Literature Survey & Project Outline Accurate methods for computing rotation-dominated flows

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The shallow water equations (SWEs) are a simplified but very effective model for incompressible fluid flow in which the fluid depth is very small compared to the horizontal length scale of the domain. Among other applications, this model has been particularly successful in modelling large-scale geophysical flows of atmospheric and oceanic currents. For these applications, models will usually include source terms to represent the Coriolis force, due to the Earth's rotation, and non-uniform bathymetry or topography. Computational difficulties arise in steady flow when these very large forces balance pressure gradients exactly. This is referred to as a geostrophic balance, and large-scale real-world currents are close to this balance at all times. Therefore, the interest of this project lies in numerical methods which work around these computational problems and remain stable for geostrophic flows.

Literature Survey

A lot of effort has been put into studying and solving hyperbolic systems of conservation laws. Several textbooks about the numerical methods considered in this project exist, including LeVeque [1992], Toro [1999] and LeVeque [2002], as well as Toro [2001] which focuses on the SWEs in particular. In addition, there is a textbook on the rotating SWEs by Zeitlin [2007], Chapter 4 of which focuses on numerical methods.

Conservation laws are systems of partial differential equations (PDEs) which, in two dimensions, can be written in the form:

$$\mathbf{q}_t + \mathbf{f}(\mathbf{q})_x + \mathbf{g}(\mathbf{q})_y = \mathbf{s}(\mathbf{q}). \tag{1}$$

Here, \mathbf{q} is a vector of density functions of conserved quantities, \mathbf{f} and \mathbf{g} are flux terms, while \mathbf{s} stands for a number of source terms. Subscripts denote partial differentiation. For the components q_i to be conserved means that the integral $\int_{-\infty}^{\infty} \int_{-\infty}^{\infty} (q_i - s_i) \, \mathrm{d}x \, \mathrm{d}y$ is independent of time. The flux terms describe how the quantities \mathbf{q} are transported through the domain. Apart from actual sources or sinks the source terms \mathbf{s} may be used to model a variety of physical and geometric effects.

Such a system is called *hyperbolic* if any linear combination of the Jacobian matrices $\partial \mathbf{f}/\partial \mathbf{q}$ and $\partial \mathbf{g}/\partial \mathbf{q}$ has real eigenvalues.

The two-dimensional SWEs are a system of three partial differential equations in three conserved quantities: the water depth, h, and the two Cartesian components of the momentum, hu and hv (where u and v are the components of the velocity). The PDEs can be written as:

$$h_t + (hu)_x + (hv)_y = 0$$

$$hu + (hu^2 + \frac{1}{2}gh^2)_x + (huv)_y = -ghB_x + fhv$$

$$hv + (huv)_x + (hv^2 + \frac{1}{2}gh^2)_y = -ghB_y - fhu,$$

where g is acceleration due to gravity, B the bathymetry (bed elevation) and f is the Coriolis coefficient. These can be obtained from the Navier–Stokes equations by assuming that the depth of the water is small compared to some significant horizontal length-scale and by depth-averaging the flow variables. For a full derivation see Dellar and Salmon [2005]. Associating this with Eq. 1, the SWEs can be written in vector form using

$$\mathbf{q} = \begin{pmatrix} h \\ hu \\ hv \end{pmatrix}, \ \mathbf{f} = \begin{pmatrix} hu \\ hu^2 + \frac{1}{2}gh^2 \\ huv \end{pmatrix}, \ \mathbf{g} = \begin{pmatrix} hv \\ huv \\ hv^2 + \frac{1}{2}gh^2 \end{pmatrix}, \ \mathbf{s} = \begin{pmatrix} 0 \\ -ghB_x + fhv \\ -ghB_y - fhu \end{pmatrix},$$

which can indeed be shown to be hyperbolic (e.g. Toro [2001], Section 3.5).

For the purpose of this project only the above bathymetry and Coriolis source terms will be considered, but more advanced treatment of shallow water systems might include further terms to model other physical effects. Examples include bed friction, surface tension and eddy viscosity. If the SWEs are discretised on an irregular grid, geometric source terms might also be used which represent properties of the grid cells.

Such hyperbolic systems can be solved using finite volume methods in which the domain is divided up into a (not necessarily regular) grid of control volumes, and the conserved quantities are discretised by assuming that they are constant across each such volume. Note that "volume" is being used in a generalised sense here — for a two-dimensional system like the SWEs, the grid cells are actually areas. The most popular such method is due to Godunov $[1959]^1$ and is simply known as Godunov's method — in fact, today a whole family of Godunov methods has been developed based on the concepts derived in this paper. It has been thoroughly studied for homogeneous hyperbolic conservation laws, where $\mathbf{s} = 0$, and is based on the integral form of such systems. In the one dimensional case, this integral form is

$$\frac{\mathrm{d}}{\mathrm{d}t} \int_{x_1}^{x_2} \mathbf{q}(x) \, \mathrm{d}x = \mathbf{f}(\mathbf{q}(x_1)) - \mathbf{f}(\mathbf{q}(x_2)),$$

where x_1 and x_2 are the boundaries of a control volume. As opposed to the differential form, this integral form admits discontinuities, like hydraulic jumps.

Due to the piecewise-constant discretisation, every cell boundary x_i is now a simple step-function. Such an initial-value problem for conservation laws is known as a Riemann problem (see Toro [1999], Section 2.2.2, LeVeque [2002], Section 3.8, or Toro and Garcia-Navarro [2007], Section 2.1). By solving each of these Riemann problems, the discretised SWEs can be solved. The other insight of Godunov's method is that each hyperbolic system has a number of characteristic waves which only propagate in certain directions,

¹The author could not obtain an English translation of this Russian paper, but the method developed by Godunov has been extensively reiterated in papers and textbooks. Therefore, the following explanation of the method is based on what the author was able to find in those secondary sources.

which allows to simplify the computation by only looking at waves that can propagate *into* each cell. This technique is known as *upwinding*.

While it is possible to solve the Riemann problem exactly for many systems (including the SWEs), this dominates the computations necessary to solve each time step. Therefore, approximate solvers have been developed, the most popular one being due to Roe [1981]. See the textbooks mentioned above for an overview of exact and approximate Riemann solvers

Note that special care has to be taken with dry states in systems where negative values for one or more conserved quantities are unphysical. The SWEs are an example of such a system, as the water depth, h, is strictly non-negative. In fact, the SWEs do not even hold for regions where the depth is zero. Therefore, wet/dry fronts have to be dealt with differently than other cell boundaries. Modelling these is particularly important in geophysical flows, where they appear on beaches of protruding topography, or during outcropping of stratified flow. See Toro [2001], Chapter 6, for exact Riemann solvers in the presence of dry states.

There are several ways to apply Godunov's method in two dimensions. In general, a numerical scheme can be developed from the equivalent two-dimensional integral form:

$$\frac{\partial}{\partial t} \iint \mathbf{q} \, \mathrm{d}S = -\oint \mathbf{n} \cdot \mathbf{h} \, \mathrm{d}\ell,$$

where the left-hand integral is over each cell area, the right-hand integral around its boundary and $\mathbf{h} = (\mathbf{f}, \mathbf{g})$.

This usually yields the best results but is also computationally more intensive. A simpler approach is to apply a dimensional split, and alternate between solving one-dimensional Riemann problems along slices in the x and y directions, respectively (Toro [2001], Section 12.3.1, LeVeque [2002], Section 19.5). For unsplit two-dimensional solvers, see Toro [1999], Chapter 16, Toro [2001], Section 12.3.2 and LeVeque [2002], chapters 19–21.

A fairly recent review of Godunov-type methods was conducted by Toro and Garcia-Navarro [2007] and an older review can be found in Sweby [2001].

The previous discussion assumed homogeneous systems. However, the interest of this project does not lie in homogeneous systems, but in conservation laws with source terms. Similar to how a dimensional splitting can be applied, traditionally, hyperbolic systems with source terms were solved by splitting the system into two parts. The homogeneous hyperbolic PDEs:

$$\mathbf{q}_t + \mathbf{f}(\mathbf{q})_x + \mathbf{g}(\mathbf{q})_y = 0.$$

And a set of ordinary different equations (ODEs) for the source terms:

$$\mathbf{q}_t = \mathbf{s}(\mathbf{q}).$$

This way, the homogeneous system can be solved using well-studied Godunov-type methods, and the source terms can be solved independently be a simple integration in time, also using established methods like Runge-Kutta (originally developed by Runge [1895] and Kutta [1901]; see Kaw et al. [2009], Sections 8.3 and 8.4 for a modern account). See Toro [2001], Section 12.2.2 or LeVeque [2002], Sections 17.2.2 to 17.5, for instance.

However, this leads to problems in steady or quasi-steady scenarios, where $\mathbf{q}_t \approx 0$, such that the flux terms and source terms are balanced at all times. According to Toro and Garcia-Navarro [2007], the first authors to recognise this were Glimm et al. [1984]. To preserve the balance, this would require a time step in the hyperbolic system to be

cancelled exactly by the corresponding time step in the source ODE. Since these terms can in principle be very large, even for balanced systems, due to different methods being employed and numerical inaccuracies, this is practically impossible, such that equilibria cannot be modelled. Furthermore, small perturbations away from equilibrium would be completely dominated by said numerical errors. These problems are particularly relevant for large-scale geophysical flows, which are usually very close to geostrophic balance—which is an instance of balanced flux and source terms—at all times.

Therefore, a lot of research was conducted over the past two decades to develop socalled *well-balanced* methods which are able to preserve these equilibria exactly. To the best of the author's knowledge, Greenberg and LeRoux [1996] were the first to use the term "well-balanced".

Subsequently, dozens of well-balanced methods have been developed, including LeV-eque [1998], Garcia-Navarro and Vazquez-Cendon [2000], Hubbard and Garcia-Navarro [2000], Burguete and García-Navarro [2001], Gascón and Corberán [2001], Rogers et al. [2001], Bale et al. [2003], Rogers et al. [2003], Audusse et al. [2004], Chinnayya et al. [2004], Liang and Borthwick [2009], Liang and Marche [2009]. The most recent articles the author could find are by Zhang et al. [2014] and Chertock et al. [submitted 2014], the latter being of particular relevance here, as their assumptions align with those made in this project. Furthermore, Section 4.4 of Zeitlin [2007] presents a long list of other well-balanced methods applicable to the rotating SWEs and refers to the method discussed in Audusse et al. [2004] and related works as "the most classical [well-balanced] method". With Bouchut [2004], there is also a textbook focusing primarily on these methods.

Due to the vast amount of publications in this area, an in-depth review of these methods is beyond the scope of this preliminary literature survey. In fact, such a review will account for a large part of the overall project. At this stage, it should only be noted that a variety approaches exist, some of which employ a change of variables that allow standard Godunov methods to be applied and others which derive completely new (unsplit) methods based on the full inhomogeneous integral form of the equations.

Finally, an additional technique should be mentioned, which can be applied independently of the chosen Riemann solver: Adaptive mesh refinement (AMR) refers to using grids of different resolutions for different patches of the domain, and changing the resolution of each patch during solving based on the current state of the solution. The motivation behind this is to improve both the quality of the solution and the runtime and memory performance of the solver. Some features of the domain, like irregular boundaries, or propagating shocks, can benefit immensely from an increased resolution to mitigate aliasing effects of the discretisation. However, using the necessary high resolution across the entire domain can result in infeasible computation times. Therefore, using patches of different resolution, we can compute the critical parts of the domain as exactly as necessary while still using a low resolution on uniform regions to greatly reduce the computational effort.

Several AMR techniques have been developed, including quad-tree-based approaches, as discussed in Rogers et al. [2001], and so-called Berger-Oliger-Colella adaptive refinement, as introduced in Berger and Oliger [1984] and Berger and Colella [1989] and further extended in Berger and LeVeque [1998]. Since a detailed analysis of these methods is beyond the scope of this project, the reader is referred to Plewa et al. [2005] for a broad overview over the research done in this field.

Project Outline

The previous survey shows that there are many orthogonal concerns to be taken into account when selecting a solver for the shallow water equations. These include:

- What grid should be used for discretisation?
- Should adaptive mesh refinement be applied?
- Which source terms should be included in the model?
- How are the source terms to be balanced, or is it even feasible to use source splitting?
- Should wet/dry fronts be modelled?
- Should multilayer flow be modelled?
- Should a dimensional split be applied?

As the focus of this project lies on balancing source terms for applications in large-scale geophysical flows, most of the technical decisions will be made in favour of simpler implementation to begin with. The only source terms that will be considered are the Coriolis forcing term and the bathymetry term. The grid will be assumed to be rectangular and regular. Adaptive mesh refinement will be used in order to deal with non-rectangular boundaries. For simplicity, dry states will be ignored, and the fluid will be assumed to consist only of a single layer. Unless it produces noticeably inferior results, dimensional splitting will be used, as unsplit approaches require multiple solvers to be written for each approach.

On the basis of these decisions the project will be carried out as follows:

Initially, existing well-balanced Godunov-type methods will be carefully reviewed with focus on their assumptions, difficulty of implementation and suitability for use with Coriolis and bathymetry source terms, in particular. About one month of time will be allocated for this in-depth review.

Subsequently, based on these criteria, some of the methods will be selected and implemented in Fortran 90, using LeVeque's framework CLAWPACK². In particular, the AMRClaw version will be used, which implements an adaptive mesh refinement algorithm detailed in Berger and LeVeque [1998]. This will allow AMR to be used with all solvers without the need for any additional implementation effort. These implementations will then be tested against model problems with known steady solutions, such as a still fluid over varying bathymetry (under rotation) and a steadily rotating vortex of fluid over topography. This phase of the project is expected to take until the end of January.

Depending on the results of these tests, the remaining time can be spent on one or more of the following three tasks:

- Attempting to improve methods which performed badly on the model problems by using more advanced techniques in areas that were simplified initially (for instance, by replacing the dimensional split with an unsplit method).
- Attempting to model additional physical effects with methods that performed well. Here, the focus would lie on including an additional source term to model bed friction, and dealing with dry states and multi-layer fluids, as these are of particular importance in large-scale geophysical flows.

²http://www.clawpack.org/

• Using well-working solvers to look for steady or quasi-steady solutions in additional problems which are known to exhibit such solutions in the non-rotating case. This includes long-range breathers and Prandtl–Meyer expansion fans.

Thus, at the end of the project will stand a comprehensive review and comparison of existing well-balanced solvers for the shallow water equations, potentially with novel results for steady solutions under strong rotation.

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