Are quasi-stationary waves and planetary waves the same?

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6 Key Points:

- List up to three key points (at least one is required)
 - Key Points summarize the main points and conclusions of the article
- Each must be 100 characters or less with no special characters or punctuation

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Abstract

A good abstract will begin with a short description of the problem being addressed, briefly describe the new data or analyses, then briefly states the main conclusion(s) and how they are supported and uncertainties.

1 Introduction

Atmospheric planetary waves are blabalbaal... In some latitudes and tiems of the year they exhibit some degree of stationatiry

At present, there exists a lack of consistency in the literature when it comes to a

2 Story

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Given a set of atmospheric fields, we define "zonal waves" (ZW; sometimes also called planetary waves) as waves observed in each individual "instantaneous" field and quasistationary waves (QS) as the resulting waves in the mean field. While these definitions depend on which are the "instantaneous field" in question and the averaging timescales used, they illustrate that, ZWs are properties of the *elements* of the set, while the QS is a property of the set as a whole. This is an important distinction with theorical and methodological implications that is not always appreciated in the literature.

As an example, Figure 1 shows the monthly seasonal cycle of amplitude of planetary waves at 60° S using monthly fields from the NCEP/NCAR reanalysis (Kalnay et al., 1996) between 1950 and 1998. The left column (\overline{ZW}) reproduces Figure 3 from Rao et al. (2004) and is computed by taking –for each month and level– the average amplitude of the 49 individual amplitudes. The right column (QS), on the other hand, is computed by taking the amplitude of the average geopotential field for each month and level.

The resulting fields convey different information. First, the amplitude of \overline{ZW} fields is always greater than the one for QS fields. This is a ^{c1}mathematical necessity that explains Rao et al. (2004)'s observation that their Wave 1 amplitude was greater than that reported by Hurrell, van Loon, and Shea (1998). Secondly, they have different annual

^{c1} elio: Deberia demostrar eso? Vale la pena una demostracion en un material suplementario?

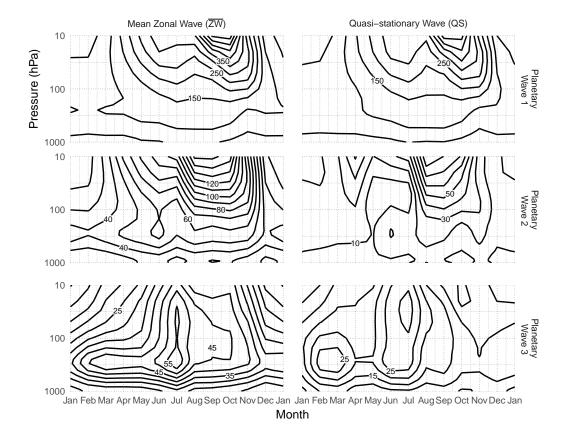


Figure 1. Seasonal cycle of amplitude of the geopotential planetary waves 1 to 3 at 60° S computed as the mean amplitude of the monthly waves (\overline{Z}) and as the amplitude of the mean wave (QS). The period of analysis is 1950 to 1998. The left column reproduces Figure 3 from Rao et al. (2004).

cycles and vertical structures. QS2, for example, has a strong minimum in the low stratosphere during the austral autumn that is not aparent in $\overline{ZW2}$. Similarly, the austral winter mid-troposferic maxmium is very well defined in $\overline{ZW3}$ but not so in QS3. Thirdly, the relative importance between each wavenumber vary. \overline{ZW} fields show, for example, a preponderance of wave 2 over 3 in almost every level and month. However, the QS3 has greater amplitude than QS2 in the first half of the year. In contrast with wave-numbers 2 and 3, $\overline{ZW1}$ and QS1 fields are very similar.

These differences are related to the degree of stationative of zonal waves and are location-dependent. Figure 2 show the same variable that Figure 1 but for 300hPa. The contrast between the northern and shouthern hemisphere is not only evident in the amplitude of the planetary waves, but also in the comparison between \overline{ZW} and QS. Spe-

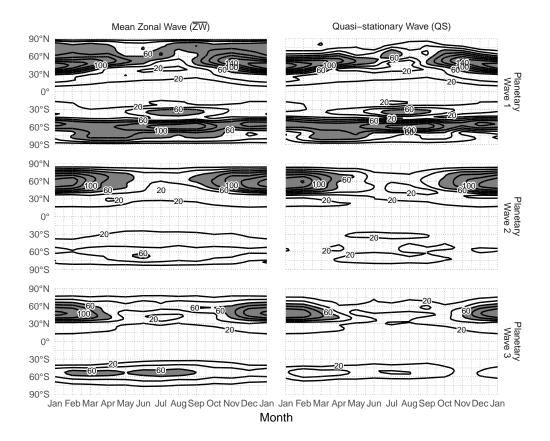


Figure 2. Seasonal cycle of amplitude of the geopotential planetary waves 2 at 300hPa computed as the mean amplitude of the monthly waves (\overline{ZW}) and as the amplitude of the mean wave (QS). From monthly NCEP/NCAR Reanalisys, 1979 to 2017.

cially for wave-numbers 2 and 3, \overline{ZW} and QS fields are very similar in the north but they have significant differences in the south.

2.1 Stationarity

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Another important consequence of the distinction between \overline{ZW} and QS is that the quotient between the two can be used as a measure of stationarity. As an analogy with the steadiness of the wind (Singer, 1967), planetary wave stationarity can be calculated as

$$S = \frac{2}{\pi} \arcsin\left(\frac{QS}{\overline{ZW}}\right) \tag{1}$$

It can be shown that S = 1 for completely stationary waves and that $E(S) = n^{-1/2 c 1}$ for completely non-stationary waves (where n is the sample size).

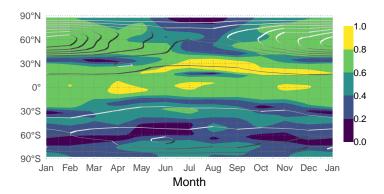


Figure 3. Seasonal cycle of stationarity of the 300hPa geopotential QS2 computed using Equation 1. From monthly NCEP/NCAR Reanalisys, 1979 to 2017.

As an example, Figure 3 shows S for QS2 computed using Equation 1. The southern hemisphere clearly shows a lower degere of QS2 stationarity than the northern hemisphere or the tropics. In the northern midlatitudes there is a seasonal cycle of stationarity that follows the seasonal cycle of \overline{ZW} (Figure 2). In the southern hemisphere, instead, the June maximum of \overline{ZW} coincides with a minimum of stationarity.

While S is used –sometimes without the arcsin transformation– in the meteorological literature in the context of wind steadiness (c1 Agregar citas?), to our knowledge this is the first time it has been applied to the study of atmospheric waves. Furthemore, its statistical properties are not well studied. For example, it can be seen that the estimation of S from a finite sample has a positive bias that is inversely proportional to the population stationarity.

2.2 QS activity

Defining quasi-stationary waves as a climatological property of a set of atmospheric fields, precludes, in principle, the possibility of quantifying a QS metric that applies to instantaneous fields. It would seem impossible to, for example, construct an timeseires

 $^{^{\}rm c1}$ elio: El factor 1/2 creo que es para el S sin el asin. El S de la ecuacion anterior tiene un factor raro de 0.688 o algo asi.

c1 elio: Text added.

of QS activity that could be use as a basis for correlations with other variables, compositions or for use in other methodologies. But there are ways of solving this issue.

One posibility is recognising that individual fields can be characterised by their degree of similarity with the climatological QS. The index produced by Raphael (2004) for the QS3 is an example. While not expressly a measure of similarity, it is sensitive to wave 3 patterns with phase close to the stationary phase. Yuan and Li (2008) use Principal Component Analisys on the meridional wind field; the spatial pattern of the leading mode is very similar to the QS3 and so a timeseries can be obtained by projecting each instantaneous field to it.

Another way of constructing a timeseires is to exploit the fact the timescale dependence of QS. By applying a running mean with a suitable windows before computing wave amplitudes, one obtains the QS wave amplitude of that window.

3 Conclusions

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The fact that zonal waves (ZW) and quasi-stationarity waves (QS) are two distinct but related phenomena has both practical and theorical implications.

Researchers should be aware of which phenomena they want to study and use the appropriate methods.

Comparison between results should also be made having this issues in mind. For example, Irving and Simmonds (2015) compare their planetary wave activity index with Raphael (2004)'s wave 3 index and conclude that the later cannot account for events with waves far removed from their climatological position. However, if we understand it as an index of QS3 similitude, then that is a feature, not a bug.

Since planetary waves are generally more stationary in the northern hemisphere, these issues are more critical in studies of the southern hemisphere.

For clarity, when reporting results that use mean amplitude of waves, researchers should always describe if they are using the mean amplitude of individual waves or the amplitude of the mean wave.

Besides those direct implications, separating ZW and QS can lead to novel levels 109 of analysis. Here, we showed a simple metric of the stationarity, but others are possi-110 ble. 111

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