

Climate change and its implications

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Lecture-1



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Course outline: 2 credit

- Introduction to climate & climate change
- Evidence of climate change
- Global climate models & future climate projections
- Implications to
 - : Biodiversity, mitigation practices
 - : Agriculture, mitigation practices
 - : Water resources, mitigation practices



QGIS-Quantum GIS

<https://qgis.org/en/site/>

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Class outline:

Earth system Components

- Atmosphere
- Oceans
- Cryosphere
- Biosphere
- Earth's crust and mantle



Class outline:

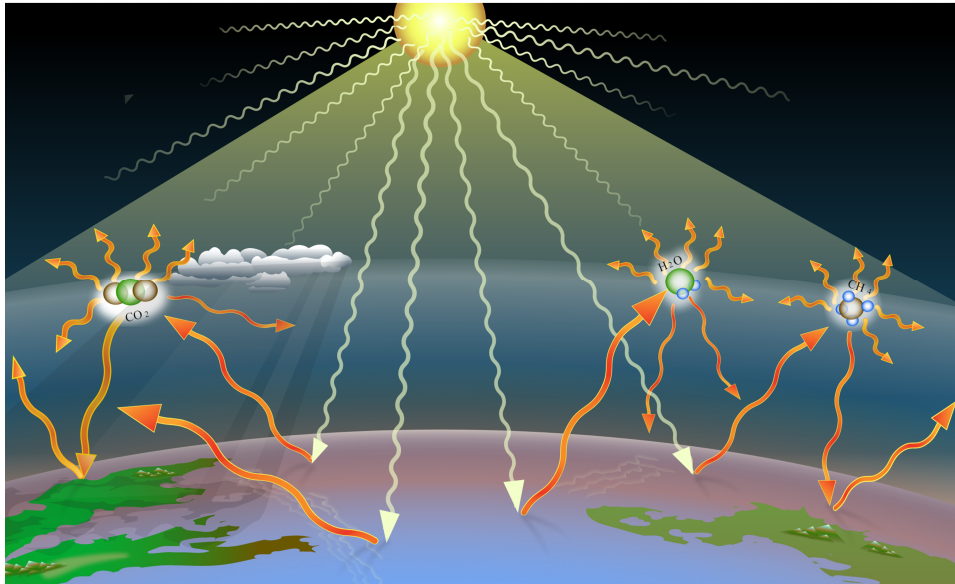
Earth system Components: 1) Atmosphere

Characteristics of Earth's atmosphere

- Optical properties
- Mass
- Composition
- Vertical structure: temperature, general circulations



Optical properties



- ▶ Earth's atmosphere is transparent to incoming solar radiation
- ▶ The outgoing radiation emitted by Earth is absorbed by the atmosphere (green house effect), and this makes the Earth's atmosphere warm
- ▶ About 22% of incoming solar rad is backscattered to space without absorption

Mass of the atmosphere



Mass of the atmosphere

- ▶ At any point on the earth's surface, the atmosphere exerts a downward force on the underlying surface due to earth's gravitational attraction.
- ▶ The downward force (i.e., the weight) of a unit volume of air with density ρ is given by,

$$F = \rho g$$

Where g is the acceleration due to gravity.

- ▶ Integrating the equ from earth's surface to the top of the atmosphere, we obtain the atmospheric pressure on the earth's surface (P_s) due to the weight (per unit area) of the air in the overlying column.

$$\text{i.e., } P_s = \int_0^{\infty} \rho g \, dz$$

- ▶ Neglect the small variation of g with lat, long, and height, we can take the mean value g , which is equal to $9.807 \, \text{m}^2/\text{s}$, we can take it outside the integral.

Then $P_s = g \int_0^\infty \rho \, dz$,

which is again equal to $= g m$

where m is the vertically integrated mass per unit area of the overlying air

Exercise 1

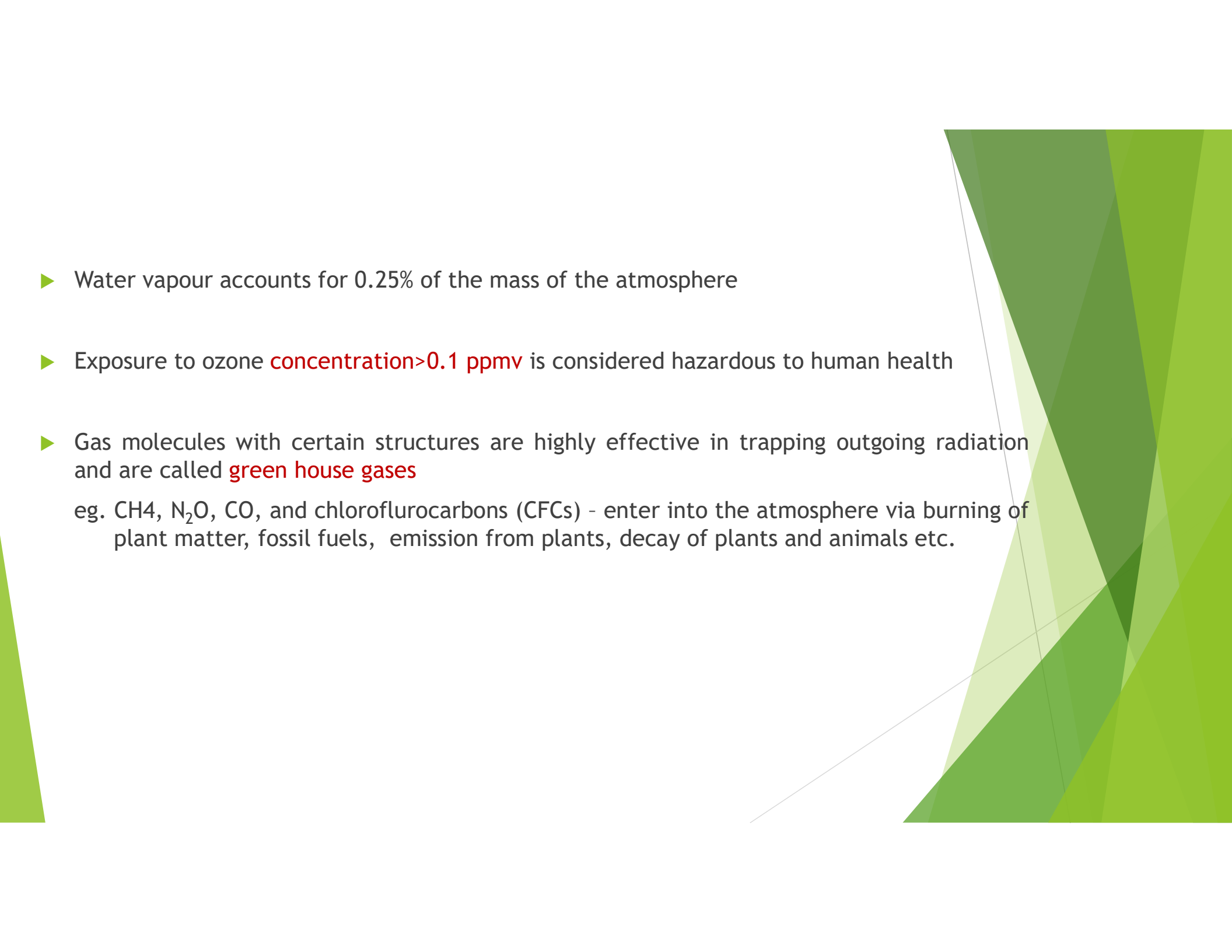
- The globally averaged surface pressure is 985 hPa. Estimate the mass of the atmosphere.

Earth's radius = 6.37×10^6 m



Chemical composition

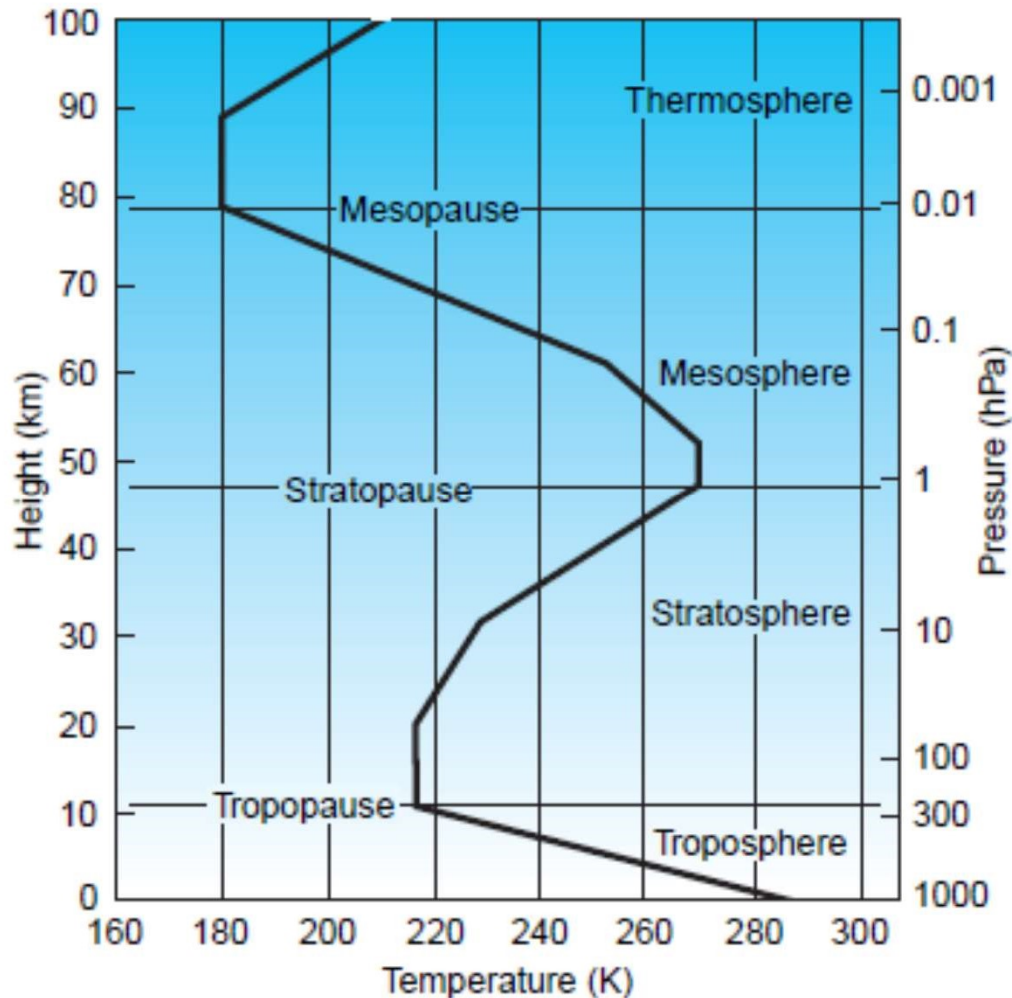
Constituent ^a	Molecular weight	Fractional concentration by volume
Nitrogen (N ₂)	28.013	78.08%
Oxygen (O ₂)	32.000	20.95%
Argon (Ar)	39.95	0.93%
Water vapor (H₂O)	18.02	0–5%
Carbon dioxide (CO₂)	44.01	380 ppm
Neon (Ne)	20.18	18 ppm
Helium (He)	4.00	5 ppm
Methane (CH₄)	16.04	1.75 ppm
Krypton (Kr)	83.80	1 ppm
Hydrogen (H ₂)	2.02	0.5 ppm
Nitrous oxide (N₂O)	56.03	0.3 ppm
Ozone (O₃)	48.00	0–0.1 ppm

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- The background of the slide features abstract, overlapping green geometric shapes, primarily triangles and polygons, in various shades of green, creating a modern, layered effect on the right side.
- ▶ Water vapour accounts for 0.25% of the mass of the atmosphere
 - ▶ Exposure to ozone **concentration > 0.1 ppmv** is considered hazardous to human health
 - ▶ Gas molecules with certain structures are highly effective in trapping outgoing radiation and are called **green house gases**
eg. CH₄, N₂O, CO, and chlorofluorocarbons (CFCs) - enter into the atmosphere via burning of plant matter, fossil fuels, emission from plants, decay of plants and animals etc.

Vertical structure of atmosphere



Vertical structure

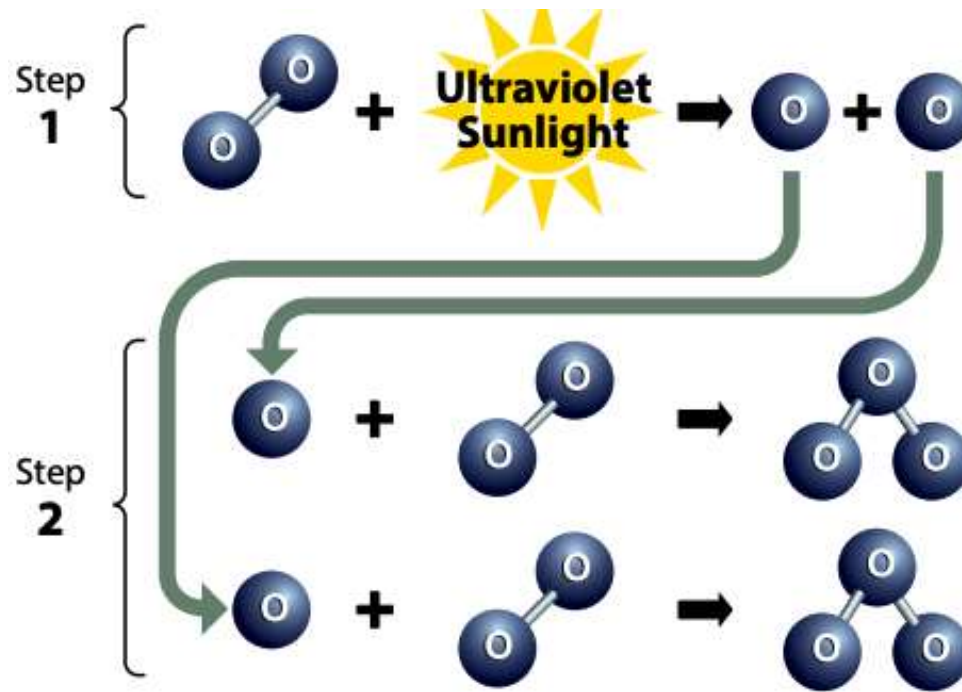


- Tropo-(turning or changing/vertical mixing) - sphere
- Temperature decreases with height ($\sim 6.5^{\circ}\text{C}$)
- Tropospheric air accounts for the 80% of the mass of the atmosphere

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- Strato-(layered)-sphere, vertical mixing is prohibited due to the increase of temp with height
 - Residence time of particles are longer
 - Air is extremely dry and ozone rich
 - They absorb the UV from the spectrum
 - This increases the temperature

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- Meso-(inbetween)-sphere: temperature decreases with height

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- Thermosphere-temp increases with height due to the absorption of solar rad, and lots of ionization processes occurs



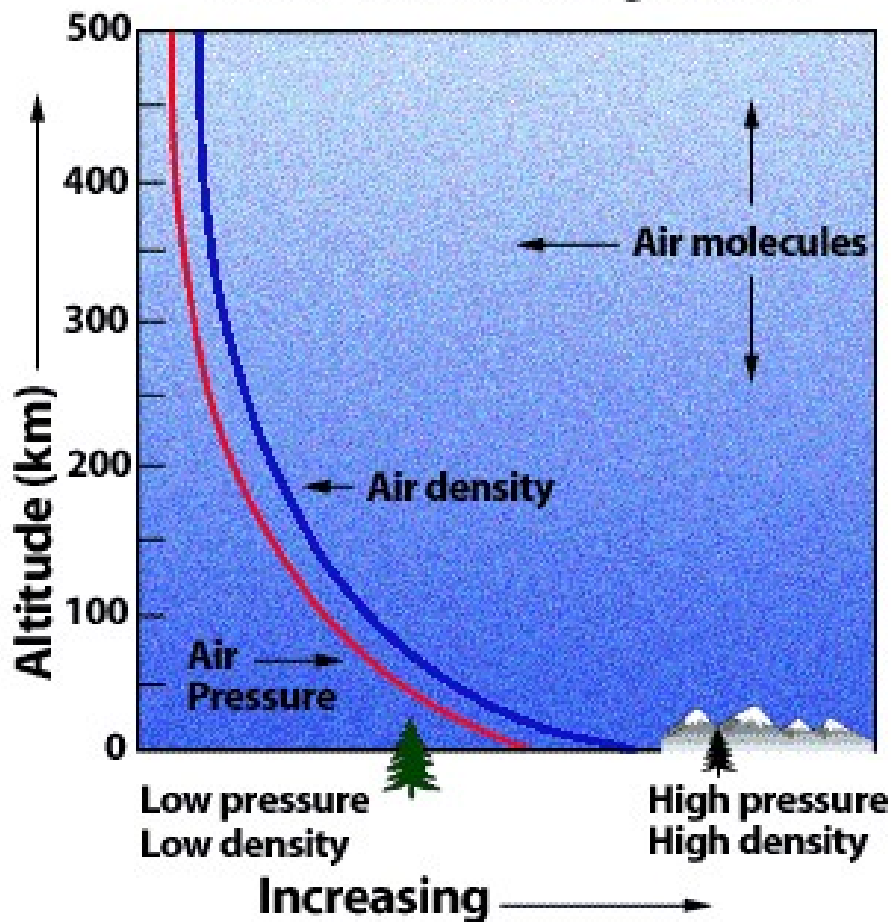
Ozone formation in stratosphere



Anvil cloud (Cloud flattening) at the tropopause

- All the weather and climate activities are under the tropopause

Both air pressure and air density decrease with increasing altitude.



Density of air @ sea level is 1.25 kg/m^3

Pressure at any height:

$$p \approx p_0 e^{-z/H}$$

P_0 is the pressure @ sea level (reference level)

H - scale height; e-folding depth (height at which pressure becomes $1/e$ times p_0), 7 to 8 Km

$$\ln \frac{p}{p_0} \approx -\frac{z}{H}$$

$$z = H \ln(p_0/p)$$