nevertheless, the choice of these parameters is not too critical with respect to the main features of the resulting model.

As shown in the application examples, the software can work for large data sets (for instance, coming from InSAR data), and for reduced sets of sequential data (for instance, coming from a small GPS monitoring network). It can work for fully 3D data, or also for 1D or 2D data (using e.g, GNSS, InSAR, or levelling data). For both application examples, geometry of the resulting models offers interesting features to evaluate the deformation causative phenomenon.

This software consists of two executable programs, and can be freely downloaded from the web page of the journal or asking to the corresponding author. The first program (**ConfigPAF.exe**) allows the user to set (in a dialog mode) the values for the inversion parameters, and for the geometrical configuration of the small cells. The second program (**InverPAF.exe**) works with the results of the former program, and does not require any user involvement (display of intermediate or final result is optional). It can be incorporated into a larger software tool for monitoring at a volcanic observatory

364 and even can run in an automatic way in real time, giving also very useful results (see Cannavò et al.,
365 2015). The source codes, user manual, executable files and a test example (including input data and
366 results) can be obtained on request to the corresponding author, under an open source license.

367 **Acknowledgments**

368 This research has been supported by the Spanish Ministry of Economy and Competitiveness grant
369 ESP2013-47780-557 C2-1-R and the EU 7th FP MED-SUV project (contract 308665). It is a
370 contribution to the Moncloa Campus of International Excellence.

371 **References.**

- 372 Allard, P., Behncke, B., D'Amico, S., Neri, M., and Gambino, S. (2006) Mount Etna 1993–2005:
373 anatomy of an evolving eruptive cycle, *Earth-Sci. Rev.*, 78, 85–114,
374 doi:10.1016/j.earscirev.2006.04.002.
- 375 Aloisi, M., Bonaccorso, A., and Gambino, S. (2006) Imaging compositive dike propagation (Etna,
376 2002 case), *J. Geophys. Res.*, 111, B06404, doi:10.1029/2005JB003908.
- 377 Battaglia, M., and Hill, D.P. (2009), Analytical modeling of gravity changes and crustal deformation
378 at volcanoes: The Long Valley caldera, California, case study, *Tectonophysics*, 471, 45-57, doi:
379 10.1016/j.tecto.2008.09.040.
- 380 Branca, S., Coltelli, M., and Groppelli, G. (2011) Geological evolution of a complex basaltic
381 stratovolcano: Mount Etna, Italy, *Ital. J. Geosci.*, 130, 306–317, 2011. doi: 10.3301/IJG.2011.13.
- 382 Camacho, A. G., González, P. J., Fernández, J., & Berrino, G. (2011) Simultaneous inversion of
383 surface deformation and gravity changes by means of extended bodies with a free geometry:
384 Application to deforming calderas. *J. Geophys. Res.*, 116, B10401. doi.org/10.1029/2010JB008165

385 Camacho, A. G., J. C. Nunes, E. Ortiz, Z. França, and R. Vieira (2007), Gravimetric determination of
386 an intrusive complex under the Island of Faial (Azores): Some methodological improvements,
387 *Geophys. J. Int.*, 171, 478–494, doi:10.1111/j.1365-246X.2007.03539.x.

388 Cannavò, F., Arena, A., & Monaco, C. (2015a). Local geodetic and seismic energy balance for shallow
389 earthquake prediction. *Journal of Seismology*, 19(1), 1-8, doi:10.1007/s10950-014-9446-z.

390 Cannavò, F.; Camacho; A.G.; González, P.J.; Mattia, M.; Puglisi, G; and Fernández, J. (2015b) Real
391 Time Tracking of Magmatic Intrusions by means of Ground Deformation Modeling during
392 Volcanic Crises, *Scientific Reports*, 5:10970. doi: 10.1038/srep10970

393 Charco, M., Fernández, J., Luzón, F., Tiampo, K.F., and Rundle, J.B. (2007), Some insights into
394 topographic, elastic a self-gravitation interaction in modelling ground deformation and gravity
395 changes in active volcanic areas, *Pure Appl. Geophys.*, 164, 865-878, doi 10.1007/s00024-004-
396 0190-y

397 Dzurisin D. (2003), A comprehensive approach to monitoring volcano deformation as a window on
398 the eruption cycle, *Rev. Geophys.*, 41(1), 1001, doi:10.1029/2001RG000107.

399 Lisowski, M. (2007), Analytical volcano deformation source models, in *Volcano Deformation*, chap.
400 8, pp. 279-304, Springer Praxis, Chichester, U.K.

401 Mogi, K. (1958) Relations between the eruption of various volcanoes and the deformation of the
402 ground surface around them. *Bull. Earthquake Res. Inst. Univ. Tokyo*, 36, 99-134.

403 Rymer, H and Williams-Jones, G. (2000). Volcanic eruption prediction: Magma chamber physics from
404 gravity and deformation measurements. *Geophys. Res. Lett.*, 27-16, 2389-2392. doi:
405 10.1029/1999GL011293

406 Samsonov, S. V., Tiampo, K. F., Camacho, A. G., Fernández, J., and González, P. J. (2014)
407 Spatiotemporal analysis and interpretation of 1993–2013 ground deformation at Campi Flegrei,
408 Italy, observed by advanced DInSAR. *Geophys. Res. Lett.*, 41, 6101–6108 (2014). doi:10.1002/
409 2014GL060595.

410 Willians, C.A., and Wadge, G. (1998), The effects of topography on magma chamber deformation
411 models: Application to Mt. Etna and radar interferometry, *Geophys. Res. Lett.*, 25, 1549-1552.