

1 Dawn-dusk asymmetries in the auroral particle precipitation and their modulations by substorms
2 Simon Wing¹, Jay R. Johnson², and Enrico Camporeale³

³ ¹Applied Physics Laboratory, The Johns Hopkins University, Laurel, Maryland, USA

⁴Princeton Plasma Physics Laboratory, Princeton University, Princeton, New Jersey, USA

⁵Center for Mathematics and Computer Science (CWI), Amsterdam, The Netherlands

Abstract. Auroral particle precipitation exhibits dawn-dusk asymmetries that reflect the asymmetries in the particle populations, waves, and processes in the magnetosphere. The diffuse auroral electrons can be observed mainly in 22:00 – 09:00 MLT, which coincides much with the spatial distribution of the whistler-mode chorus waves that have been shown to be the predominant mechanism for pitch-angle scattering magnetospheric electrons into the loss cone. On the other hand, the monoenergetic auroral electrons can be observed at dusk-midnight sector. The monoenergetic electrons are magnetospheric electrons that have gone through a quasi-static parallel electric field in the upward field-aligned current regions. The broadband auroral electrons can be found mostly at 22:00 – 02:00 MLT where a peak in the Poynting flux of Alfvén waves is observed. Alfvén waves are known to cause broadband acceleration of electrons. There may be a connection between monoenergetic and broadband electrons in that the low frequency Alfvén wave–electron interaction can result in monoenergetic electron signature. Substorms increase the power of the diffuse, monoenergetic, and broadband electron aurora by 310%, 71%, and 170%, respectively. The duration of the substorm cycle for monenergetic and broadband auroral is ~5 hr, but it is larger than 5 hr for diffuse auroral electrons.

22

23

24 **Keywords:** particle precipitation, diffuse aurora, broadband aurora, monoenergetic aurora, ion
25 aurora, substorm cycle.

26 **Index terms:** 2704, 2790, 2407, 2455, 2483

27 **Key points:** (1) Broadband aurora increases at substorm onset, but it decreases rapidly 15 min
28 after onset. (2) Monoenergetic aurora start increasing approximately 75 min before the substorm
29 onset. (3) Monoenergetic and broadband electron aurorae complete the substorm cycle within 5
30 hr. (4) Diffuse electron aurora peak at 1 hr after the substorm onset.

31

32 **1. Introduction**

33 The Earth's magnetic field lines converge in the northern and southern polar regions.
34 Ions and electrons in the magnetosphere follow the magnetic field. Downgoing particles, that are
35 closely field-aligned, precipitate into the high latitude ionosphere. As a result, the particle
36 precipitation detected within the auroral oval can serve as a window to the magnetosphere.

37 In the magnetosphere, the magnetic field strength, and the ion and electron temperatures
38 increase with decreasing distance from Earth. Where the magnetic field is sufficiently strong or
39 the particle (ion and electron) temperatures are sufficiently high, typically at distance $r < 10 - 12$
40 R_E , the curvature and gradient drifts, which are temperature dependent, can play a significant
41 role in the transport of ions and electrons [e.g., *Wang et al.*, 2004]. The curvature and gradient
42 drifts move hotter ions westward and hotter electrons eastward, leading to dawn-dusk
43 asymmetries in the near-Earth plasma sheet and the inner magnetosphere. The ion temperatures
44 are higher at dusk-midnight sector than at midnight-dawn sector during the quiet time and
45 substorm growth phase [e.g., *Spence and Kivelson*, 1993; *Wing and Newell*, 2002; *Wing et al.*,
46 2005; *Johnson and Wing*, 2009; *Wang et al.*, 2001; 2006; 2011]. The electron temperatures
47 exhibit the opposite asymmetries with the temperature being higher at the midnight-dawn sector
48 than at the dusk-midnight sector. These dawn-dusk asymmetries can also manifest in the auroral
49 ion and electron precipitation in the ionosphere [e.g., *Newell et al.*, 2010; *Wing et al.*, 2013].

50 The dawn-dusk asymmetries of the auroral particle precipitation can also be modulated
51 by substorms, which inject into and energize particles at the inner magnetosphere. There are
52 generally three phases of a substorm: growth, expansion, and recovery [e.g., *Tanskanen et al.*,
53 2002; *Lui*, 1991]. The growth phase typically begins in the quiescent period at the time of the
54 southward turning of the interplanetary magnetic field (IMF) and ends at the onset of the

expansion phase (commonly referred to as the substorm onset), which is an interesting topic in its own right (e.g., *Lyons et al.*, [1997], *Birn and Hones* [1981], *McPherron* [1991], *Hsu and McPherron* [2004], *Meng and Liou* [2004], *Rostoker* [2002], *Lui* [2004], *Angelopoulos et al.* [2008], *Haerendel* [2010], *Sergeev et al.* [2012], *Johnson and Wing* [2014]). During the growth phase, the auroral oval expands equatorward, the aurora and the electrojet gradually intensify, the plasma sheet thins, and the magnetospheric magnetic field lines stretch (become tail-like), as the solar wind energy is stored in the magnetotail. During the expansion phase the auroral oval brightens and expands poleward, eastward, and westward, the westward electrojet significantly increases, and the magnetic field configuration in the inner plasma sheet changes rapidly from the stretched tail-like configuration to a more dipolar configuration. The expansion phase is followed by the recovery phase, during which the magnetosphere returns back to its original undisturbed state. The start of the recovery phase is usually signaled by the waning of the substorm aurora and weakening of the westward electrojet. The recovery phase ends when the magnetosphere reaches its normal undisturbed state.

In the present paper, the dawn-dusk asymmetries in the auroral ion and electron precipitation and how these asymmetries change throughout the substorm phases are presented. The time scales for the substorm cycle and each substorm phase are examined. We categorize the electron precipitation into three categories based on their spectra: diffuse, monoenergetic, and broadband auroral electrons [e.g., *Newell et al.*, 2009; 2010; *Wing et al.*, 2013]. These are described in Sections 2 – 4. The diffuse auroral electrons are most likely downgoing field-aligned plasma sheet electrons that precipitate in the ionosphere. The electrons in the loss cone (the field-aligned electrons) are replenished by the pitch angle scattering resulting primarily from electron interactions with the very low frequency (VLF) whistler-mode chorus waves [e.g.,

78 *Thorne, 2010; Reeves et al., 2009; Summers et al., 1998; Ni et al., 2011*]. When precipitating
79 electrons exhibit an intense and broad energy spectrum, they are classified as broadband. This
80 electron population is believed to primarily result from acceleration by Alfvén waves [*Chaston et*
81 *al., 2002; 2003; 2008*], which are often observed around the time of substorm dipolarization
82 [*Lessard et al., 2006*]. Monoenergetic electron spectra, on the other hand, likely indicate the
83 presence of a parallel electric field that accelerates the electrons downward and are usually
84 related to the global upward currents. The monoenergetic electrons may also result from the
85 electron interaction with low frequency Alfvén waves/ballooning modes that accelerate electrons
86 [e.g., *Pritchett and Coroniti, 2010; Damiano and Johnson, 2012*]. Precipitating ions are
87 magnetospheric ions that have been pitch angle scattered into the loss cone by current sheet
88 scattering [e.g., *Speiser, 1965; Lyons and Speiser, 1982; Sergeev et al., 1983; 1993*] and/or
89 interaction with electromagnetic ion cyclotron (EMIC) waves [e.g., *Jordanova et al., 2001;*
90 *Sergeev et al., 2015*]. Thus, these ions and the three types of electron spectra can provide
91 information about the particle populations, waves, and processes in the magnetosphere.

92

93 **2. Diffuse electron aurora**

94 Figure 1 presents energy flux maps from diffuse auroral electrons obtained from the
95 SSJ4/5 particle detector on board of DMSP F12 – F16 satellites in the interval 1996 – 2007. The
96 SSJ4/5 can detect ions and electrons with energy range of 30 eV – 30 keV. The DMSP satellites
97 are in sun-synchronous, nearly circular polar orbits at approximately 845 km altitude. The
98 SSJ4/5 detector apertures always point toward local zenith, which means that at high latitudes,
99 only highly field-aligned precipitating particles well within the atmospheric loss cone are
100 observed. The substorm onset times are taken from two substorm databases: (1) the Polar UVI

101 substorm database [Liou *et al.*, 1997; 2001] and (2) the IMAGE substorm database [Frey *et al.*,
102 2004]. Only isolated substorms that are separated by at least 5 hr from other substorms are used
103 to construct the figure. Out of 4861 substorm events in the original combined dataset, 1677
104 events, or about 34%, fall into this isolated substorm category. The diffuse electrons typically
105 have a Kappa distribution in the magnetotail [Christon *et al.*, 1991; Kletzing *et al.*, 2003].
106 However, in order not to leave out any electrons, the electrons that are not classified as either
107 monoenergetic or broadband (see Sections 3 and 4) are also classified as diffuse.

108 Figure 1 shows that the diffuse auroral electron energy flux increases around substorm
109 onset, consistent with the finding in Newell *et al.* [2010]. However, Figure 1 also shows that
110 after onset, the energy flux continues to increase and remains at an elevated level for at least 2 hr
111 after onset, reaching a maximum at about 1 hr after onset (this can be seen more easily in Figure
112 7, which is discussed in Section 6). Moreover, it appears that at and after onset, the increase of
113 the diffuse electron energy flux is confined approximately in the sector spanning 22:00–09:00
114 MLT. This dawn-dusk asymmetry is discussed further in Section 8.

115 The diffuse electrons are the magnetospheric electrons that precipitate into the loss cone
116 in the ionosphere. As the plasma sheet electrons $\mathbf{E} \times \mathbf{B}$ convect earthward, they also curvature
117 and gradient drift eastward toward dawn. The field-aligned component of these electrons are
118 quickly lost through the lost cone, but they are replenished by pitch-angle scattering. A leading
119 mechanism for pitch-angle scattering is VLF whistler-mode chorus wave–electron interactions
120 [e.g., Thorne, 2010; Reeves *et al.*, 2009; Summers *et al.*, 1998; Camporeale, 2015; Camporeale
121 *et al.*, 2015]. Upper band chorus waves provide the dominant scattering process for electrons
122 100 eV – 2 keV, while lower band chorus waves are most effective for scattering electrons >
123 ~2keV [Ni *et al.*, 2011]. Electrostatic electron cyclotron harmonic (ECH) waves can also pitch

124 angle scatter electrons, but *Ni et al.* [2011] found that ECH wave scattering rates are at least one
125 order of magnitude smaller than those by whistler-mode chorus waves. Studies have shown that
126 whistler-mode chorus waves are excited in the region spanning premidnight to noon, which
127 includes the region where the diffuse electrons are observed, at 22:00–09:00 MLT. This can be
128 seen in Figure 2, which shows the whistler-mode chorus wave and diffuse electron spatial
129 distributions. Around 09:00 MLT, the diffuse electron flux decreases, which may suggest that
130 the whistler-mode chorus waves start weakening. In the magnetosphere, the electrons continue
131 to drift eastward, circling the earth, but they are only observed in the ionosphere as diffuse
132 electrons when and where there are whistler-mode chorus waves to pitch-angle scatter them.
133 Because the bounce motion is generally faster than the drift motion, the signatures of the waves
134 in the ionosphere can be mapped to the magnetosphere at the same MLT. Whistler mode chorus
135 waves are important for the study of radiation belts because they have been found to contribute
136 to both energy and pitch angle diffusion [*Li et al.*, 2007; *Horne and Thorne*, 2003].

137

138 **3. Broadband electron aurora**

139 The precipitating electrons are classified as broadband electrons if three or more energy
140 channels have differential energy flux (dJ_E/dE) $> 2.0 \times 10^8$ (eV cm $^{-2}$ s $^{-1}$ sr $^{-1}$ eV $^{-1}$). There are
141 some caveats, which are listed in *Newell et al.* [2010] and are not repeated here. Figure 3
142 displays the broadband electron energy flux maps for 1 hr before onset through 1 h and 45 min
143 after onset in the same format as in Figure 1. The database and methodology used to construct
144 Figure 3 are also similar as those used to construct Figure 1 and are briefly discussed in Section
145 2. The broadband aurora rises after onset in the 21:00–02:00 MLT interval and appears to wane
146 or to start waning approximately 15 min after onset.

147 Broadband aurora is characterized by precipitating electrons having a broad energy
148 spectrum. Such precipitation is thought to result from electron interaction with dispersive Alfvén
149 waves when their frequency is comparable to the electron transit time [*Chaston et al.*, 2002,
150 2003, 2007; 2008; *Damiano et al.*, 2015]. Typical Poynting fluxes in Alfvén waves at high
151 latitude have been shown sufficient to account for 30-35% of auroral luminosity [*Wygant et al.*,
152 2002; *Keiling et al.*, 2002; 2003]. Figure 4 shows that the statistical local time distribution of
153 Alfvén waves is similar to that of the broadband electrons.

154 The association of Alfvén waves, substorm, and broadband electrons has been made by
155 *Lessard et al.* [2006]. The paper notes a connection between dipolarization events observed in
156 the magnetotail and dispersive Alfvén waves observed above the ionosphere, which are
157 associated with the broadband electron precipitation. Observations of Pi1B (irregular bursty
158 pulsations with periods from 1 to 40 sec) were detected by GOES 9, FAST, and at the ground in
159 conjunction with a substorm. While GOES detected compressional magnetic field fluctuations
160 along with dipolarization at geosynchronous orbit, FAST (which was conjugate to GOES)
161 detected shear Alfvén waves as a broad band ELF wave spectrum. The ratio of $\delta E/\delta B$ for the
162 waves was consistent with Doppler-shifted dispersive Alfvén waves that have been reported
163 [*Stasiewicz et al.*, 2000; *Chaston et al.*, 2008], suggesting that compressional waves mode
164 convert to dispersive shear Alfvén waves in this region. These same waves were also observed
165 by ground-based magnetometers on conjugate field lines.

166 Because transfer of energy by Alfvén waves is most efficient when the perpendicular
167 wavelength is small [*Hasegawa*, 1976; *Lysak and Song*, 2003; *Damiano et al.*, 2007], it is
168 additionally necessary that there be cascade of energy from large scales to small scales [*Chaston*
169 *et al.*, 2008]. Cross-scale coupling may result from linear phase mixing in inhomogeneity

170 [Lysak and Song, 2011; *Camporeale et al.*, 2012], nonlinear wave-wave cascade [*Vasconez et*
171 *al.*, 2014; *Schekochihin et al.*, 2009 and references therein] or by nonlinear wave-particle
172 interactions [*Damiano and Johnson*, 2012].

173 Electrons with broadband energy distribution are consistent with acceleration in a time-
174 varying parallel electric field [*Chaston et al.*, 2002] that is associated with small-scale dispersive
175 Alfvén waves. Electrons can be resonantly trapped in the wave potential of an Alfvén wave
176 pulse [*Kletzing*, 1994] typically leading to the development of a velocity-dispersed beam in front
177 of the pulse [*Watt et al.*, 2005]. At lower altitude, the electrons escape the potential well and
178 precipitate into the ionosphere as an energy-dispersed population [*Watt and Rankin*, 2009].

179 In the transient response models [e.g., *Nishida*, 1979; *Kan et al.*, 1982; *Hull et al.*, 2010],
180 the magnetospheric reconfiguration and diversion of the cross-tail currents by the current wedge
181 are communicated to the ionosphere by Alfvén waves. The broadband aurora that results from
182 the initial substorm pulse may be expected to last a few Alfvén bounce periods because Alfvén
183 waves damp kinetically on electrons absorbing most of the wave energy after a few reflections
184 via wave-particle interactions [*Lysak and Song*, 2003; *Damiano and Johnson*, 2012] and/or Joule
185 dissipation in the ionosphere [*Hull et al.*, 2010].

186 As shown in Figure 3, broadband auroral power peaks following onset and remains
187 elevated for about 15 minutes, consistent with the decay time of Alfvén waves. As shown later
188 in Sections 6 and 7, the broadband auroral power reaches its half maximum value in ~40 min
189 from the time it reaches its maximum value (the end of the expansion phase), which gives a
190 measure of how fast the auroral power and Alfvén waves decay. After the rapid decay of the
191 main Alfvén waves in the interval 15–45 min after onset, there seems to be residual Alfvén
192 waves that slowly decay starting at approximately 45 min after onset (as can be seen later in

193 Figure 7b).

194

195 **4. Monoenergetic electron aurora**

196 The algorithm for classifying monoenergetic electrons is as follows. (1) Identify the
197 differential energy flux peak, and subsequently look at the drops one and two energy channels
198 above and below the peak. If the differential energy flux drops to 30% or less of the peak within
199 these two steps (at energies above and below the peak), then the event is considered
200 monoenergetic. (2) The differential energy flux must be above 1.0×10^8 (eV cm $^{-2}$ s $^{-1}$ sr $^{-1}$ eV $^{-1}$)
201 at the peak channel. (3) If either the average energy is below 80 eV or the differential energy
202 flux peak is below 100 eV, the spectrum is not considered “accelerated.” Such events may result
203 from spacecraft charging (low acceleration potentials are excluded by this rule).

204 The monoenergetic auroral electron energy flux map can be constructed using the same
205 dataset and methodology described in Section 2. Figure 5 shows the monoenergetic auroral
206 electron energy flux in the same format as in Figures 1 and 3. Many of its features are quite
207 different from those in the diffuse and broadband aurorae. The monoenergetic aurora is
208 concentrated mainly in the dusk-midnight sector. The monoenergetic auroral electron energy
209 flux starts increasing approximately 1 hr and 15 min before onset and increases more
210 significantly at substorm onset, reaching a maximum at 15 min after onset. This can be more
211 clearly seen in Figure 7c and discussed in Section 6.

212 The monoenergetic electron energy flux reaches a minimum 1 hr and 15 min before onset
213 (as can be seen later in Figure 7c) and increases thereafter, suggestive of a correlation between
214 monoenergetic electron precipitation and the growth phase magnetic field configuration. As the
215 tail stretches during the growth phase, field-aligned currents intensify [McPherron, 1972;

216 *Watanabe and Iijima, 1993; Wing and Sibeck, 1997; Tsyganenko and Sibeck, 1994; Tsyganenko*
217 *et al., 1993; Zanetti and Potemra, 1986*]. In the upward field-aligned current regions, the
218 current-voltage relationship implies that the parallel potential drop would increase in order to
219 maintain higher currents by drawing more electrons downward [Knight, 1973]. Hence, an
220 increase in the monoenergetic electron aurora may simply be an indicator of elongation of the
221 tail that occurs during the growth phase. Figure 5 shows that the monoenergetic electrons can be
222 observed mainly in the dusk-midnight sector. This would be consistent with the increase in the
223 upward region-1 field-aligned current (R1) in the dusk-midnight sector during the growth phase.
224 In the midnight-dawn sector, R1 may also increase, but here R1 is downward and so fewer
225 monoenergetic electrons would be expected. Region-2 field-aligned current (R2) at the
226 midnight-dawn sector flows upward, but few monoenergetic electrons are observed at this
227 location. This may result from the higher electron density on the downside than on the duskside
228 magnetosphere due to the eastward curvature and gradient drifts of the electrons.

229 Another possibility for the increase in monoenergetic precipitation around substorm onset
230 is the development of low-frequency waves that accelerate electrons, but do not lead to global
231 instability. One such possibility is the kinetic-balloonning/interchange mode discussed by
232 *Pritchett and Coroniti [2010]*, which operates in a stretched-tail configuration with a minimum
233 in B_z . These modes are thought to be associated with interchange heads, which generate auroral
234 streamers and contribute to the monogenergetic electron precipitation. Many studies have shown
235 that an auroral streamer is a fast flow signature in the ionosphere [e.g., *Nakamura et al., 2001*;
236 *Sergeev et al., 2004*]. Fast flows have been attributed to reconnection leading to flux tubes
237 having depleted total entropy (S), which initiate unstable growth of ballooning and interchange
238 instabilities resulting in earthward propagation of flux tubes [e.g., *Birn et al., 2009; 2011*;

239 *McPherron et al., 2011; Wing and Johnson, 2009; 2010; Wolf et al., 2009*]. Fast flows have also
240 been attributed to current disruption leading to field-line collapse [*Lui, 1994; Wolf et al., 2009*].

241 Fast flows are also observed following substorm onset, in the expansion and even
242 recovery phases. For example, *Baumjohann et al.*, [1999] shows that the earthward fast flow
243 occurrence rate peaks between 0 and 60 min after onset depending on the GSM X location of the
244 fast flow. The superposition of these peaks may give the broad peak in the monoenergetic
245 auroral power seen in the interval 15–60 min after onset in Figure 7c. These fast flows can
246 launch low frequency global Alfvén waves that are associated with monoenergetic precipitation
247 [*Damiano and Johnson, 2012*] (the low frequency Alfvén wave connection to monoenergetic
248 electrons is further explored in Section 9). Regardless of how they are formed, fast flows have
249 been observed more frequently in the dusk-midnight than the midnight-dawn sector in the tail
250 [e.g., *McPherron et al., 2011*]. This dawn-dusk asymmetry has also been seen in RCM-
251 Equilibrium (RCM-E) simulation and has been attributed to the ion westward curvature and
252 gradient drifts [e.g., *Zhang et al., 2008*]. The dawn-dusk asymmetry is discussed further in
253 Section 8.

254

255 **5. Ion Aurora**

256 Precipitating ions are magnetospheric ions that have been pitch angle scattered into the
257 loss cone by current sheet scattering [e.g., *Speiser, 1965; Lyons and Speiser, 1982; Sergeev et al.,*
258 *1983; 1993; Wing et al., 2005*] and/or interaction with electromagnetic ion cyclotron (EMIC)
259 waves [e.g., *Jordanova et al., 2001; Sergeev et al., 2015*]. Figure 6 shows aurora ion energy flux
260 in the same format as Figures 1, 3, and 5. Apparently, ion aurora too is enhanced after onset, but
261 in comparison with electron aurorae of any type, the enhancement is smaller (note the different

262 color scale with respect to the previous figures). The enhancement appears to continue for at
263 least 1 hr and 45 min after onset as shown in Figure 6. Near the end of the growth phase, i.e., Δt
264 = –30 to 0 min, the ion energy flux peaks at 18:00–21:00 MLT sector, which can be attributed to
265 the ion curvature and gradient westward drifts toward dusk [e.g., *Spence and Kivelson*, 1993;
266 *Wing and Newell*, 2002; *Wing et al.*, 2005; *Johnson and Wing*, 2009; *Wang et al.*, 2001; 2006;
267 2011].

268 There is also a second peak at 02:00–05:00 MLT that can be seen throughout the
269 substorm cycle (sometimes the two peaks are close together that they can't be easily
270 distinguished). *Wing et al.* [2013] shows that the ion pressure exhibits a second peak at the same
271 location. These ion pressure and energy flux maxima can be attributed to a peak in the ion
272 number flux (not shown). The pressure peak around 02:00–05:00 MLT in the plasma sheet has
273 been previously observed and attributed to a density peak during active times [e.g., *Korth et al.*,
274 1999; *Wing and Newell*, 1998] and during growth phase [*Wing et al.*, 2007; *Wing and Johnson*,
275 2009]. The dawn density enhancement during high magnetic activity may result from the cold
276 solar wind ion entry on the dawn flank and flow stagnation when enhanced $\mathbf{E} \times \mathbf{B}$ and corotation
277 are nearly cancelled by the curvature and gradient drifts [e.g., *Friedel et al.*, 2001]. Another
278 possible mechanism also presumes solar wind ion entry along the dawn flank, but additionally,
279 an enhanced $\mathbf{E} \times \mathbf{B}$ pushing the solar wind origin ions closer to Earth where the flux tube volume
280 is smaller and hence ions have higher density [*Wang et al.*, 2003]. Figure 6 shows that the dawn
281 peak persists after onset. This may be associated with substorm injection and flow stagnation.
282

283 **6. Quantifying the nightside electron auroral power increase by substorms**

284 We examine the evolution of the nightside particle precipitation power quantitatively

285 from 2 hr before onset to 3 hr after. The solid line in Figure 7a shows the diffuse electron
286 auroral power for the entire nightside, 18:00–06:00 MLT (nightside hemispheric power), whereas
287 the dashed line is for the midnight-dawn sector, 00:00–06:00 MLT. The procedure used to
288 calculate these quantities is similar the one used in *Newell et al.* [2010]. The nightside diffuse
289 auroral power increases at onset and continues to increase after onset, reaching a maximum at 1
290 hr after onset. The nightside diffuse auroral power spans a huge range during the substorm
291 cycle, from approximately 1.2 GW at 15 min before onset to 4.9 GW at about 1 hr after onset.
292 Therefore, substorm-led magnetospheric reconfiguration typically increases the diffuse auroral
293 electron power by 310%. The nightside diffuse electron auroral power is dominated by the
294 power in the midnight-dawn sector. As can be seen from Figure 7a, the diffuse auroral power for
295 midnight-dawn sector (dashed line) constitutes approximately 70–80% of the power for the
296 entire nightside (solid line). The green asterisk marks the position of half maximum power,
297 which is based on the maximum power taken at the end of the expansion phase and the baseline
298 power taken at 15 min before onset.

299 In Figure 7b, the solid line plots the broadband auroral power integrated for the entire
300 nightside, 18:00–06:00 MLT, whereas the dashed line plots the power integrated from 21:00 to
301 02:00 MLT. Comparing the two curves, we can see that the power at 21:00–02:00 MLT
302 constitutes approximately 50–75% of the total nightside power. Substorms increase the
303 integrated power from 0.56 GW at 15 min before onset to 1.5 GW at 15 min after onset, which
304 represents a 170% increase. The broadband aurora peaks at 15 min after onset.

305 In Figure 7c, the solid line plots the evolution of the nightside electron monoenergetic
306 auroral power obtained by integrating from 18:00 to 06:00 MLT. The monoenergetic auroral
307 power reaches a minimum at 1 hr and 15 min before onset. The monoenergetic power increases

308 drastically at onset. After onset, the power continues to increase, but only for a short time,
309 reaching the maximum at 15 min after onset. The nightside monoenergetic auroral power is
310 dominated by the power in the dusk-midnight sector, 18:00–24:00 MLT, which is plotted as the
311 dashed line in Figure 7c. From the comparison of the solid and dashed lines, it can be seen that
312 approximately 60–75% of the monoenergetic nightside auroral power come from the dusk-
313 midnight sector. The substorm increases the monoenergetic auroral electron power by 71% from
314 1.05 GW at 15 min before onset to 1.8 GW at 15 min after onset.

315

316 **7. The durations of substorm cycle and phases from electron precipitation perspective**

317 Previous studies have reported greatly varying durations for the substorm cycle and
318 phases [e.g., *Huang et al.*, 2003; *Tanskanen et al.*, 2002; *Bargatze et al.*, 1999; *Baker et al.*,
319 1994; *McPherron et al.*, 1986; *Baker et al.*, 1981; *Pulkkinen et al.*, 1994]. Some of these
320 differences may be attributed, at least partly, to the differences in the instrumentations used to
321 make observations and the locations where the observation are made. These rich and diverse
322 observations help advance our understanding of the substorm processes. Here, we attempt to
323 estimate timescales of the substorm phases for the three different types of electrons using the
324 substorm events obtained from satellite UV images [*Liou et al.*, 1997; 2001; *Frey et al.*, 2004].
325 We define the minimum auroral power before onset as the start of the growth phase and the
326 maximum auroral power after onset as the end of the expansion phase/start of the recovery
327 phase.

328 The monoenergetic and broadband electron auroral powers complete the entire substorm
329 cycle approximately within 5-hr. That is, 3 hr after onset, these auroral powers finally reach
330 roughly the values to those at 2 hr before onset. On the other hand, the diffuse electron auroral

331 power appears to require more than 5 hr to complete the cycle.

332 The duration of the substorm cycle has been reported to be approximately 2–3 hr [e.g.,
333 *Huang et al.*, 2003] or 4 hr [e.g., *Tanskanen et al.*, 2002] based on substorm onsets determined
334 by the ground magnetic field observations or indices derived from these observations such as
335 AE, AL or IL. So, the durations of the substorm cycle for all electron aurorae appear larger than
336 or at the upper end of the range of the values obtained from ground magnetic field observations.
337 For the remainder of this section, we examine the duration of each substorm phase for the three
338 types of electrons.

339 The monoenergetic and broadband electron aurora growth phases start at approximately 1
340 hr and 15 min before onset, when the power reaches the minimum. The growth phase signature
341 for the diffuse electron aurora is not so clear. The minimum at 15 min before onset in the diffuse
342 auroral power in Figure 7a is not likely the start of the growth phase. All types of electron and
343 ion aurorae increase substantially at the substorm onset. So, the growth phase ends roughly at
344 the same time for all electron and ion aurorae. In other words, the substorm onsets obtained
345 from optical observations seem to agree with those obtained from the particle precipitation. The
346 1 hr and 15 min duration of the growth phase for the monoenergetic and broadband electron
347 aurorae is at the upper end of the range of the growth phase obtained from ground magnetic field
348 observations [e.g., *Bargatze et al.*, 1999; *Huang et al.*, 2003].

349 The end of the expansion phase can be defined as the time when the maximum power is
350 reached. The duration of the diffuse aurora expansion phase is longer than that of the other two
351 types of electron aurorae. The duration of the expansion phase of the diffuse aurora is 1 hr,
352 whereas that of the monoenergetic and broadband aurorae is only about 15 min. It is interesting
353 to note that the recovery phase onset for the electron diffuse aurora, ~1 hr after substorm onset, is

354 comparable to the start of the plasma sheet recovery reported in *Baker et al.* [1994]. They
355 reported that plasma sheet recoveries, e.g., expansion of plasma sheet, reduction of cross-tail
356 current, etc., can occur with a delay ranging 10–120 min after substorm onset with a median
357 delay of 45 min. However, their substorm onsets were determined from the ground magnetic
358 field observations. The short expansion phase duration in the broadband aurora may result from
359 the quick damping of the Alfvén waves as discussed in Section 3.

360 It is a challenge to determine the end of the recovery phase, primarily because it is hard to
361 determine when the quiet time state is reached and what the quiet time power ought to be. For
362 the broadband and monoenergetic electron aurorae, the powers at 3 hr after onset are
363 approximately the same as those at 2 hr before onset. The declining power in the interval 2–1 hr
364 before the substorm onset in the broadband and monoenergetic electron aurorae may suggest that
365 some of the points in this interval come from the recovery phase of the preceding substorm.
366 Assuming that (1) the minimum power before the substorm onset is the baseline for the quiet
367 time power and (2) many points in the interval 2–1 hr before onset could also be in the interval
368 3–4 hr after onset, given our criterion for isolated substorms and so the durations of the
369 monoenergetic and broadband electron aurora recovery phases can be estimated to be 3 hr and 30
370 min. It is harder to determine the recovery time for the diffuse electron aurora. Perhaps, some or
371 many of the points in the interval 0 – 2 hr before onset may actually come from the recovery
372 phase of the preceding substorm. This result may suggest that the recovery phase duration for
373 the diffuse electron aurora could be more than 4 hr, but in order to get a better estimate, one
374 would have to use substorms that are separated by more than 7 or 8 hr.

375 The recovery durations of all three types of the electron precipitation are larger than the
376 0.5–2 hr recovery duration obtained from the ground magnetic field observations [e.g., *Bargatze*

377 *et al.*, 1999; McPherron *et al.*, 1986; Baker *et al.*, 1981; Huang *et al.*, 2003; Horwitz, 1985].
378 Pulkkinen *et al.* [1994] reported that the recovery period of the geosynchronous magnetic field
379 and energetic particle observations is on the order of 1–3 hr. They attributed the long recovery
380 of the near-Earth magnetic field to the effect of the developing ring current. However, it is not
381 clear whether this can explain the long recovery period of the precipitating electrons, some of
382 which map farther out than geosynchronous orbit.

383 Table 1 summarizes the time scales for the substorm cycles and phases for the three types
384 of electrons.

385

386 **8. The electron auroral dawn-dusk asymmetry modulations by substorms**

387 Previous Sections present and discuss the electron energy fluxes for all three types of
388 electron aurorae. The dawn-dusk asymmetries in diffuse and monoenergetic electron aurorae
389 can be seen prominently in the electron energy fluxes in Figures 1, 2, and 5. In this section, we
390 examine more quantitatively the dawn-dusk asymmetries and how these asymmetries are
391 modulated by substorms.

392 Figure 8a plots the ratio of dawn (18:00 – 24:00 MLT) to dusk (24:00 – 06:00 MLT)
393 auroral power for monoenergetic (red), broadband (blue), and diffuse (green) electrons. The
394 monoenergetic electron auroral power is larger at dawn than at dusk and this asymmetry, as
395 measured by the dawn/dusk power ratio, increases after substorm onset. The asymmetry does
396 not return to its growth phase value until ~135 min after onset. The increase of the asymmetry
397 after the substorm onset may be related to the increase of Alfvén wave activities at premidnight
398 after onset as discussed in Sections 4 and 9.

399 The broadband electron auroral power does not show significant dawn-dusk asymmetry.

400 This can be seen by the dawn/dusk power ratio (blue line) that hovers around one in Figure 8a.

401 The diffuse electron auroral power has the opposite asymmetry from that of
402 monoenergetic electron auroral power with the power at dawn larger than that at dusk (the
403 dawn/dusk power ratio < 1). At 0 – 30 min after substorm onset, the dawn/dusk power ratio
404 increases a bit, suggesting a smaller dawn-dusk asymmetry. This asymmetry can be seen more
405 clearly in Figure 8b. The smaller dawn-dusk asymmetry after onset may be attributed to the
406 plasma sheet electron injection and energization that occurs over a wide local time after onset,
407 including at premidnight. The dawn/dusk power ratio minimum at –45 min before onset may be
408 due to the interference from the long recovery phase of the preceding substorm (> 4 hr).

409

410 **9. Is there any link between broadband and monoenergetic electrons?**

411 As shown in Table 1 and Figures 7b and 7c, the substorm dynamics of broadband and
412 monoenergetic electrons are more similar to each other than to those of diffuse electrons. For
413 example, both broadband and monoenergetic electron powers peak 15 min after onset whereas
414 diffuse electron power peaks 1 hr after onset. There may be a link between the mechanisms for
415 broadband and monoenergetic electrons.

416 Substorms increase the Alfvén wave activities in the magnetotail [e.g., *Lessard et al.*,
417 2006]. The high frequency Alfvén wave interaction with electrons lead to time-varying parallel
418 electric fields that accelerate electrons, resulting in the broadband signature in the electrons
419 [*Chaston et al.*, 2002; 2003; 2008]. This is consistent with the broadband electron energy flux
420 increase after substorm onset around 21:00 – 02:00 MLT, as shown in Figure 7b. However,
421 Alfvén waves can also be responsible for the monoenergetic electrons. For example, *Hull et al.*
422 [2010] suggests that the Alfvén waves can lead to the formation of density cavities and quasi-

423 static parallel electric fields. The low frequency Alfvén waves can accelerate electrons that lead
424 to the monoenergetic signature [Damiano and Johnson, 2012; Pritchett and Coroniti, 2010].
425 After a few Alfvén bounce periods, the Alfvén waves damp due to electron energy absorption
426 and/or Joule dissipation [Hull *et al.*, 2010; Lysak and Song, 2003; Damiano and Johnson, 2012].
427 As a result, the broadband auroral electron power decays after a few Alfvén bounce periods (~ 15
428 min), as seen in Figure 7b. On the other hand, the low frequency Alfvén waves damp more
429 slowly, which is consistent with the slower decay of the monoenergetic electron power seen in
430 Figure 7c. The decay time scales can be illustrated with the time it takes the power to reach its
431 half maximum value from the maximum value at the end of the expansion phase. As can be seen
432 in Figure 7 and summarized in Table 1, the time it takes to reach half maximum is ~ 1.6 hr for the
433 monoenergetic electrons, but only ~ 0.42 hr for broadband electrons.

434 Figure 7b shows that after reaching its maximum value, the broadband electron power
435 initially decays rapidly and then slowly. This two stage decay of the broadband electron power
436 suggests that after the rapid decay of the main Alfvén waves in the interval 15–45 min after
437 onset, there seems to be residual Alfvén waves that slowly decay that can be seen at
438 approximately 45 min after onset. These residual Alfvén waves seem to linger on for the rest of
439 the recovery period, which has the same time scale as the recovery period for the monoenergetic
440 electrons.

441 However, a significant amount of monoenergetic electrons are likely produced by quasi-
442 static electric fields that can be attributed to mechanisms other than low frequency Alfvén waves.
443 For example, in upward field-aligned current regions, quasi-static electric fields can arise when
444 the magnetospheric electron density is too low to carry the currents [Knight, 1973]. Figures 3
445 and 5 show that monoenergetic and broadband electrons are not always observed in the same

446 region. They seem to overlap roughly in the region spanning 21:00–01:00 MLT. Westward of
447 the overlapping region, e.g., $\text{MLT} < 21:00$, monoenergetic electrons can be observed without
448 significant broadband electrons.

449

450 **10. Summary**

451 Substorms change the magnetospheric configuration, e.g., when the magnetic field lines
452 change from stretched tail-like to more dipolar configurations. During this process, a significant
453 amount of energy is released, some of which energizes precipitating particles. On the nightside,
454 ion and all three types of electron (diffuse, monoenergetic, and broadband) energy fluxes and
455 powers increase at or shortly after substorm onset. However, the energy increases differ for each
456 type of aurora. The increases are 71%, 170%, and 310% for the monoenergetic, broadband, and
457 diffuse electron auroral powers, respectively. In contrast, the ion pressure increases only by 30%
458 [Wing et al., 2013]. At the end of the expansion phase (maximum power), different types of
459 electron auroral power decays at different rates. The time it takes to reach half maximum power
460 from the maximum for the broadband, diffuse, and monoenergetic electrons are ~ 0.42 , ~ 1.2 , and
461 ~ 1.6 hr, respectively. Among the electron aurorae, the broadband aurora has the smallest power,
462 followed by the monoenergetic aurora, whereas the diffuse aurora has the largest power. The ion
463 auroral power and energy flux are comparable to those of the broadband electron aurora during
464 the growth phase, but after the onset the broadband electron auroral power and energy fluxes
465 increase much more than those of the ion aurora. Relative to their pre onset values, substorms
466 appear to energize ion aurora less than electron aurora of any kind.

467 The MLT distribution of each type of aurora also differs. The diffuse electrons and
468 auroral power concentrate mainly in 22:00–09:00 MLT, which overlap much with the whistler-

469 mode chorus waves that can pitch-angle scatter magnetospheric electrons into the loss cone.
470 However, the monoenergetic auroral power mainly from the dusk-midnight sector, 18:00-24:00
471 MLT, whereas the broadband auroral power comes mainly from the region centered at
472 premidnight, spanning roughly 21:00 to 02:00 MLT. The broadband electrons result from time
473 varying electric fields due to electron-Alfvén wave interaction. On the other hand, the
474 monoenergetic electrons are magnetospheric electrons that have gone through quasi-static
475 parallel electric field in the upward field-aligned current regions. However, there may be a
476 connection between the mechanism for monoenergetic and broadband electrons – low frequency
477 Alfvén wave-electron interaction can lead to monoenergetic electron signature.

478 After substorm onset, the dawn-dusk asymmetry increases for about 135 min in the
479 monoenergetic electron power. In contrast, after the onset, the dawn-dusk asymmetry decreases
480 for about 30 min in the diffuse electron power. In comparison, substorms do not change the
481 dawn-dusk asymmetry of the broadband electron power by much.

482 Near the end of the growth phase, i.e., $\Delta t = -30$ to -15 min, the ion energy flux peaks at
483 18:00–21:00 MLT sector, which can be attributed to the ion curvature and gradient westward
484 drifts toward dusk. There is also a second peak at 02:00–05:00 MLT, which can be attributed to
485 cold solar wind ion entry along the dawn flank, flow stagnation, and enhanced $\mathbf{E} \times \mathbf{B}$.

486 The substorm cycle time scales for electron precipitation are longer than those previously
487 reported for other magnetospheric and ionospheric parameters. The substorm complete cycle
488 duration for monoenergetic and broadband electrons is ~ 5 hr, but it is greater than 5 hr for
489 diffuse electrons.

490

491 **Acknowledgments.** Gordon Wilson and Air Force Research Lab have been helpful in acquisition

492 of DMSP SSJ/4/5 data, as has the World Data Center in Boulder, Colorado. Simon Wing
493 gratefully acknowledges support from NSF Grant AGS-1058456 and NASA Grants
494 (NNX13AE12G, NNX15AJ01G). Jay R. Johnson acknowledges support from NASA Grants
495 (NNH11AR07I, NNX14AM27G, NNH14AY20I), NSF Grants (ATM0902730, AGS-1203299),
496 and DOE contract DE-AC02-09CH11466.

497

498

References

- 499 Angelopoulos et al. (2008), Tail reconnection triggering substorm onset, *Science*, 321(5891),
500 931-935, DOI: 10.1126/science.1160495.
- 501 Baker, D. N., E. W. Hones, Jr., P. R. Higbie, and R. D. Belian (1981), Global properties of the
502 magnetosphere during a substorm growth phase: A case study, *J. Geophys. Res.*, 86, 8941–
503 8956.
- 504 Baker, D. N., T. I. Pulkkinen, E. W. Hones Jr., R. D. Belian, R. L. McPherron, and V.
505 Angelopoulos (1994), Signatures of the substorm recovery phase at high-altitude
506 spacecraft, *J. Geophys. Res.*, 99, 10,967-10,979.
- 507 Bargatze, L. F., T. Ogino, R. L. McPherron, and R. J. Walker (1999), Solar wind magnetic field
508 control of magnetospheric response delay and expansion phase onset timing, *J. Geophys.*
509 *Res.*, 104, 14,583-14,599.
- 510 Baumjohann, W., M. Hesse, S. Kokubun, T. Mukai, T. Nagai, and A. A. Petrukovich (1999),
511 Substorm dipolarization and recovery, *J. Geophys. Res.*, 104(A11), 24,995–25,000,
512 doi:10.1029/1999JA900282.
- 513 Birn, J., and E. W. Hones Jr. (1981), Three-dimensional computer modeling of dynamic
514 reconnection in the geomagnetic tail, *J. Geophys. Res.*, 86, 6802–6808,
515 doi:10.1029/JA086iA08p06802.
- 516 Birn, J., M. Hesse, K. Schindler, and S. Zaharia (2009), Role of entropy in magnetotail
517 dynamics, *J. Geophys. Res.*, 114, A00D03, doi:10.1029/2008JA014015.
- 518 Birn, J., R. Nakamura, E. V. Panov, and M. Hesse (2011), Bursty bulk flows and dipolarization
519 in MHD simulations of magnetotail reconnection, *J. Geophys. Res.*, 116, A01210,
520 doi:10.1029/2010JA016083.

- 521 Camporeale, E., G. L. Delzanno, and P. Colestock (2012), Lower hybrid to whistler mode
522 conversion on a density striation, *J. Geophys. Res.*, 117, A10315,
523 doi:[10.1029/2012JA017726](https://doi.org/10.1029/2012JA017726).
- 524 *Camporeale, E. (2015), Resonant and nonresonant whistlers-particle interaction in the radiation*
525 *belt. Geophys. Res. Lett.*, 42, 3114–3121. doi: 10.1002/2015GL063874.
- 526 Camporeale, E., & Zimbardo, G. (2015). Wave-particle interactions with parallel whistler waves:
527 nonlinear and time-dependent effects revealed by Particle-in-Cell simulations, *Phys.*
528 *Plasmas* 22, 092104, doi.org/10.1063/1.4929853.
- 529 Chaston, C., et al. (2008), Turbulent heating and cross-field transport near the magnetopause
530 from THEMIS, *Geophys. Res. Lett.*, 35, L17S08, doi:10.1029/2008GL033601.
- 531 Chaston, C. C., C. W. Carlson, J. P. McFadden, R. E. Ergun, and R. J. Strangeway (2007), How
532 important are dispersive Alfvén waves for auroral particle acceleration?, *Geophys. Res.*
533 *Lett.*, 34, L07101, doi:10.1029/2006JG029144.
- 534 Chaston, C. C., J. W. Bonnell, C. W. Carlson, J. P. McFadden, R. E. Ergun, and R. J. Strangeway
535 (2003), Properties of small-scale Alfvén waves and accelerated electrons from FAST, *J.*
536 *Geophys. Res.*, 108, 8003, doi:10.1029/2002JA009420.
- 537 Chaston, C. C., J. W. Bonnell, L. M. Peticolas, C. W. Carlson, J. P. McFadden, and R. E. Ergun
538 (2002), Driven Alfvén waves and electron acceleration, *Geophys. Res. Lett.*, 29, 1535.
- 539 Christon, S. P., D. J. Williams, D. G. Mitchell, C. Y. Huang, and L. A. Frank (1991), Spectral
540 Characteristics of Plasma Sheet Ion and Electron Populations during Disturbed
541 Geomagnetic Conditions, *J. Geophys. Res.*, 96(A1), 1–22, doi:10.1029/90JA01633.
- 542 Damiano, P. A., A. N. Wright, R. D. Sydora, and J. C. Samson (2007), Energy dissipation via
543 electron energization in standing shear Alfvén waves, *Phys. Plasmas*, 14, 062,904.

- 544 Damiano, P. A., and J. R. Johnson (2012), Electron acceleration in a geomagnetic Field Line
545 Resonance, *Geophys. Res. Lett.*, 39, L02102, doi:10.1029/2011GL050264.
- 546 Damiano, P. A., J. R. Johnson, and C. C. Chaston (2015), Ion temperature effects on magnetotail
547 Alfvén wave propagation and electron energization, *J. Geophys. Res.*, 120, 5623–5632,
548 doi:[10.1002/2015JA021074](https://doi.org/10.1002/2015JA021074).
- 549 Frey, H. U., S. B. Mende, V. Angelopoulos, and E. F. Donovan (2004), Substorm onset
550 observations by IMAGE-FUV, *J. Geophys. Res.*, 109, A10304,
551 doi:10.1029/2004JA010607.
- 552 Friedel, R. H. W., H. Korth, M. G. Henderson, and M. F. Thomsen (2001), Plasma sheet access
553 to the inner magnetosphere, *J. Geophys. Res.*, 106, 5845-5858.
- 554 Haerendel, G. (2010), Equatorward moving arcs and substorm onset, *J. Geophys. Res.*, 115,
555 A07212, doi:[10.1029/2009JA015117](https://doi.org/10.1029/2009JA015117).
- 556 Hasegawa, A., Particle acceleration by MHD surface wave and formation of the aurora, *J.
557 Geophys. Res.*, 81, 5083, 1976.
- 558 Hodges, J. L., Jr., and E. L. Lehmann (1967), *J. Am. Stat. Ass.*, 62, 319, 926–931.
- 559 Horne, R. B., and R. M. Thorne (2003), Relativistic electron acceleration and precipitation
560 during resonant interactions with whistler-mode chorus, *Geophys. Res. Lett.*, 30, 1527,
561 doi:[10.1029/2003GL016973](https://doi.org/10.1029/2003GL016973), 10.
- 562 Horwitz, J. L. (1985), The substorm as an internal magnetospheric instability: Substorms and
563 their characteristic time scales during intervals of steady interplanetary magnetic field, *J.
564 Geophys. Res.*, 90, 4164–4170.
- 565 Huang, C.-S., G. D. Reeves, J. E. Borovsky, R. M. Skoug, Z. Y. Pu, and G. Le (2003), Periodic
566 magnetospheric substorms and their relationship with solar wind variation, *J. Geophys.*

- 567 Res., 108(A6), 1255, doi:10.1029/2002JA009704.
- 568 Hull, A. J., M. Wilber, C. C. Chaston, J. W. Bonnell, J. P. McFadden, F. S. Mozer, M. Fillingim,
569 and M. L. Goldstein (2010), Time development of field-aligned currents, potential drops,
570 and plasma associated with an auroral poleward boundary intensification, *J. Geophys.*
571 Res., 115, A06211, doi:10.1029/2009JA014651.
- 572 Hsu, T.-S., and R. L. McPherron (2004), Average characteristics of triggered and nontriggered
573 substorms, *J. Geophys. Res.*, 109, A07208, doi:10.1029/2003JA009933.
- 574 Johnson, J. R., and S. Wing (2009), Northward interplanetary magnetic field plasma sheet
575 entropies, *J. Geophys. Res.*, 114, A00D08, doi:10.1029/2008JA014017.
- 576 Johnson, J. R., and S. Wing (2014), External versus internal triggering of substorms: An
577 information-theoretical approach, *Geophys. Res. Lett.*, 41, 5748–5754,
578 doi:[10.1002/2014GL060928](https://doi.org/10.1002/2014GL060928).
- 579 Jordanova, V. K., C. J. Farrugia, R. M. Thorne, G. V. Khazanov, G. D. Reeves, and M. F.
580 Thomsen (2001), Modeling ring current proton precipitation by electromagnetic ion
581 cyclotron waves during the May 14–16, 1997, storm, *J. Geophys. Res.*, 106(A1), 7–22,
582 doi:10.1029/2000JA002008.
- 583 Kan, J. R., D. U. Longenecker, and J. V. Olson (1982), A transient response model of Pi 2
584 pulsations, *J. Geophys. Res.*, 87(A9), 7483–7488, doi:10.1029/JA087iA09p07483.
- 585 Keiling, A., J. R. Wygant, C. Cattell, W. Peria, G. Parks, M. Temerin, F. S. Mozer, C. T. Russell,
586 and C. A. Kletzing (2002), Correlation of Alfvén wave Poynting flux in the plasma sheet at
587 4–7 RE with ionospheric electron energy flux, *J. Geophys. Res.*, 107, 1132,
588 doi:10.1029/2001JA900140.
- 589 Keeling, A., J. R. Wygant, C. A. Cattell, F. S. Mozer, and C. T. Russell (2003), The global

- 590 morphology of wave Poynting flux: Powering the aurora, *Science*, 299, 383-386,
591 DOI: 10.1126/science.1080073.
- 592 Kenney, J. F., and E. S. Keeping (1951), Mathematics of Statistics part two, 2nd edition, pp. 370–
593 371, D. V. Nostrand Co. Inc., Princeton, NJ, USA.
- 594 Kletzing, C. A. (1994), Electron Acceleration by Kinetic Alfvén Waves, *J. Geophys. Res.*,
595 99(A6), 11,095–11,103, doi:10.1029/94JA00345.
- 596 Kletzing, C. A., J. D. Scudder, E. E. Dors, and C. Curto (2003), The auroral source region:
597 plasma properties of the high altitude plasma sheet, *J. Geophys. Res.*, 108, 1360,
598 doi:10.1029/2002JA009678.
- 599 Knight, L. (1973), Parallel electric fields, *Planet. Space Sci.*, 21, 741-750.
- 600 Korth, H., M. F. Thomsen, J. E. Borovsky, and D. J. McComas (1999), Plasma sheet access to
601 geosynchronous orbit, *J. Geophys. Res.*, 104, 25,047-25,061.
- 602 Lessard, M. R., E. J. Lund, S. L. Jones, R. L. Arnoldy, J. L. Posch, M. J. Engebretson, and K.
603 Hayashi (2006), Nature of Pi1B pulsations as inferred from ground and satellite
604 observations, *Geophys. Res. Lett.*, 33, L14108, doi:10.1029/2006GL026411.
- 605 Li, W., Y. Y. Shprits, and R. M. Thorne (2007), Dynamic evolution of energetic outer zone
606 electrons due to wave-particle interactions during storms, *J. Geophys. Res.*, 112, A10220,
607 doi:[10.1029/2007JA012368](https://doi.org/10.1029/2007JA012368).
- 608 Liou, K., P. T. Newell, C.-I. Meng (1997), A. T. Y. Lui, M. Brittnacher, and G. Parks, Dayside
609 auroral activity as a possible precursor of substorm onsets: a survey using POLAR UVI
610 imagery, *J. Geophys. Res.*, 102, 19835-19844.
- 611 Liou, K., P. T. Newell, D. G. Sibeck, C.-I. Meng, M. Brittnacher, and G. Parks (2001),
612 Observation of IMF and seasonal effects in the location of auroral substorm onset, *J.*

- 613 *Geophys. Res.*, 106, 5799.
- 614 Lui, A. (1994), Mechanisms for the substorm current wedge, in Substorms 2: Proceedings of the
615 Second International Conference on Substorms, Fairbanks, Alaska, March 7–11, 1994,
616 edited by J. R. Kan, J. D. Crave, and S.-I. Akasofu, pp. 195–203, Univ. of Alaska,
617 Fairbanks, Alaska.
- 618 Lui, A. T. Y. (1991), Extended consideration of a synthesis model for magnetospheric
619 substorms, in *Magnetospheric Substorms*, pp. 43-60, *Geophys. Monog.*, 64, edited by J. R.
620 Kan, T. A. Potemra, S. Kokubun, and T. Iijima, AGU, Washington D.C.,
- 621 Lui, A. T. Y. (2004), Potential plasma instabilities for substorm expansion onsets, *Space Sci.
622 Rev.*, 113, 127 – 206, doi:10.1023/B:SPAC.0000042942.00362.4e.
- 623 Lyons, L. R. and T. W. Speiser (1982), Evidence for current sheet acceleration in the
624 geomagnetic tail, *J. Geophys. Res.*, 87(A4), 2276–2286, doi:10.1029/JA087iA04p02276.
- 625 Lyons, L. R., G. T. Blanchard, J. C. Samson, R. P. Lepping, T. Yamamoto, and T. Moretto
626 (1997), Coordinated observations demonstrating external substorm triggering, *J. Geophys.
627 Res.*, 102, 27,039–27,052, doi:10.1029/97JA02639.
- 628 Lysak, R. L., and Y. Song (2003), Kinetic theory of the Alfvén wave acceleration of auroral
629 electrons, *J. Geophys. Res.*, 108, 8005, doi:10.1029/2002JA009406.
- 630 Lysak, R. L., and Y. Song (2011), Development of parallel electric fields at the plasma sheet
631 boundary layer, *J. Geophys. Res.*, 116, A00K14, doi:10.1029/2010JA016424, [printed
632 117(A1), 2012].
- 633 McPherron, R. L. (1972), Substorm related changes in the geomagnetic tail: The growth phase,
634 *Planet. Spa. Sci.*, 29, 1521–1539.
- 635 McPherron, R. L., T.-S. Hsu, J. Kissinger, X. Chu, and V. Angelopoulos (2011), Characteristics

- 636 of plasma flows at the inner edge of the plasma sheet, *J. Geophys. Res.*, 116, A00133,
637 doi:10.1029/2010JA015923.
- 638 McPherron, R. L., T. Terasawa, and A. Nishida (1986), Solar wind triggering of substorm
639 expansion onset, *J. Geomag. Geoelectr.*, 38, 1089–1108.
- 640 McPherron, R. (1991), Physical processes producing magnetospheric substorms and magnetic
641 storms, in *Geomagnetism*, edited by J. Jacobs, pp. 593–739, Academic Press, London, U.
642 K.
- 643 Meng, C.-I., and K. Liou (2004), Substorm timings and time scales: A new aspect, *Space Sci.*
644 *Rev.*, 113, 41 – 75, doi:10.1023/B:SPAC.0000042939.88548.68.
- 645 Nakamura, R., W. Baumjohann, M. Brittmacher, V. A. Sergeev, M. Kubyshkina, T. Mukai, and
646 K. Liou (2001), Flow bursts and auroral activations, *J. Geophys. Res.*, 106(A6), 10,777–
647 10,789.
- 648 Newell, P. T., A. R. Lee, K. Liou, S.-I. Ohtani, T. Sotirelis, and S. Wing (2010), Substorm cycle
649 dependence of various types of aurora, *J. Geophys. Res.*, 115, A09226,
650 doi:10.1029/2010JA015331.
- 651 Newell, P. T., T. Sotirelis, and S. Wing (2009), Diffuse, monoenergetic, and broadband aurora:
652 the global precipitation budget, *J. Geophys. Res.*, 114, A09207,
653 doi:10.1029/2009JA014326.
- 654 Ni, B., R. M. Thorne, N. P. Meredith, R. B. Horne, and Y. Y. Shprits (2011), Resonant scattering
655 of plasma sheet electrons leading to diffuse auroral precipitation: 2. Evaluation for whistler
656 mode chorus waves, *J. Geophys. Res.*, 116, A04219, doi:[10.1029/2010JA016233](https://doi.org/10.1029/2010JA016233).
- 657 Ni, B., R. M. Thorne, Y. Y. Shprits, and J. Bortnik (2008), Resonant scattering of plasma sheet
658 electrons by whistler-mode chorus: Contribution to diffuse auroral precipitation, *Geophys.*

- 659 Res. Lett., 35, L11106, doi:[10.1029/2008GL034032](https://doi.org/10.1029/2008GL034032).
- 660 Nishida, A. (1979), Possible origin of transient dusk-to-dawn electric field in the nightside
661 magnetosphere, *J. Geophys. Res.*, 84(A7), 3409–3412, doi:10.1029/JA084iA07p03409.
- 662 Pritchett, P. L., and F. V. Coroniti (2010), A kinetic ballooning/interchange instability in the
663 magnetotail, *J. Geophys. Res.*, 115, A06301, doi:10.1029/2009JA014752.
- 664 Pulkkinen, T. I., D. N. Baker, P. K. Toivanen, R. J. Pellinen, R. H. W. Friedel, and A. Korth
665 (1994), Magnetospheric field and current distributions during the substorm recovery phase,
666 *J. Geophys. Res.*, 99, 10,955-10,966.
- 667 Reeves, G. D., A. Chan, and C. Rodger (2009), New Directions for Radiation Belt Research,
668 *Space Weather*, 7, S07004, doi:10.1029/2008SW000436.
- 669 Rostoker, G. (2002), Identification of substorm expansive phase onsets, *J. Geophys. Res.*
670 Lett., 107(A7), doi:[10.1029/2001JA003504](https://doi.org/10.1029/2001JA003504).
- 671 Schekochihin, A. A., S. C. Cowley, W. Dorland, G. W. Hammett, G. G. Howes, E. Quataert, and
672 T. Tatsuno (2009), Astrophysical Gyrokinetics: Kinetic and Fluid Turbulent Cascades in
673 Magnetized Weakly Collisional Plasmas, *Astrophysical Journal Supplemental Series*, 182,
674 310-377, doi:10.1088/0067-0049/182/1/310.
- 675 Sergeev, V.A., K. Liou, P. T. Newell, S.-I. Ohtani, M. R. Hairston, and F. Rich (2004), Auroral
676 streamers: Characteristics of associated precipitation, convection and field-aligned
677 currents, *Ann. Geophys.*, 22, 537–548.
- 678 Sergeev, V. A., M. Malkov, and K. Mursula (1993), Testing the isotropic boundary algorithm
679 method to evaluate the magnetic field configuration in the tail, *J. Geophys. Res.*, 98, 7609-
680 7620.
- 681 Sergeev, V. A., E. M. Sazhina, N. A. Tsyganenko, J. A. Lundblad and F. Soraas (1983), Pitch-

- angle scattering of energetic protons in the magnetotail current sheet as the dominant source of their isotropic precipitation into the nightside ionosphere, *Planet. Space Sci.*, 31, 1147–1155.
- Sergeev, V. A., V. Angelopoulos, and R. Nakamura (2012), Recent advances in understanding substorm dynamics, *Geophys. Res. Lett.*, 39, L05101, doi:[10.1029/2012GL050859](https://doi.org/10.1029/2012GL050859)
- Sergeev, V. A., S. A. Chernyaeva, S. V. Apatenkov, N. Y. Ganushkina, and S. V. Dubyagin (2015), Energy–latitude dispersion patterns near the isotropy boundaries of energetic protons, *Ann. Geophys.*, 33, 1059–1070, doi:10.5194/angeo-33-1059-2015.
- Spence, H. E., and M. G. Kivelson (1993), Contributions of the low-latitude boundary layer to the finite width magnetotail convection model, *J. Geophys. Res.*, 98(A9), 15487–15496, doi:[10.1029/93JA01531](https://doi.org/10.1029/93JA01531).
- Speiser, T. W. (1965), Particle trajectories in model current sheets: 1. Analytical solutions, *J. Geophys. Res.*, 70(17), 4219–4226, doi:10.1029/JZ070i017p04219.
- Stasiewicz, K., P. Bellan, C. Chaston, C. Kletzing, R. Lysak, J. Maggs, O. Pokhotelov, C. Seyler, P. Shukla, L. Stenflo, A. Streltsov, and J. E. Wahlund (2000), Small scale Alfvénic structure in the aurora, *Space Science Rev.*, 92, 423.
- Summers, D., R. M. Thorne, and F. Xiao (1998), Relativistic theory of wave-particle resonant diffusion with application to electron acceleration in the magnetosphere, *J. Geophys. Res.*, 103(A9), 20,487–20,500, doi:10.1029/98JA01740.
- Tanskanen, E., T. I. Pulkkinen, and H. E. J. Koskinen (2002), Substorm energy budget during low and high solar activity: 1997 and 1999 compared, *J. Geophys. Res.*, 107(A6), 1086, doi:10.1029/2001JA900153.
- Thorne, R. M. (2010), Radiation belt dynamics: The importance of wave-particle interactions,

- 705 *Geophys. Res. Lett.*, 37, L22107, doi:10.1029/2010GL044990.
- 706 Tsyganenko, N. A., and D. G. Sibeck (1994), Concerning Flux Erosion from the Dayside
707 Magnetosphere, *J. Geophys. Res.*, 99(A7), 13,425–13,436, doi:10.1029/94JA00719.
- 708 Tsyganenko, N. A., D. P. Stern, and Z. Kaymaz (1993), Birkeland Currents in the Plasma Sheet,
709 *J. Geophys. Res.*, 98(A11), 19,455–19,464, doi:10.1029/93JA01922.
- 710 Vasconez, C. L., Valentini, F., Camporeale, E., & Veltri, P. (2014). Vlasov simulations of kinetic
711 Alfvén waves at proton kinetic scales. *Phys. Plasmas*, 21(11), 112107.
- 712 Wang, C.-P., L. R. Lyons, M. W. Chen, and R. A. Wolf (2001), Modeling the quiet time inner
713 plasma sheet protons, *J. Geophys. Res.*, 106(A4), 6161–6178, doi:[10.1029/2000JA000377](https://doi.org/10.1029/2000JA000377).
- 714 Wang, C-P, L. R. Lyons, M. W. Chen, R. A. Wolf, and F. R. Toffoletto (2003), Modeling the
715 inner plasma sheet protons and magnetic field under enhanced convection, *J. Geophys.
716 Res.*, 108 (A2), 1074, doi:10.1029/2002JA0009620.
- 717 Wang, C.-P., L. R. Lyons, M. W. Chen, and F. R. Toffoletto (2004), Modeling the transition of the
718 inner plasma sheet from weak to enhanced convection, *J. Geophys. Res.*, 109, A12202,
719 doi:[10.1029/2004JA010591](https://doi.org/10.1029/2004JA010591).
- 720 Wang, C.-P., L. R. Lyons, J. M. Weygand, T. Nagai, and R. W. McEntire (2006), Equatorial
721 distributions of the plasma sheet ions, their electric and magnetic drifts, and magnetic
722 fields under different interplanetary magnetic field B_z conditions, *J. Geophys. Res.*, 111,
723 A04215, doi:[10.1029/2005JA011545](https://doi.org/10.1029/2005JA011545).
- 724 Wang, C.-P., M. Gkioulidou, L. R. Lyons, R. A. Wolf, V. Angelopoulos, T. Nagai, J. M. Weygand,
725 and A. T. Y. Lui (2011), Spatial distributions of ions and electrons from the plasma sheet to
726 the inner magnetosphere: Comparisons between THEMIS-Geotail statistical results and
727 the Rice convection model, *J. Geophys. Res.*, 116, A11216, doi:[10.1029/2011JA016809](https://doi.org/10.1029/2011JA016809).

- 728 Watanabe, M., and T. Iijima (1993), Substorm Growth Phase on the Magnetotail, *J. Geophys.*
729 *Res.*, 98(A10), 17,299–17,316, doi:10.1029/93JA01474.
- 730 Watt, C. E. J., and R. Rankin (2009), Electron trapping in shear Alfvén waves that power the
731 aurora, *Phys. Rev. Lett.*, 102, 045002.
- 732 Watt, C. E. J., R. Rankin, I. J. Rae, and D. M. Wright (2005), Self-consistent electron
733 acceleration due to inertial Alfvén wave pulses, *J. Geophys. Res.*, 110, A10S07,
734 doi:10.1029/2004JA010877.
- 735 Wing, S., M. Gkioulidou, J. R. Johnson, P. T. Newell, and C.-P. Wang (2013), Auroral particle
736 precipitation characterized by the substorm cycle, *J. Geophys. Res. Space*
737 *Physics*, 118,1022–1039, doi:[10.1002/jgra.50160](https://doi.org/10.1002/jgra.50160).
- 738 Wing, S., and J. R. Johnson (2010), Introduction to special section on Entropy Properties and
739 Constraints Related to Space Plasma Transport, *J. Geophys. Res.*, 115, A00D00,
740 doi:10.1029/2009JA014911.
- 741 Wing, S., and J. R. Johnson (2009), Substorm entropies, *J. Geophys. Res.*, 114, A00D07,
742 doi:10.1029/2008JA013989.
- 743 Wing, S., and D. G. Sibeck (1997), The effects of interplanetary magnetic field Z-component and
744 the solar wind dynamic pressure on the magnetospheric magnetic field line, *J. Geophys.*
745 *Res.*, 102, 7207–7216.
- 746 Wing, S., J. W. Gjerloev, J. R. Johnson, and R. A. Hoffman (2007), Substorm plasma sheet ion
747 pressure profiles, *Geophys. Res. Lett.*, 34, L16110, doi:10.1029/2007GL030453.
- 748 Wing, S., J. R. Johnson, P. T. Newell, and C.-I. Meng (2005), Dawn-dusk asymmetry in the
749 northward IMF plasma sheet, *J. Geophys. Res.*, 110, A08205, doi:10.1029/2005JA011086.
- 750 Wing, S., and P. T. Newell (2002), 2D plasma sheet density profile for northward and southward

- 751 IMF, *Geophys. Res. Lett.*, 29(9), 10.1029/2001GL013950.
- 752 Wing, S., and P. T. Newell (1998), Central plasma sheet ion properties as inferred from
753 ionospheric observations, *J. Geophys. Res.*, 103, 6785–6800.
- 754 Wolf, R. A., Wan, X. Xing, J. Zhang, and S. Sazykin (2009), Entropy and plasma sheet transport,
755 *J. Geophys. Res.*, 114, A00D05, doi:10.1029/2009JA014044.
- 756 Wygant, J. R., et al. (2002), Evidence for kinetic Alfvén waves and parallel electron energization
757 at 4–6 RE altitudes in the plasma sheet boundary layer, *J. Geophys. Res.*, 107, 1201,
758 doi:10.1029/2001JA900113.
- 759 Zanetti, L. J., and T. Potemra (1986), The relationship of Birkeland and ionospheric current
760 systems to the interplanetary magnetic field, in *Solar Wind-Magnetosphere Coupling*,
761 edited by Y. Kamide and J. A. Slavin, pp. 547–562, Terra Sci., Tokyo.
- 762 Zhang, J.-C., R. A. Wolf, S. Sazykin, and F. R. Toffoletto (2008), Injection of a bubble into the
763 inner magnetosphere, *Geophys. Res. Lett.*, 35, L02110, doi:10.1029/2007GL032048.
- 764

765
766
767
768
Table

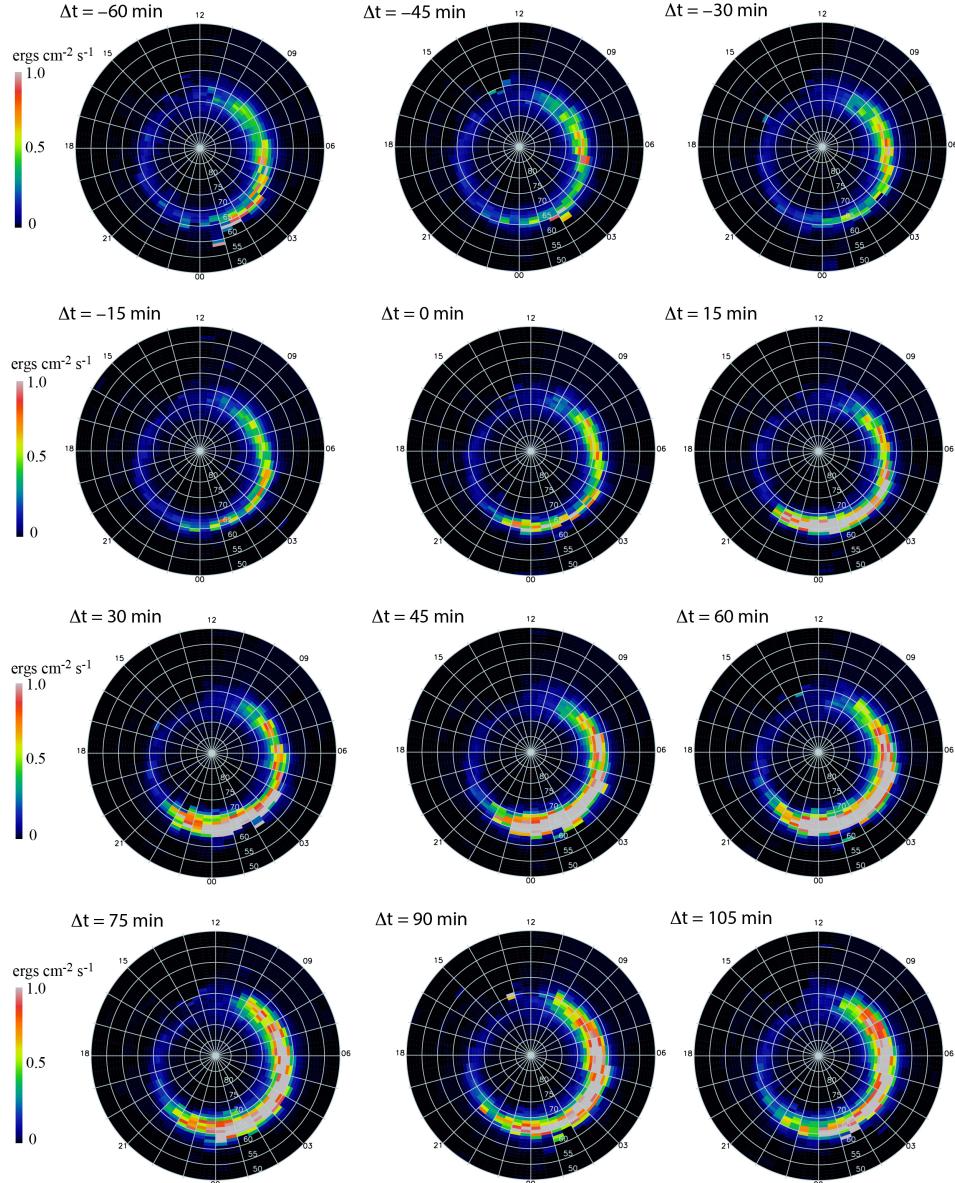
	total substorm cycle	growth phase	expansion phase	recovery phase	Δt at half max
diffuse electrons	> 5hr	?	1 hr	>4 hr?	~1.2 hr
broadband electrons	~ 5 hr	1.25 hr	0.25 hr	~3.50 hr	~0.42 hr
monoenergetic electrons	~ 5 hr	1.25 hr	0.25 hr	~3.50 hr	~1.6 hr

769 Table 1. Summary of the time scales of the substorm cycles and phases from electron
770 precipitation perspective. The growth and recovery phases for the diffuse electrons are hard to
771 determine (see text). Δt at half max denotes the time it takes to reach half maximum (green
772 asterisks in Figure 7) from the maximum auroral power at end of the expansion phase, which
773 gives a measure of how quickly the auroral power decays after reaching the maximum value.
774
775

776
777
778

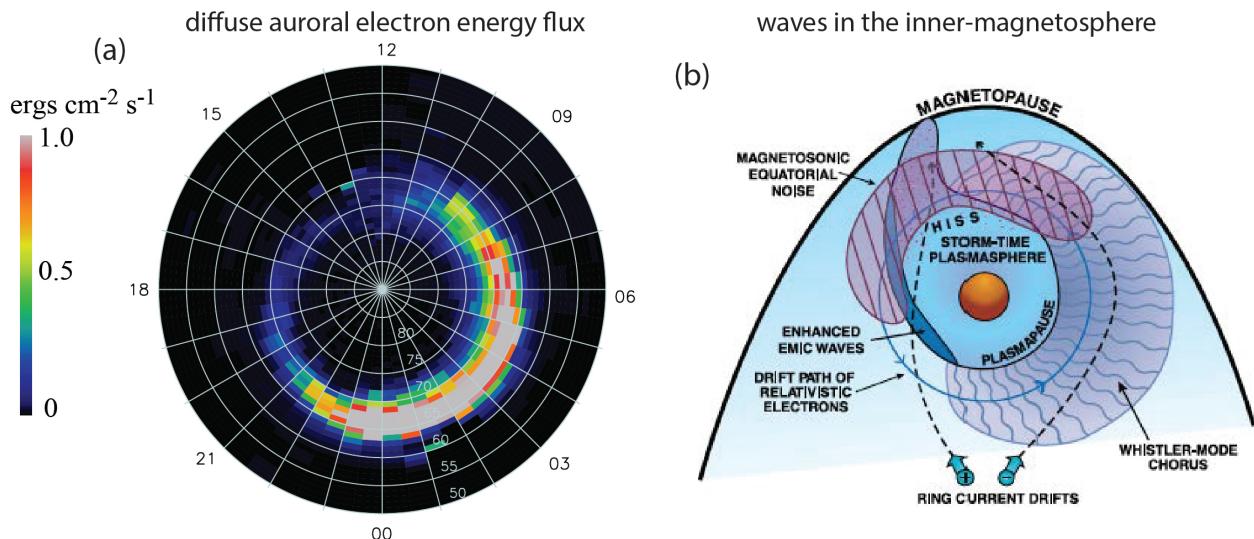
Figures

Diffuse auroral electron energy flux



779
780
781
782
783
784
785
786
787

Figure 1. Diffuse auroral electron energy flux from 1 hr before to 1 hr and 45 min after the substorm onset. Each map shows the median energy flux of 48 MLT by 40 MLat bins at 15 min time steps centered at the time labeled. The substorm onset occurs at $\Delta t = 0$ min. The MLat range is 50° - 90° , with data from the two hemispheres combined. (from Wing et al. [2013]). 1677 substorm events were considered for the construction of this and other figures in this paper.



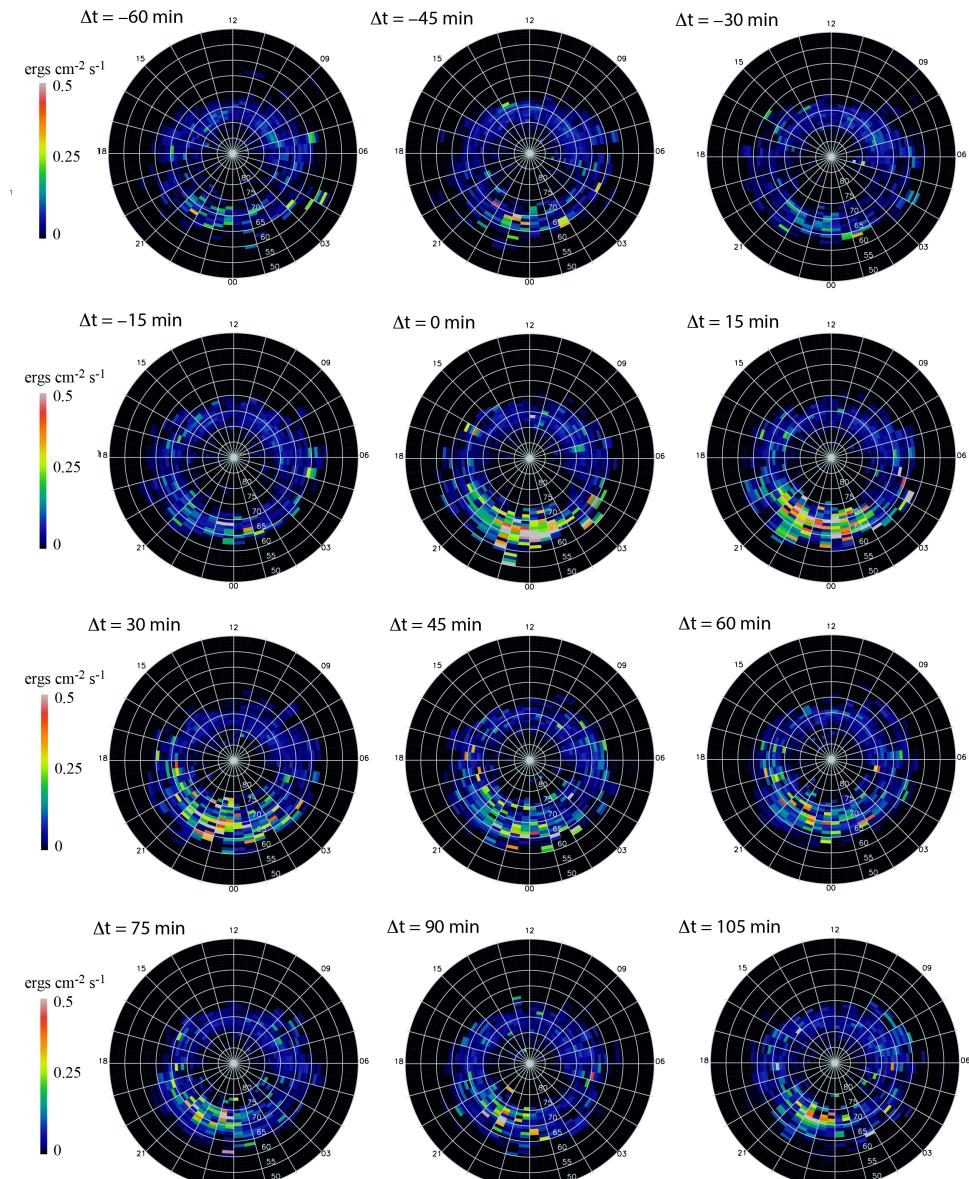
788

789

790 Figure 2. Diffuse auroral electron spatial distribution is similar to that of whistler-mode chorus
 791 waves. (a) diffuse auroral electron energy flux 60 min after substorm onset (from Figure 1). (b)
 792 Schematic showing spatial distribution of important waves in the inner magnetosphere (from
 793 Thorne [2010]).

794

Broadband auroral electron energy flux

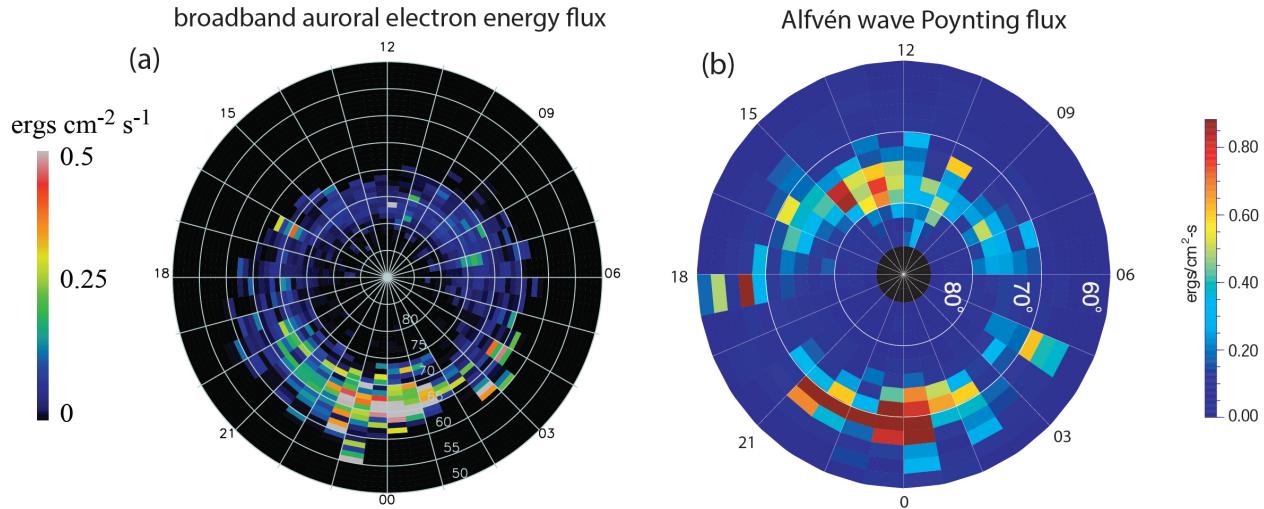


795

796

797 Figure 3. Broadband auroral electron energy flux from 1 hr before to 1 hr and 45 min after the
798 substorm onset in the same format as in Figure 1. The substorm onset occurs at $\Delta t = 0 \text{ min}$.
799 (from Wing et al. [2013])

800



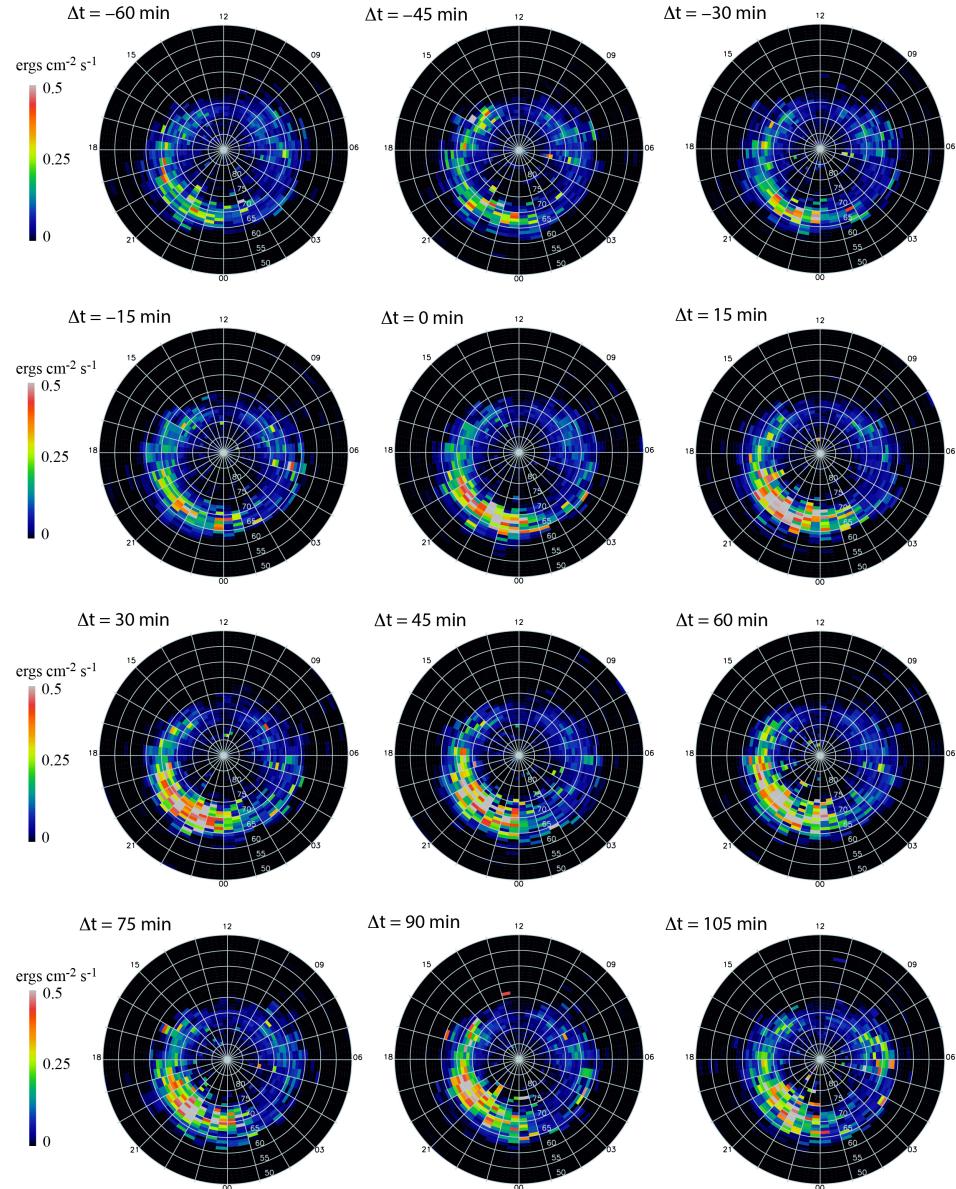
801

802

803 Figure 4. Broadband auroral electrons have similar spatial distribution with that of Alfvén waves
 804 in the magnetosphere. (a) Broadband auroral electron energy flux near substorm onset (from
 805 Figure 3). (b) Average Alfvén wave Poynting flux flowing toward the Earth at high-altitude as
 806 observed by Polar satellite at 25,000 to 38,000 km obtained from an interval when large Alfvén
 807 waves were present during auroral and substorm activity (from Keiling et al. [2003]).

808

Monoenergetic auroral electron energy flux



809

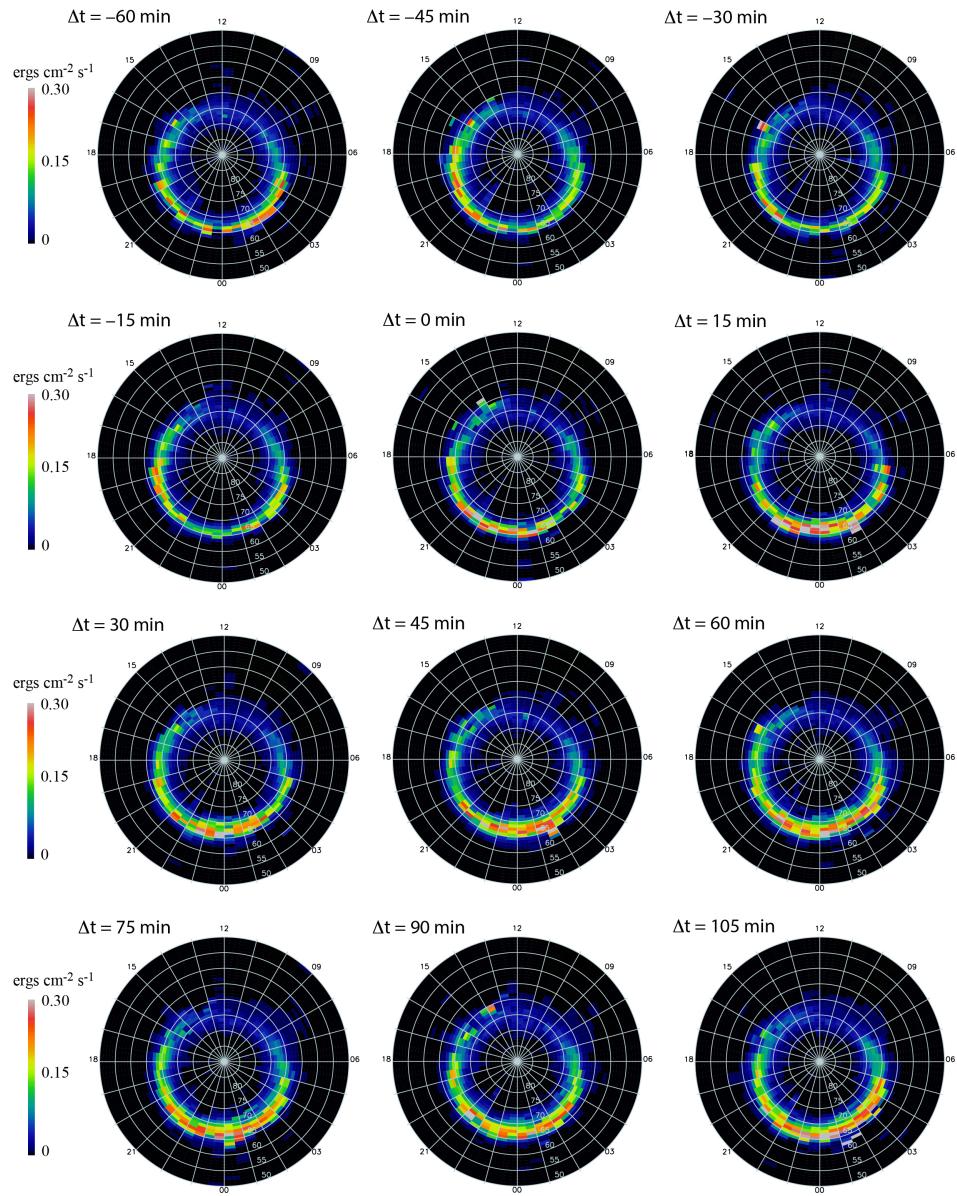
810 Figure 5. Monoenergetic auroral electron energy flux from 1 hr before to 1 hr and 45 min after
 811 the substorm onset in the same format as in Figure 1. The substorm onset occurs at $\Delta t = 0$ min.
 812 (from Wing et al. [2013])

813

814

815

Auroral ion energy flux



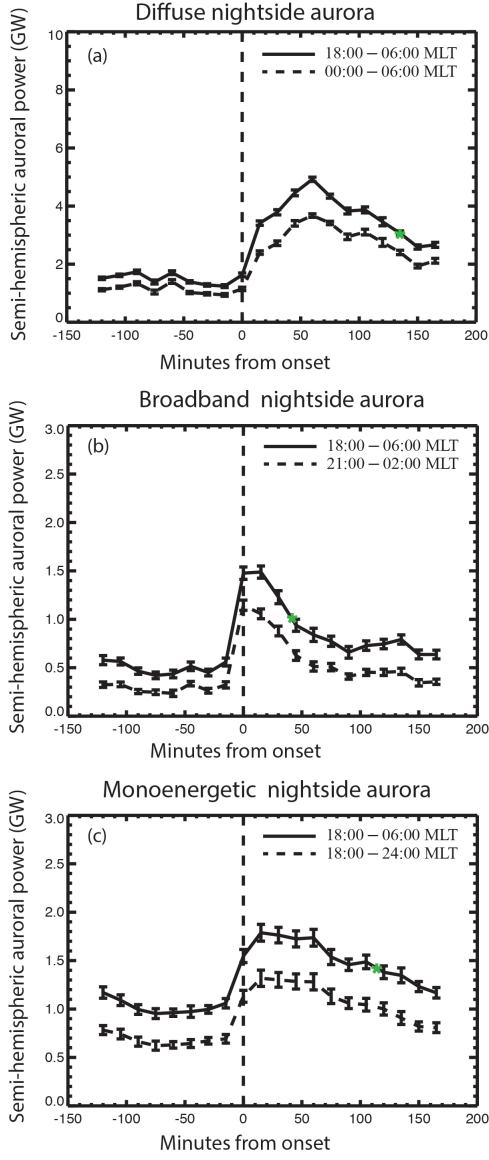
816

817 Figure 6. Auroral ion energy flux from 1 hr before to 1 hr and 45 min after the substorm onset in
 818 the same format as in Figure 1. The substorm onset occurs at $\Delta t = 0$ min. (from Wing et al.
 819 [2013])

820

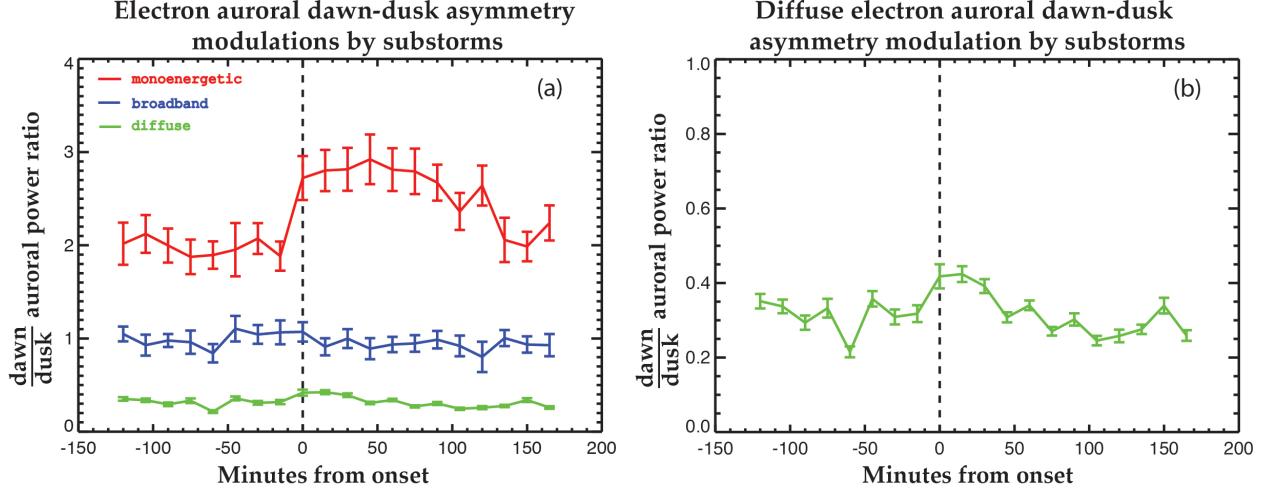
821

Electron auroral power increases by substorms



822

823 Figure 7. The nightside semihemispheric auroral electron power spanning the interval 2 hr to 3 hr
 824 before to 3 hr after substorm onset for (a) diffuse, (b) broadband, and (c) monoenergetic electrons. The solid line shows
 825 the entire nightside power obtained from integrating the electron powers in all bins in 18:00–06:00 MLT
 826 and 50°–90° MLat. The electron power in each bin is computed by multiplying the physical surface area
 827 of each grid by the median energy flux in each bin in Figures 1, 3, and 5. For example, each image in
 828 Figure 1 contributes one point on the solid line in (a). The error bars are the standard deviations of the
 829 median [e.g., Kenney and Keeping, 1951; Hodges and Lehmann, 1967]. (a) The dashed line shows the
 830 midnight–dawn powers obtained from integrating the electron power in all bins in 00:00–06:00 MLT
 831 and 50°–90° MLat. (b) The dashed line shows the power obtained by integrating all powers in all bins in
 832 21:00–02:00 MLT and 50°–90° MLat. (c) the dashed line shows the dusk–midnight sector power obtained
 833 from integrating the powers in all bins 18:00–24:00 MLT and 50°–90° MLat. The time it takes to reach
 834 half maximum power (green asterisks) from the end of the expansion phase (maximum power) for the
 835 diffuse (a), broadband (b) and monoenergetic (c) electrons are ~1.2, ~0.42, and 1.6 hr, respectively.
 836 (adapted from Wing et al. [2013]).



837
838 Figure 8. Electron auroral power dawn-dusk asymmetries modulations by substorms. (a) the ratio of
839 dawn (18:00 – 2400 MLT) to dusk (24:00 – 06:00 MLT) auroral power for monoenergetic (red),
840 broadband (blue), and diffuse (green) electrons. The monoenergetic electron auroral power is larger at
841 dawn than at dusk and this asymmetry increases after substorm onset. The asymmetry does not get back
842 to its growth phase value until \sim 135 min after onset. In contrast, broadband electron auroral power does
843 not show much dawn-dusk asymmetry. The diffuse electron auroral power has the opposite asymmetry
844 from that of monoenergetic electron auroral power (the ratio < 1). At 0 – 30 min after substorm onset, the
845 diffuse electron auroral power dawn-dusk asymmetry decreases a bit. This can be seen more clearly in
846 (b), which plots the same dawn/dusk auroral power ratio as in (a), except that the scale of the Y-axis is
847 smaller.